OBSERVATIONS OF HIGHLY OXIDISED MOLECULES AND PARTICLE NUCLEATION IN THE ATMOSPHERE OF BEIJING

James Brean¹, Roy M. Harrison¹,†, Zongbo Shi¹
David C.S. Beddows¹, W. Joe F. Acton²
C. Nicholas Hewitt², Freya A. Squires³ and James Lee³

¹Division of Environmental Health and Risk Management, School of Geography, Earth and Environmental Sciences
University of Birmingham
Edgbaston, Birmingham B15 2TT
United Kingdom

²Lancaster Environment Centre
Lancaster University, Lancaster LA1 4YQ
United Kingdom

³National Centre for Atmospheric Science
Wolfson Atmospheric Chemistry Laboratory
University of York, York YO10 5DD
United Kingdom

* To whom correspondence should be addressed.
Tele: +44 121 414 3494; Fax: +44 121 414 3709; Email: r.m.harrison@bham.ac.uk

†Also at: Department of Environmental Sciences / Center of Excellence in Environmental Studies, King Abdulaziz University, PO Box 80203, Jeddah, 21589, Saudi Arabia
ABSTRACT

Particle nucleation is one of the main sources of atmospheric particulate matter by number, with new particles having great relevance for human health and climate. Highly oxidised multifunctional organic molecules (HOMs) have been recently identified as key constituents in the growth, and, sometimes, in initial formation of new particles. While there have been many studies of HOMs in atmospheric chambers, flow tubes and clean environments, analyses of data from polluted environments are scarce. Here, measurements of HOMs and particle size distributions down to small molecular clusters are presented alongside volatile organic compounds (VOC) and trace gas data from a campaign in June 2017, in Beijing. Many gas phase HOMs have been characterised and their temporal trends and behaviours analysed in the context of new particle formation. The HOMs identified have a comparable degree of oxidation to those seen in other, cleaner, environments, likely due to an interplay between the higher temperatures facilitating rapid hydrogen abstractions and the higher concentrations of NOx and other RO2 terminators ending the autoxidation sequence more rapidly. Our data indicate that alkylbenzenes, monoterpenes, and isoprene are important precursor VOCs for HOMs in Beijing. Many of the C5 and C10 compounds derived from isoprene and monoterpenes have a slightly greater degree of average oxidation state of carbon compared to those from other precursors. Most HOMs except for large dimers have daytime peak concentrations, indicating the importance of OH chemistry in the formation of HOMs, as O3 tends to be lower on days with higher HOM concentrations; similarly, VOC concentrations are lower on the days with higher HOM concentrations. The daytime peaks of HOMs coincide with the growth of freshly formed new particles, and their initial formation coincides with the peak in sulfuric acid vapours, suggesting that the nucleation process is sulfuric acid-dependent, with HOMs contributing to subsequent particle growth.
1. INTRODUCTION

Atmospheric particle nucleation, or the formation of solid or liquid particles from vapour phase precursors is one of the dominant sources of global aerosol by number, with primary emissions typically dominating the mass loadings (Tomasi et al., 2016). New particle formation (NPF) or the secondary formation of fresh particles is a two-step process comprising of initial homogeneous nucleation of thermodynamically stable clusters and their subsequent growth. The rate of growth needs be fast enough to out-compete the loss of these particles by coagulation and condensation processes in order for the new particles to grow, and hence NPF is a function of the competition between source and sink (Gong et al., 2010). New particle formation has been shown to occur across a wide range of environments (Kulmala et al., 2005). The high particle load in urban environments was thought to suppress new particle formation until measurements in the early 2000s (McMurry et al., 2000; Shi et al., 2001; Alam et al., 2003), with frequent occurrences observed even in the most polluted urban centres. NPF events in Beijing occur on about 40% of days annually, with the highest rates in the spring (Wu et al., 2007, 2008; Wang et al., 2016). Chu et al. (2019) review many studies of NPF which have taken place in China and highlight the need for long-term observations and mechanistic studies.

NPF can lead to the production of cloud condensation nuclei (CCN) (Wiedensohler et al., 2009; Yu and Luo, 2009; Yue et al., 2011; Kerminen et al., 2012) which influences the radiative atmospheric forcing (Penner et al., 2011). A high particle count, such as that caused by nucleation events, has been shown to precede haze events in environments such as Beijing (Guo et al., 2014). These events are detrimental to health and quality of life. The sub-100 nm fraction of particles to which new particle formation contributes to is often referred to as the ultrafine fraction. Ultrafine particles (UFPs) pose risks to human health due to their high number concentration. UFPs exhibit gas-like behaviour and enter all parts of the lung before penetrating into the bloodstream (Miller et al., 2017). They can initiate inflammation via oxidative stress responses, progressing conditions such as
atherosclerosis and initiating cardiovascular responses such as hypertension through to myocardial infarction (Delfino et al., 2005; Brook et al., 2010).

Highly oxidised multifunctional molecules (HOMs), organic molecules with O:C ratios >0.6, are the result of atmospheric autooxidation and have recently been subject to much investigation, in part because the extremely low volatilities arising from their high O:C ratios favour their condensation into the particulate phase. HOMs are most well characterised as the product of oxidation of the biogenic monoterpenoid compound α-pinene (Riccobono et al., 2014; Tröstl et al., 2016; Bianchi et al., 2017). Although globally, biogenic volatile organic compound (BVOC) concentrations far exceed anthropogenic volatile organic compound (AVOC) concentrations, in the urban environment the anthropogenic fraction is far more significant. Formation of HOMs from aromatic compounds has been demonstrated in laboratory studies and these have been hypothesised to be large drivers of NPF in urban environments (Wang et al., 2017; Molteni et al., 2018; Qi et al., 2018). The formation of HOMs through autoxidation processes begins with the reaction of VOCs with OH, O3 or NO3; formation of a peroxy radical (RO2•) is followed by rapid O2 additions and intra-molecular hydrogen abstractions (Jokinen et al., 2014; Rissanen et al., 2014; Kurtén et al., 2015). Furthermore, generation of oligomers from stabilised Criegee intermediates arising from short chain alkenes has been hypothesised as a contributor of Extremely Low Volatility Organic Compounds (ELVOCs) and Low Volatility Organic Compounds (LVOCs) (Zhao et al., 2015). The low volatilities of these molecules arise from their numerous oxygen-containing functionalities, and this allows them to make a significant contribution to early stage particle growth where other species cannot due to the Kelvin effect (Tröstl et al., 2016), although the contribution of HOMs to the initial molecular clusters is still debated (Kurtén et al., 2016; Elm et al., 2017; Myllys et al., 2017).

Recent technological advances have facilitated insights into the very first steps of nucleation which were previously unseen, with mass spectrometric techniques such as the Atmospheric Pressure
Interface Time of Flight Mass Spectrometer (API-ToF) and its chemical ionisation counterpart (CI-API-ToF) allowing for high mass and time resolution measurements of low volatility compounds and molecular clusters. Diethylene glycol based particle counters, such as the Particle Size Magnifier (PSM) allow for measurements of particle size distributions down to the smallest molecular clusters nearing 1 nm. Recent chamber studies have elucidated the contribution of individual species to particle nucleation, ammonia and amines greatly enhancing the rate of sulfuric acid nucleation (Kirkby et al., 2011; Almeida et al., 2013). In these studies, HOMs have been identified, formed through autoxidation mechanisms (Schobesberger et al., 2013; Riccobono et al., 2014; Ehn et al., 2014). These are key to early particle growth (Tröstl et al., 2016) and can nucleate even in the absence of sulfuric acid in chambers (Kirkby et al., 2016) and in the free troposphere (Rose et al., 2018). In this paper, we report the results of HOM and particle size measurements during a summer campaign in Beijing, China.

2. DATA AND METHODS

2.1. Sampling Site

Sampling was performed as part of the Air Pollution and Human Health in a Developing Megacity (APHH-Beijing) campaign, a large international collaborative project examining emissions, processes and health effects of air pollution. For a comprehensive overview of the programme, see Shi et al. (2019). All sampling was conducted across a one month period at the Institute for Atmospheric Physics (IAP), Chinese Academy of Sciences, Beijing (39°58.53’N, 116°22.69’E). The sampling was conducted from a shipping container, with sampling inlets 1-2 metres above ground level, the nearest road being 30 metres away. Meteorological parameters (wind speed, wind direction, relative humidity (RH) and temperature) were measured at the IAP meteorological tower, 20 metres away from the sampling site, 30 metres from the nearest road at a height of 120 metres. Data was continuously taken from the CI-API-ToF during a two week period, but due to data losses
only five days of data is presented here. Particle size distribution measurements were taken during a
33 day period from 24/05/2017 – 26/06/2017.

2.2 Chemical Ionisation Atmospheric Pressure Interface Time of Flight Mass Spectrometry

The Aerodyne Nitrate Chemical Ionisation Atmospheric Pressure Interface Time of Flight Mass Spectrometer (CI-APi-ToF) was used to make measurements of neutral oxidised organic compounds, sulfuric acid and their molecular clusters at high time resolution with high resolving power. The ionization system charges molecules by adduct formation, such as in the case of organic compounds with two or more hydrogen bond donor groups (Hyttinen et al., 2015), or proton transfer in the case of strong acids like sulfuric acid. Hydroxyl or hydroperoxyl functionalities are both common hydrogen bond donating groups, with hydroperoxyl being the more efficient hydrogen bond donor (Møller et al., 2017). This instrument has been explained in great detail elsewhere (Junninen et al., 2010; Jokinen et al., 2012), but briefly the front end consists of a chemical ionisation system where a 10 LPM sample flow is drawn in through the 1 metre length 1” OD stainless steel tubing opening. A secondary flow was run parallel and concentric to this sample flow, rendering the reaction chamber effectively wall-less. A 3 SCCM flow of a carrier gas (N\textsubscript{2}) is passed over a reservoir of liquid HNO\textsubscript{3}, entraining vapour which is subsequently ionised to NO\textsubscript{3}– via an X-ray source. This flow is then guided into the sample flow. The nitrate ions will then charge molecules either by clustering or proton transfer. The mixed flows travelling at 10 LPM enter the critical orifice at the front end of the instrument at 0.8 LPM and are guided through a series of differentially pumped chambers before reaching the ToF analyser. Two of these chambers contain quadrupoles which can be used to select greater sensitivity for certain mass ranges, and the voltages across each individual chamber can be tuned to maximise sensitivity and resolution for ions of interest. Mass spectra are taken at a frequency of 20 kHz but are recorded at a rate of 1 Hz. All data analysis was carried out in the Tofware package in Igor Pro 6 (Tofwerk AG, Switzerland). A seven
point mass calibration was performed for every minute of data, and all data was normalised to
signal at 62, 80 and 125 m/Q to account for fluctuations in ion signal, these masses representing
NO$_3^-$, H$_2$ONO$_3^-$ and HNO$_3$NO$_3^-$ respectively. Typical values for calibration coefficients range from
10$^9$-10$^{10}$ molecules cm$^{-3}$ from these normalised data (Kürten et al., 2012), producing peak sulfuric
acid concentrations in the range of 10$^6$ molecules cm$^{-3}$. From the very limited periods with
simultaneous data for SO$_2$, OH radical and condensation sink, it was possible to calculate H$_2$SO$_4$
concentrations of 10$^3$ to 10$^5$ molec cm$^{-3}$, in which range the calibration constant was 7.0 ± 1.6 x 10$^8$
cm$^{-3}$ which fits well with that expected for this concentration range (Kürten et al., 2012). The
nitrate-water cluster is included as the presence of many nitrate-water clusters of the general
formula (H$_2$O)$_x$(HNO$_3$)$_y$NO$_3^-$ were found, where $x = (1, 2, 3... 20)$ and $y = (0, 1)$. No sensitivity
calibration was performed for these measurements, and so all values are reported in normalised
signal intensity. Due to the high resolving power of the CI-API-ToF system (mass resolution of
3500 m/dm and mass accuracy of 20 ppm at 288 m/Q), multiple peaks can be fit at the same unit
mass and their molecular formulae assigned. These peaks follow the general formula C$_x$H$_y$O$_z$N$_w$
where $x = 2$–20, $y = 2$–32, $z = 4$–16 and $w = 0$–2, spanning from small organic acids like oxalic and
malonic acid through to large dimers of oxidised monoterpene RO$_2$ radicals such as C$_{20}$H$_{31}$O$_9$N.
Beyond 500 m/Q, peak fitting and assignment of compositions becomes problematic as signal
decreases, mass accuracy decreases, and the total number of chemical compositions increases, so
peaks above the C$_{20}$ region have not been assigned, and a number of peaks have been unassigned
due to this uncertainty (Cubison and Jimenez, 2015). As proton transfer mostly happens with acids,
and nearly all HOM molecules will be charged by adduct formation it is possible to infer the
uncharged formula; therefore all HOMs from here onwards will be listed as their uncharged form.
2.3. Size Distribution Measurements

Two Scanning Mobility Particle Sizer (SMPS) instruments measured particle size distributions at 15 minute time resolution, one LongSMPS (TSI 3080 EC, 3082 Long DMA, 3775 CPC, TSI, USA) and one NanoSMPS (3082 EC, 3082 Nano DMA, 3776 CPC, TSI, USA) measuring the ranges 14-615 nm and 4-65 nm respectively. A Particle Size Magnifier (A10, Airmodus, FN) linked to a CPC (3775, TSI, USA) measured the sub-3 nm size fraction. The PSM was run in stepping mode, operating at four different saturator flows to vary the lowest size cut-off of particles that it will grow (this cut-off is technically a point of 50% detection efficiency) of <1.30, 1.36, 1.67 and 2.01 nm. The instrument switched between saturator flows per 2.5 minutes, giving a sub-2.01 nm size distribution every 10 minutes. The data was treated with a moving average filter to account for jumps in total particle count, and due to the similar behaviour of the two upper and two lower size cuts, these have been averaged to two size cuts at 1.30 and 1.84 nm.

2.4. Calculations

The condensation sink (CS) was calculated from the size distribution data as follows:

\[ CS = 4\pi D \sum \beta_{m,d_p} d_p' N_{d_p'} \]  

(1)

where D is the diffusion coefficient of the diffusing vapour (assumed sulfuric acid), \( \beta_m \) is a transition regime correction (Kulmala et al., 2012), \( d_p' \) is particle diameter, and \( N_{d_p'} \) is the number of particles at diameter \( d_p' \).

2.5. Other Measurements

Measurements of the classical air pollutants were measured on the same site, and have been reported in the campaign overview paper (Shi et al., 2019). \( \text{SO}_2 \) was measured using a 43i \( \text{SO}_2 \) analyser (ThermoFisher Scientific, USA), \( \text{O}_3 \) with a 49i \( \text{O}_3 \) analyser (ThermoFisher Scientific,
USA) and NO, with a 42i-TL Trace NO, analyser (ThermoFisher Scientific, USA), and a T500U CAPS NO₂ analyser (Teledyne API, USA). VOC mixing ratios were measured using a Proton Transfer Reaction-Time of Flight-Mass Spectrometer (PTR-ToF 2000, Ionicon, Austria).

3. RESULTS AND DISCUSSION

3.1. Characteristics of Sampling Period

A total of five days of CI-API-ToF data were collected successfully, from 2017/06/21 midday through 2017/06/26 midday. New particle formation events were observed on 24th June in the late afternoon and 25th June at midday. Some nighttime formation of molecular clusters was seen earlier in the campaign, as were several peaks to the 1.5 – 100 nm size range, likely from pollutant plumes containing freshly nucleating condensable materials. The trace gases, O₃, SO₂, NO and NO₂ are plotted in the Figure S1. O₃ shows mid-afternoon peaks, around ~120 ppb on the first two days of the campaign, and 50-70 ppb for the latter days. SO₂ shows a large peak, reaching 4 ppb on 22/06 but <1 ppb for the rest of campaign. NO shows strong mid-morning rush hour related peaks, declining towards midday due to being rapidly consumed by O₃. NO₂ shows large traffic related peaks. The sulfuric acid signal across this period as measured by NO₃ CI-API-ToF showed strong midday peaks, with signal highest on 24/06/2017 and 25/06/2017. The meteorological data are shown in Figure S2 alongside condensation sink (CS). The conditions were generally warm and humid, with temperature reaching its maximum on 25/06/2017, with a peak hourly temperature of 31°C. High temperatures were seen on 21/06 and 24/06 also, of 30°C and 26°C respectively.

3.2. Gas Phase HOM Chemistry

3.2.1. Bulk chemical properties

For the peaks that have had chemical formulae assigned, oxidation state of carbon, or OSₖ, can be used to describe their bulk oxidation chemistry. OSₖ is defined as (Kroll et al., 2011)
This does not account for the presence of nitrate ester groups, which has been accounted for previously by subtracting five times the N:C ratio (Massoli et al., 2018), under the assumption that all nitrogen containing functionality is in the form of nitrate ester (RONO₂) groups. In Beijing, multiple sources of nitrate-containing organic compounds are seen, in the forms of amines, nitriles and heterocycles. The variation of oxidation state with carbon number (Cₙ) without correction for nitrate esters is plotted in Figure 1. The average oxidation state of carbon in this dataset tends to decrease with an increase to Cₙ, highest where Cₙ = 5, attributable both to high O:C and peak area for the peak assigned to C₅H₁₀N₂O₈ at m/Q 288. Cₙ = 5 also shows the greatest distribution of oxidation states, likely due to the high ambient concentration of isoprene and therefore its many oxidation products being of high enough signal for many well resolved peaks to be seen in this dataset. It is worth noting that some of the ions plotted here may not form through peroxy radical autoxidation, such as C₅H₁₀N₂O₈, which may be a second-generation oxidation product of isoprene under high NOₓ (Lee et al., 2016). Cₙ = 10 and 15 also see a small increase to average oxidation number compared to their neighbours. The lower oxidation state of the larger products is likely a function of two things. First and foremost, any autoxidation mechanism must undergo more steps in order for a larger molecule to reach an equivalent O:C ratio with a smaller one, and the equivalent O:C ratio is ultimately less likely to be reached before the radical is terminated (Massoli et al., 2018). Secondly, the lower vapour pressures of these larger products will lead to their partitioning into the condensed phase more readily than the smaller, thus they are more rapidly lost (Mutzel et al., 2015).

The degrees of OSₙ observed here are similar to those seen in other environments such as during the SOAS campaign in 2013 in southern United States, characterised by low NO/NO₂ and high temperatures, where campaign averages of 0.3 ppb, 0.4-0.5 ppb, and 25°C respectively were
measured, although an additional parameter to account for nitrogen containing VOCs is included in the calculation (Massoli et al., 2018). The OS$_c$ observed in Beijing is also higher than that seen in the boreal forest environment of Hyytiälä, despite extremely low NO$_x$ concentrations, likely due to low temperature conditions dominating in those conditions (Schobesberger et al., 2013). These relatively similar degrees of oxidation to those seen in other, cleaner, environments are likely due to an interplay between the higher temperatures facilitating rapid hydrogen abstractions (Crounse et al., 2013; Quéléver et al., 2019) and the higher concentrations of NO$_x$, HO$_2^\cdot$, and other RO$_2^\cdot$ molecules terminating the autoxidation sequence more efficiently (Praske et al., 2018, Rissanen, 2018, Garmash et al., 2019).

A mass defect plot is shown in Figure 2, which shows nominal mass plotted against mass defect for all peaks in this dataset. Mass defect is defined as the ion mass minus integer mass. This is shown for two separate daytime periods, one where nucleation was not occurring and HOM concentrations are lower (10:30 – 12:00 23/06/2017) and one where nucleation was occurring under high HOM concentrations (10:30 – 12:00 25/06/2017). The band of lower mass defect is characterised by a number of large peaks with high signal, for example, at $m/Q$ 436 the ion (C$_2$H$_7$N)$_2$(H$_2$SO$_4$)$_2$HSO$_4^-$.

The upper component of the mass defect is dominated by organic compounds, the upper end of more positive mass defect is occupied by molecules with more $^1$H (mass defect 7.825 mDa) and $^{14}$N (mass defect 3.074 mDa). The end of less positive mass defect has lower $^1$H and more $^{16}$O (mass defect -5.085 mDa); alternatively put, the mass defect reflects the variation in OS$_c$. The organic components with more positive mass defects will be more volatile than their lower mass defect counterparts as they will contain fewer oxygen functionalities (Tröstl et al., 2016, Stolzenburg et al., 2018). These higher volatility products may still contribute to larger size particle growth. The more negative mass defect components will be those of greater O:C and therefore lower volatility, LVOCs, and the yet larger and more oxidised components, ELVOCs (Tröstl et al., 2016). During the nucleation period, the signal intensity for the species in the upper band of more negative mass
defect have the most marked increase in concentration, with significantly less difference >500 m/Q. This region 200-400 m/Q will contain most of the >C$_5$ monomer HOMs seen in this dataset.

### 3.2.2. Diurnal trends of HOMs

Temporal trends of HOMs in the urban atmosphere can reveal their sources and behaviour in the atmosphere. Most of the HOM species peak in the daytime. These species all follow a similar diurnal trend, as shown in Figure 3. Both the concentrations of O$_3$ and OH are high during the summer period in Beijing (although the nitrate chemical ionisation technique is not sensitive to all OH oxidation products (Berndt et al., 2015)). Figure S1 shows the time series of concentrations of NO which is considered a dominant peroxy radical terminator of particular importance in the polluted urban environment (Khan et al., 2015). Radicals such as HO$_2$ and RO$_2$ also typically peak during daytime. The HOM components peaking in the daytime are presumed to be the oxidation products of a mixture of anthropogenic and biogenic components, such as alkylbenzenes, monoterpenes and isoprene. The oxidation of monoterpenes, specifically the monoterpene $\alpha$-pinene, has been the subject of extensive study recently, with the O$_3$-initiated autoxidation sequence being the best characterised (Ehn et al., 2014; Jokinen et al., 2014; Kurtén et al., 2015; Kirkby et al., 2016); ozonolysis of $\alpha$-pinene opens the ring structure and produces a RO$_2$ radical (Kirkby et al., 2016). In the case of aromatics, OH addition to the ring and the subsequently formed bicyclic peroxy radical is the basis for the autooxidation of compounds such as xylenes and trimethylbenzenes (Molteni et al., 2018; Wu et al., 2017).

The identified compounds have been roughly separated into several categories, each of these plotted in Figure 3. Figure 3a shows the separation of components into non-nitrogen containing HOMs, and nitrogen containing HOMs, or organonitrates (ONs). The ON signal is much higher than that of the HOM, attributable in part to a few ions of high signal, such as the isoprene organonitrate C$_5$H$_{10}$N$_2$O$_8$. A few similar structural formulae are seen (C$_5$H$_{10}$N$_2$O$_6$, C$_5$H$_{11}$NO$_6$, C$_5$H$_{11}$NO$_7$, etc),
some of which have been identified as important gas phase oxidation products of isoprene under high NOx conditions (Xiong et al., 2015), and their contribution to SOA has been explored previously (Lee et al., 2016). A high nitrophenol signal is also seen, C6H5NO3. The signal for HOM compounds is less dominated by a few large ions. The prevalence of ON compounds points towards the important role of NOx as a peroxy radical terminator, with the probability for the RO2 + NOx reaction to produce nitrate ester compounds increasing with the size of the RO2 molecule (Atkinson et al., 1982). The NOx concentrations in urban Beijing are approximately a factor of 10 higher than seen at the Hyytiälä station in Finland as reported by Yan et al. (2016), and hence it is expected to be a more significant peroxy radical terminator.

Despite the very large fluxes of anthropogenic organic pollutants in Beijing, biogenic emissions are still an important source of reactive VOCs in the city, with abundant isoprene oxidation products observed (see above), as well as monoterpane monomers (C10H16O9, C10H15O9N) and some dimer products (C20H30O11, C20H31O11N). The time series of the signals of all C5, C10 and C20 molecules is plotted in Figure 3b, with C5 species assumed to be isoprene dominated, C10 and C20 assumed to be monoterpene dominated. Signals for isoprene oxidation products are higher, with abundant isoprene nitrate and dinitrate products. C10 products show similar behaviour, with, for example, several C10H15OxN x = 5−9 compounds seen. The C20 signal intensities are low, and follow the general formula C20HxOyNz, where x = 26−32, y = 7−11 and z = 0−2; in Figure 3 the signal for C20 compounds has been multiplied by a factor of 50 for visibility. The low signals reflect the lack of RO2 cross reactions necessary for the production of these accretion products.

Other identified peaks are plotted in Figure 3c. The C2−C4 components are summed together, these being small organic acids such as malonic acid and oxalic acid, as well as products such as C4H7O6N. Malonic acid is the most prominent here, seen both as an NO3− adduct (C3H4O4NO3−) and a proton transfer product (C3H3O4+), at a ratio of around 2:3. Measurements of particle phase
dicarboxylic acids in cities typically show greater concentrations of oxalic acid than malonic (Ho et al., 2016), and these acids are primarily produced in the aqueous phase (Bikkina et al., 2014). Primary sources of dicarboxylic acid include fossil fuel combustion (Kawamura and Kaplan, 1987) and biomass burning (Narukawa et al., 1999), which are both plentiful in urban Beijing. The C₆-C₉ components are assumed to be dominated by oxidation products of alkylbenzenes such as C₈H₁₂O₅, although fragments of other compounds, i.e., monoterpenes, can also occupy this region (Isaacman-Vanwertz et al., 2018). It is assumed the majority of the signal for these peaks come from alkylbenzenes. This assumption is supported by the relative signal intensity ratios of the oxygen numbers of monomer C₈H₁₂O₈ compounds being similar to those seen for xylene oxidation products in previous work (Molteni et al., 2018). The largest fraction, C₁₁ through C₁₈, includes the larger compounds, oxidation products of larger aromatics, or products of the cross reaction of smaller RO₂ radicals. Here they are grouped without more sophisticated disaggregation as they all follow much the same time series, species such as C₁₁H₁₁O₈N following the same temporal trends as C₁₅H₁₆O₉ and C₁₆H₂₄O₁₂.

Nearly all ions with the exception of the larger compounds attributed to the cross reaction of C₁₀ monomers follow similar temporal patterns, with the majority of peaks occurring in the daytime. This reflects the importance of the concentration of atmospheric oxidants. Some selected oxidation products are plotted against their precursor VOCs in Figure 4. The concentration of isoprene is plotted against the signal of a nitrate HOM product, C₅H₉NO₆ (Xiong et al., 2015; Lee et al., 2016), while monoterpenes are plotted against C₁₀H₁₆O₉ (Ehn et al., 2014; Berndt et al., 2016; Yan et al., 2016; Kirkby et al., 2016; Massoli et al., 2018), and C₂-benzenes against C₆H₁₂O₆ (Molteni et al., 2018; Wang et al., 2017). The first half of the time series shows little correlation between the VOC species and the resultant oxidation products, while isoprene, monoterpenes and C₂-benzenes follow their usual diurnal cycles, isoprene having the most distinct with a strong midday peak. The latter two days, however, show similar and coinciding peaks in both the VOCs and HOMs - HOMs show
afternoon peaks on both days, and an initial shelf on the final half day. The C$_5$H$_9$NO$_6$ peak follows some of the peaks of the isoprene, but not all (e.g., morning shelf of isoprene on 24/06).

Concentrations of isoprene do not seem to determine directly the signal of HOM, as the day with the lowest isoprene of all is the day with highest C$_5$H$_9$NO$_6$. The C$_{10}$H$_{16}$O$_9$ trace has coincidental peaks with the monoterpene trace also, including two 4-hour separated simultaneous peaks on 25/06. The peaks in the concentrations of C$_2$-benzenes are nearly synchronous with the peaks in C$_8$H$_{12}$O$_6$, for which the data exhibit a strong mid afternoon peak likely due to the lack of an efficient ozonolysis reaction pathway; the main oxidant of C$_2$-benzenes is the OH radical. Trends of both C$_3$ benzenes and their HOMs are much the same as C$_2$ benzenes as discussed above, pointing to similar sources and oxidation chemistries. The concentration of precursor VOC is likely a driving force in the identity and quantity of various HOM products, but not the sole determinant, as while there are simultaneous peaks of VOCs and HOMs, both the condensation sink and oxidant concentrations also influence HOM product signals.

The first half of campaign measurements is marked by an episode of low HOM signals. A diurnal cycle still exists but it is weak. The radiation intensity was significantly lower on these prior days than it was on the 24th. No data is available for the final period of measurement. Ozone is higher on the prior measurement days with lower HOM signals (see Figure S1). Little agreement is seen between VOC concentration and HOM signals on these days. The condensational sinks are roughly similar to those on days of higher HOM concentrations, but temperature and solar radiation are much lower. HOM formation is largely dependent upon VOC concentration, oxidant concentration (which will be lower if solar radiation is lower, especially in the case of OH, the main oxidant of aromatic species especially), and temperature (as H-shifts are highly temperature dependent) (Quéléver et al., 2019), as well as losses by RO$_2$ termination before a molecule can become HOM, and losses to condensational sink. The low HOM concentration is likely due to these lower temperatures, and weaker solar radiation not facilitating HOM formation.
The C20 compounds plotted in Figure 3b show no strong diurnal sequence, contrasting with other HOMs. We can presume that all C20 compounds identified are the result of the reaction of two monoterpenoid C10RO2 radicals, a reasonable assumption as all identified C20 species follow the general formula outlined for these reactions (C20H28−32O6−16). The formation of C20 dimers is dependent upon two processes, initial oxidation of monoterpenes, and RO2−RO2 termination. Initial oxidation is contingent upon oxidant concentration, which is highest in the daytime, and RO2−RO2 termination is contingent upon the probability of the molecular collision between the RO2− molecules occurring before other radical termination (i.e., RO2−NOx, or RO2−HO2). There is likely a strong diurnal sequence in the dominant RO2− termination mechanisms across the day period, and the combination of the two factors discussed above results in there being no strong diurnal trend in these molecules. A lower oxidant concentration at night results in less RO2− molecules, but less NO and HO2− results in a greater chance for those RO2− molecules to dimerise (Rissanen, 2018, Garmash et al., 2019). As the levels of NOx in Beijing fall, the peroxy radical termination reactions will be less probable compared to continued autoxidation (Praske et al., 2018), and it is expected that more oxidised HOM products will be seen with lower volatilities and therefore a greater potential contribution to earlier stage particle formation and growth.

3.3. New Particle Formation

Nearly all the signal intensity in the CI-API-ToF instrument arises from molecules charged by NO3−, therefore plotting the unit mass resolution data (the data gained by integrating over the entire area at each m/Q integer) against time describes simply the evolution of oxidised organic molecules, acids and their molecular clusters both with each other and stabilising amine species. This is done in Figure 5. As the signal intensity varies by factors of 10 from mass to mass, each value has been normalised so they have maxima at 1. This has been done separately for two days for clarity, as the signal intensity also varies from day to day. PSM data for these two days is plotted in Figure 5 also, with both total particle count >1.30 nm in black and the number difference between the lower and
upper size cuts (1.30 and 1.84 nm) in blue, which shows the number of particles between these sizes. The relationship between mass and electrical mobility diameter can be defined thus (Tammet, 1995),

\[ d_e = \left( \frac{6m}{\pi \rho} \right)^{\frac{1}{3}} + d_g \]  

where \( d_e \) is the electrical mobility diameter of the cluster or particle, \( m \) is the mass of the cluster or particle expressed in kg, \( \rho \) is the density and \( d_g \) is the effective gas diameter, determined to be 0.3 nm for smaller particles (Larriba et al., 2011). We can use this to draw a comparison between the PSM and CI-API-ToF measurements. If a density of 1.2 g cm\(^{-3}\) is assumed, then once molecular clusters reach the >400 \( m/Q \) range, they will be seen in the lowest size cut of the PSM, or >700 \( m/Q \) if a density of 2.0 g cm\(^{-3}\) is assumed. A full table of densities is provided in the Supplementary Information.

A burst in the signal seen by the CI-API-TOF occurs first in the late morning in the top panel of Figure 5, and this is at the same time as peaks begin to rise in the identified HOMs (see Figure 3). Here, the PSM is not available due to an instrumental fault until 16:00; however, at that point, an elevation to particle count and a large elevation to cluster count can be seen. Moving into the evening period, the mass contour shows peaks to larger masses >400 \( m/Q \). This is likely dimerised compounds and products of NO\(_3\) chemistry with little contribution to newly forming particles, but still sensitive to chemical ionisation by NO\(_3^-\). Many of these peaks cannot be assigned due to uncertainties in the structural formula assignment for higher mass peaks, as the number of possible dimerised compounds is many, being the combination of most possible RO\(_2\) radicals. Graphically, these are over-represented in Figure 5 due to the normalisation, their signals (especially >500 \( m/Q \)) are much lower than the signals <400 \( m/Q \).
The second day plotted in the lower panel of Figure 5 (25/06/2017) shows a strong afternoon peak to the HOMs (for most HOMs, stronger than that on the day prior). Particle formation is shown in the PSM data. A strong midday peak to particle number is seen with two distinct peaks to cluster count. These two peaks are not coincidental with the two peaks to HOM signal (i.e., nitrogen-containing HOMs in Figure 3a peaking at 11:00 and 16:00). Sulfuric acid, however, does peak synchronously with the particle number count. Sulfuric acid is plotted across the contour plot in Figure 6, where PSM data is also shown in the bottom panel. The peak to CI-APi-TOF mass signal, visible in Figure 5 occurs at around 12:00/13:00, peaks in the PSM cluster count occur at 10:00 and 13:00. Peaks in mass up to 550 m/Q are seen in the CI-APi-ToF at 13:00. Assuming the density of these species is ≤1.6 g cm\(^{-3}\) then these will be suitably sized to be grown in the PSM saturator.

These newly formed particles then go on to grow and contribute significantly to the larger particle count (Figure S3). As initial particle formation coincides with sulfuric acid signal peaks and before HOM signals peak, it can be assumed on these days, the HOM contribution to the initial particle formation is modest.

There is recent strong evidence to suggest that the driving force of the earliest stages of particle formation in urban Shanghai is from sulfuric acid and C\(_2\)-amines (Yao et al., 2018), and the coincidental peaks of sulfuric acid with new particles as seen in Figure 6 suggest a similar behaviour. Dimethylamine (DMA) can efficiently stabilise the sulfuric acid clusters (Almeida et al., 2013). Here, few larger sulfuric acid-DMA clusters were visible in the dataset, as seen in the work by Yao et al., 2018, although five sulfuric acid-dimethylamine (SA-DMA) ions were observed, the others were likely too low in signal to be confidently resolved from their neighbouring peaks; however, clusters of up to 4 sulfuric acid molecules and 3 dimethylamine molecules were seen, with similar diurnal trends to sulfuric acid. The scarcity of SA-DMA clusters is likely due to instrumental conditions, rather than their absence in the atmosphere. The nitrate chemical ionisation system tends to evaporate amine compounds upon charging, and as specific voltage-tuning setups
can lend themselves towards preservation or breakage of molecular clusters, the signal for larger
sulfuric acid clusters was also very weak. The formation of HOM-sulfuric acid clusters is unlikely
under atmospheric conditions (Elm et al., 2017) and few of these were observed. Signals of HOMs
seem to coincide with later particle growth; it can be expected that HOM molecules make a more
significant contribution to particle growth than to early particle formation, with the largest and most
oxidised being involved in early growth, and the smaller and less oxidised contributing to later
growth as the necessary vapour pressure properties become less demanding.

4. CONCLUSIONS

The average degree of HOM oxidation in Beijing is comparable with that seen in other
environments. Rapid intramolecular hydrogen shifts during autoxidation due to the higher
temperatures are probably offset by the frequent termination reactions due to high NOx
concentrations. $O_{5c}$ values seem to be marginally higher for biogenic species.

The temporal trend of nearly every HOM shows afternoon or evening maxima. Both $O_3$ and OH
have high daytime concentrations and these likely drive the initial oxidation steps. The species
arising from alkylbenzene precursors show sharper afternoon peaks, probably since their oxidation
is OH dominated. Many of the rest of the peaks, coming from largely BVOC precursors show
broader daytime peaks, being influenced by $O_3$ also. There seems to be no direct link between VOC
concentrations and HOM signals, with days of lower precursor VOC sometimes having higher
HOM signals and vice versa.

Initial particle formation coincides with peak sulfuric acid signals, while the growth of the particles
correlates more closely with the signals of HOMs. This is very similar to behaviour observed in a
study of NPF in Shanghai which was attributed to sulfuric acid-dimethylamine-water nucleation
with condensing organic species contributing to particle growth (Yao et al., 2018), and this is
further backed up by numerous SA-DMA clusters present in this dataset. The freshly formed particles grow and contribute significantly to total particle loading. This is visible when the unit mass CI-API-ToF data is plotted as a contour plot, and further to this is visible in the PSM data, with bursts to both total number count >1.30 nm and the number of molecular clusters between 1.30 and 1.84 nm. As NO$_3$ levels fall in Beijing due to traffic emission control measures being enforced it is likely that autoxidation will become increasingly significant in the new particle formation processes. The number of molecules detected by the NO$_3$ CIMS is undoubtedly many more than have had formulae assigned here, but to identify more requires a more sophisticated data deconvolution.

DATA ACCESSIBILITY

Data supporting this publication are openly available from the UBIRA eData repository at https://doi.org/10.25500/edata.bham.00000304

AUTHOR CONTRIBUTIONS

The study was conceived and planned by RMH and ZS. DCSB and JB set up and operated the main instrumental measurements, and JB prepared the first draft of the paper and responded to comments from RMH and ZS. CNH and WJA contributed the hydrocarbon data and provided comments on the draft manuscript, and ES and JL contributed the gas phase pollutant data.

COMPETING INTERESTS

The authors have no conflict of interests.

ACKNOWLEDGMENTS

This work was part of the APHH-Beijing programme funded by the UK Natural Environmentl Research Council (NE/N007190/1) and the Natural Sciences Funding Council of China. It was
additionally facilitated by the National Centre for Atmospheric Science ODA national capability programme ACREW (NE/R000034/1), which is supported by NERC and the GCRF. We thank Professor X.M Wang from the Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Brian Davison from Lancaster University and Ben Langford, Eiko Nemitz, Neil Mullinger and other staff from the Centre for Ecology and Hydrology, Edinburgh for assistance with the VOC measurements and associated infrastructure.
REFERENCES


Measurements of Highly Oxidized Gas-Phase Molecules during the Southern Oxidant and Aerosol

10 nm atmospheric particles: implications for nucleation mechanisms, Philosophical Transactions

Vesey, A. T., Fokkens, P. H., Boere, A. J. F., Krystek, P., Campbell, C. J., Hadoke, P. W.,


Molteni, U., Bianchi, F., Klein, F., Haddad, I. E., Frege, C., Rossi, M. J., Dommen, J., and
Baltensperger, U.: Formation of highly oxygenated organic molecules from aromatic compounds,

Mutzel, A., Poulain, L., Berndt, T., Inumma, Y., Rodigast, M., Böge, O., Richters, S., Spindler, G.,
Sipila, M., Jokinen, T., Kulmala, M., Herrmann, H.: Highly oxidized multifunctional organic
compounds observed in tropospheric particles: A field and laboratory study, Environ. Sci. Techn.,

Myllys, N., Olenius, T., Kurtén, T., Vehkamäki, H., Riipinen, I., and Elm, J.: Effect of Bisulfate,

Narukawa, M., Kawamura, K., Takeuchi, N. and Nakajima, T.: Distribution of dicarboxylic acids
and carbon isotopic ratios in aerosols from 1997 Indonesian forest fires, Geophys. Res. Lett.,

Penner, J. E., Xu, L., and Wang, M.: Satellite methods underestimate indirect climate forcing by

Praske, E., Otkjær, R. V., Crounse, J. D., Hethcox, J. C., Stoltz, B. M., Kjaergaard, H. G., and
Wennberg, P. O.: Atmospheric autoxidation is increasingly important in urban and suburban North
America, PNAS, 115, 64-69, 2018.

contributions to new particle formation and growth: comparison of boreal forest in Finland and a

Quéléver, L. L. J., Kristensen, K., Normann Jensen, L., Rosati, B., Teiwes, R., Daellenbach, K. R.,
Peräkylä, O., Roldin, P., Bossi, R., Pedersen, H. B., Glasius, M., Bilde, M., and Ehn, M.: Effect of
temperature on the formation of highly oxygenated organic molecules (HOMs) from alpha-pinene

Riccobono, F., Schobesberger, S., Scott, C., Dommen, J., Ortega, I., Rondo, L., Almeida, J.,
Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E., Duplissy, J.,
Ehrhart, S., Flagan, R., Francin, A., Hansel, A., Junninen, H., Kajos, M., Keskinen, H., Kupc, A.,
Kürten, A., Kvashin, A., Laaksonen, A., Lehtipalo, K., Makhmuto, V., Mathot, S., Nieminen, T.,


Lawler, M., Lehtipalo, K., Leiminger, M., Mai, H., Mathot, S., Mentler, B., Molteni, U., Nie, W.,
Nieminen, T., Nowak, J., Ojdanic, A., Onnela, A., Passananti, M., Petäjä, T., Quéléver, L.,
Rissanen, M., Sarnela, N., Schallhart, S., Tauber, C., Tomé, A., Wagner, R., Wang, M., Weitz, L.,
Wimmer, D., Xiao, M., Yan, C., Ye, P., Zha, Q., Baltensperger, U., Curtius, J., Dommen, J., Flagan,
R., Kulmala, M., Smith, J., Worsnop, D., Hansel, A., Donahue, N., Winkler, P., Rapid growth of
organic aerosol nanoparticles over a wide tropospheric temperature range, PNAS, 115, 9122-9127,
2018.

Tammet, H.: Size and mobility of nanometer particles, clusters and ions, J. Aerosol Sci., 26, 459-

Tomasi, C., Fuzzi, S., Kokhanovsky, A.: Atmospheric Aerosols: Life Cycles and Effects on Air
Quality and Climate, Wiley, 2016.

Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C.,
Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J.,
Adamov, A., Almeida, J., Bernhammer, A. K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S.,
Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T.,
Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Kräpfl, M., Kürten, A., Laaksonen, A.,
Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F.,
Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schoebesberger, S., Sengupta, K., Sipilä,
M., Smith, J. N., Steiner, G., Tomè, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D.,
Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen,
J., Worsnop, D. R., Donahue, N. M., and Baltensperger, U.: The role of low-volatility organic
compounds in initial particle growth in the atmosphere, Nature, 533, 527-531,
https://doi.org/10.1038/nature18271, 2016.

Wang, S., Wu, R., Berndt, T., Ehn, M., and Wang, L.: Formation of Highly Oxidized Radicals and
Multifunctional Products from the Atmospheric Oxidation of Alkylbenzenes, Environ. Sci. Technol.,
51, 8442-8449, 2017.

Wang, Z., Wu, Z., Yue, D., Shang, D., Guo, S., Sun, J., Ding, A., Wang, L., Jiang, J., Guo, H., Gao,

Wiedensohler, A., Cheng, Y. F., Nowak, A., Wehner, B., Achtelt, P., Berghof, M., Birmili, W., Wu,
increase of cloud condensation nucleus activity by secondary aerosol formation and condensation:
A case study for regionalair pollution in northeastern China, Journal of Geophysical Research

distribution in the urban atmosphere of Beijing, China, Atmos. Environ., 42, 7967-7980, 2008.

Wiedensohler, A., and Zhang, Y.: Thermodynamic properties of nanoparticles during new particle


**FIGURE LEGENDS:**

**Figure 1** Oxidation state of carbon calculated as two times the oxygen to carbon ratio minus the hydrogen to carbon ratio against carbon number for (colored) individual ions and (blue circles) signal weighted average for each carbon number. Area and colour are both proportional to the peak area for each ion.

**Figure 2** Mass defect plot of fitted mass spectral peaks between 100-600 mass units on (a) 10:30 – 12:00 23/06/2017, a non nucleation day, and (b) 10:30 - 12:00 25/06/2017, a nucleation day. Mass defect can be defined as the mass - integer mass. The size of point is proportional to the signal intensity. As $^1$H has a positive mass defect (1.007276 Da), the upward trend along the horizontal indicates increasing carbon chain length, and differences at similar masses are due to increasing oxygen functionality, clustering with species such as sulfuric acid (negative mass defect) and ammonia (positive mass defect), as $^{16}$O and $^{32}$S have negative mass defects (15.9949 and 31.9721 Da respectively), while $^{14}$N has a positive mass defect at 14.0031 Da. The two large peaks seen at 201 and 288 m/Q are the nitrophenol-nitrate cluster and a C5H10N2O8-nitrate cluster respectively.

**Figure 3** Summed time series of the normalised signals of (A) all non-nitrogen containing HOMs and all organonitrates identified, (B) C5, C10 and C20 components, assumed to be dominated by isoprene, monoterpenes monomer and monoterpenes dimers, signal for C20 multiplied 50 times to fit scale, and (C) summed C6 - C9 components, and summed C11 - C18 components, assumed to be dominated by alkylbenzenes and other larger components respectively.

**Figure 4** Time series for the whole sampling campaign for the concentrations of (left axis) VOCs as measured by PTR-ToF and (right axis) a selected HOM product associated with that precursor.

**Figure 5** Normalised unit mass NO3- CI-APi-ToF signal intensity on 24/06/2017 (A) and 25/06/2017 (B). Each individual unit mass was normalised to a maximum of 1. Each period is normalised separately so the individual signal maxima on each day are visible. The graph is plotted between 200-600 mass units, with every 10 mass units averaged for simplicity. On the secondary axis is plotted PSM data, both total particle count $> 1.30$ nm (black trace) and total clusters between 1.30 and 1.84 nm (blue trace). Data is plotted at 1 hour time resolution.

**Figure 6** SMPS + PSM contour plot for two nucleation days on 24/06/2017 and 25/06/2017. Data in bottom panel is from the PSM instrument, top panel from NanoSMPS, units in colour bar are $\log_{10}$ (dN/logD$_p$) for N in cm$^{-3}$. Points signify normalised sulfuric acid concentration (right axis) as measured by CI-APi-ToF.
Figure 1. Oxidation state of carbon plotted against carbon number for (colored) individual ions and (purple triangles) signal weighted average for each carbon number. Area and colour are both proportional to the peak area for each ion.
Figure 2. Mass defect plot of fitted mass spectral peaks between 100-600 mass units on (a) 10:30 – 12:00 23/06/2017, a non nucleation day, and (b) 10:30 -12:00 25/06/2017, a nucleation day. Mass defect can be defined as the mass - integer mass. The size of point is proportional to the signal intensity. As $^1$H has a positive mass defect (1.007276 Da), the upward trend along the horizontal indicates increasing carbon chain length, and differences at similar masses are due to increasing oxygen functionality, clustering with species such as sulfuric acid (negative mass defect) and ammonia (positive mass defect), as $^{16}$O and $^{32}$S have negative mass defects (15.9949 and 31.9721 Da respectively), while $^{14}$N has a positive mass defect at 14.0031 Da. The two large peaks seen at 201 and 288 m/Q are the nitrophenol-nitrate cluster and a C$_5$H$_{10}$N$_2$O$_8$-nitrate cluster respectively.
Figure 3. Summed time series of the normalised signals of (A) all non-nitrogen containing HOMs and all organonitrates identified, (B) C5, C10 and C20 components, assumed to be dominated by isoprene, monoterpane monomer and monoterpane dimers, signal for C20 multiplied 50 times to fit scale, and (C) summed C6 - C9 components, and summed C11 - C18 components, assumed to be dominated by alkylbenzenes and other larger components respectively.
**Figure 4.** Time series for the whole sampling campaign for the concentrations of (left axis) VOCs as measured by PTR-ToF and (right axis) a selected HOM product associated with that precursor.
**Figure 5.** Normalised unit mass NO$_3^-$ Cl-APi-ToF signal intensity on 24/06/2017 (A) and 25/06/2017 (B). Each individual unit mass was normalised to a maximum of 1. Each period is normalised separately so the individual signal maxima on each day are visible. The graph is plotted between 200-600 mass units, with every 10 mass units averaged for simplicity. On the secondary axis is plotted PSM data, both total particle count $>$1.30 nm (black trace) and total clusters between 1.30 and 1.84 nm (blue trace). Data is plotted at 1 hour time resolution.
**Figure 6.** SMPS + PSM contour plot for two nucleation days on 24/06/2017 and 25/06/2017. Data in bottom panel is from the PSM instrument, top panel from NanoSMPS, units in colour bar are $\log_{10}(dN/d\log D_p)$ for N in cm$^{-3}$. Points signify normalised sulfuric acid concentration (right axis) as measured by CI-API-ToF.