Response to Referee Comment 1 (RC1) on “Rapid transition in winter aerosol composition in Beijing from 2014 to 2017: response to clean air actions” by H. Li et al.

In past publications, the authors have reported on the effects of emission controls on China’s air quality. This paper focuses on how those emission changes affect aerosol chemical composition. The authors present a comprehensive analysis involving measurements and models to assess the role of just emissions, apart from transport and meteorological differences. The results are interesting and the topic appropriate for this journal, however some details were missing and in many cases the analysis or description of processes was not of sufficient technical detail, such as the thermodynamic analysis. Also, strong conclusions are made with limited analysis to support them. For example, a major finding seems to be that these changes in emissions altered the aerosol formation mechanisms. It is stated in the Abstract and Conclusions that lower SO2 suppressed rapid sulfate formation through heterogeneous reactions. But it is not clear that this is really shown or tested in a quantitative way, instead it seems to be mostly speculation, based mainly on this hypotheses being consistent with the observations. Line 324 simply states: the results here imply that: : : Could not a chemical transport model, or maybe a simpler 0-D or 1D box model be run to further test this hypothesis? Overall, my recommendation is that the authors check closely if their reported findings are truly supported in the manuscript by quantitative analyses, and if only based on a consistency with expectations, that this clearly be stated. More detailed comments are provided below.

We would like to thank Rodney Weber for giving the constructive and helpful comments and suggestions, especially for the discussions on thermodynamic analysis. In the revised manuscript, more technical details about the thermodynamic analysis have been added according to the comments. One finding that the decrease in SO2 emissions suppressed the rapid sulfate formation through heterogeneous reactions was speculated based on the ambient observations. We found that compared to the fast SO2-to-sulfate formation starting from a RH threshold of ~50% in 2014, the promptly increased sulfate formation through heterogeneous reactions was observed to delay to a higher RH of 70% in 2017. Therefore, this is one hypothesis based on the consistency with ambient observations. We have clearly stated it in the Abstract, Sect. 3.2.4, and Conclusions in the manuscript.

In the following, we will answer the comments point by point.

Specific Comments:

Line 59-60: I would think atmospheric chemical reactions (secondary aerosol) and deposition would also be a major contributor to PM2.5 composition.

Agreed. We slightly modified the sentence to “The chemical composition of PM2.5 is mainly affected by the following factors: precursor emissions, meteorological conditions, atmospheric chemical reactions, and regional transport and deposition”.

Line 136 to 137, in giving the ambient data vs model comparison, state the integration time, ie 24 hr average data?

It has been clearly stated that the data are 24-hour averages.

ISORROPIA calculations: In this paper the model is run without considering nonvolatile cations. Maybe a few words should be added why this is ok, ie, it may be argued that for PM1 this is reasonable. As another example, the nitrate considered in the paper is all semivolatile nitrate (ie, NH4NO3), but it is possible that nonvolatile nitrate also exists in the ambient PM2.5 (eg, Ca(NO3)2). Thus if Ca2+ was considered in the thermodynamic calculations, it could affect predicted pH and NO3- concentrations. Most of this Ca2+ would likely be in the 1 to 2.5 um range, and since the comparison between PM1 and PM2.5 mass is reasonable, the authors could argue that it is not a large contribution. Also, I suggest the authors specifically note what RH range was used in the thermodynamic calculations, many of the assumptions, such as no separate organic/inorganic phases,
etc., may be less likely at lower RH (say <40 to 50%). Line 231-232 notes that the observed RH was about 33 to 34% in the winter of 2017. This a very low RH to comfortably run ISORROPIA under the metastable assumption without some test on the reliability of the results.

Indeed, including nonvolatile cations in ISORROPIA calculations would influence the model results. But as the reviewer said, nonvolatile cations, i.e., Na⁺, K⁺, Ca²⁺, Mg²⁺, mainly exists in the size range of 1.0 to 2.5 µm in particles and has a minor contribution to PM₁₀. Therefore, nonvolatile cations are not considered in pH calculation in this study. A previous study by Song et al. (2018) showed that including nonvolatile cations in ISORROPIA calculations did not significantly change the particle pH. Discussions about the effects of nonvolatile cations on pH calculations have been added in the manuscript “The effects of nonvolatile cations (i.e., Na⁺, K⁺, Ca²⁺, Mg²⁺) are not considered in this study because the fraction of nonvolatile cations in PM₁₀ in Beijing is generally negligible compared to SO₄²⁻, NO₃⁻, and NH₄⁺ (Sun et al., 2014). Although nonvolatile nitrate may exist in ambient particles as Ca(NO₃)₂; and Mg(NO₃)₂, Ca²⁺ and Mg²⁺ are mainly abundant at sizes above 1 µm (Zhao et al., 2017). In addition, the mixing state of PM₁₀ nonvolatile cations with SO₄²⁻, NO₃⁻, and NH₄⁺ remains to be investigated (Guo et al., 2016, 2017). Previous studies showed that including the nonvolatile cations in ISORROPIA-II does not significantly affect the pH calculations unless the cations become important relative to anions (Guo et al., 2016; Song et al., 2018). The sensitivity test for Beijing winter conditions suggested that with nonvolatile cations, the predicted pH values increase by about 0.1 units.”

We agree with the reviewer that RH ranges influence the liquid or solid phases of atmospheric aerosols. So far, there are no observational data showing whether aerosols are in a metastable or stable state in Beijing wintertime (Song et al., 2018). According to previous studies, at low RH, especially when RH < 20% or 30%, aerosols are less likely to be in a completely liquid state (Fountoukis and Nenes, 2007; Guo et al., 2016, 2017). Therefore, we exclude periods with RH < 30% in this study. After that, an average RH value of 50% is now used in the thermodynamic calculations assuming that aerosols were in metastable states. We also did a sensitivity study assuming that solid phases are present. For that case, over 88% of the data resulted in pH values approximating 7.6 with few variations, which is unrealistic. After the correction, the average pH values for year 2013, 2014, and 2017 are 4.5, 4.8, and 5.3, respectively. The results indicate a moderately acidic condition for aerosols in Beijing in winter, consistent with previous studies (Guo et al., 2017; Liu et al., 2017; Song et al., 2018). The correction did not change the trend of pH variation from 2013 to 2017 because the reduced sulfate concentration played a dominant role in pH variation. The corresponding explanations and corrections have been added and updated in the manuscript. “Up to now, there are no observational data showing whether aerosols are in a metastable or stable state in Beijing in winter (Song et al., 2018). According to previous studies, at low RH (RH < 20% or 30%), aerosols are less likely to be in a completely liquid state (Fountoukis and Nenes, 2007; Guo et al., 2016, 2017). Therefore, periods with RH < 30% were excluded in this study.” “During 2013-2017, the average particle pH varied from 4.5 to 5.3, with a significant decrease in sulfate concentration, resulting in a more neutral atmospheric environment. The pH values here agree reasonably with previous ISORROPIA-II calculations, showing that fine particles are moderately acidic in northern China during wintertime (Guo et al., 2017; Liu et al., 2017; Song et al., 2018).”

Figures in the manuscript showing the results of thermodynamic analysis have also been updated:
Figure 9. Sensitivity of (a) AWC and (b) particle pH to the mass concentrations of particulate sulfate and nitrate. The stars indicate the average winter conditions for the years 2013, 2014, and 2017.

Figure 10. Sensitivity of the ammonium partitioning ratio to the mass concentrations of particulate sulfate and nitrate. The stars indicate the average winter conditions for the years 2013, 2014, and 2017.

In the experimental section, there was no discussion of measurement of HNO₃ or NH₃ (gas species), yet ISORROPIA was run such that these data are needed, i.e., run in forward mode. More information is needed in the paper on how the model was run without these critical gas phase species.

To investigate how the variations in particulate nitrate and sulfate concentrations affect aerosol properties, this study used ISORROPIA-II to generate the contour plots in Fig. 9 and Fig. 10. ISORROPIA-II was run in the forward mode, which calculates the equilibrium partitioning with the total concentration of both gas and particle phase species. Previous study shows that the forward mode is less sensitive to measurement error than the reverse mode (Hennigan et al., 2015). To run the model, a selected sulfate concentration with the average temperature, RH, and total ammonia concentration (NH₃ + NH₄⁺) during the winters of 2014 and 2017 was input to ISORROPIA-II, where the total nitrate concentration (HNO₃ + NO₃⁻) was left as the free variable. The gaseous HNO₃ and NH₃ concentrations were not directly measured in this work. To estimate the NH₃ concentration, an empirical equation derived based on long-term measurements in winter in Beijing was
applied, \( \text{NH}_3 \text{ (ppb)} = 0.34 \times \text{NO}_x \text{ (ppb)} + 0.63 \) (Meng et al., 2011). On average, the \( \text{NH}_3 \) concentration was estimated to be around \( 14.0 \mu g \text{ m}^{-3} \) during the winters of 2014 and 2017 in Beijing, consistent with previous observations in the same season (Meng et al., 2011; Zhao et al., 2016; Zhang et al., 2018). For gaseous \( \text{HNO}_3 \),

the total \( \text{NO}_3^- \) concentration (\( \text{HNO}_3^-+\text{aerosol NO}_3^- \)) varying from 0.2 to 75 \( \mu g \text{ m}^{-3} \) was used as the input. More discussions about the consideration of gaseous \( \text{HNO}_3 \) and \( \text{NH}_3 \) concentrations have been added in the manuscript as follows: “The gaseous \( \text{HNO}_3 \) and \( \text{NH}_3 \) concentrations were not directly measured during our campaign. But long-term measurements in Beijing showed that gaseous \( \text{NH}_3 \) concentration correlated well with \( \text{NO}_x \) concentration in winter (Meng et al., 2011). Therefore, the empirical equation derived from Meng et al. (2011), \( \text{NH}_3 \text{ (ppb)} = 0.34 \times \text{NO}_x \text{ (ppb)} + 0.63 \), was applied to estimate the gaseous \( \text{NH}_3 \) concentration. On average, the \( \text{NH}_3 \) concentration was approximated to be \( 14.0 \mu g \text{ m}^{-3} \) during the winters of 2014 and 2017, consistent with previous observations in the same season of Beijing (Meng et al., 2011; Zhao et al., 2016; Zhang et al., 2018). The total nitrate concentration, including both gaseous \( \text{HNO}_3 \) and particulate nitrate, varied from 0.2 to 75 \( \mu g \text{ m}^{-3} \) for the sensitivity study.”

Line 175 or in Table 1 title state this is observational data (not model).

It has been clearly stated in the title of Table 1.

Line 258, what does Until 2017 mean? These changes were completed by 2017?

It means that these changes have been completed by the end of 2017. To make it more clearly, we changed “Until 2017” to “By the end of 2017”.

Line 311, does not tell the complete story. There are publications, see below, that show the predictions of Cheng et al. 2016 and Wang et al. 2016 are likely not correct due to their incorrect calculation of fine particle pH and that this proposed heterogeneous chemistry is highly sensitive to pH. This counter argument should also be noted here in this paper for completeness.


Guo, H., R. J. Weber, and A. Nenes (2017), High levels of ammonia do not raise fine particle pH sufficiently to yield nitrogen oxide-dominated sulfate production, Sci. Reports, 7(12109), DOI:10.1038/s41598-41017-11704-41590.


Agreed. The corresponding argument has been noted in the manuscript as follows: “Recently, studies have found that \( \text{SO}_2 \) oxidation by \( \text{NO}_2 \) in aerosol water with near neutral aerosol acidity, which is usually ignored in current model simulations, plays an important role in the persistent formation of sulfate during haze events in northern China (B. Zheng et al., 2015; Cheng et al., 2016; Wang et al., 2016). Others pointed out that regardless of the high \( \text{NH}_3 \) levels, aerosols are always moderately acidic in northern China, and there are probably other alternative formation pathways contributing to fast sulfate production in haze pollution (Guo et al., 2017b; Liu et al., 2017; Song et al., 2018).”

Line 318-319, define SOR and NOR (ie, what does the acronym stand for?)

The SOR and NOR stand for sulfur oxidation ratio and nitrogen oxidation ratio, respectively. This has been clearly clarified in the manuscript as “The sulfur oxidation ratio (SOR) and nitrogen oxidation ratio (NOR) were further estimated as the molar ratio of sulfate to the sum of sulfate and \( \text{SO}_2 \) and the molar ratio of nitrate to the sum of nitrate and \( \text{NO}_x \), respectively, to quantify the degree of \( \text{SO}_2 \) and \( \text{NO}_x \) oxidations (Zheng et al., 2015; Li et al., 2016).”
Lines 339 to 344. Although this discussion tends to follow the Seinfeld and Pandis discussion of sulfate/nitrate/ammonium interactions, it is largely based on weak intuitive arguments and not a rigorous thermodynamic discussion. It is suggested that these types of statements be avoided. Below are possible other ways to discuss the interplay between sulfate, nitrate and ammonium and gas species, nitric acid and ammonia:

Change: Particulate nitrate in PM2.5 is mainly formed through the neutralization of HNO3 with NH3 to something like: Semivolatile PM2.5 particulate nitrate is formed through the partitioning of HNO3 to the particle phase, which is more favored at higher aerosol pH. pH is affected by gas phase NH3 concentrations, where higher NH3 generally leads to higher pH and so possibly more particulate nitrate.

Thanks for the suggestions. We have revised the text accordingly.

Change: Nitrate formation was also affected by the competition for available NH3 between sulfate and nitrate. In the atmosphere, NH3 prefers to react first with H2SO4 to form ammonium sulfate due its stability. To: Because sulfate is nonvolatile, when it is a significant fraction of the aerosol mass it has a dominant influence on aerosol pH, making the aerosol acidic (low pH). In contrast ammonium and nitrate are semivolatile and so their particle-phase concentrations depend on the meteorological conditions (T, RH) their corresponding gas phase concentrations, (NH3 and HNO3 respectively) and aerosol pH. For example, at high sulfate and moderate NH3 concentrations the aerosol can be too acidic for partitioning of HNO3, but at higher NH3, or if sulfate concentrations drop sufficiently (or RH increases), particle pH will increase and can reach a point at which HNO3 partitioning can occur and nitrate aerosol formed. Lower T also favors partitioning to the particle phase through Henry’s law constants.

Thanks for the suggestions. We have modified the text accordingly.

As for the last line, when RH>60% maybe an additional explanation for the trend in Fig8d is that as RH increases liquid water levels increase resulting in higher pH, which allows more nitrate to partition to the aerosol through a feedback loop, as is discussed later in the paper. That is, the increase in NOR may be due to more than just heterogeneous nitrate production.

Agreed. This additional explanation has been added as “In addition, as RH increases, the AWC increases accordingly, resulting in higher aerosol pH. This allows more semivolatile nitrate to partition to the particle phase through a feedback loop, thus favoring the formation of particulate nitrate.”

Line 366-369, In Fig S9a provide a reference for the calculation of epsilon(NO3-). This pH of 3 at which the sensitivity of epsilon(NO3-) changes, as found in this work, was discussed in detail by Guo et al., ACP 2018, which should be cited.

A reference for the calculation of $\varepsilon$ (NO$_3^-$) has been added. The reference of Guo et al., ACP 2018, has been cited.

Finally, there are a number of studies reporting pH in Beijing over different periods. Some of these did analysis to test the thermodynamic model predictions, which was not done here. A summary of these papers and comparison to pH reported in this paper is warranted to support this analysis.

According to the comments above regarding the thermodynamic analysis, summarization of previous studies reporting pH in Beijing and comparison with results in this study have been included in the manuscript. In addition, more information were added: “Previous studies showed that including the nonvolatile cations in ISORROPIA-II does not significantly affect the pH calculations unless the cations become important relative to anions (Guo et al., 2016; Song et al., 2018). The sensitivity test for Beijing winter conditions suggested that with nonvolatile cations, the predicted pH values increase by about 0.1 units.”
References

Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient thermodynamic equilibrium model for K+·Ca2+·Mg2+·NH4+·Na+·SO42−·NO3−·Cl−·H2O aerosols, Atmos Chem Phys, 7, 4639-4659, 2007.


Guo, H., Weber, R. J., and Nenes, A.: High levels of ammonia do not raise fine particle pH sufficiently to yield nitrogen oxide-dominated sulfate production, Scientific Reports, 7, 12109, 10.1038/s41598-017-11704-0, 2017.


Response to Referee Comment 2 (RC2) on “Rapid transition in winter aerosol composition in Beijing from 2014 to 2017: response to clean air actions” by H. Li et al.

This study combines in-situ measurements of PM1 composition in Beijing during the winters of 2014 and 2017 and the simulation results from a regional chemical transport model to investigate the impacts that the clean air actions in China have had on aerosol chemistry. The relative contributions of anthropogenic emissions, meteorological conditions, and regional transport to the changes in aerosol composition in Beijing are also investigated. This is an interesting work and provides a timely and relevant analysis of a current problem – how aerosol pollution in Beijing responded to the implementation of the Air Pollution Prevent and Control Policy. Overall, the manuscript is well written and fits well within ACP’s aims and scope. However, the current version may require substantial revision before publication can be considered. A major shortcoming in this manuscript is that the Methods section is very much lacking of important technical details, for example on aerosol source contribution analysis and CMT model performance, thereby raises doubts about the credibility of the results. Moreover, there are some inconsistencies in the results and discussions which could cause concerns about the quality of the data, representativeness of the findings, or validity of the conclusions. Additionally, the figure legends and captions are often too brief to make the figures understandable.

We thank the reviewer for the positive feedback and helpful comments. In the revised manuscript, more technical details in the Method section have been added. In the following, we will answer the comments point by point.

Detailed comments:

Line 44-46, how much did air pollutants reduce between 2013 – 2017 in the Beijing area?

Emission changes of different air pollutants in Beijing with the implementation of the clean air actions was analyzed in details in Section 3.2.2 and Figure 6.

Section 2.2. is cursory and provides very little information on the organic aerosol source apportionment analysis. Details must be provided on how the PMF/ME-2 analysis of the ACSM data was performed, what data treatments were implemented, and how the solution conditions (eg number of factors, a value, fpeak) were selected and evaluated.

For the ME-2 analysis of the ACSM organic data, the mass spectra and error matrices were prepared based on the procedures given by Ulbrich et al. (2009) and Zhang et al. (2011). Detailed evaluations of different solutions were provided in the Supplement (Figs. S2-11), including the factor time series, mass spectra, and diurnal patterns with different a values. Corresponding changes can be found in the revised manuscript and Supplement.

In Section 2.3. more information should be provided on the performance of the CMT model at simulating PM2.5 composition and how the modeling results compare to observations for 2014 and 2017 winters separately.

As the reviewer suggested, we added more information about the evaluation of the CMT model performance in Section 2.3 and the Supplement (Table S1-2, Fig. S12).

Section 2.5 lists three assumptions about aerosol properties in the ISORROPIA modeling. Several references are cited and claimed to support these assumptions for this study. But upon reading the references more closely, they don’t seem so as the references either talked about aerosols from different locations or under different meteorological conditions, or simply did not provide direct evidence on aerosol physical states. In fact, given the wintry weather (very low RH and T) and intense local emissions in Beijing, it is hard to believe that internal mixing and single aqueous phase were the prevailing aerosol conditions relevant to this work.
When running the ISORROPIA model, it is assumed that aerosols are internally mixed and composed of a single aqueous phase. Previous study in Beijing during wintertime showed that ISORROPIA predictions with these assumptions were in good agreement with observations (Liu et al., 2017). Up to now, there are no observational data showing whether aerosols are in a metastable (only liquid) or stable (solid plus liquid) state in Beijing in winter (Song et al., 2018). At low RH (RH < 20% or 30%), aerosols are less likely to be in a completely liquid state (Fountoukis and Nenes, 2007; Guo et al., 2016, 2017). Therefore, periods with RH < 30% were excluded in this study and aerosol solutions were assumed to be metastable. Particle liquid-phase separations can occur between the inorganic and organic components. It remains unclear how the phase separations influence pH values. A recent laboratory study suggested that the pH value of the organic-rich fraction under phase separation is about 0.4 units higher than that for a fully mixed aqueous phase (Dallemagne et al., 2016; Song et al., 2018).

Line 168 – 171, there was no mentioning of NH3 and HNO3 measurements in 2.1., but it is mentioned here that the values of NH3+NH4 and NO3+HNO3 were input into the model. Where did the NH3 and HNO3 data come from?

In the revised manuscript, discussions were added about how the gas-phase NH3 and HNO3 were considered in this study. “The gaseous HNO3 and NH3 concentrations were not directly measured during our campaign. But long-term measurements in Beijing showed that gaseous NH3 concentration correlated well with NOx concentration in winter (Meng et al., 2011). Therefore, the empirical equation derived from Meng et al. (2011), NH3 (ppb) = 0.34 × NOx (ppb) + 0.63, was applied to estimate the gaseous NH3 concentration. On average, the NH3 concentration was approximated to be 14.0 μg m⁻³ during the winters of 2014 and 2017, consistent with previous observations in the same season of Beijing (Meng et al., 2011; Zhao et al., 2016; Zhang et al., 2018). The total nitrate concentration, including both gaseous HNO3 and particulate nitrate, varied from 0.2 to 75 μg m⁻³ for the sensitivity study.”

Line 110, missing numbers after “0.” What “a” values were used?

For the ME-2 analysis, an optimal solution of four factors with the a value of 0 was accepted.

Line 167 -168, what does “the transition in aerosol composition” mean in this sentence?

The transition in aerosol composition refers to the variations in nitrate and sulfate concentrations. During the winters of 2014-2017, the mass concentrations of both sulfate and nitrate decreased due to the implementation of the clean air actions. The reduction in sulfate concentration was higher than that in nitrate concentration. In this study, we investigated how the variations in nitrate and sulfate concentrations influence particle properties.

Line 186 – 188, the Xu et al. study was also conducted in Beijing in the winters of similar years, but the nitrate to sulfate ratios reported there were much lower than in this study. Normal measurement uncertainty could not explain such large discrepancies (more than a factor of two in difference). Was it due to measurement artifacts or does it suggest some issues with the representativeness of the measurement data? What’s the implication for the validity of the conclusions presented in this paper?

The study by Xu et al. (2019b) performed aerosol measurements in Beijing from the mid of November to the mid of December in 2014 and 2016. In this work, field measurements were conducted in Beijing from 6 December to 27 February in the year of 2014 and from 11 December to 2 February in the year of 2017. For the same year of 2014, ambient observations were performed in different months in the study by Xu et al. (2019b) and in this work. Different emission intensities and different meteorological conditions influence the formation of particulate nitrate and sulfate in different months. In this work, one of the conclusions is that the ratio of nitrate/sulfate was higher in winter 2017 than in winter 2014 with the implementation of the clean air actions. The decrease in nitrate concentration was larger in the winter of 2017 than in the winter of 2014. This is consistent with the conclusions by Xu et al. (2019b).

Line 205 – 206, organics were higher than sulfate and nitrate and in both 2014 and 2017, so calling Beijing aerosol pollution being “sulfate-driven” or “nitrate-driven” does not seem logical.
The driving factors in this work refers to the species whose mass fractions increased with the increase of aerosol loadings, consistent with the discussions in previous studies (Wang et al., 2014, Li et al., 2017, 2018; Sun et al., 2018; Xu et al., 2019a). The driving factors favored the development of haze pollution and the increase of aerosol mass concentrations. While organics composed a high fraction of aerosols in both 2014 and 2017, the mass fraction of organics in aerosols decreased with the increase of aerosol loadings.

Line 259, for coal usage reduction in addition to quoting the absolute amount, it would be also interesting in knowing the relative amount of reduction.

The total coal usage for coal-fired boilers is around 20.4 million tons in Beijing in 2013. By the end of 2017, the coal use was reduced by more than 17 million tons in Beijing due to emission controls of coal-fired boilers. In other way, the coal usage in Beijing was reduced by ~83% for coal-fired boilers compared to 2013.

Figure 2, according to the diurnal profiles, HOA concentration in 2014 was 2-3 times lower than 2017. If HOA is representative of emissions from transportation, is this level of decrease consistent with the decrease in emissions according to emission inventory? Moreover, the decrease of BC concentration from 2014 to 2017 was between 30-40% but the reduction of total combustion POA (sum of CCOA, BBOA and HOA) was close to 70%. This would suggest some very large, probably unrealistic, changes in the combustion emission factors.

Compared with 2014, the HOA concentration in 2017 was reduced by around 50%. Consistently, emissions of primary organic carbon from transportation decreased by around 30% in Beijing from 2014 to 2017 according to the emission inventory. Compared to 2014, both BC concentration and POA concentration were reduced largely in 2017. This is not only caused by the changes in the combustion emission factors via the more advanced control technology but also contributed by the reduced usage of traditional fuels, i.e., coal and biomass, through the application of clean energy.

Line 289 – 292, the comparisons of PM1 concentrations between different air trajectory classes do not logically lead to a conclusion about how much Beijing aerosol was influenced by polluted air masses transported from surrounding areas. Beijing has local pollution sources which could cause high PM events as well. In fact, in the paragraph immediately beneath, the authors reported that CMT simulation indicates that regional transport contributes to only 30 -40% of PM2.5 in Beijing.

Back trajectory analysis shows the air masses paths as they move through time and space before they arrive at the receptor site. Different aerosol concentrations for different air masses indicate the influence of regional transport. For example, due to the high emissions of air pollutants and sever aerosol pollutions in the southern surrounding areas of Beijing, air masses transported from the south of Beijing usually showed higher PM1 concentrations than the air masses from others regions. According to back trajectory analysis, Beijing was less influenced by polluted air masses from surrounding areas in 2017. This is consistent with the results of CMT model simulations, which showed that the lower PM2.5 concentration from regional transport contributed around 40% to PM2.5 reduction in Beijing in 2017. The results are not contradictory to the main conclusion that emission reductions in Beijing and its surrounding regions played a dominant role in air quality improvement during 2014-2017.

Section 3.2.4, 2nd paragraph, what are the rationales for using SO4/BC and NO3/BC ratios in the analysis? Sulfate is a secondary species with formation time scales usually much longer than the emission time scales of BC. So the physical meaning of SO4/BC ratio is unclear.

The absolute concentration of sulfate and nitrate in the atmosphere is not only controlled by atmospheric chemical reactions but also influenced by boundary layer developments. By using SO4/BC and NO3/BC in the analysis, we can see more clearly how the secondary formation of sulfate and nitrate varied compared to primary emissions.

Section 3.2.4, 2nd paragraph, nitrate/(nitrate+NOx) is not a proper index for the oxidation ratio of nitrogen. Discussions related to NOR should be either removed or revised.
According to the ISORROPIA calculations in this study, the fraction of particulate nitrate in total nitrate was higher than 0.99 for the average winter conditions in both 2014 and 2017. Therefore, it is meaningful to use nitrate/(nitrate+NO\textsubscript{x}) to evaluate the oxidation ratio of nitrogen without the consideration of gas-phase HNO\textsubscript{3}.

For the discussions of the relationship between SOR and RH, it is important to point out that RH was measured locally at the sampling site but sulfate was mostly formed on a regional scale, ie, in air masses upwind of Beijing. Is it valid to assume that in-situ RH measurement data are representative of the RH conditions in the air masses where sulfate was formed?

The locally measured RH was largely influenced by the RH of the air masses transported from the upwind direction. For example, the RH in Beijing is usually low when the dry northwesterly/northeasterly air masses arrive. When Beijing is influenced by the humid southerly air masses, the RH in Beijing is high (An et al., 2019). Therefore, in-situ RH measurements are representative of the RH conditions in the air masses where sulfate was formed. Relationship between sulfate and the locally measured RH has been discussed a lot by previous studies (Zheng et al., 2015; Li et al., 2017; Fang et al., 2019).

Line 330-331, does the higher SOR in 2017 than 2014 necessarily demonstrate “a higher sulfate production rate in 2017”? A relatively larger contribution from background air masses could also lead to higher SOR.

With the implementation of the clean air actions, SO\textsubscript{2} emissions were reduced in both Beijing and the surrounding regions. According to back trajectory analysis, Beijing was less influenced by polluted air masses from regional transport in 2017 and sulfate contributed a smaller fraction in the air masses transported from surrounding areas. Therefore, the higher SOR in 2017 than 2014 indicates the higher sulfate production rate in 2017.

Line 544-545, this citation is incomplete

Updated.

Figures in the supplementary materials are fuzzy, need to use better resolution.

The fuzzy figures in the Supplement have been updated to better resolution.

Figure S6d, explain how to read the figure and the meanings of N-E, W-N, E-S, S-W?

The diurnal pattern in Figure S6d showed the diurnal variations in the contributions of different wind directions. The meanings of N-E, W-N, E-S, and S-W were already explained in the manuscript as “from north to east (N-E; 0° ≤ WD < 90°), east to south (E-S; 90° ≤ WD < 180°), south to west (S-W; 180° ≤ WD < 270°), and west to north (W-N; 270° ≤ WD ≤ 360°)”.

Figure S7, what do the color bars stand for?

The color bars indicate the variations of PM\textsubscript{2.5} concentration, which was labelled in the figure.

References


Response to Short Comment 1 (SC1) on “Rapid transition in winter aerosol composition in Beijing from 2014 to 2017: response to clean air actions” by H. Li et al.

(a) Overview The authors used fine particle data collected in Beijing during the winters of 2014 and 2017 to understand changes in the chemical composition of the PM over a period when particle mass decreased significantly. The results here are consistent with past data, showing a large decrease in PM mass, but also revealing that this is largely because of decreases in sulfate, organics, and unidentified components, as a result of decreasing local and regional emissions. The authors find that meteorology also contributed to lower PM concentrations in 2017 but that this was a smaller factor.

This is an interesting paper with a number of complementary components. On the other hand, there are a number of issues that need to be addressed, as described below. The thermodynamic modeling seems problematic and needs more clarification and information.

We would like to thank Cort Anastasio for giving the constructive and helpful comments and suggestions. In the following, we will answer the comments point by point.

(b) Major Points 1a. Figure 2 and lines 190 – 191: The text states that “All SIA species showed similar diel trends in the two winters: : :.”. For some species this is true, but for many of the species, the diel variability is much lower in 2017 than in 2014, suggesting that while more local sources dominated in 2014, regional sources are more important in 2017. Take CCOA as an example: in 2014, the concentration varied by a factor of 3, while in 2017 the variation was probably only a factor of 1.5 or less. Can the change in diel variability be used as an indicator of local vs. regional pollution?

SIA species refer to sulfate, nitrate, and ammonium. For the SIA species, they showed similar diel trends in the winters of 2014 and 2017. For many of the other species, especially the primary species, their diel variability was much lower in 2017 than in 2014. However, the changes in the diel variability of these primary species can not be easily used as an indicator of local vs. regional pollution. For example, previous studies reported that coal combustion sources in Beijing were mainly attributed to regional transport from surrounding areas (Wang et al., 2011; Shang et al., 2018). Therefore, we could not rely on the diel variability of CCOA to provide information of local vs. regional pollution.

1b. Of course boundary layer height is a major influence on concentrations as well. I suggest you add a plot of CO as a tracer for BL height. Does plotting the ratio of PM component/CO help disentangle chemistry and BL height? These might be useful supplemental plots.

The ambient variations of PM components are not only controlled by atmospheric chemical reactions but are also influenced by boundary layer developments. According to the comment, we analyzed the CO-scaled concentrations for PM components to eliminate the influence of different dilution/mixing conditions. The corresponding diurnal plots are shown below. By plotting the ratio of PM component/CO, we can more clearly see the photochemical production of sulfate, nitrate, ammonium, and OOA during daytime. These diurnal plots are included in the supplementary information as Fig. S15.
2. Figure 8 and its corresponding text. (i) Why normalize panels (a) and (b) to BC concentrations? Since both the numerator and denominator are changing between 2014 and 2017, the ratio seems less effective as a normalizing factor. What do the plots look like if sulfate and nitrate are not normalized by BC? (ii) The data are very noisy and so the authors have chosen to plot the median values (as the solid circles). But what does it look like if the mean and standard deviation are plotted instead? Are there any significant differences between RH values for a given year? If the authors want to stay with median, they should at least show some interquartile ranges or other measures of variability about the median. (iii) Line 323: “: : :the starting point of SOR growth was clearly delayed in 2017: : :”. Given the spread in the data, this is speculative since the median has no associated uncertainty or variability. Can this difference be tested statistically? (iv) Lines 330 – 335. The authors explain that the higher sulfate / (sulfate + SO2) ratio (i.e., SOR) in 2017 at low RH results from less of an oxidant limitation. But isn’t an alternative explanation that a higher fraction of the sulfur pollution in Beijing is from regional transport, which would have a higher SOR since it is more aged? (v) The nitrate / (nitrate + NO2) ratio (i.e., NOR) discussion in lines 339 – 344 seems to be missing a few points. First, NOR is a poor measure for the extent of NOx oxidation since it considers only particulate nitrate and not gas-phase HNO3. Under acidic particle conditions (as seems typical in 2014), most N(V) will be gas-phase HNO3, a result of sulfate acidity driving out the particle nitrate. But this is not accounted for if NOR is calculated only with particulate nitrate, as is done here. Running the e-AIM thermodynamic model on the average 2014 data shows that the gas-phase HNO3 concentration is over an order of magnitude higher than particulate nitrate. Thus N(V) / NOx + N(V), with both phases considered for N(V), would be a better measure of NOx oxidation than NOR. ISOROPPIA could be run to examine HNO3(g) in the two years. (vi) As a consideration for future work, it would be enormously helpful to have gas-phase measurements of NH3 and/or HNO3 to constrain the modeled aerosol pH.
(i) BC is an aerosol species coming from only primary emissions and its variations well reflect the influence of atmospheric mixing conditions. The ratio of sulfate and nitrate to BC can better represent the contributions from chemical reactions. Plots not normalized by BC are shown below.

(ii) Figures with the mean and standard deviations are displayed below. They show no significant differences from the results with median values. The use of median values can to some extent avoid the bias by the odd values.

(iii) In this study, there were around 3-month and 2-month hourly data points for the winters of 2014 and 2017, respectively. To make the comparison between 2014 and 2017 more visible, we checked the mean and median
values. Both of them indicated that compared to the winter of 2014, the starting point of SOR growth with RH was delayed in the winter of 2017.

(iv) With the implementation of the clean air actions, SO\textsubscript{2} emissions were reduced in both Beijing and the surrounding regions. Compared to 2014, Beijing was less influenced by polluted air masses from regional transport, which has been analyzed in detail in Sect. 3.2.3. Therefore, the increase of SOR in 2017 at lower RH is more likely to be caused by the weakened oxidant limitation on SO\textsubscript{2} oxidations.

(v) According to the ISORROPIA calculations in this study, the fraction of particulate nitrate in total nitrate was higher than 0.99 for the average winter conditions in both 2014 and 2017. Therefore, the calculation of NOR without the consideration of gas-phase HNO\textsubscript{3} has no significant influence on the discussion of NO\textsubscript{x} oxidation.

(vi) We agree that the gas-phase measurements of NH\textsubscript{3} and HNO\textsubscript{3} are important for the calculation of aerosol pH. As we stated in the response to Rodney Weber, gas-phase NH\textsubscript{3} and HNO\textsubscript{3} were considered in the ISORROPIA calculations in this work.

3. Figure 9 and its corresponding text. (i) This figure indicates that the aerosol pH in 2014 was approximately 5, but this is inconsistent with the average composition indicated in Table 1. Using the averages for 2014 in Table 1, assuming protons make up the missing positive charge (which I assume was also done in ISORROPIA) and using a relative humidity of 60\% in the e-AIM model results in a pH of -0.04 assuming a proton activity coefficient of 1; the pH is -0.7 if the e-AIM activity coefficient (4.22) is used. These results were done allowing solids to form, but it shouldn’t change significantly if solids were not allowed to form. While the pH undoubtedly varies between samples, it is hard to believe the typical pH is near 5 given the ion imbalance in the data. (ii) Is the range of pH values given in line 361 (pH 5.0 to 6.2) from their work? Why is it for 2013 – 2017, rather than just the two winters of this study? It’s not clear that the authors did much with the ISORROPIA data: why not report ALW and pH for every sample? This would be very useful information. (ii) The 2017 average data in Table 1 has a net positive charge, requiring a nonsensical negative concentration for protons that is of similar magnitude (though smaller) than the proton concentration from the 2014 data. This prevents use of a thermodynamic model and suggests that there are other charged species, likely organics, which are affecting the ion balance and the pH calculation. This should be discussed. Were the authors able to calculate ALW and pH values for the 2017 data? (iii) Under the acidic conditions I calculated for 2014, the value of NO\textsubscript{3} / (HNO3 + NO\textsubscript{3}-) is very small, in contrast to the values described in the manuscript. (v) Similarly, I calculate a NH\textsubscript{4}+/NH\textsubscript{x} ratio near unity in 2014, in contrast to the small values reported in the text.

(i) According to ISORROPIA calculations, the average pH was around 4.8 in the winter of 2014, consistent with the results from previous studies. Many researches have shown that aerosols are moderately acidic in Beijing during wintertime (Guo et al., 2017; Liu et al., 2017; Song et al., 2018). In this study, the ISORROPIA was run in the forward mode. In addition to aerosol chemical compositions, gas-phase NH\textsubscript{3} and HNO\textsubscript{3} were also input into the model. Since only conditions with RH > 30\% were considered, aerosols were assumed to be in a metastable state. Details about the ISORROPIA calculations were provided in the revised manuscript. A few studies have demonstrated that ion balance and equivalent ratios can not be used to predict the acidity of particles, especially under ammonia-rich conditions (Hennigan et al., 2015; Guo et al., 2016; Song et al., 2018). Song et al. (2018) showed that pH values calculated using the forward mode of thermodynamic models are not sensitive to the ion balance.

(ii) The pH values reported in the manuscript were calculated by this study. It has already been stated in the manuscript that “Data for the winter of 2013 were acquired from Sun et al. (2016).” While the ALW and pH were calculated for almost every sample, we did not report them in this study. The research focus of section 3.3 is to investigate how the clean air actions influence particle properties in different years.

(ii) As we stated above, the ion balance is not a good criteria to evaluate particle acidity, especially under ammonia-rich conditions. In addition to aerosol chemical compositions, gas-phase NH\textsubscript{3} and HNO\textsubscript{3} were also input into the thermodynamic model to calculate the pH values.
For the winter conditions of 2014, the fraction of particulate nitrate in total nitrate was higher than 0.99 according to ISORROPIA calculations. This is consistent with the results of Guo et al. (2018), which showed that the fraction of particulate nitrate in total nitrate was approaching 1 under Beijing winter conditions when aerosol pH was above 3.

(v) A $\text{NH}_4^+$/NH$_x$ ratio near unity means that almost all NH$_x$ remains in the particle phase, which is not possible in the winter of Beijing in 2014. Previous observations showed abundant existence of gas-phase NH$_3$ in urban Beijing during wintertime (Meng et al., 2011; Zhao et al., 2016). The ISORROPIA predictions by Guo et al. (2018) demonstrated that with the gas-phase NH$_3$ concentration above 5 μg m$^{-3}$, NH$_4^+$/NH$_x$ ratio stays lower than 0.4.

(c) Minor Points 1. Throughout the manuscript, it would be much better (and more consistent with the other inorganic nomenclature in the paper) to use “Cl$^-$” as the abbreviation for chloride rather than “Chl”.

Changed.

2. lines 183-184: This sentence is unclear. Clearer wording would be “The decreases in particulate nitrate and ammonium during this period were 1.3 and 1.5 g m$^{-3}$, respectively.”

Changed.

3. Figure 1. The second line of the figure caption is unclear. These are not “rates”, but rather decreases in concentrations, yes? Also, three significant figures in the decreases seems one too many.

The second sentence in the caption of Figure 1 were changed to “The decreases in the mass concentrations of different components from 2014 to 2017 are as follows”.

4. line 188: It is odd that the Xu nitrate/sulfate ratios are approximately half of the values in the current work. Why this discrepancy?

The measurements in winter 2014 in this study were conducted from the start of December to the end of February while those by Xu et al. (2019) were from the mid of November to the mid of December. Different emission intensities and different meteorological conditions influence the formation of particulate nitrate and sulfate in different months.

5. lines 203-204: Sulfate is a minor fraction of haze in 2014 as well as 2017, so this statement isn’t very useful. It would be better to indicate the sulfate contributions to each of the pollution classifications in 2017.

In the winter of 2014, the fraction of sulfate in total aerosol was more than 10% for all pollution levels. Therefore, it is not a minor fraction. Compared to 2014, sulfate comprised a smaller fraction of haze development in 2017. This has been stated in the revised manuscript.

6. lines 204–205: This description of nitrate contribution is for 2017?

Yes. This has been clearly stated in the revised manuscript.

7. line 208: State the 2014 and 2017 OA concentrations along with the overall decline.

The OA concentrations in the winters of 2014 and 2017 have been stated in the revised manuscript.

8. line 209: “The contribution from HOA was 2.6 g m$^{-3}$: : :”. This is not as clear as it should be: better to state something like “HOA decreased by 2.6 g m$^{-3}$: : :”.

The sentence has been changed to “The decrease in the mass concentration of HOA was 2.6 μg m$^{-3}$”.

9. lines 240 – 241: It is not clear what cases A, B, C, and D refer to. Are these panels of Fig. S7?

The explanations of cases A, B, C, and D was provided in the experimental method, section 2.3.
10. lines 258-259: This is unclear. Should “Until 2017: : :” be “By 2017: : :”?

Changed.

11. line 298: The 57.9% is not a “rate” and likely has one too many significant figure.

It has been changed to “a decrease of 57.9% in its mass concentration”.

(d) Recommendation I recommend that the paper be accepted after major revisions to address the points above.

References


Guo, H., Weber, R. J., and Nenes, A.: High levels of ammonia do not raise fine particle pH sufficiently to yield nitrogen oxide-dominated sulfate production, Scientific Reports, 7, 12109, 10.1038/s41598-017-11704-0, 2017.


Rapid transition in winter aerosol composition in Beijing from 2014 to 2017: response to clean air actions

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Abstract. The clean air actions implemented by the Chinese government in 2013 have led to significantly improved air quality in Beijing. In this work, we combined the in-situ measurements of the chemical components of submicron particles (PM₁) in Beijing during the winters of 2014 and 2017 and a regional chemical transport model to investigate the impact of clean air actions on aerosol chemistry and quantify the relative contributions of anthropogenic emissions, meteorological conditions, and regional transport to the changes in aerosol chemical composition from 2014 to 2017. We found that the average PM₁ concentration in winter in Beijing decreased by 49.5% from 2014 to 2017 (from 66.2 μg m⁻³ to 33.4 μg m⁻³). Sulfate exhibited a much larger decline than nitrate and ammonium, which led to a rapid transition from sulfate-driven to nitrate-driven aerosol pollution during the wintertime. Organic aerosol (OA), especially coal combustion OA, and black carbon also showed large decreasing rates, indicating the effective emission control of coal combustion and biomass burning. The decreased sulfate contribution and increased nitrate fraction were highly consistent with the much faster emission reductions in sulfur dioxide (SO₂) due to phasing out coal in Beijing compared to reduction in nitrogen oxides emissions estimated by bottom-up inventory. The chemical transport model simulations with these emission estimates reproduced the relative changes in aerosol composition and suggested that the reduced emissions in Beijing and its surrounding regions played a dominant role. The variations in meteorological conditions and regional transport contributed much less to the changes in aerosol concentration and its chemical composition during 2014-2017 compared to the decreasing emissions. Finally, we observed speculated that changes in precursor emissions also possibly altered the aerosol formation mechanisms based on ambient observations. The observed explosive growth of sulfate at a relative humidity (RH) greater than 50% in 2014 was delayed to a higher RH of 70% in 2017, which was likely caused by the suppressed sulfate formation through heterogeneous reactions due to the decrease of SO₂ emissions. The decreased SO₂ emissions suppressed the rapid formation of secondary sulfate through heterogeneous reactions. The observed explosive growth of sulfate at a relative humidity (RH) greater than 50% in 2014 was delayed to a higher RH of 70% in 2017. Thermodynamic simulations showed that the decreased sulfate and nitrate concentrations have lowered the aerosol water content, particle acidity, and ammonium particle fraction. The results in this study demonstrated the response of aerosol chemistry to the stringent clean air actions and identified that the anthropogenic emission reductions are a major driver, which could help to further guide air pollution control strategies in China.

1 Introduction

Beijing, the capital of China, is one of the most heavily polluted cities in the world (Lelieveld et al., 2015), and it frequently experiences severe and persistent haze pollution episodes in winter (Guo et al., 2014). For example, in January 2013, the daily
concentration of ambient particles with an aerodynamic diameter less than 2.5 μm (PM$_{2.5}$) reached a record high of 569 μg m$^{-3}$ in Beijing (Ferreri et al., 2018), which was over 20 times higher than the World Health Organization standard (25 μg m$^{-3}$ for daily average PM$_{2.5}$). As a complex mixture of many different components, ambient aerosols have a range of chemical compositions and originate from various emission sources and formation processes in the atmosphere (Seinfeld and Pandis, 2012). The adverse effects of aerosols on visibility (Pui et al., 2014), climate (IPCC, 2013), and human health (Pope et al., 2009) are intrinsically related to the chemical composition of particles.

To tackle severe aerosol pollution, the Chinese State Council implemented the Air Pollution Prevention and Control Action Plan (denoted as clean air actions) in September 2013, which is the most stringent pollution mitigation policy ever in China. As a consequence, China’s anthropogenic emissions have declined by 59% for SO$_2$, 21% for NO$_x$, 32% for organic carbon (OC), and 28% for black carbon (BC) during 2013-2017 (Zheng et al., 2018). The annual average PM$_{2.5}$ concentration in Beijing decreased by 35.6% from 2013 to 2017, reaching 58 μg m$^{-3}$ in 2017. Combining the bottom-up emission inventory and chemical transport model simulations, our recent study (Cheng et al., 2019) quantified the relative contributions of meteorological conditions, emission reductions from surrounding regions, and emission reductions from local sources to the decrease in PM$_{2.5}$ concentration in Beijing during 2013-2017. While changes in meteorological conditions partially explained air quality improvement in Beijing in 2017, local and regional emission controls played major roles. In addition, the aerosol chemical composition is expected to change correspondingly due to the rapid reductions in precursor emissions, which is not well understood yet because the chemical components of PM$_{2.5}$ are not measured by China’s monitoring network. A few studies have examined the change in aerosol composition in Beijing after 2013, including a semicontinuous measurement of carbonaceous aerosols during 2013-2018 (Ji et al., 2019) and an aerosol mass spectrometry study comparing aerosol composition and size distribution between 2014 and 2016 (Xu et al., 2019). However, neither performed a comprehensive assessment of all the main factors affecting aerosol concentration and its composition. A deep understanding of how the aerosol composition has changed since the clean air actions were activated and the possible linkage between them is urgently needed.

The chemical composition of PM$_{2.5}$ is mainly affected by three following factors: precursor emissions, meteorological conditions, atmospheric chemical reactions, and regional transport and deposition patterns. Emissions are typically the main driver of aerosol composition changes. During 2005-2012, the sulfate concentration in China decreased, while the nitrate concentration increased, which was caused by the considerable reduction in SO$_2$ emissions but limited control of NO$_x$ (Geng et al., 2017). Based on the measurements of organic aerosol (OA) composition in Beijing, a larger decrease in secondary OA than primary OA was found during the 2014 Asia-Pacific Economic Cooperation summit due to the strict emission controls (Sun et al., 2016). Meteorological conditions affect aerosol composition by changing emissions, chemical reactions, and transport and deposition processes (Mu and Liao, 2014). For example, increases in relative humidity (RH) enhance the secondary formation of sulfate through heterogeneous reactions (Zheng et al., 2015; Cheng et al., 2016), and decreases in temperature favor particulate nitrate formation by facilitating gas-to-particle partitioning (Pye et al., 2009; Li et al., 2018). With chemical transport model simulations in China for the years 2004-2012, Mu and Liao (2014) demonstrated that due to the large variations in meteorological parameters in North China, all aerosol species showed large corresponding interannual variations. Furthermore, aerosol characteristics in Beijing are influenced by regional transport from adjacent polluted regions. Polluted air masses from the southern regions contributed more secondary inorganic aerosols (SIAs) than primary aerosols in Beijing (Zhang et al., 2014; Du et al., 2018). Following our previous work (Cheng et al., 2019), the main objective of this study is to investigate the impact of clean air actions on changes in aerosol chemical composition from 2014 to 2017. With both the in-situ observations of aerosol species in Beijing during the winters of 2014 and 2017 and model simulations for the corresponding periods, this work provides the opportunity for a detailed evaluation of the underlying drivers. First, changes in aerosol characteristics are illustrated for inorganics and organics
by comparing aerosol measurements in 2014 and 2017. Then, the relative importance of different factors in varying aerosol composition is assessed by combining direct observations and model simulations, including synoptic conditions, emission changes, regional transport and formation mechanisms. Last, we show that the transition in aerosol characteristics influenced particle properties, such as aerosol water content (AWC) and particle acidity, which in turn affects secondary aerosol formation.

2 Experimental methods

2.1 Ambient sampling and instrumentation

Online aerosol measurements were performed in urban Beijing during the winters of 2014 (from 6 December 2014 to 27 February 2015) and 2017 (from 11 December 2017 to 2 February 2018). The sampling site is located on the roof of a three-story building on the campus of Tsinghua University (40.0° N, 116.3° E), which is surrounded by school and residential areas. No major industrial sources are situated nearby. An Aerodyne Aerosol Chemical Speciation Monitor (ACSM) was deployed for the real-time chemical observations of nonrefractory PM$_1$ (NR-PM$_1$), including organics, sulfate, nitrate, ammonium, and chloride. A detailed description of the instrument can be found in Ng et al. (2011a). The mass concentration of BC in PM$_1$ was measured using a multigangle absorption photometer (MAAP, model 5012; Petzold and Schönlinner, 2004). In addition, the total PM$_{2.5}$ mass was simultaneously recorded with a PM-712 monitor based on the β-ray absorption method (Kimoto Electric Co., Ltd., Japan). For gaseous species, the mixing ratios of SO$_2$, NO$_x$, CO, and O$_3$ were monitored by a suite of commercial gas analyzers (Thermo Scientific). The meteorological parameters, including temperature, RH, wind speed (WS), and wind direction (WD), were obtained from an automatic meteorological observation instrument (MILOS520, VAISALA Inc., Finland).

2.2 ACSM data analysis

The ACSM data were analyzed using the standard analysis software within Igor Pro (WaveMetrics, Inc., Oregon USA). Default relative ionization efficiencies (RIEs) were applied to organics (1.4), nitrate (1.1), and chloride (1.3), while the RIEs of ammonium and sulfate were experimentally determined through calibrations with pure ammonium nitrate and ammonium sulfate, respectively. A composition-dependent collection efficiency (CE) algorithm was used to account for the incomplete detection of aerosol particles (Middlebrook et al., 2012). As shown in Fig. S1, the total measured PM$_1$ mass (NR-PM$_1$ plus BC) correlated well with the PM$_{2.5}$ obtained from PM-712 ($r^2 = 0.80$ and 0.87 for 2014 and 2017, respectively). On average, PM$_1$ accounted for 68% and 80% of the total PM$_{2.5}$ in Beijing during the winters of 2014 and 2017.

The ACSM provides unit-mass-resolution mass spectra of submicron particles, facilitating source apportionment via factor analysis. In this study, positive matrix factorization (PMF) was implemented to resolve OA into various sources using a multilinear engine (ME-2; Paatero, 1999) via the SoFi toolkit (Source Finder; Canonaco et al., 2013). The so-called a value approach allows for the introduction of a priori factor profile or time series to reduce the rotational ambiguity and obtain a unique solution. The spectra and error matrices of organics were pretreated based on the procedures given by Ulbrich et al. (2009) and Zhang et al. (2011). Ions up to larger than m/z 120 were not considered in this study given the interferences of the internal standard of naphthalene at m/z 127-129, and the low signal-to-noise ratio of larger ions. For the winter of 2014, a reference hydrocarbon-like OA (HOA) profile from Ng et al. (2011b) was introduced into the ME-2 analysis to constrain the model performance, varying a values from 0 to 1. Following the guidelines by Canonaco et al. (2013) and Crippa et al. (2014), four factors was finally accepted, with an a value of 0. Detailed evaluation of the factor time series, mass spectra, and diurnal patterns with different a values can be found in the Supplement (Figs. S2-10). Figure S112 shows the source apportionment results
with three primary factors, i.e., HOA, coal combustion OA (CCOA), and biomass burning OA (BBOA), and one secondary factor, oxygenated OA (OOA). For the 2017 dataset, the mass spectral profiles of HOA, CCOA, and BBOA from the ME-2 analysis of 2014 were adopted to constrain the model performance. Similarly, a four-factor solution with HOA, BBOA, CCOA, and OOA was selected for the winter of 2017, which allowed a better comparison of the OA sources between 2014 and 2017.

### 2.3 WRF-CMAQ model

The Weather Research and Forecasting (WRF) model, version 3.8, and the Community Multiscale Air Quality (CMAQ) model, version 5.1, were applied to evaluate the impact of meteorological changes, regional transport and emission variations on the PM$_{2.5}$ concentration in Beijing in winter. The simulated area was designed as three nested domains, and the innermost area covered Beijing and its surrounding regions (including Tianjin, Hebei, Shanxi, Henan, Shandong and Inner Mongolia), with a horizontal resolution of 4 km × 4 km. The simulated period basically followed the observation time, which covered December 2014 – February 2015 and December 2017 – February 2018. A one-month spin-up was applied in each simulation.

The WRF model is driven by the National Centers for Environmental Prediction Final Analysis (NCEP-FNL) reanalysis data, which then provided the meteorological fields for the CMAQ model. We used CB05 and AERO6 as the gas and particulate matter chemical mechanisms, respectively. The in-line windblown dust and photolytic rate calculation modules were also adopted to improve the simulation. The chemical initial and boundary conditions originated from the interpolated outputs of the Goddard Earth Observing System with chemistry (GEOS-Chem) model (Bay et al., 2001).

The anthropogenic emission inventory for Beijing was taken from the Beijing Municipal Environmental Monitoring Center (BMEMC), which was documented and analyzed in Cheng et al. (2019), while the emission inventory outside Beijing was provided by the Multi-resolution Emission Inventory for China (MEIC) (http://www.meicmodel.org; Zheng et al., 2018) and the MIX emission inventory for the other Asian countries (M. Li et al., 2017). The biogenic emissions were obtained by the Model of Emission of Gases and Aerosols from Nature (MEGAN v2.1); however, open biomass burning was not considered in this work. Detailed model configurations and validations can be found in Cheng et al. (2019), and the simulated results well reproduced the temporal and spatial distributions and variations in PM$_{2.5}$ in Beijing and its surrounding areas. In this study, we evaluated the model performance by comparing the simulated PM$_{2.5}$ concentrations and compositions in Beijing with observation data. The hourly observed PM$_{2.5}$ concentrations were collected from the Beijing Municipal Environmental Protection Bureau; and the observed PM$_{2.5}$ compositions came from the Surface PARTiculate mAtter Network (SPARTAN, www.spartan-network.org). We also compared the simulated PM$_{2.5}$ compositions with the observed PM$_1$ species from this work. The average simulated PM$_{2.5}$ in Beijing decreased from 91.5 (winter of 2014) to 52.5 (winter of 2017) μg m$^{-3}$, with a total decrease of 39 μg m$^{-3}$, while the observed PM$_{2.5}$ varied from 81.9 to 40.6 μg m$^{-3}$, decreasing by 41.3 μg m$^{-3}$. Generally, The Pearson correlation coefficients between the simulated and observed PM$_{2.5}$ (24-hour averages) in Beijing were 0.81 (winter of 2014) and 0.78 (winter of 2017). The time series comparison and detailed monthly descriptive statistic of the observed and CMAQ-simulated PM$_{2.5}$ concentrations can be found in Fig. S12 and Table S1. For PM$_{2.5}$ compositions, the Pearson correlation coefficients between simulated PM$_{2.5}$ and observed PM$_1$ components were all above 0.7 (Table S2a), indicating that the model simulations could well reproduce the species variations. Detailed comparisons of the simulated and observed PM$_{2.5}$ components were listed in Table S2b.

We designed six simulation cases to investigate the impact of meteorological and emission variations. Two base cases were driven by the actual emission inventory and meteorological conditions in the winter of 2014 (case A) and winter of 2017 (case B). Cases C and D were designed to quantify the impact of meteorological changes; case C was simulated with the emissions in 2014 and meteorological conditions of 2017, while case D used the 2017 emissions and 2014 meteorological conditions. Therefore, the differences between A and C or between B and D show the influence of meteorological conditions, and the differences between A...
and $D$ or between $B$ and $C$ correspond to the contributions of emission variations. We used the averaged differences as the final impacts. Cases $E$ and $F$ were developed to evaluate the effect of regional transport on PM$_{2.5}$ variations in Beijing in the winter of 2014 ($E$) and winter of 2017 ($F$). In these two cases, the emissions in Beijing were set to zero, while the regional emissions remained at the actual level. The balances between $A$ and $E$ or between $B$ and $F$ represent the contributions of regional transport to the PM$_{2.5}$ concentration in Beijing during the corresponding periods.

### 2.4 Clustering analysis of back trajectories

The Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model was conducted to calculate the back trajectories of air masses arriving in Beijing during the observation periods in 2014 and 2017. The meteorological input was downloaded from the National Oceanographic and Atmospheric Administration (NOAA) Air Resource Laboratory Archived Global Data Assimilation System (GDAS) (ftp://arlftp.arlhq.noaa.gov/pub/archives/). Each trajectory was run for three days, with a time resolution of 1 hour, and the initialized height was 100 m above ground level. In total, 2108 and 1292 trajectories were obtained for the winters of 2014 and 2017, respectively. Based on the built-in clustering calculation, the trajectories were then classified into different groups to represent the main airflows influencing the receptor site. Finally, the optimal 5-cluster and 7-cluster solutions were adopted for the winters of 2014 and 2017, respectively. Details are shown in Fig. S13.

### 2.5 ISORROPIA-II equilibrium calculation

The ISORROPIA-II thermodynamic model was used to investigate the effects of particle chemical composition on aerosol properties, i.e., particle pH, AWC, and the partitioning of semivolatile species (Fountoukis and Nenes, 2007). The model computes the equilibrium state of an NH$_4^+$-SO$_4^{2-}$-NO$_3^-$-Cl$^-$-Na$^+$-Ca$^{2+}$-K$^+$-Mg$^{2+}$-H$_2$O inorganic aerosol system with its corresponding gases (Fountoukis and Nenes, 2007). When running the ISORROPIA-II model, it is assumed that the bulk PM$_4$ or PM$_{2.5}$ properties have no compositional dependence on particle size, and aerosols are internally mixed and composed of a single aqueous phase. The validity of these assumptions has been evaluated by a number of studies in various locations (Guo et al., 2015; Weber et al., 2016; M.X. Liu et al., 2017; Li et al., 2018).

The model was run in the forward mode by assuming that aerosol solutions were metastable. The forward mode calculates the gas-particle equilibrium partitioning with the total concentrations of both gas and particle phase species. Compared to the reverse mode using only aerosol phase compositions, calculations with the forward mode are affected much less by the measurement errors (Hennigan et al., 2015; Guo et al., 2017a; Song et al., 2018). Particle water associated with OA was not considered in this study given its minor effects. M. X. Liu et al. (2017) showed that organic matter (OM)-induced particle water accounted for only 5% of the total AWC in Beijing. Up to now, there are no observational data showing whether aerosols are in a metastable (only liquid) or stable (solid plus liquid) state in Beijing in winter (Song et al., 2018). According to previous studies, at low RH (RH < 20% or 30%), aerosols are less likely to be in a completely liquid state (Fountoukis and Nenes, 2007; Guo et al., 2016, 2017). Therefore, periods with RH < 30% were excluded in this study. The effects of nonvolatile cations (i.e., Na$^+$, K$^+$, Ca$^{2+}$, Mg$^{2+}$) are not considered in this study because the fraction of nonvolatile cations in PM$_1$ in Beijing is generally negligible compared to SO$_{2}^{2-}$, NO$_2^-$, and NH$_4^+$ (Sun et al., 2014). Although nonvolatile nitrate may exist in ambient particles as Ca(NO$_3$)$_2$ and Mg(NO$_3$)$_2$, Ca$^{2+}$ and Mg$^{2+}$ are mainly abundant at sizes above 1 µm (Zhao et al., 2017). In addition, the mixing state of PM$_1$ nonvolatile cations with SO$_{2}^{2-}$, NO$_2^-$, and NH$_4^+$ remains to be investigated (Guo et al., 2016, 2017). Previous studies showed that including the nonvolatile cations in ISORROPIA-II does not significantly affect the pH calculations unless the cations become important relative to anions (Guo et al., 2016; Song et al., 2018). The sensitivity test for Beijing winter conditions suggested that with nonvolatile cations, the predicted pH values increase by about 0.1 units.
In this study, the transition in aerosol composition was mainly reflected in the variations in nitrate and sulfate concentrations. For the analysis of the sensitivity of aerosol properties to particle composition, a selected sulfate concentration combined with the average temperature, RH, and total ammonia concentration (NH$_3$ + NH$_4^+$) during the winters of 2014 and 2017 was input into the ISORROPIA-II model, where the total nitrate concentration (HNO$_3$ + NO$_3^-$) was left as the free variable. The gaseous HNO$_3$ and NH$_3$ concentrations were not directly measured during our campaign. But long-term measurements in Beijing showed that gaseous NH$_3$ concentration correlated well with NO$_x$ concentration in winter (Meng et al., 2011). Therefore, the empirical equation derived from Meng et al. (2011), NH$_3$ (ppb) = 0.34 × NO$_x$ (ppb) + 0.63, was applied to estimate the gaseous NH$_3$ concentration. On average, the NH$_3$ concentration was approximated to be 14.0 μg m$^{-3}$ during the winters of 2014 and 2017, consistent with previous observations in the same season of Beijing (Meng et al., 2011; Zhao et al., 2016; Zhang et al., 2018). The total nitrate concentration, including both gaseous HNO$_3$ and particulate nitrate, varied from 0.2 to 75 μg m$^{-3}$ for the sensitivity study.

3 Results and discussions

3.1 Overall variations in aerosol characteristics from 2014 to 2017

Figures S14 and S15 display the temporal variations in meteorological parameters, trace gases, and aerosol species during the two winter campaigns, with the average values shown in Table 1. Compared to the frequently occurring haze episodes in the winter of 2014, more clean days with lower PM$_1$ concentrations were observed in the winter of 2017. On average, the PM$_1$ concentrations were 66.2 μg m$^{-3}$ and 33.4 μg m$^{-3}$ during the winters of 2014 and 2017, respectively. The large reduction in PM$_1$ concentration reflects the effectiveness of pollution abatement strategies. Satellite-derived estimates also showed an evident decrease in PM$_{2.5}$ concentration in North China in recent years (Gui et al., 2019).

3.1.1 Changes in SIA characteristics

Sulfate, nitrate, and ammonium are the dominant components in SIA$_{s}$ and are generally recognized as ammonium sulfate and ammonium nitrate in PM$_{2.5}$. With the implementation of clean air actions, sulfate underwent the largest decline in the mass concentration among all SIA species (from 7.7 μg m$^{-3}$ to 2.8 μg m$^{-3}$ during 2014-2017). The decreases in particulate contributions of nitrate and ammonium to PM$_{1}$ mass reduction during this period were 1.3 μg m$^{-3}$ and 1.5 μg m$^{-3}$, respectively. Different changes in the mass concentration of SIA species led to variations in the PM$_{1}$ chemical composition. As illustrated in Fig. 1, nitrate exhibited an increasing mass fraction in PM$_1$ from 18% to 30%, whereas the mass contribution of sulfate decreased from 12% to 8%. Correspondingly, the mass ratio of nitrate/sulfate increased from 1.4 in 2014 to 3.5 in 2017. Based on the measurements in Beijing from November to December, Xu et al. (2019) also observed a higher nitrate/sulfate ratio in 2016 (1.36) than in 2014 (0.72). Similar annual variations in aerosol chemical composition were found in North America over 2000-2016, with an increased proportion of nitrate and a decreased contribution of sulfate (van Donkelaar et al., 2019). The diurnal cycles of SIA$_{s}$ are displayed in Fig. 2. All SIA species showed similar diel trends in the two winters, with increasing concentrations after noon due to enhanced photochemical processes and peak concentrations at night caused by a lower boundary layer height. The CO-scaled diurnal plots for SIA species are shown in Fig. S16 to eliminate the influence of different dilution/mixing conditions. However, the absolute variations in the SIA mass concentration differed greatly between 2014 and 2017. While the mass concentration of sulfate decreased by a factor of 2.3 in 2017, nitrate and ammonium showed much smaller reductions of 15-40% in their mass concentrations throughout the day.

Previous studies have concluded that the dramatically enhanced contribution of sulfate was a main driving factor of winter haze pollution in China (Wang et al., 2014; Wang et al., 2016; H. Y. Li et al., 2017). However, with the emission mitigation efforts, the
role of SIA species in aerosol pollution changed significantly. Aerosol pollution was classified into three categories in this study: clean (PM$_1$ ≤ 35 μg m$^{-3}$), slightly polluted (35 < PM$_1$ ≤ 115 μg m$^{-3}$), and polluted (PM$_1$ > 115 μg m$^{-3}$). The contributions of different pollution levels and the PM$_1$ chemical compositions at each pollution level are shown in Fig. 3 for the winters of 2014 and 2017. While the polluted level accounted for 38% of the observation period in the winter of 2014, only 14% of the observation period was recognized as being polluted in the winter of 2017. In 2014, the mass fraction of sulfate in PM$_1$ was 16.1% during clean periods. With the increase in pollution level, the contribution of sulfate increased from 10.6% in slightly polluted periods to 13.6% in polluted periods, while the mass fraction of nitrate decreased. In contrast, sulfate comprised a minor smaller fraction of haze development in 2017 compared to 2014. It was nitrate that exhibited a substantially increased mass fraction at higher PM$_1$ loadings in the winter of 2017. From clean to polluted periods, the nitrate contribution to PM$_1$ increased from 22.6% to 34.9%. These results demonstrated that aerosol pollution in Beijing has gradually changed from sulfate-driven to nitrate-driven in recent years.

3.1.2 Changes in OA characteristics

In response to the strict emission controls, the mass concentration of organics declined by ~18.5 μg m$^{-3}$ from 2014 (30.4 μg m$^{-3}$) to 2017 (11.9 μg m$^{-3}$), which was mainly caused by OOA (~6.8 μg m$^{-3}$) and CCOA (~6.0 μg m$^{-3}$). The contribution from decrease in the mass concentration of HOA was 2.6 μg m$^{-3}$, which was associated with the strengthened controls on vehicle emissions. BBOA decreased by 3.2 μg m$^{-3}$ because the use of traditional biofuels, such as wood and crop residuals, was forbidden in Beijing by the end of 2016. Generally, the concentrations of all OA factors declined substantially throughout the day in 2017. For primary factors, the reductions in their mass concentrations were much higher at night than during the day (Fig. 2). Compared to 2014, CCOA decreased by a factor of 4-5 at night in 2017 and a factor of 1.5 during the day.

Overall, the mass fraction of organics in PM$_1$ declined from 49% to 36% over the period (Fig. 1). The source apportionment results demonstrated that coal combustion was largely accountable for the reduced contribution of organics. During 2014-2017, the mass fraction of CCOA in the total OA decreased from 27% to 18%. Reports from the Beijing Municipal Environmental Protection Bureau (MEPB) also revealed that the contribution of coal combustion to aerosol pollution showed a large decrease during 2013-2017. The decline in CCOA was largely driven by the reduced emissions of organics from coal combustion with the implementation of clean air actions. In contrast, the mass contribution of OOA in the total OA increased from 41% to 49% during 2014-2017. OOA is formed in the atmosphere through various oxidation reactions of volatile organic compounds (VOCs). From 2013 to 2017, VOCs emissions decreased by approximately half in Beijing but remained constant in the surrounding regions. Large amounts of OOA brought to Beijing via regional transport weakened the efforts of local emission cuts. Therefore, stronger emission controls of VOCs need to be placed in both local Beijing and adjacent areas in the future.

3.2 Factors affecting aerosol characteristics from 2014 to 2017

3.2.1 Meteorological conditions

To evaluate the influence of weather conditions on air quality improvement, we compared the daily changes in meteorological parameters during the winters of 2014 and 2017 (Fig. S176). Compared to 2014, the temperature in 2017 was slightly lower throughout the whole day, which may have facilitated gas-particle conversion for semivolatile species, such as ammonium nitrate. Although the RH was similar between 2014 and 2017 during the daytime, the nighttime RH in 2017 was slightly higher than that in 2014, which was favorable for the heterogeneous reactions of secondary species. On average, the observed RH was 29.6% in the winter of 2014 and 33.9% in the winter of 2017. Diurnal cycles of WS showed that the WS in winter of 2017 was somewhat higher, implying beneficial conditions for the dispersal of air pollutants. To illustrate the variations in WD, the observed data were classified into four groups: from north to east (N-E; 0° ≤ WD < 90°), east to south (E-S; 90° ≤ WD < 180°), south to west (S-W;
180° ≤ WD < 270°), and west to north (W-N; 270° ≤ WD ≤ 360°). As displayed in Fig. S17d, the winters of 2014 and 2017 were both dominated by W-N and N-E, which usually bring clean air masses. After noon, the contribution of winds from S-W started to increase. According to previous studies, southerly winds arriving in Beijing generally carry higher levels of air pollutants from the southern regions (Sun et al., 2006; Zhao et al., 2009).

Simulations with the WRF-CMAQ model helped to assess the relative importance of meteorology for changes in aerosol concentration and chemical composition. The effects of meteorology were quantified by comparing cases A and C or cases B and D. The differences between A and D or B and C reflected the effectiveness of emission control. For the total PM$_{2.5}$ concentration, the simulation results clearly demonstrated that variations in meteorology from 2014 to 2017 had a much lower influence on the PM$_{2.5}$ reduction than the changes in air pollutant emissions (Fig. S182). On average, changes in weather conditions resulted in a PM$_{2.5}$ decrease of 9.6 µg m$^{-3}$, which explained 24.8% of the total PM$_{2.5}$ reduction. These results suggest that meteorological variations are far from sufficient to explain PM$_{2.5}$ abatement during 2014-2017. In terms of aerosol composition, we compared the simulated results of cases B and D and found that meteorological changes from 2014 to 2017 had a negligible influence on the chemical composition of PM$_{2.5}$ (Fig. 4). Therefore, we conclude that weather conditions in 2017 marginally favored air quality improvement in Beijing, and emission reductions in air pollutants played a dominant role in the variations in aerosol concentration and composition.

### 3.2.2 Emission changes

According to both the observations (Fig. 1) and simulation results (Fig. 5a), sulfate and organics experienced the largest decreases among different components in Beijing from 2014 to 2017, which is consistent with the considerable emission reductions in SO$_2$ and primary OC in local Beijing and its surrounding regions (Fig. 6; i.e., Tianjin, Hebei, Shandong, Henan, Shanxi, and Inner Mongolia). Comparatively, the wintertime nitrate concentration showed the lowest reduction during 2014-2017, which was expected from the smaller emission cut of NO$_x$ in Beijing and its surrounding areas.

Based on the bottom-up emission inventories (Zheng et al., 2018; Cheng et al., 2019), SO$_2$ emissions decreased by 79.9% in Beijing during 2014-2017, mainly due to the effective control of coal combustion sources and the optimization of the energy structure. Until the end of 2017, all coal-fired power units were shut down, and small coal-fired boilers with capacities of <7 MW were eliminated in Beijing, which reduced coal use by more than 17 million tons. In addition, most of the clustered and highly polluted enterprises and factories were phased out during this period. These control measures remarkably reduced SO$_2$ emissions from power and industry sectors. Enhanced energy restructuring was also implemented in the residential sector. During 2013-2017, more than 2 million tons of residential coal was replaced by cleaner natural gas and electricity, involving 900,000 households in Beijing. Apart from coal burning, the use of traditional biomass, such as wood and crops, was thoroughly forbidden in Beijing by the end of 2016. The strict governance of residential fuel also made substantial contributions to the BC and OC emission reductions in Beijing, which decreased by 71.2% and 59.9%, respectively, during 2014-2017. In comparison, NO$_x$ showed a lower emission reduction of 38.1% from 2014 to 2017 in Beijing. The decline in NO$_x$ emissions was mainly caused by the strengthened emission control of on-road and off-road transportation, the shutdown of all coal-fired power plants, and the application of low-nitrogen-burning (LNB) technologies in industrial boilers. However, due to the insufficient end-of-pipe control of widespread gas-fired facilities and the rapid increase in the vehicle population (the number of vehicles in Beijing increased by nearly 10% during 2013-2017), the NO$_x$ emission reduction in Beijing was not as significant as the SO$_2$ emission reduction.

In adjacent regions, SO$_2$ emissions decreased by 50.6% from 2014 to 2017, while NO$_x$ emissions showed a much smaller reduction of 15.2%. Comparatively, the energy structure adjustments in surrounding areas were less intense than those in Beijing. Emission reductions in SO$_2$ and NO$_x$ in surrounding regions were mainly attributed to ultralow power plant emissions and the reinforced
end-of-pipe control of key industries. Because of the looser emission standards for vehicles and the lack of vehicle management, control measures on transportation in adjacent regions were highly insufficient for NOx emission reduction compared with those in Beijing. Overall, the observed transition in PM1 chemical composition with increasing nitrate contribution and decreasing sulfate fraction was in agreement with the emission changes in their precursors.

3.2.3 Regional transport

Variations in regional weather patterns and emission changes in air pollutants in surrounding regions influenced the effect of regional transport on aerosol characteristics in Beijing. Statistical analysis of air mass trajectories was performed using the HYSPLIT model. Based on the clustering technique, back trajectories were classified into groups of similar length and curvature to identify the main airflows affecting the site. The five-cluster solution and seven-cluster solution were adopted for the winters of 2014 and 2017, respectively. The PM1 mass concentration and mass composition for each cluster are shown in Fig. S198. For a better comparison between 2014 and 2017, clusters were further grouped into two categories according to PM1 loadings. Clusters arriving in Beijing when the local PM1 concentration was less than 35 μg m⁻³ were recognized as clean clusters, while clusters with PM1 concentrations greater than 35 μg m⁻³ were defined as polluted clusters. As displayed in Fig. 7, the average PM1 concentration in local Beijing was 114 μg m⁻³ in 2014 when the polluted clusters arrived, which was much higher than that in 2017 (74 μg m⁻³). While the contribution of polluted clusters in 2014 was 47%, polluted air masses transported from surrounding regions influenced Beijing approximately 20% of the time in 2017. The results here indicate that compared to 2014, Beijing was less influenced by polluted air masses transported from surrounding areas in 2017 during the wintertime, which benefited air quality improvement. In addition, air masses in 2017 brought more nitrate and less sulfate to Beijing than those in 2014.

The WRF-CMAQ model simulations showed that the contributions of regional transport to the PM2.5 concentration in Beijing were 31.4 μg m⁻³ and 19.0 μg m⁻³ in the winters of 2014 and 2017, respectively (Fig. 5b). Although the proportion of regional transport (relative to the total PM2.5 concentration in Beijing) remained at approximately 35% in the two winters (34.4% in the winter of 2014 and 36.4% in the winter of 2017), the absolute amount decreased by 39.6%. This result further supported that less PM2.5 transported from surrounding regions indeed helped with PM2.5 abatement in Beijing. Compared with 2014, the variations in PM2.5 components due to regional transport (Fig. 5b) in 2017 were basically consistent with the total aerosol composition changes that were observed (Fig. 1) and simulated (Fig. 5a) in Beijing. Sulfate had the most notable decrease, with a decreasing rate of 57.9% in its mass concentration, and the regional transport of OM and BC decreased by over 38%. The significant reduction in sulfate was mainly attributed to the effective SO2 emission controls in the surrounding regions, such as the special emission limits for power plants and the innovation of industrial boilers. The decreasing rate of regional transport OM was obviously lower than the total change, suggesting that the local emission controls of VOCs and primary OM in Beijing had a dominant contribution to the decrease in OM. The reduction in nitrate from regional transport was much smaller than that in other components. This was not only due to the insufficient NOx emission controls in the surrounding areas but also the relatively rich ammonium environment in North China, which might have weakened the effects of NOx reductions. Therefore, the collaborative reductions in NOx and NH3 are important for future air pollution control strategies (Liu et al., 2019).

3.2.4 Formation mechanisms

From a traditional viewpoint, sulfate formation mainly includes SO2 oxidation by OH in the gas phase and SO2 oxidation in cloud droplets by H2O2 and O3 in the aqueous phase (Seinfeld and Pandis, 2012). This is actually the case for global sulfate production (Roelofs et al., 1998). The formation rate of sulfate through aqueous reactions is typically much faster than that through gas-phase oxidations. Recently, studies have found that the heterogeneous oxidation of SO2 oxidation by NO2 in aerosol water with near
neutral aerosol acidity, which is usually ignored in current model simulations, plays an important role in the persistent formation of sulfate during haze events in northern China (B. Zheng et al., 2015; Cheng et al., 2016; Wang et al., 2016). However, others pointed out that regardless of the high NH$_3$ levels, aerosols are always moderately acidic in northern China, and there are probably other alternative formation pathways contributing to fast sulfate production in haze pollution (Guo et al., 2017b; Liu et al., 2017; Song et al., 2018). As the SO$_2$ emissions decreased substantially with the clean air actions, the substantial decrease in SO$_2$ emissions currently, the importance of heterogeneous chemistry in sulfate formation is highly uncertain.

To shed light on this query, the formation of sulfate and nitrate with increasing RH was compared between 2014 and 2017. As displayed in Fig. 8, the SO$_2$/BC ratio was much lower in 2017 than in 2014, especially at a higher RH, indicating greatly weakened sulfate formation in 2017 compared to primary BC emissions. NO$_2$/BC showed little difference between 2014 and 2017. The sulfur oxidation ratio (SOR) and nitrogen oxidation ratio (NOR) of sulfur and nitrogen were further estimated as SOR (the molar ratio of sulfate to the sum of sulfate and SO$_2$) and NOR (the molar ratio of nitrate to the sum of nitrate and NO$_2$), respectively, to quantify the degree of SO$_2$ and NO$_2$ oxidations (Zheng et al., 2015; Li et al., 2016). Median values were used for comparison between 2014 and 2017 to avoid bias caused by outliers. When the RH>50%, SOR started to increase significantly with the enhancement in RH in 2014, which was consistent with previous observations in Beijing in 2013 (G. J. Zheng et al., 2015). A year-long study in Beijing from 2012 to 2013 also revealed that a rapid increase in SOR was found at a RH threshold of ~45% (Fang et al., 2019). However, the starting point of SOR growth was clearly delayed in 2017, with a higher RH of 70%. Considering the decrease in the SO$_2$ mixing ratio from 15.5 ppb in the winter of 2014 to 2.8 ppb in the winter of 2017 (Table 1), the results here imply we speculated that with the large reduction in gaseous precursors, the rapid formation of sulfate through heterogeneous reaction is more difficult to occur. In addition to emission reduction, reduced regional transport from southern polluted regions in 2017 helped to lower SO$_2$ concentrations in Beijing. Previous studies have revealed the positive feedback between aerosols and boundary layers, as high aerosol loadings could decrease the boundary layer height and further increase aerosol concentrations (Petäjä et al., 2016; Z. Li et al., 2017). With a lower PM$_{2.5}$ concentration in 2017, the interactions between aerosols and the boundary layer were weakened, which in turn also favored a decrease in the SO$_2$ concentration. At a lower RH, the SOR in 2017 (~0.14) was unexpectedly higher than that in 2014 (~0.06), demonstrating a higher sulfate production rate in 2017. Similar results have been observed over the eastern United States, where a considerable decrease in SO$_2$ resulted in the more efficient formation of particulate sulfate during wintertime (Shah et al., 2018). Combining airborne measurements, ground-based observations, and GEOS-Chem simulations, Shah et al. (2018) explained that sulfate production in winter is limited by the availability of oxidants and particle acidity. At lower concentrations of precursor gases, the oxidant limitation on SO$_2$ oxidation weakened, leading to a higher formation rate of sulfate.

Semivolatile PM$_{2.5}$ nitrate is formed through the partitioning of HNO$_3$ to the particle phase, which is more favored at higher aerosol pH. Aerosol pH is affected by gas phase NH$_3$ concentrations, where higher NH$_3$ generally leads to higher pH and therefore possibly more particulate nitrate. Aerosol pH is also influenced by the abundance of particulate sulfate (Seinfeld and Pandis, 2012). Sulfate is nonvolatile in the atmosphere. When sulfate is a significant fraction of aerosol mass, it has a dominant influence on aerosol pH, making aerosol acidic (low pH). In contrast, ammonium and nitrate are mainly semivolatile in the atmosphere. The particle-phase concentrations of ammonium and nitrate depend on the meteorological conditions (i.e., temperature and RH), their corresponding gas phase concentrations (NH$_3$ and HNO$_3$ respectively), and aerosol pH. For example, at high sulfate and moderate NH$_3$ concentrations, aerosols can be too acidic for the partitioning of HNO$_3$ to particle phase. However, at higher RH, or higher NH$_3$ concentrations, or if sulfate concentrations drop sufficiently, particle pH will increase and can reach a point at which HNO$_3$ partitioning to the particle phase occurs and particulate nitrate is formed. Lower temperature also favors HNO$_3$ partitioning to the particle phase through Henry’s law constants. Particulate nitrate in PM$_{2.5}$ is mainly formed through the neutralization of HNO$_3$ with NH$_3$. HNO$_3$ is produced by NO oxidation via OH during the day and the hydrolysis reaction of dinitrogen pentoxide (N$_2$O$_5$)
...ions (temperature enhancement in the atmosphere, \( \text{NH}_3 \) prefers to react first with \( \text{H}_2\text{SO}_4 \) to form ammonium sulfate due to its stability. When excess \( \text{NH}_3 \) is available, ammonium nitrate is formed (Seinfeld and Pandis, 2012). With the decrease in sulfate concentration in 2017, some \( \text{NH}_3 \) was freed up to react with HNO\(_3\). This may have also facilitated the formation rate of nitrate. When RH>60%, NOR increased substantially in 2014 and 2017, indicating the importance of heterogeneous reactions in nitrate production. In addition, as RH increases, the AWC increases accordingly, resulting in higher aerosol pH. This allows more semivolatile nitrate to partition to the particle phase through a feedback loop, thus favoring the formation of particulate nitrate.

### 3.3 Influence of the transition in aerosol characteristics on particle properties

According to thermodynamic calculations, various aerosol properties were affected by changes in aerosol characteristics associated with clean air actions. As shown in Fig. 9a, nitrate and sulfate play key roles in determining the AWC in PM\(_{2.5}\). The decreasing mass concentrations of nitrate and sulfate result in a lower AWC. Similar observations have been reported previously across northern China, revealing that nitrate and sulfate are dominant anthropogenic inorganic salts driving AWC (Wu et al., 2018). With the clean air actions enacted, the mass concentrations of nitrate and sulfate decreased during 2013-2017, leading to an average decline in AWC from 12.048 to 8.5 \( \mu \text{g m}^{-3} \). Data for the winter of 2013 were acquired from Sun et al. (2016). The reduced AWC further helped air quality improvement by lowering the ambient aerosol mass and enhancing visibility. Because aqueous-phase reactions contribute largely to sulfate formation in winter, the decrease in AWC decelerated the formation of sulfate. In addition, the lower AWC slowed down the uptake coefficient of \( \text{N}_2\text{O}_5 \) for heterogeneous processing, thereby suppressing the formation of particulate nitrate.

Figure 9b displays the effects of nitrate and sulfate concentrations on particle acidity. Particle acidity is largely driven by the mass concentration of sulfate and is less sensitive to the variation in nitrate. Particle pH substantially decreases with increasing sulfate concentration. Ding et al. (2019) suggested that sulfate is one of the common driving factors influencing particle acidity in Beijing across all four seasons. In contrast, more particulate nitrate leads to a slightly higher pH by increasing the particle liquid water and diluting aqueous H\(^+\) concentrations. Through the comparison of pH predictions among various locations worldwide, Guo et al. (2018) also found that a higher particle pH was generally associated with higher concentrations of nitrate. During 2013-2017, the average particle pH varied from 4.55 to 3.362, with a significant decrease in sulfate concentration, resulting in a more neutral atmospheric environment. The pH values here agree reasonably with previous ISORROPIA-II calculations, showing that fine particles are moderately acidic in northern China during wintertime (Guo et al., 2017a; Liu et al., 2017; Song et al., 2018; Ding et al., 2019). When pH > 5.0, aqueous-phase productions of sulfate are dominated by \( \text{SO}_2 \) oxidation with \( \text{H}_2\text{O}_2 \), \( \text{O}_3 \), and \( \text{NO}_2 \) under haze conditions in Beijing (Cheng et al., 2016). The sulfate oxidation rates by \( \text{O}_3 \) and \( \text{NO}_2 \) increase with increasing particle pH. Therefore, a more neutral atmosphere would favor aqueous-phase sulfate formation in Beijing. Particle acidity also influences the gas-particle partitioning of nitrate. The rising particle pH would result in a higher fraction of particulate nitrate (\( \varepsilon (\text{NO}_3^-) = \frac{[\text{NO}_3^-]}{[\text{HNO}_3]+[\text{NO}_3^-]} \) (Guo et al., 2016). Figure S209a displays the variation in \( \varepsilon (\text{NO}_3^-) \) as a function of particle pH under typical Beijing winter conditions (temperature of approximately 0\(^\circ\)C). With a particle pH below 3, \( \varepsilon (\text{NO}_3^-) \) increases sufficiently with the enhancement in particle pH. However, when the particle pH is larger than 3, \( \varepsilon (\text{NO}_3^-) \) remains relatively stable (approaching 1), consistent with previous findings by Guo et al. (2018). From 2013 to 2017, with the particle pH remaining above 3 in Beijing, no clear change in \( \varepsilon (\text{NO}_3^-) \) was observed (Fig. S209b).
The variations in nitrate and sulfate concentrations also affected the gas-particle partitioning of total ammonium (NH₄ = NH₄⁺ + NH₃). As expected, the decreased concentrations of nitrate and sulfate led to a reduction in the ammonium particle fraction (∈ (NH₄⁺) = NH₄⁺/NH₄; Fig. 10). From 2013 to 2017, ∈ (NH₄⁺) in Beijing always stayed below 0.43, indicating that most ammonium existed in the gas phase. Therefore, a minor reduction in NH₄ would not be sufficient for air quality improvement. Guo et al. (2018) revealed that for winter haze conditions in Beijing, an approximate 60% decrease in NH₄ was required to achieve an effective reduction in PM₂.₅. Due to the close linkage between ammonia emissions and agricultural activities, it may be difficult to attain substantial ammonia reduction in China.

4 Conclusions

This study investigated the variations in aerosol characteristics in Beijing during the winters of 2014 and 2017 by combining the online measurements of aerosol chemical composition with a comprehensive model analysis of meteorological conditions, anthropogenic emissions, and regional transport. The average PM₁ concentration decreased from 66.2 μg m⁻³ in the winter of 2014 to 33.4 μg m⁻³ in the winter of 2017, with decreasing concentrations of organics, sulfate, nitrate, and ammonium by 18.5 μg m⁻³, 4.9 μg m⁻³, 1.3 μg m⁻³, and 1.5 μg m⁻³, respectively. These changes reduced the mass fractions of organics and sulfate from 59% to 36% and from 13% to 9%, respectively, whereas increased the nitrate contribution from 19% to 32%. Consequently, the winter haze pollution changed from sulfate-driven to nitrate-driven in Beijing from 2014 to 2017, implicating the increasing role of nitrate in aerosol pollution.

The chemical transport model simulations suggest that the rapidly declining emissions in Beijing and its adjacent regions account for ~75% of PM₂.₅ abatement in Beijing, and the remaining portion can be explained by the favorable weather conditions in 2017. The faster reductions in SO₂ emissions compared to NOₓ emissions are in line with the decreased sulfate contribution and increased nitrate fraction in observed aerosols, and the model simulations with these emission estimates can reproduce the relative changes in aerosol composition. Regional transport contributed moderately to the variations in aerosol concentration and its chemical composition, with less polluted air masses transported from surrounding regions to Beijing in the winter of 2017. The air masses were observed to have brought more nitrate and less sulfate to Beijing. Furthermore, The fast SO₂-to-sulfate conversion through heterogeneous reactions was observed to increase promptly at a RH threshold of ~50% in 2014, while a higher RH of 70% was observed in 2017. Based on these ambient observations, the suppressed sulfate formation during wintertime was possibly caused by the considerable decrease in SO₂ emissions suppressed the rapid formation of sulfate during wintertime. The fast SO₂-to-sulfate conversion through heterogeneous reactions was observed to increase promptly at a RH threshold of ~50% in 2014, while a higher RH of 70% was observed in 2017.

Thermodynamic calculations showed that the decreased sulfate and nitrate concentrations in 2017 caused a lower AWC in PM₂.₅, which further decreased the ambient aerosol mass and weakened the formation rates of sulfate and nitrate through aqueous-phase reactions. Particle acidity displayed a decline during 2014-2017, mostly driven by the declining sulfate concentration. In turn, the more neutral ambient environment would favor the aqueous oxidation of sulfate in Beijing. Analysis of the ammonium particle fraction indicated that most ammonium in Beijing existed in the gas phase. Therefore, increased efforts are needed to achieve an effective reduction in particle ammonium in the future.
Author contributions

QZ and KH conceived the study. HL conducted the field measurements and carried out the data analysis. JC provided the emission data and performed the model simulations. BZ participated the data analysis. HL, JC and QZ wrote the paper with inputs from all coauthors.

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Table 1 Summary of the average meteorological parameters, mixing ratios of gaseous species, and mass concentrations of the PM$_1$ chemical components observed during the winters of 2014 and 2017.

<table>
<thead>
<tr>
<th>Sampling period</th>
<th>2014 winter</th>
<th>2017 winter</th>
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<tbody>
<tr>
<td><strong>Meteorological parameters</strong></td>
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<tr>
<td>T (°C)</td>
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<tr>
<td>RH (%)</td>
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<tr>
<td>WS (m s$^{-1}$)</td>
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<td>SO$_2$ (ppb)</td>
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<td>NO$_2$ (ppb)</td>
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<td>CO (ppm)</td>
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<tr>
<td>O$_3$ (ppb)</td>
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<td><strong>Aerosol species (μg m$^{-3}$)</strong></td>
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</table>
[Diagram of pie charts showing the percentage distribution of various components in 2014 and 2017 for different categories: (a) OM, SO_4^{2-}, NO_3^-, NH_4^+, Chl, BC; (b) OM, SO_4^{2-}, NO_3^-, NH_4^+, Cl^-; (c) HOA, CCOA, BBOA, OOA; (d) HOA, CCOA, BBOA, OOA. The percentages for each component are indicated in the charts.]
Figure 1. Average chemical compositions of PM$\text{\textsubscript{1}}$ and OA in (a, c) winter of 2014 and (b, d) winter of 2017. The decreasing rates of decreases in the mass concentrations of different components from 2014 to 2017 are as follows: 60.9% for organics, 64.1% for sulfate, 11.6% for nitrate, 21.7% for ammonium, 50.0% for chloride, and 37.5% for BC.
Figure 2. Average diurnal cycles of different aerosol species in the winter of 2014 (solid line) and winter of 2017 (dashed line).
Figure 3. Aerosol chemical composition at different pollution levels in the (a) winter of 2014 and (b) winter of 2017. The contributions of each pollution level are shown at the top of each bar.
Figure 4. The average PM$_{2.5}$ chemical composition simulated by the WRF-CMAQ model for the observation periods in 2017: (a) base scenario with the 2017 emissions and the 2017 meteorological conditions; (b) simulation with the 2017 emissions and 2014 meteorological conditions.
Figure 5. Simulated concentrations of PM$_{2.5}$ and its chemical components during the observation periods of 2014 and 2017: (a) total changes in Beijing and (b) changes due to regional transport.
Figure 6. Changes in the anthropogenic emissions of SO₂, NOₓ, NH₃, VOCs, primary PM₂.₅, BC, and OC in (a) Beijing and (b) its surrounding regions from 2014 to 2017.
(a) Clean: 1, 2
Polluted: 3, 4, 5

(b) Clean: 1, 2, 3, 4, 5
Polluted: 6, 7
Figure 7. Comparison of the air masses arriving in Beijing between 2014 and 2017. (a) and (b) show the clustering analysis of the back trajectories in the winters of 2014 and 2017, respectively, with pie charts displaying the contributions of the clean and polluted air masses. The stacked bar charts on the right show the average aerosol compositions for the clean and polluted clusters.
Figure 8. Variations in (a) $\text{SO}_4^{2-}/\text{BC}$, (b) $\text{NO}_3^-/\text{BC}$, (c) SOR, and (d) NOR plotted against increasing RH. The data are also binned according to RH values, with the median value shown for each bin.
Figure 9. Sensitivity of (a) AWC and (b) particle pH to the mass concentrations of particulate sulfate and nitrate. The stars indicate the average winter conditions for the years 2013, 2014, and 2017.
Figure 10. Sensitivity of the ammonium partitioning ratio to the mass concentrations of particulate sulfate and nitrate. The stars indicate the average winter conditions for the years 2013, 2014, and 2017.
Supplementary information

Rapid transition in winter aerosol composition in Beijing from 2014 to 2017: response to clean air actions

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Table S1. Monthly descriptive statistic of the comparison between the observational PM$_{2.5}$ concentrations and CMAQ-simulated PM$_{2.5}$ concentrations in Beijing during the winters of 2014 and 2017. Concentration values are all reported in μg m$^{-3}$.

<table>
<thead>
<tr>
<th>Year</th>
<th>Month</th>
<th>Mean_obs</th>
<th>Mean_sim</th>
<th>R</th>
<th>MB</th>
<th>ME</th>
<th>RMSE</th>
<th>NMB(%)</th>
<th>NME(%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>2014</td>
<td>12</td>
<td>58.75</td>
<td>77.87</td>
<td>0.91</td>
<td>19.12</td>
<td>24.53</td>
<td>39.57</td>
<td>13.73</td>
<td>31.24</td>
</tr>
<tr>
<td>2015</td>
<td>1</td>
<td>97.76</td>
<td>104.05</td>
<td>0.75</td>
<td>6.29</td>
<td>32.27</td>
<td>34.84</td>
<td>15.14</td>
<td>37.29</td>
</tr>
<tr>
<td></td>
<td>2</td>
<td>93.36</td>
<td>92.66</td>
<td>0.86</td>
<td>-0.70</td>
<td>21.26</td>
<td>31.47</td>
<td>15.14</td>
<td>37.29</td>
</tr>
<tr>
<td>2017</td>
<td>12</td>
<td>43.35</td>
<td>47.80</td>
<td>0.90</td>
<td>4.45</td>
<td>19.86</td>
<td>25.17</td>
<td>12.01</td>
<td>28.06</td>
</tr>
<tr>
<td>2018</td>
<td>1</td>
<td>33.87</td>
<td>55.76</td>
<td>0.79</td>
<td>21.89</td>
<td>35.97</td>
<td>36.41</td>
<td>29.45</td>
<td>44.45</td>
</tr>
<tr>
<td></td>
<td>2</td>
<td>50.14</td>
<td>54.28</td>
<td>0.83</td>
<td>4.14</td>
<td>34.55</td>
<td>32.83</td>
<td>14.08</td>
<td>27.39</td>
</tr>
</tbody>
</table>

Table S2 (a). Correlations between simulated PM$_{2.5}$ compositions and observed PM$_{1}$ compositions in this study.

<table>
<thead>
<tr>
<th>Period</th>
<th>OM</th>
<th>SO$_4^{2-}$</th>
<th>NO$_3$</th>
<th>NH$_4^+$</th>
<th>BC</th>
</tr>
</thead>
<tbody>
<tr>
<td>2014 Winter</td>
<td>0.83</td>
<td>0.76</td>
<td>0.82</td>
<td>0.84</td>
<td>0.76</td>
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<tr>
<td>2017 Winter</td>
<td>0.72</td>
<td>0.81</td>
<td>0.89</td>
<td>0.90</td>
<td>0.73</td>
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</tbody>
</table>
Table S2 (b). Comparison of the simulation and observations of PM$_{2.5}$ compositions. The observation data was collected from the SPARTAN, Beijing Site (40.01° N, 116.207° E) (https://www.spartan-network.org/beijing-china); while the simulated data was extracted from the same grid where the Beijing Site located in the third domain. Concentration values are all reported in μg m$^{-3}$.

<table>
<thead>
<tr>
<th>year</th>
<th>Date_range (dd/mm)</th>
<th>PM$_{2.5}$</th>
<th>SIA</th>
<th>SO$_4^{2-}$</th>
<th>NO$_3^-$</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Obs</td>
<td>Sim</td>
<td>MB</td>
<td>Obs</td>
<td>Sim</td>
</tr>
<tr>
<td>2014</td>
<td>01/12-28/12</td>
<td>64.12</td>
<td>70.31</td>
<td>6.19</td>
<td>15.14</td>
</tr>
<tr>
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<td>22/01-27/02</td>
<td>95.38</td>
<td>97.64</td>
<td>2.26</td>
<td>18.55</td>
</tr>
<tr>
<td>2017</td>
<td>27/11-12/12</td>
<td>39.14</td>
<td>42.70</td>
<td>3.56</td>
<td>14.21</td>
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</tbody>
</table>

<table>
<thead>
<tr>
<th>year</th>
<th>Date_range (dd/mm)</th>
<th>NH$_4^+$</th>
<th>RM/OM</th>
<th>BC</th>
<th>OTHER</th>
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</thead>
<tbody>
<tr>
<td></td>
<td>Obs</td>
<td>Sim</td>
<td>MB</td>
<td>Obs</td>
<td>Sim</td>
</tr>
<tr>
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<td>01/12-28/12</td>
<td>3.17</td>
<td>4.46</td>
<td>1.29</td>
<td>30.64</td>
</tr>
<tr>
<td>2015</td>
<td>22/01-27/02</td>
<td>4.61</td>
<td>5.23</td>
<td>0.62</td>
<td>30.00</td>
</tr>
<tr>
<td>2017</td>
<td>27/11-12/12</td>
<td>3.20</td>
<td>2.35</td>
<td>-0.85</td>
<td>14.25</td>
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</tbody>
</table>

$^a$SIA represents the total amount of secondary inorganic aerosol, that the sum of SO$_4^{2-}$, NO$_3^-$ and NH$_4^+$.

$^b$RM represents the residual matter in SPARTAN observations. There is no direct organic composition in SPARTAN, however RM in SPARTAN measurements is dominated by organics, and has been used to evaluate the organic mass simulations (Weagle et al., 2018). OM represents the organic mass in CMAQ simulations.

$^c$BC represents the black carbon.

$^d$Other represents the rest of the PM$_{2.5}$ compositions, such as crustal matter, sea salt.
Figure S1. Correlation of total PM$_1$ concentration (NR-PM$_1$ plus BC) with PM$_{2.5}$ concentration during the winter of (a) 2014 and (b) 2017.
Figure S2. Factor profiles of HOA with different $a$ values of ME-2 analysis.
Figure S3. Factor profiles of CCOA with different $a$ values of ME-2 analysis.
Figure S4. Factor profiles of BBOA with different $\alpha$ values of ME-2 analysis.
Figure S5. Factor profiles of OOA with different $\alpha$ values of ME-2 analysis.
Figure S6. Time series of HOA with different $a$ values of ME-2 analysis.
Figure S7. Time series of CCOA with different $a$ values of ME-2 analysis.
Figure S8. Time series of BBOA with different $a$ values of ME-2 analysis.
Figure S9. Time series of OOA with different $a$ values of ME-2 analysis.
Figure S10. Diurnal cycles of (a) HOA, (b) CCOA, (d) BBOA, and (d) OOA with different $a$ values of ME-2 analysis.
Figure S112. Source apportionment results of OA during the winter of 2014. (a) Mass spectra of HOA, CCOA, BBOA, and OOA. (b) Time series of different OA factors and their corresponding tracers. The correlation coefficients of OA factors with the tracers are also shown. (c) Diurnal cycles of OA factors. (d) The average fractional pie chart of OA factors to total OA. (e) The average diurnal mass contributions of OA factors to total OA.
Figure S12. Comparison of observed (red) and CMAQ-simulated (blue) daily mean PM$_{2.5}$ concentrations over Beijing during the winters of 2014 (a) and 2017 (b). Observation data was obtained and averaged from 12 national observation stations in Beijing. Simulated concentrations were extracted from the grids corresponding to the station locations.
Figure S13. Back trajectories of air masses arriving in Beijing in winter 2014 and winter 2017 with unique colors for different clusters. The average percentage of each cluster is shown.
Figure S14. Temporal variations of (a, b) meteorological parameters, (c, d) gaseous species, and (e) aerosol species during the winter of 2014.
Figure S15. Temporal variations of (a, b) meteorological parameters, (c, d) gaseous species, and (e) aerosol species during the winter of 2017.
Figure S16. Average diurnal cycles of CO-scaled aerosol species in the winter of 2014 (solid line) and winter of 2017 (dashed line).
Figure S176. Diurnal cycles of meteorological parameters in 2014 and 2017. (a-c) The average variations of temperature, relative humidity (RH), and wind speed (WS). The shaded areas indicate 25th and 75th percentiles. (d) The distribution of different wind directions through the day. The lines and the stacked areas represent the year of 2014 and 2017, respectively.
Figure S187. WRF-CMAQ simulated PM$_{2.5}$ concentration under different scenarios for the observation periods of 2014 and 2017.
Figure S198. The mass concentrations of aerosol species and their fractional contributions in total PM$_1$ for different clusters in (a, b) winter 2014 and (c, d) winter 2017. The average PM$_1$ concentration of each cluster is shown on the top of the bar.
Figure S209. (a) Variation of $\epsilon (NO_3^-)$ as a function of particle pH. (b) Sensitivity of the fraction of particulate nitrate to the mass concentrations of particulate sulfate.