The size resolved cloud condensation nuclei (CCN) activity and its prediction based on aerosol hygroscopicity and composition in the Pearl Delta River (PRD) Region during wintertime 2014

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Abstract. A hygroscopicity-tandem differential mobility analyzer (H-TDMA), a scanning mobility CCN analyzer (SMCA), and an aerodyne high resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS) were used to respectively measure the hygroscopicity, condensation nuclei activation, and chemical composition of aerosol particles at the Panyu site in the Pearl River Region during wintertime 2014. The distribution of the size-resolved cloud condensation nuclei (CCN) at four supersaturations (SS=0.1%, 0.2%, 0.4%, and 0.7%) and the aerosol particle size distribution were obtained by the SMCA. The hygroscopicity parameter κ (κ_{CCN}, κ_{H-TDMA}, and κ_{AMS}) was respectively
calculated based upon the SMCA, H-TMDA, and AMS measurements. The results showed that the 
\( \kappa_{\text{H-TDMA}} \) value was slightly smaller than the \( \kappa_{\text{CCN}} \) one at all diameters and for particles larger than 100 
nm and the \( \kappa_{\text{AMS}} \) value was significantly smaller than the others (\( \kappa_{\text{CCN}} \), and \( \kappa_{\text{H-TDMA}} \)), which could be 
attributed to the underestimated hygroscopicity of the organics (\( \kappa_{\text{org}} \)). The activation ratio (AR) 
calculated from the growth factor – probability density function (GF-PDF) without surface tension 
correction was found to be lower than that from the CCN measurements, due most likely to the 
uncorrected surface tension (\( \sigma_{\text{s/a}} \)) that did not consider the surfactant effects of the organic compounds.

We demonstrated that better agreement between the calculated and measured AR could be obtained by 
adjusting \( \sigma_{\text{s/a}} \). Various schemes were proposed to predict the CCN number concentration (\( N_{\text{CCN}} \)) based 
on the H-TDMA and AMS measurements. In general, the predicted \( N_{\text{CCN}} \) agreed reasonably well with 
the corresponding measured ones using different schemes. For the H-TDMA measurements, the \( N_{\text{CCN}} \) 
value predicted from the real time AR measurements was slightly smaller (~6.8%) than that from the 
activation diameter (\( D_{\text{a}} \)) method due to the assumed internal mixing in the \( D_{\text{a}} \) prediction. The \( N_{\text{CCN}} \) 
values predicted from bulk chemical composition of PM\(_1\) were higher (~11.5%) than those from 
size-resolved composition measured by the AMS because a significant fraction of PM\(_1\) was composed 
of inorganic matter. The \( N_{\text{CCN}} \) values calculated from AMS measurement were under-predicted at 0.1% 
and 0.2% supersaturations, which could be due to underestimate of \( \kappa_{\text{org}} \) and overestimate of \( \sigma_{\text{s/a}} \). For 
SS=0.4% and 0.7%, slight over-predicted \( N_{\text{CCN}} \) values were found because of the internal mixing 
assumption. Our results highlight the need for accurately evaluating the effects of organics on both the 
hygroscopic parameter \( \kappa \) and the surface tension \( \sigma \) in order to accurately predict CCN activity.
Introduction

Aerosol particles can directly impact global climate by scattering and absorbing solar radiation (Stocker, 2013), while they can influence cloud formation, life time and optical properties by acting as cloud condensation nuclei (CCN), indirectly exerting climatic forcing on the Earth’s atmosphere. In general, aerosol particles increase the CCN concentration and hence cause cooling effects on the global radiation balance. However, to what extent aerosol particles contribute to the radiation forcing is still highly uncertain (Stocker, 2013). It is hence important to measure chemical composition and properties of aerosol particles in order to assess their abilities of acting as CCN and contribution to cloud formation, further facilitating our understanding of the impacts of atmospheric aerosols on regional and global climate.

The extent to which aerosol particles can affect cloud formation is dependent on their fraction that can be activated to become CCN. This fraction of activation is termed as CCN activity that is determined by the chemical composition, sizes, and the water saturation ratio of the particles (Farmer et al., 2015).

The size-dependent saturation ratio \( S \) can be calculated from the Köhler equation (Köhler, 1936):

\[
S = a_w \exp \left( \frac{4 \sigma_s/a M_w}{RT \rho_w D} \right)
\]  

where \( a_w \) is the water activity in solution, \( \sigma_s/a \) is surface tension of the solution/air interface, \( M_w \) is the mole weight of water, \( R \) is the universal gas constant, \( T \) is temperature in Kelvin, and \( D \) is the diameter of the droplet. The \( a_w \) represents Raoult effect, which means that the activation potential increases with the concentration of the solution. The term \( \exp \left( \frac{4 \sigma_s/a M_w}{RT \rho_w D} \right) \) represents Kelvin effect, which relates the surface curvature to the saturation vapor pressure of the droplet. The activation
potential increases with increase of the droplet diameter or decrease of surface tension $\sigma_{s/a}$ and the
$\sigma_{s/a}$ value is sensitive to the organic surfactant effect. The two important parameters, the water activity
($a_w$) and surface tension ($\sigma_{sw}$), are dependent on the composition of the aerosol particles, assuming
those particles have the same properties as their corresponding bulk solutions. The effects of organics
on the CCN activity have been extensively investigated; however, many outstanding questions still
remain. Sorjamaa et al. (2004) suggested that the partitioning of surfactants had to be considered when
evaluating the Kelvin effect and the Raoult effect. According to their experimental results, the
surfactant partitioning could alter the Raoult effect and that the change is large enough to depress CCN
activity. However, another experiment conducted by Engelhart et al. (2008) revealed that the organics
in aged monoterpane aerosols could depress surface tension by about 0.01 N m$^{-1}$ and hence increase
CCN activity. Ovadnevaite et al. (2017) also presented observational and theoretical evidences that the
decrease of surface tension could prevail over the Raoult effect, which led to the increase of CCN
activity. Salma et al. (2006) isolated humic-like substances (HULIS) from PM2.5 fraction aerosol
samples and investigated the surface tension properties of the HULIS pure solutions. The results show
that thermodynamic equilibrium on surface could only be reached after several hours. Because the
depression of surface tension was controlled by diffusion of surfactants from the bulk of the droplet to
its surface, the extent of the actual decrease of surface tension was hence kinetically limited. A hybrid
model proposed by Petters and Kreidenweis (2013) was used to predict the effects of surfactants on the
CCN activity. The model predicted strong effects of the surfactants on ternary systems where common
ions were present. However, due to the limited measurement techniques, the available laboratory data
were still not sufficient to support this prediction and more solid data were needed to validate the
surfactant effects on the CCN activity.

The CCN activity can be characterized by the hygroscopicity parameter $\kappa$ that was initially proposed by Petters and Kreidenweis (2007). Aerosol hygroscopicity represents the ability of the particles to grow by absorbing water vapor from the atmosphere and the extent to which the particles are hygroscopic can be evaluated by the $\kappa$ values, which can be determined from the H-TDMA or CCNe measurements.

The $\kappa$ values were measured worldwide extensively either in the field measurements or in the laboratory experiments and depending on the organic content of the particles, a wide range of $\kappa$ values were reported in the literatures. Cerully et al. (2011) showed that the $\kappa$ values measured in 2007 by Flow-Streamwise Thermal Gradient CCN Chamber (CFSTGC) ranged mostly between 0.1 and 0.4 in a forest environment in Finland. Hong et al. (2014) obtained the average $\kappa$ values of 0.15 (110 nm) and 0.28 (102 nm) measured by H-TDMA at the same site in 2010. Chang et al. (2010) used an AMS to measure aerosol chemical composition and a mole ratio of atomic oxygen to atomic carbon (O/C) at a rural site in Canada. They reported a relationship between the $\kappa$ values of organics and the O/C ratio as $\kappa_{\text{org}}=(0.29 \pm 0.05)\times(O/C)$. Tritscher et al. (2011) conducted smog chamber experiments for measurements of the $\kappa$ values of aging secondary organic aerosols and they found that the $\kappa$ was a sensitive indicator of the SOA properties.

Although the $\kappa$ values were reported under different environments in many locations, only a few studies were conducted to measure $\kappa$ in the Pearl River Delta (PRD) region (Cheung et al. 2015; Schurman et al. 2017). Jiang et al. (2016) compared the $\kappa$ values between wintertime (0.18-0.22) and
summertime (0.17-0.21) in Guangzhou. Cai et al. (2017) reported the κ values of about 0.4-0.6 and 0.2-0.3 measured by the H-TDMA respectively in Cape Hedo (Japan) and in Guangzhou (China). Alternatively, the average κ values can be predicted by the ZSR mixing rule (Zdanovskii, 1948; Stokes and Robinson, 1966) which is based on the chemical composition of the aerosol particles from the AMS measurements. Liu et al. (2014) reported the κ values of 0.22 to 0.32 using the ZSR mixing rule, consistent with the values (0.25 to 0.34) based on the H-TDMA measurements.

Once the κ values were determined, they could then be employed to predict the CCN activity that was characterized by two important parameters: activation diameter ($D_{50}$) and activation ratio (AR). Until now, the CCN activity (thus the above three parameters) can be determined using the following three methods:

1. The combination of Cloud Condensation Nuclei counter (CCNc) and Scanning Mobility Particle Sizer (SMPS). The CCN number was measured by the CCNc at different supersaturation ratios (SS, typically 0.05% ~ 1%). Meanwhile, the $D_{50}$ and size-resolved activation ratios could be measured by combining the CCNc with a differential mobility analyzer (DMA) and a condensation particle counter (CPC) (Moore et al., 2011; Deng et al., 2011), referred to as Scanning Mobility CCN Analysis (SMCA) based on measurements from a SMPS (DMA+CPC) and a CCNc. This method can measure the size-resolved CCN distributions at a high time resolution (Moore et al., 2010) and has been applied in lab experiments (Asaawuku et al., 2009) and filed campaigns (Moore et al., 2008) to measure CCN activity.

2. The ZSR method based on chemical composition measurements. The CCN concentrations were
inverted from the chemical composition and the size distribution of the aerosol particles measured respectively from the aerosol mass spectrometer (AMS) and SMPS (Moore et al., 2012; Meng et al., 2014). The $\kappa$ was then calculated from the ZSR mixing rule. In general, the particles were assumed to be internally mixed, which might lead to a large uncertainty (up to 80%) in predicting $N_{\text{CCN}}$ in some cases (Wang et al., 2010).

(3) The H-TDMA method. The size-resolved CCN distribution and activation ratios could be determined from hygroscopocity and size distribution measured using the H-TDMA (Good et al., 2010; Wu et al., 2013). The H-TDMA measured the distribution of hygroscopic growth factor ($G_f$) at a fixed relative humidity for a selected diameter of aerosol particles. Väisänen et al. (2016) reported that the measured $N_{\text{CCN}}$ with the H-TDMA agreed well with that from in-cloud prediction, where the sample was collected from a tower approximately 224 m above the surrounding lake level. On the other hand, Chan (2008) attributed differences in $\kappa$ from HTDMA and CCN measurements to sparingly soluble organics that did not easily deliquesce in the former measurements.

The PRD region is one of the most economically invigorating regions in China. This region is subjected to severe air pollution due to intensive human activities and insufficient pollution control measures. High particle loading leads to both visibility degradation and large cooling effects due to decrease of solar radiation. During wintertime, high concentrations of fine particles also cause severe haze events that pose health risk on people at the regional scale. It is hence an ideal location to investigate the influence of local anthropogenic emissions on the particles properties. However, there is still lack of understanding on the relationship between the CCN activity and its controlling factors (e.g.,
chemical composition and hygroscopicity of aerosol particles), hindering policy-makers to propose effective measures for climate-related policy-making.

In this study, we used the SMCA, H-TDMA, and HR-ToF-AMS to respectively measure CCN activity, hygroscopicity, and chemical composition. We reported the relationship between CCN activity and hygroscopicity/chemical composition of aerosol particles in the PRD region, where only a few studies on such relationship were available in the literature. The measurements were performed during wintertime 2014 (November and December). The CCN properties were predicted based on the combined SMCA, H-TDMA and HR-ToF-AMS measurements. The methods employed to predict the CCN concentrations were evaluated and the impact of organics on CCN concentrations was discussed.

2 Experiments and data analysis

2.1 Measurement site

The field measurements were conducted at the Chinese Meteorological Administration (CMA) Atmospheric Watch Network (CAWNET) Station in Panyu, Guangzhou, China, during wintertime 2014 (November and December). The Panyu Station is located at the center of the PRD region and at the top of Dazhengang Mountain (23°00’N, 113°21’E) with an altitude of about 150 m. No significant local emission sources were around the site. Detailed description of the measurement site and instruments (i.e., the HTDMA and the AMS) can be found elsewhere (Cai et al., 2017; Qin et al. 2017).

2.2 Instrumentation
2.2.1 Aerosol hygroscopicity measurements

Size-resolved aerosol hygroscopicity and particle number size distribution (PNSD) were measured by a H-TDMA which was developed by Tan et al. (2013). The hygroscopicity data were only available in November due to the failure of the H-TDMA during December. An aerosol sampling port equipped with a PM$_{1.0}$ impactor inlet was used during the measurement period. Ambient sampling flow first passed through a Nafion dryer (Model PD-70T-24ss, Perma Pure Inc., USA) to achieve a RH of <10%. We considered the particles to be dry when the RH values were less than 10%. The particles were subsequently charged by a neutralizer (Kr85, TSI Inc.) and size-selected by a differential mobility analyzer (DMA1, Model 3081L, TSI Inc.). The mono-disperse particles with a specific diameter ($D_0$) were then introduced into a Nafion humidifier (Model PD-70T-24ss, Perma Pure Inc., USA) under a fixed RH of (90 ± 0.44) %. Another differential mobility analyzer (DMA2, Model 3081L, TSI Inc.) and a condensation particle counter (CPC, Model 3772, TSI Inc.) were used to measure the number size distribution of the humidified particles ($D_p$). Thus, growth factor (Gf) of the particles can be calculated:

$$G_f = \frac{D_p}{D_0}$$

(2)

During the campaign, we selected five dry mobility diameters (40, 80, 110, 150, and 200 nm) for the H-TDMA measurements. The measurements were performed continuously except for regular calibration of the instrument. We used standard polystyrene latex spheres and ammonium sulfate to perform the DMA calibration to ensure the instrument to function normally.
2.2.2 Size-resolved CCN activity measurements

Size-resolved CCN spectra and activation ratios were measured with the SMCA initially proposed by Moore et al. (2010). In this work, the SMCA consisted of a CCNc-100 (DMT Inc.), a differential mobility analyzer (DMA, Model 3081L, TSI Inc.) and a condensation particle counter (CPC, Model 3787, TSI Inc.). In the SMCA system, the combined DMA and CPC were used as a scanning mobility particle sizer (SMPS) during the measurements. The dry particles after the Nafion dryer were neutralized by the Kr85 neutralizer and were subsequently classified by the DMA. The mono-disperse particles were split into two streams: one to the CPC for measurement of total particle number concentration (N_{CN}) and another to the CCNc-100 for measurements of the CCN number concentration. The aerosol and CPC flow rate was both 1.0 L min^{-1} for the DMA and the CPC (0.5 L min^{-1} makeup flow and 0.5 L min^{-1} sample flow), respectively. The CCNc-100 drew another aerosol flow rate of 0.5 L min^{-1}. The SMCA was protocoled to measure particles at a mobility diameter range of 10 - 400 nm. The supersaturation in the CCNc-100 was set to be 0.1%, 0.2%, 0.4% and 0.7% respectively for each measurement cycle. The CCNc-100 was regularly calibrated with ammonium sulfate particles at the four SS (0.1%, 0.2%, 0.4%, and 0.7%). Previous studied showed that different parameterizations in the Köhler theory can retrieve different critical supersaturations (Rose et al., 2008; Wang et al., 2017).

When performing the CCNc calibration, we assumed the density and molecular weight of ammonium sulfate to be 1770 kg m^{-3} and 0.132141 kg mol^{-1}, respectively. We also set the temperature and the pressure to 298.15 K and 1026 hPa, respectively. A temperature gradient ΔT of about 3-8 K in the CCNc column was also used in the calibrations. Similarly, the DMA was calibrated with standard polystyrene latex spheres before and after the campaign for quality assurance and control.
2.2.3 Aerosol chemical composition measurements

An Aerodyne high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS) was employed in the campaign to measure non-refractory PM$_1$ chemical composition (bulk and size-resolved) including sulfate, nitrate, ammonium, chloride, and organics. The refractory components such as black carbon, sea salts and crustal species cannot be measured by this instrument. Detailed description of the HR-ToF-AMS can be found elsewhere (DeCarlo et al., 2006; Jimenez et al., 2003). Here only a brief description relevant to the measurements is given. The instrument was operated in three modes (pToF, V, and W mode). Particle size distribution could be obtained based on time-of-flight of the particles in pToF mode. In V and W modes, the resolving power of the mass spectrometer was approximately 2000 and 4000, respectively. The instrument collected alternatively 5-min average mass spectra for the V + pToF modes and the W mode. The monodisperse pure ammonium nitrate (NH$_4$NO$_3$) particles selected by a DMA (400 nm) were used weekly in the ionization efficiency (IE) calibration. Background signals were obtained daily for about 30 minutes by introducing filtered ambient air with a HEPA filter in the sample flow. Before and after the measurement, the sampling flow rate was calibration with a Gilian gilibrator. We also generated PSL (Duke Scientific) and ammonium nitrate particles in a size range of 178–800 nm to calibrate the pToF size. Note that the mass concentrations were too low for particle diameters smaller than 65 nm and data for those particles were hence discarded in this study. A more detailed description of the AMS performance during the measurements can be found in Qin et al. (2017) and Cai et al. (2017).
The AMS measured size-resolved chemical composition of particles in vacuum aerodynamic diameter ($D_{\text{a}}$). It is hence necessary to convert aerodynamic diameter to mobility diameter in order to compare the AMS data and the SMCA data. We adopted the equation derived by DeCarlo et al. (2004) to do the conversion. Here we assume a density of 1700 kg m$^{-3}$ for particles measured by the AMS (DeCarlo et al., 2004).

2.3 Data processing and methodology

2.3.1 Hygroscopicity

Due to the effects of diffusing transfer function, the measured distribution function (MDF) given by H-TDMA is only a skewed and smoothed integral transform of the actual growth factor probability density function (Gf-PDF) of the particles (Gysel et al., 2009). Here the TDMAfit algorithm (Stolzenburg and McMurry, 2008) was applied to narrow the uncertainties caused by the diffusion broadening. The TMDAfit algorithm describes the Gf-PDF as a combination of several (usually smaller than three) lognormal distribution functions, in which the parameters of each mode are considered as mean Gf, standard deviation, and number fraction. The detailed data inversion process of the H-TDMA instrument can be found in Tan et al. (2013). Note that we include multiply charged correction for the SMCA, SMPS and H-TDMA data when the data were inverted so that the contributions of the multiply charged particles were accounted for all the measured particle data.

As mentioned in the introduction, the CCN activity can be represented by a widely used hygroscopicity parameter $\kappa$ (Petters and Kreidenweis, 2007). According to the $\kappa$-Köhler theory, for a known
temperature, κ and Gf can be related via eq. 3 (Petters and Kreidenweis, 2007):

\[
\kappa = (G f^3 - 1) \left[ \frac{1}{R H} \exp \left( \frac{A \sigma_s/a M_w}{R T \rho_w D} \right) - 1 \right]
\]  

(3)

where \( \rho_w \) is the density of water (about 997.04 kg m\(^{-3} \) at 298.15 K), \( M_w \) is the molecular weight of water (0.018 kg mol\(^{-1} \)), \( \sigma_s/a \) is the surface tension of the solution/air interface and here pure water is tentatively assumed for the solution (\( \sigma_s/a = 0.0728 \) N m\(^{-1} \) at 298.15 K), \( R \) is the universal gas constant (about 8.31 J mol\(^{-1} \)K\(^{-1} \)), \( T \) is thermodynamic temperature in Kelvin (298.15 K), and \( D \) is the particle diameter (in meter).

2.3.2 CCN activation

The \( N_{CN} \) and \( N_{CCN} \) data were respectively measured by the SMPS and the CCNc-100 and they were used to calculate the size-resolved CCN activation ratios (AR) which was defined as the ratio of \( N_{CCN} \) to \( N_{CN} \) at each particle size. The activation ratio can be obtained by fitting the ratio with the sigmoidal function with respect to \( D_p \):

\[
\frac{N_{CCN}}{N_{CN}} = \frac{B}{1 + \left( \frac{D_p}{D_{50}} \right)^C}
\]

(4)

where \( D_p \) is the particle dry diameter, \( B \), \( C \) and \( D_{50} \) are fitting coefficients that represent the asymptote, the slope, and the inflection point of the sigmoid, respectively (Moore et al., 2010). A steep activation curve is associated with a small \( C \) value. \( D_{50} \) is also called the critical diameter or the activation diameter, that is, the diameter at which 50% of the particles are activated at a specific SS.

Alternatively, the hygroscopicity parameter κ can be calculated from the critical saturation ratio (Sc) and \( D_{50} \) from the following equation (Petters and Kreidenweis, 2007):
\[ \kappa = \frac{4A^3}{27D_0^2(\ln \delta c)^2}, \quad A = \frac{4\sigma_s d M_w}{RT \rho_w} \]  

(5)

3 2.3.3 CCN prediction based on H-TDMA and AMS measurements

The \( N_{\text{CCN}} \) can be predicted based on either the aerosol hygroscopicity data (measured by the H-TDMA) or the AMS data. Figure 1 is the schematic diagram of the four approaches we followed to predict \( N_{\text{CCN}} \) based on the above two measured datasets. In the first approach (I in Fig. 1), the mixing state and size dependence were taken into account. We assumed the critical hygroscopicity parameter \( \kappa_{\text{critical}} \) to be a function of the particle diameter and the supersaturation ratio (denoted as \( \kappa_{\text{critical}}(D_p, SS) \)). The \( \kappa_{\text{critical}} \) was hence defined as the point at which all the particles were activated at a specific diameter and a specific SS. Here we measured hygroscopicity using the H-TDMA at five dry diameters and the CCN concentrations at four SS. We calculated the \( \kappa_{\text{critical}}(D_p, SS) \) using eq.5 for a known diameter and SS. A particle with a \( \kappa \) value higher than \( \kappa_{\text{critical}}(D_p, SS) \) was considered to be activated as an CCN (Fig. 1a) and the shadow area represented the particles which can be activated as CCN for a known diameter and SS. The activation ratio for a specific diameter at a specific SS was obtained by integrating the \( \kappa \)-PDF for \( \kappa > \kappa_{\text{critical}}(D_p, SS) \).

This approach is similar to the one employed in Kammermann et al. (2010), however, we used the size-resolved activation ratio (AR\(_{\text{SR}}\)) to calculate the \( N_{\text{CCN}} \). The AR\(_{\text{SR}}\) was determined by fitting the AR\((D_p,SS)\) to the diameter \( D_p \) using eq. 4 for the five measured diameters (Fig. 1d). Thus, the calculated \( N_{\text{CCN}} \) using the activation ratio can be expressed as (Fig. 1e):

\[ N_{\text{CCN}}(SS) = \int_0^{\infty} AR_{\text{SR}}(D_p, SS)N_{\text{CN}}(D_p)dDp \]  

(6)
In the second approach (II in Fig. 1), the particles were assumed to be internally mixed. The $D_{50}$ was determined by fitting the AR(Dp,SS) to the diameter Dp (Fig. 1d). The $N_{CCN}$ was obtained by integrating the cloud nuclei concentration for particles larger than $D_{50}$ based on the particle size distribution (Fig. 1f), according to the following equation (eq. 7):

$$N_{CCN}(SS) = \int_{D_{50}}^{\infty} N_{CN}(Dp) dDp$$  

(7)

In the third and fourth approaches (III and IV in Fig. 1), the particles were also assumed to be internally mixed. We then calculated the $\kappa$ value according to the ZRS rule (eq. 8) based on the AMS measurements.

$$\kappa = \sum_1 \varepsilon_i \kappa_i$$  

(8)

where $\varepsilon_i$ is the volume fraction of each component in the particles, $\kappa_i$ is the $\kappa$ value of each component.

The AMS only provided the ion concentrations during the measurements, while the ZSR rule required the volume fraction and hygroscopicity of each component. A simplified ion pairing scheme developed by Gysel et al. (2007) was used to reconstruct the $\text{NH}_4^+$, $\text{SO}_4^{2-}$ and $\text{NO}_3^-$ measured by the AMS:

$$n_{\text{NH}_4\text{NO}_3} = n_{\text{NO}_3^-}$$

$$n_{\text{H}_2\text{SO}_4} = \max(0, N_{\text{SO}_4^{2-}} - n_{\text{NH}_4^+} + n_{\text{NO}_3^-})$$

$$n_{\text{NH}_4\text{HSO}_4} = \min\left(2n_{\text{SO}_4^{2-}} - n_{\text{NH}_4^+} + n_{\text{NO}_3^-}, n_{\text{NH}_4^+} - n_{\text{NO}_3^-}\right)$$

$$n_{(\text{NH}_4)^2\text{SO}_4} = \max\left(n_{\text{NH}_4^+} - n_{\text{NO}_3^-} - n_{\text{SO}_4^{2-}}, 0\right)$$

$$n_{\text{HNO}_3} = 0.$$

(9)
where $n$ denotes the number of moles of each component (i.e., $\text{NH}_4^+$, $\text{SO}_4^{2-}$, and $\text{NO}_3^-$). Here we used the ADDEM proposed by Topping et al. (2005) to calculate the $\kappa$ values of the inorganic species and those of the organics were tentatively assumed to be 0.1 (Meng et al., 2014). Table 1 lists the $\kappa$ values of the relevant species used in the study based on the calculations and the above assumption.

Here instead of being determined from fitting of $\text{AR}_{SR}$ to $D_p$ used in the second approach, the $D_{50}$ was calculated from the above $\kappa$ values using eq. 5. In the third approach, the $\kappa$ values were size-resolved because the chemical composition of the particles was size dependent (Fig. 1b). In the fourth approach, the particles were assumed to have the same chemical composition and hygroscopicity as those in PM$_1$ (Fig. 1c). The $N_{\text{CCN}}$ was then predicted using eq. 7 (Figs. 1 g and h).

### 3 Results and discussion

#### 3.1 Overview

Table 2 summarizes the observed CCN activity during the campaign. Overall, the average $N_{\text{CCN}}$ at 0.1, 0.2, 0.4, and 0.7% SS were about 3100, 5100, 6500, and 7900 cm$^{-3}$, respectively. The average activation ratios (AR) at the above four SS were 0.26, 0.41, 0.53 and 0.64, respectively. The average $D_{50}$ at the above four SS were 156, 107, 78 and 58 nm, respectively. The $N_{\text{CCN}}$ at 0.7% SS was respectively lower than those of the previous measurements (10731 cm$^{-3}$ at 0.67% SS) in July 2006 in Guangzhou (Rose et al., 2010), but much higher than those measured (2085 cm$^{-3}$ at 0.7% SS) in May 2011 in Hong Kong (Meng et al., 2014), while the $N_{\text{CCN}}/N_{\text{CN-tot}}$ was lower than those from the previous measurements (at 0.67% SS) in Guangzhou (0.59, Rose et al., 2010) and similar to those from the
measurements (at 0.7% SS) in Hong Kong (0.64, Meng et al., 2014). The $D_{50}$ was larger than that in the previous measurements in Guangzhou (49 nm) and in Hong Kong (47 nm), due to the lower particle hygroscopicity in Guangzhou. The differences of the $\kappa_{CCN}$ values between the two measurements (0.21 in this winter campaign vs 0.28 during the summer season in Guangzhou both at 0.7% SS) suggested that the particles in the summer were in general more hygroscopic and hence were more readily activated than those in the winter, implying different chemical composition of the particles between the two distinct seasons. Mochida et al. (2010) measured the size-resolved CCN activity in Cape Hedo, a remote marine site rarely affected by anthropogenic emissions. The results showed that the $D_{50}$ at 0.1% SS in Cape Hedo was about 130 nm, much larger than that in Guangzhou, leading to higher hygroscopicity of atmospheric particles in Cape Hedo than that in Guangzhou.

Figure 2 shows the average mass fraction of NR-PM$_1$ bulk composition and size-resolved (64-731 nm) composition. The organics was dominant in the bulk NR-PM$_1$ (50%), followed by sulfate (26%) and nitrate (12%) (Fig. 2a). The mass fraction of the organics decreased with the size (Fig. 2b), from 73% at 64 nm to 42% at 397 nm. The mass fraction of organics at 397 nm was close to that of NR-PM$_1$ bulk, due to the fact that the PM$_1$ mass is dominated by particles in a diameter range of 200~500 nm (Tan et al., 2016). In comparison, the dominant NR-PM$_1$ species observed in Hong Kong were sulfate (51.0%) and organics (28.2%) (Lee et al., 2013), significantly different from our measurements, due probably to different origins of the dominant air masses between the two seasons. The measurement site in Guangzhou was impacted predominantly by the air mass from north, where straw burning contributes to a high mass fraction of organics matter (Cao et al., 2008).
Figure 3 shows the $\kappa$ values based respectively on the CCN ($\kappa_{\text{CCN}}$), AMS ($\kappa_{\text{AMS}}$), and H-TDMA ($\kappa_{\text{H-TDMA}}$) measurements, along with the measured particle number size distribution (PNSD, 10-400 nm) during the campaign. The shadow area represents the interquartile range of the PNSD. A distinct peak at around 90 nm was observed from the PNSD (Fig. 3). The $\kappa_{\text{AMS}}$ was calculated based on the size-resolved chemical composition, assuming the particles are internally mixed. At 0.7% SS, the $D_{50}$ was about 58 nm. Hence no $\kappa_{\text{AMS}}$ was reported at this SS since we only measured particle composition above 63 nm using the AMS in this study. The $\kappa$ values were shown in the interquartile range, with the largest variation from the CCN measurements (Fig. 3). As shown in Fig. 3, the difference between $\kappa_{\text{CCN}}$ and $\kappa_{\text{H-TDMA}}$ is statistically insignificant at all employed diameters, while the one between $\kappa_{\text{AMS}}$ and $\kappa_{\text{CCN}}$ became statistically significant at larger sizes of the particles. The $\kappa_{\text{H-TDMA}}$ values were lower than those of the corresponding $\kappa_{\text{CCN}}$ at most of the particle sizes, consistent with the previous observation (Pajunoja et al., 2015). This was probably due to the facts that particles contain a certain fraction of low solubility composition, such as secondary organic aerosols (SOA), contributing differently to hygroscopic growth and CCN activation. The available AMS data (Fig. 3) show that the $\kappa_{\text{AMS}}$ values were lower than the corresponding $\kappa_{\text{CCN}}$ and $\kappa_{\text{H-TDMA}}$ values at all size ranges and the differences become larger with increasing particle sizes. This was probably due to underestimated hygroscopicity in the organic composition when using the AMS data, since we assumed a $\kappa$ value of 0.1 for all organics at all particle sizes. The hygroscopicity increased with particle diameters due to aerosol aging which increased the hygroscopic organic contents. Previous studies showed that the $\kappa_{\text{org}}$ values of larger particles are lower than those for smaller particles (Lance et al., 2013; Zhao et al. 2015) and
hygroscopicity of organics is often found to be related to its chemical composition ($f_{44}$ or O/C) in both
field and laboratory studies (Chang et al. 2010; Massoli et al., 2010; Lambe et al., 2011; Mei et al.,
2013, and others reference therein). We showed that the $f_{44}$ increased with the particle size from the
AMS data (Fig. S1). Note that the $f_{44}$ for particle diameters smaller than 100 nm was discarded due to
the data quality. The results indicate that the degree of oxidation of the organics was higher for larger
size particles and the hygroscopicity for larger particles is higher (Chang et al., 2010). The measured
$k_{\text{mean}}$ values fall in a range of 0.22-0.30 for the particle sizes of 40-200 nm measured by H-TDMA in
this study. The other aerosol hygroscopicity measurement in PRD (Jiang et al., 2016) reported the $k_{\text{mean}}$
values ranging from 0.18 to 0.22 in 2012 winter season and 0.17 to 0.21 in 2013 summer season,
suggesting an increase of the aerosol hygroscopicity, which might result from an increasing mass
fraction of nitrate in recent years (Zhang et al., 2015; Itahashi et al., 2018), although the fraction
decrease of less hygroscopic compounds is not as significant as the fraction increase of the nitrate.
However, the fraction of the non-hydroscopic compounds (i.e. EC) decreases more rapidly than the
organic compounds.

Figure 4 shows the activation ratios (AR) measured by SMCA at four supersaturation ratios (0.1, 0.2,
0.4, 0.7%) for particles below 300 nm. The activation curves obtained in this study were segmented
into three sections: a steady rise at low ARs, a middle sharp increase, and a plateau at almost 100% AR.
We defined the steepness as the rate at which the AR increased with the particle sizes. Figure 4 shows
the steepness increased with the SS, indicating that the curves became steeper with the SS and a larger
variation of the $D_{50}$ was expected. In addition, the CCN activity was more sensitive to particle
diameters at higher SS, which can be seen from partial derivative of $\kappa_{\text{critical}}$ by $D_{50}$ (eq. 10):

$$\frac{\partial \kappa_{\text{critical}}}{\partial D_{50}} = -\frac{4A^3}{9D_{50}^3 (\ln \gamma)^2}$$  (10)

For a certain SS, the $\kappa_{\text{critical}}$ value became more sensitive to $D_{50}$ with decrease of the $D_{50}$. Meanwhile, a high SS usually led to a low $D_{50}$. Therefore, the AR would vary with $D_p$ more readily at higher SS and the curve would become steeper. A higher SS allowed a smaller particle to be activated and the activation curve became steeper, and vice versa for a lower SS.

The steepness of activation curve was also associated with the heterogeneity of aerosol chemical composition, that was, a steeper activation curve meant that aerosol particles had higher similarity in hygroscopicity. A bimodal distribution (peaks at about 1-1.1, and 1.5-1.7 Gf) of the Gf-PDFs was observed along the Gf coordinate at all the five sizes of the particles measured by H-TDMA in this study (Fig. 5), corresponding respectively to the less- and more-hygroscopic modes. Larger size particles contain higher fractions of more-hygroscopic inorganics matters which lead to the increase of Gf of more-hygroscopic mode. The less-hygroscopic mode usually represents externally mixed black carbon or fresh organics. Thus the less-hygroscopic mode for larger size particles more likely represents the external mixing non-hygroscopic black carbon with a Gf value of 0.8-1.1, indicating that the Gf of less-hygroscopic mode decreased and that of more-hygroscopic mode increased with the particle diameter (Fig. 5). Since less-hygroscopic particles were usually associated with externally mixed black carbon (BC) or fresh organics and more-hygroscopic particles usually represent the inorganics matters or BC coated with inorganics matters (internally mixed). The decrease of peak area
of less-hygroscopic mode and the increase of more-hygroscopic mode indicate that the number fraction of less-hygroscopic particles decreased while the more-hygroscopic particles fraction increased. Thus, the particles became more internally mixed. Here a parameter $\sigma$ is introduced to illustrate the deviation of Gf-PDF (Gysel et al., 2009):

$$G_{f\text{mean}} = \int_0^\infty Gf c(Gf)dGf$$

$$\sigma = \left( \int_0^\infty (Gf - G_{f\text{mean}})^2 c(Gf)dGf \right)^{\frac{1}{2}}$$

where the $c(Gf)$ denotes Gf-PDF and $G_{f\text{mean}}$ denotes number weighted mean $Gf$. The $\sigma$ was employed as a measure of the spread of Gf-PDF which represents the heterogeneity of aerosol chemical composition (Sjogren et al., 2008; Liu et al., 2011). A small $\sigma$ indicated that the heterogeneity of aerosol chemical composition was low and aerosol particles had higher similarity in hygroscopicity.

The parameter $C$ determined the shape of the activation curve which was segmented into steep and smooth parts. A small $C$ value means a steep activation curve and vice versa. Here an activation curve was assumed to be steep when the $C$ values are lower than the lower quartile of all the $C$ values, while the activation curve was considered to be smooth when the $C$ values are higher than the upper quartile of all the $C$ values. Table 3 summarizes the $\sigma$ values of Gf-PDF for the corresponding steep and smooth activation curve at the four supersaturations. In general, the $\sigma$ increased with the diameter, indicating that larger particles had higher heterogeneity of aerosol chemical composition. Meanwhile, the $\sigma$ values for smooth curve were generally higher than the $\sigma$ values for steep curve. The results implied that the shapes of activation curves were related to the heterogeneity of aerosol chemical composition. Cai et al. (2017) compared the Gf-PDF between Guangzhou and Cape Hedo and the results showed that only more-hygroscopic (MH) particles were observed in Cape Hedo, indicating that
atmospheric particles tend to be more internally mixed in Cape Hedo than in Guangzhou. Meanwhile, atmospheric particles in Guangzhou have a higher degree of external mixing affected by more anthropogenic emissions, which in turn affect the CCN activity.

3.2 Impact of organics on CCN activity

Figure 6 shows the relationship between the $D_{50}$ obtained from the SMCA measurements and the size-resolved mass fractions of organics ($f_{org}$) at three supersaturation ratios (0.1%, 0.2%, and 0.4% SS). In general, the $D_{50}$ increased with $f_{org}$ at the three SS, with a slope of 127, 66, 21, and a fitting coefficient ($R^2$) of 0.47, 0.31, 0.1 at 0.1%, 0.2%, and 0.4% SS, respectively. The particles usually became less hygroscopic with increase of the organic fractions ($f_{org}$), which then required larger particles to be activated. At lower SS, better correlations were found between the $f_{org}$ and the $D_{50}$ because the $D_{50}$ was more sensitive to hygroscopicity (The activation ratios increase more slowly with particle sizes at lower SS as shown in Fig. 4). It was hence more obvious at lower SS that the modification of the particle hygroscopicity caused by the change of the mass fraction of organics matter could greatly modify the $D_{50}$ which might further affect the CCN activity. At higher SS, according to eq. 5, particles were more easily activated as CCN and the change of particles hygroscopicity would not significantly alter the CCN activity.

Organics can affect the CCN activity via two opposite ways: they can decrease the CCN activity by increasing the less hygroscopic organic fraction of the particles and thus increase the $D_{50}$ as shown in Fig. 6; they can also increase the CCN activity by decreasing the surface tension of the particles.
latter effect has been demonstrated experimentally. For example, an increase of CCN activity was observed when organics were added to sulfate ammonium (Engelhart et al., 2008). In this study, we investigated the impacts of organics on CCN activity through adjusting the value of surface tension until the calculated AR values based on H-TDMA measurements agree with those obtained from SMCA measurements (measured AR). The calculated AR values were systematically lower than the corresponding measured ones because here the surface tension of bulk pure water (0.072 N m$^{-1}$) was assumed when calculating the AR from the H-TDMA measurements (Fig. 7). Note that the surface tension is not the only factor that determines the AR and other factors such as the sparingly soluble compounds in the particles may contribute to the AR, although they are currently not understood. Previous studies found that the hygroscopicity of the particles measured by the H-TDMA could be lower than that measured by the CCNc (Chan et al., 2008; Pajunoja et al. 2015; Petters et al., 2009; Hansen et al., 2015; Hong et al., 2014) which might be attributed to low soluble compounds in the particles. The deviation of the calculated AR from the measured AR is probably dependent on the degree of dissolution of particles and the oxidative state of the organics in the particles.

The surface tension of a nanoparticle was substantially different from that of its bulk solution due to the curvature effect (Ahn et al., 2010; Bogdan, 1997). The effects of size and composition on the surface tension were currently not well understood. Here we proposed an approach to evaluate the impact of organics on the surface tension ($\sigma_{sa}$) based on the fraction change of the calculated AR to the measured AR. We defined this fraction change ($\delta_{AR}$) as a function of surface tension, diameter, and supersaturation:
\[
\delta_{AR}(\sigma_{s/a}, Dp, SS) = \frac{AR_m(Dp, SS) - AR_c(\sigma_{s/a}, Dp, SS)}{AR_m(Dp, SS)} \times 100\%
\]  

where \( AR_m(Dp, SS) \) is the measured AR for a certain diameter and SS, \( AR_c(\sigma_{s/a}, Dp, SS) \) is the calculated AR for a certain diameter, SS, and \( \sigma_{s/a} \). We excluded particles at the size of 200 nm because they were easily activated even at 0.1% SS and the \( \delta_{AR} \) was expected to be independent of \( \sigma_{s/a} \). Here the \( \sigma_{s/a} \) value varied between 0.03 and 0.072 N m \(^{-1}\) (surface tension of pure water). Figure 8 shows the \( \delta_{AR} \) as a function of \( \sigma_{s/a} \) for the four particle diameters (40, 80, 110, 150 nm). The \( \delta_{AR} \) decreased with increase of the \( \sigma_{s/a} \) for all given particle sizes, changing more rapidly for smaller particles (i.e., from 200% to -100% for 40 nm) than bigger particles (i.e., from 20% to -10% for 150 nm). The \( R^2 \) between measured AR and predicted AR for a certain diameter and four supersaturations at \( \sigma_{s/a} = 0.072 \) N m \(^{-1}\) were 0.35, 0.93, 0.95 and 0.91, respectively. The \( \delta_{AR} \) values reached zero when the \( \sigma_{s/a} \) was set to be about 0.054 N m \(^{-1}\) for 40, 80, and 110 nm particles, and 0.062 N m \(^{-1}\) for 150 nm particles, with a \( R^2 \) of 0.88, 0.94, 0.94 and 0.88 respectively. As a compromise, here we adopt a \( \sigma_{s/a} \) value of 0.058 N m \(^{-1}\) (denoted as \( \sigma_{s/a}^* \)) to predict AR. This \( \sigma_{s/a}^* \) value increased significantly the \( R^2 \) compared to that based on pure water assumption (0.072 N m \(^{-1}\)) for 40 nm particles, while it was reasonable well for other sizes of particles (80, 110, 150 nm). The AR was then recalculated using the \( \sigma_{s/a}^* \) value and the prediction was significantly improved (Fig. 9). The results demonstrated that partitioning of organics into aerosol particles would decrease their surface tension. Therefore, the pure water assumption for surface tension would lead to high uncertainties when it applied to predict the activation ratios of the aerosol particles at a certain size. Note that we did not consider the effects of individual organics due to the limited data from the chemical composition measurements. How chemical composition affects the surface tension of the particles is yet to be investigated.
3.3 The $N_{CCN}$ prediction

3.3.1 The $N_{CCN}$ prediction based on the H-TDMA measurements

In this study, we used several approaches to predict the $N_{CCN}$ based on the H-TDMA measurements, from either the activation curve or the $D_{50}$. Table 4 summarizes the methods that were used to predict the $N_{CCN}$, along with the slope and $R^2$ between the predicted and the measured values. The mixing state of the aerosol particles is an important parameter in determining the $N_{CCN}$. The prediction of $N_{CCN}$ using activation curve means the $N_{CCN}$ was calculated based on eq. 6. Meanwhile the prediction of $N_{CCN}$ using the $D_{50}$ means that the $N_{CCN}$ was calculated based on eq. 7 and the $D_{50}$ was determined from fitting the size-resolved activation ratio by eq. 4. The activation curve represented actual mixing state, while the $D_{50}$ approach assumed that all particles were internally mixed. Scheme 5 in Table 4 was the method based on the activation curve with the new $\sigma_{a/s}$ (0.058 N m$^{-1}$). Eq. 6 and Eq. 7 were respectively used to calculate the $N_{CCN}$ following schemes 1, 2, 5, and the rest of the schemes. Scheme 5 (real time activation curve using $\sigma_{a/s}$) provided the best $N_{CCN}$ predicted value (closest to the measured one), followed by scheme 3 (real time $D_{50}$) > scheme 4 (average $D_{50}$) > scheme 1 (real time activation curve) > scheme 2 (average activation curve). The $R^2$ values for all the approaches were in general high (around 0.93). The CCN prediction based on scheme 2 led to the largest underestimation over the measured values. In general, the real time data (schemes 1 and 3) gave better predicted $N_{CCN}$ than the corresponding average data (schemes 2 and 4).

Figure 10 shows the correlation between the measured $N_{CCN}$ and the predicted $N_{CCN}$ from schemes 1-5.
at the four SS. For schemes 1-4, the predicted $N_{CCN}$ values were found to be the largest deviation from the corresponding measured ones at 0.1% SS among all the approaches, probably due to the pure water assumption for surface tension ($\sigma_{sw} = 0.072 \text{N m}^{-1}$). Meanwhile, because the CCN activity was sensitive to hygroscopicity of the particles at low SS, the uncertainties of hygroscopicity data would lead to large errors in the prediction of CCN. As discussed in the previous section, the $D_{50}$ was more sensitive to the $\sigma_{sw}$ at lower supersaturations, leading to a large deviation of the $N_{CCN}$ from the measured value. The best agreement between the calculated AR and the measured AR was seen using scheme 5 as the slopes at the four SS were close to 1 (Fig. 10q-t).

3.3.2 The $N_{CCN}$ prediction based on AMS measurements

We proposed five approaches based on H-TDMA measurements to predict the $N_{CCN}$ in the previous section. Alternatively, we can calculate the $N_{CCN}$ based on AMS measurements. Here we proposed four methods based on either size-resolved chemical composition or bulk PM$_1$ chemical composition from the AMS measurements (Table 5). Here we assumed that the particles were internally mixed and the median $\kappa_{AMS}$ obtained from bulk composition was 0.28, higher than those from size-resolved composition (0.24-0.26 in Fig. 3), probably due to a higher mass fraction of inorganic matters in bulk NR-PM$_1$ (Fig. 2). We excluded the size-resolved data at 0.7% SS due to their poor quality. Note that the impact on the calculated $\kappa_{AMS}$ values and the predicted $N_{CCN}$ was minor using the $\kappa$ value (0.53) of ammonium sulfate from Petters and Kreidenweis (2007). For example, the $\kappa_{AMS}$ values slightly increased from 0.27 to 0.28 at 0.1% SS; the slopes for scheme 6, 8 and 9 in Table 5 slightly increased from 0.9859 to 0.9898, 0.9721 to 0.9834, and 0.9742 to 0.9973, respectively, while the one for scheme
7 did not change. Figure 11 shows the correlation between the measured and predicted $N_{CCN}$ from schemes 6-9. The $N_{CCN}$ was under-predicted at 0.1% SS and was over-predicted at 0.7% SS. We proposed three potential factors that might impact $N_{CCN}$ prediction based on AMS measurements. (1) The assumed $\kappa_{org}$ values were probably underestimated for particles larger than 100 nm, leading to the underestimated $N_{CCN}$ at low SS. As shown in Fig. 3, the predicted $\kappa$ shows a larger deviation from the measured value for a larger particle. The $D_{50}$ values were more sensitive to particle hygroscopicity at lower SS as discussed in the previous section. (2) The pure water assumption for surface tension. As we have shown in the previous section, the $\sigma_{s/a}$ values for the aerosol particles were found to be much smaller than the $\sigma_{s/a}$ for pure water (0.072 N m$^{-1}$). As a result, the pure water assumption for surface tension led to the $N_{CCN}$ underestimation. In addition, again the $D_{50}$ was more sensitive to $\sigma_{s/a}$ at the low SS. (3) The exclusion of black carbon (BC) particles and the mixing state assumption. The BC particles were known to be non-hygrosopic and had a low CCN activity. During the campaign period, the average BC concentration was about 5.91 µg/m$^3$ which account for 7 % in PM$_{2.5}$. The assumption of no BC particles would lead to the overestimation of $N_{CCN}$. Here, the particles were assumed to be internally mixed in the AMS measurements. This would lead to an overestimation of the $N_{CCN}$ when the ambient particles tend to be externally mixed (Wang et al., 2010; Sánchez Gácita et al., 2017). However, the internal mixing assumption seems to play a minor role in predicting the $N_{CCN}$ at 0.1% SS since the particles at about 140-180 nm tend to be internally mixed as shown in Fig. 5. In this case, the $\kappa_{org}$ assumption and the pure water assumption played more important roles than the mixing state assumption at low SS (i.e., 0.1% SS). Figure 11 shows significant underestimation of $N_{CCN}$ at 0.1% SS (panels a, e, i, m), while more or less comparable to the measured $N_{CCN}$ at higher SS (i.e. 0.2%, 0.4%,...
0.7%). The difference between the $\kappa_{\text{AMS}}$ and $\kappa_{\text{CCN}}$ became smaller and the corresponding $D_{50}$ value decreased with the increase of the SS so that the impacts of the $\kappa_{\text{org}}$ assumption and the pure water assumption became minor with the increase of the SS. Instead, the internal mixing state assumption would play a more important role in the prediction (Meng et al. 2014). As shown in Fig. 5, the peak height and area of the less-hygrosopic mode became larger for the smaller size particles (i.e., 40 nm particles), implying that small particles were likely to be externally mixed, that is, the non or less hygroscopic species including BC and insoluble organics were less likely coated with inorganics salts. Hence the internal mixing assumption could lead to an overestimated $N_{\text{CCN}}$.

As discussed above, the two important parameters ($\kappa_{\text{org}}$ and $\sigma_{\text{s/a}}$) had significant impacts on the $N_{\text{CCN}}$ prediction. We denoted $\kappa_{\text{org}}^*$ and $\sigma_{\text{s/a}}^*$ as important representations respectively for hygroscopicity and surface tension contributed from organics. We also pointed out that the $\kappa_{\text{org}}$ was dependent on the particle size and hence here we further assumed the $\kappa_{\text{org}}^*$ values to be 0.15 and 0.1 respectively for particles larger and smaller than 100 nm. Note that we previously assumed the $\kappa_{\text{org}}$ to be 0.1 for all particle sizes. Here we gave an example of the improvements at 0.1% SS when the $\kappa_{\text{org}}$ and $\sigma_{\text{s/a}}$ values were respectively replaced with the $\kappa_{\text{org}}^*$ and $\sigma_{\text{s/a}}^*$ ones (Fig. 12). The $\kappa_{\text{AMS}}$ value calculated at 0.1% SS based on $\kappa_{\text{org}}^*$ was 0.288, very close to the corresponding $\kappa_{\text{CCN}}$ value (0.30), indicating that an improvement was made for the $N_{\text{CCN}}$ prediction when including the $\kappa_{\text{org}}^*$ value. The $N_{\text{CCN}}$ prediction could be greatly improved when include both $\sigma_{\text{s/a}}^*$ and $\kappa_{\text{org}}^*$ in the calculation. For example, the underestimate of $N_{\text{CCN}}$ decrease from 44% (Fig. 11a) to 4% (Fig. 12b). In addition, we also investigated the effects of the $\sigma_{\text{s/a}}$ values in a range of 0.054 to 0.062 N m$^{-1}$ as discussed in section 3.2. The shadow
area in Fig. 12b represents the variation of linear fit between the measured and predicted \( N_{CCN} \). An under- and over-estimated value of 16% (slope=0.84) and 8% (slope=1.08) was obtained for the predicted \( N_{CCN} \) to the measured \( N_{CCN} \) using a \( \sigma_{s/a} \) value of 0.054 and 0.062 N m\(^{-1}\) respectively, indicating that the predicted \( N_{CCN} \) agreed reasonably with the measured ones when the \( \sigma_{s/a} \) values between 0.054 and 0.062 N m\(^{-1}\) were used in this study. We conclude that the predicted \( N_{CCN} \) can agree better with the measured one when including both \( \sigma_{s/a}^* \) and \( \kappa_{org}^* \) in the calculation at low SS.

4 Summary and Conclusions

The CCN activity is an important parameter that determines the extent to which atmospheric particles can influence cloud formation. It is hence essential to predict CCN activity so that a quantitative assessment of atmospheric particles on cloud formation can be made. While numerous studies were performed to investigate the CCN activity under different atmospheric conditions around the world, only a few of them were made in the PRD region in China. In this study, several advanced instruments (i.e., the SMCA, AMS and H-TDMA) were used to respectively measure CCN activity, chemical composition and hygroscopicity in PRD during wintertime 2014. A variety of schemes were proposed to determine the CCN activity based on the measurements. Here two important properties were considered when evaluating the CCN activity: the hygroscopic parameter \( \kappa \) and the surface tension of the particles. Three methods (i.e., the SMCA, the AMS+ZSR, and the H-TDMA) were employed to calculate the \( \kappa \) values based on our measurements. The results show that the deviation between \( \kappa_{AMS} \) and \( \kappa_{CCN} \) became larger at low supersaturation ratios, indicating that the organic components in larger size particles were more aged and hygroscopic. The activation curve became smoother at the low SS,
which could be partly attributed to the higher heterogeneity of chemical composition for larger particles.

In general, the Gf-PDF measured by H-TDMA exhibited a bimodal distribution with a less-hygroscopic mode and a more-hygroscopic mode. The less-hygroscopic mode was more significant at smaller diameters, indicating a more external mixing for smaller particles, while the more-hygroscopic mode increased with diameter and became broader, implying higher hygroscopicity and more complex chemical composition for larger particles. The shapes of activation curve were related to the $\sigma$ values of the Gf-PDF. The higher $\sigma$ values suggest the higher heterogeneity of chemical composition and smooth activation curve. A $k$ value of 0.22-0.30 measured by H-TDMA was obtained for 40-200 nm particles in this study during the measurement period, larger than those previously measured in the PRD region, indicating an increasing mass fraction of nitrate in recent years.

Organic compounds could influence CCN activity through modifying the hygroscopicity and surface tension of the particles. The impacts of organics on CCN activity were also investigated in this study. The increase of organic mass fraction in the particles could lead to the decrease of the aerosol hygroscopicity and hence increase the $D_{50}$, especially at low supersaturation. In addition, organics could decrease the surface tension $\sigma_{s/a}$. This could lead to the underestimated CCN activity if pure water solution is assumed when inverting the H-TDMA data. We evaluated the impact of the surface tension on the activation ratios over a wide range of $\sigma_{s/a}$ values (0.03-0.07 N m$^{-1}$) for several measured size particles (40, 80, 110, and 150 nm) and found that a $\sigma_{s/a}$ value of 0.058 N m$^{-1}$ was the best fit between predicted AR and measured AR, which could then be used to predict the CCN activity in the PRD region. Based on the hygroscopicity and chemical composition measured in this study, we
proposed several schemes to predict the CCN activity. Overall, the predicted $N_{\text{CCN}}$ agreed well with the measure one. The slope and $R^2$ of $N_{\text{CCN}}$ predicted from average data was similar to the $N_{\text{CCN}}$ predicted from real time data. The $N_{\text{CCN}}$ obtained from H-TDMA measurement was under-predicted, if pure water assumption was used and better agreement with the measured values can be achieved by using the adjusted $\sigma_{s/a}$ (i.e., $\sigma_{s/a}^* = 0.058 \text{ N m}^{-1}$). Similarly, the $N_{\text{CCN}}$ predicted from AMS measurement was underestimated at low supersaturations and overestimated at high supersaturations, due to an assumption of fixed 0.1 for $\kappa_{\text{org}}$ and the external mixing state. Better predicted CCN concentrations can be obtained by using $\sigma_{s/a}^*$ and $\kappa_{\text{org}}^*$ in the calculation, especially at low supersaturations. For high supersaturations, the effect of internal mixing assumption should be taken into consideration. We concluded that better agreement between predicted and measured CCN concentrations could be achieved by taking the effects of organic into account on the hygroscopicity, surface tension, and the mixing state of the particles. More work on the roles of organics on the CCN activity is obviously needed in order to better understand the impacts of atmospheric particles on cloud formation and hence climate.

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Table 1. The κ values of the related species in the study.

<table>
<thead>
<tr>
<th>Species</th>
<th>κ</th>
</tr>
</thead>
<tbody>
<tr>
<td>NH$_4$NO$_3$</td>
<td>0.58$^a$</td>
</tr>
<tr>
<td>NH$_4$HSO$_4$</td>
<td>0.56$^a$</td>
</tr>
<tr>
<td>H$_2$SO$_4$</td>
<td>0.90$^a$</td>
</tr>
<tr>
<td>(NH$_4$)$_2$SO$_4$</td>
<td>0.48$^a$</td>
</tr>
<tr>
<td>Organics</td>
<td>0.10$^b$</td>
</tr>
</tbody>
</table>

$^a$ κ of inorganics compounds are derived from ADDEM (Topping et al., 2005)

$^b$ κ of organics is from Meng et al. (2014)
<table>
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<tr>
<th>SS</th>
<th>0.1%</th>
<th>0.2%</th>
<th>0.4%</th>
<th>0.7%</th>
</tr>
</thead>
<tbody>
<tr>
<td>Max</td>
<td>15165</td>
<td>19989</td>
<td>25964</td>
<td>26208</td>
</tr>
<tr>
<td>$N_{\text{CCN}}$ (#/cm$^3$) Min</td>
<td>258</td>
<td>361</td>
<td>408</td>
<td>502</td>
</tr>
<tr>
<td>Mean±STD</td>
<td>3103±1913</td>
<td>5095±2972</td>
<td>6524±3783</td>
<td>7913±4234</td>
</tr>
<tr>
<td>Max</td>
<td>0.68</td>
<td>0.75</td>
<td>0.89</td>
<td>0.94</td>
</tr>
<tr>
<td>$N_{\text{CCN}}/N_{\text{CN, tot}}$ Min</td>
<td>0.06</td>
<td>0.10</td>
<td>0.19</td>
<td>0.28</td>
</tr>
<tr>
<td>Mean±STD</td>
<td>0.26±0.10</td>
<td>0.41±0.14</td>
<td>0.53±0.15</td>
<td>0.64±0.13</td>
</tr>
<tr>
<td>Max</td>
<td>268.90</td>
<td>194.04</td>
<td>145.28</td>
<td>97.17</td>
</tr>
<tr>
<td>$D_{50}$ (nm) Min</td>
<td>112.47</td>
<td>76.60</td>
<td>43.50</td>
<td>24.21</td>
</tr>
<tr>
<td>Mean±STD</td>
<td>156.02±19.48</td>
<td>106.66±16.99</td>
<td>77.96±14.86</td>
<td>58.45±10.68</td>
</tr>
</tbody>
</table>
**Table 3.** The average of $\sigma$ values of Gf-PDF measured by H-TDMA for five diameters for the steep and smooth activation ratio at four supersaturations.

<table>
<thead>
<tr>
<th>SS(%)</th>
<th>0.1</th>
<th>0.2</th>
<th>0.4</th>
<th>0.7</th>
</tr>
</thead>
<tbody>
<tr>
<td>Dp(nm)</td>
<td>Steep</td>
<td>Smooth</td>
<td>Steep</td>
<td>Smooth</td>
</tr>
<tr>
<td>40</td>
<td>0.13</td>
<td>0.17</td>
<td>0.12</td>
<td>0.17</td>
</tr>
<tr>
<td>80</td>
<td>0.16</td>
<td>0.20</td>
<td>0.14</td>
<td>0.20</td>
</tr>
<tr>
<td>110</td>
<td>0.17</td>
<td>0.21</td>
<td>0.15</td>
<td>0.21</td>
</tr>
<tr>
<td>150</td>
<td>0.19</td>
<td>0.22</td>
<td>0.17</td>
<td>0.23</td>
</tr>
<tr>
<td>200</td>
<td>0.20</td>
<td>0.23</td>
<td>0.19</td>
<td>0.24</td>
</tr>
</tbody>
</table>
Table 4. The overview of different schemes used in the $N_{CCN}$ prediction based on H-TDMA measurements.

<table>
<thead>
<tr>
<th>Scheme</th>
<th>Method</th>
<th>Slope</th>
<th>$R^2$</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>Real time activation curve</td>
<td>0.8275</td>
<td>0.93</td>
</tr>
<tr>
<td>2</td>
<td>Average activation curve</td>
<td>0.8183</td>
<td>0.93</td>
</tr>
<tr>
<td>3</td>
<td>Real time $D_{50}$</td>
<td>0.8869</td>
<td>0.93</td>
</tr>
<tr>
<td>4</td>
<td>Average $D_{50}$</td>
<td>0.8738</td>
<td>0.93</td>
</tr>
<tr>
<td>5</td>
<td>Real time activation curve using $\sigma_s$</td>
<td>0.9377</td>
<td>0.93</td>
</tr>
</tbody>
</table>
Table 5. The overview of methods used in the $N_{CCN}$ prediction based on AMS measurements.

<table>
<thead>
<tr>
<th>Scheme</th>
<th>Method</th>
<th>Slope</th>
<th>$R^2$</th>
</tr>
</thead>
<tbody>
<tr>
<td>6</td>
<td>Real time bulk composition</td>
<td>0.9859</td>
<td>0.91</td>
</tr>
<tr>
<td>7</td>
<td>Average bulk composition</td>
<td>1.0108</td>
<td>0.91</td>
</tr>
<tr>
<td>8</td>
<td>Real time size-resolved composition</td>
<td>0.9721</td>
<td>0.87</td>
</tr>
<tr>
<td>9</td>
<td>Average size-resolved composition</td>
<td>0.9742</td>
<td>0.86</td>
</tr>
</tbody>
</table>
FIGURE CAPTIONS

Fig. 1. A schematic representation of \( N_{CCN} \) prediction based on the H-TDMA and the AMS measurements. The \( N_{CCN} \) can be predicted based on the fitted activation ratio (approach I) and the \( D_{50} \) (approach II) both obtained from the H-TDMA measurement, the size-resolved composition (approach III) and the bulk PM\(_1\) composition (approach IV) both obtained from the AMS measurements. Panel (a) is the representation of calculating the activation ratio for a specific diameter and SS and the shadow area represents the particles which can be activated as CCN; (b) and (c) are the representations of the \( \kappa \) values obtained respectively from size-resolved chemical composition and bulk chemical composition; (d) is the representation of fitting the activation ratio to the particle diameter \( D_p \) (red dot); (e), (f), (g), and (h) are the representations of predicting the \( N_{CCN} \) using the four approaches respectively and the shadow area represents the particles which can be activated as CCN.

Fig. 2. The mass fraction of the bulk NR-PM\(_1\) composition (a) and the mass fraction of the size-resolved composition (b).

Fig. 3. The median and interquartile PNSD, \( \kappa \) obtained from H-TDMA, AMS and CCN measurement during the campaign. The \( \kappa \) was pointed against their corresponding median \( D_{50} \) (SMCA and AMS) or measured diameter (H-TDMA). Dot points represent the median value and the bars represent the interquartile range. The blue, red, and green represent \( \kappa_{CCN} \), \( \kappa_{AMS} \) and \( \kappa_{HTDMA} \) respectively.

Fig. 4. The sized resolved activation ratios measured by the SMCA at four different supersaturations.

Note that the curves were fitted according to the SMCA measurements.

Fig. 5. The Gf-PDF as a function of Gf measured by H-TDMA for the five diameters.

Fig. 6. The relationship between size-resolved mass fraction of organics and \( D_{30} \) at three
supersaturations. The red, blue, and green dots and line represent 0.1%, 0.2%, and 0.4% SS respectively.

Fig. 7. The predicted activation ratio based on H-TDMA measurement vs. the measured activation ratio at 0.1%, 0.2%, 0.4% and 0.7% SS for 40, 80, 110, 150 and 200 nm particles.

Fig. 8. The relative deviation between predicted AR and measured AR at different assumed \( \sigma_{sa} \); the color code represents \( R^2 \) between calculated AR and measured AR.

Fig. 9. The predicted activation ratio using new surface tension assumption (\( \sigma_{sa}^{*} \)) based on H-TDMA measurement vs. the measured activation ratio at 0.1%, 0.2%, 0.4% and 0.7% SS for 40, 80, 110, 150 and 200 nm particles.

Fig. 10. The relationship between measured \( N_{CCN} \) and predicted \( N_{CCN} \) based on scheme 1, 2, 3, 4 and 5. The black lines represent 1:1 lines.

Fig. 11. The relationship between measured \( N_{CCN} \) and predicted \( N_{CCN} \) based on scheme 6, 7, 8 and 9. The black lines represent 1:1 lines.

Fig. 12. The relationship between measured \( N_{CCN} \) and predicted \( N_{CCN} \) at SS 0.1% based on size-resolved chemical composition using \( \kappa^{*}_{org} \) (a) and \( \kappa^{*}_{eg} \) and \( \sigma_{sa}^{*} \) (b). The shadow area represents the variation of the linear fit using the \( \sigma_{sa} \) values from 0.054 to 0.062 N m\(^{-1}\).
Based on H-TMDA measurement

I. Activation Ratio
   - a. Size Dependence
   - d. Mixing State

II. $D_{50}$
   - √
   - ×

Based on AMS measurement

III. Size-resolved Composition
   - b. 
   - e. 
   - g. 

IV. Bulk PM$_2.5$ Composition
   - c. 
   - f. 
   - h. 

Fig. 1.
Fig. 2.
Fig. 3.
Fig. 4.
Fig. 5.
Fig. 6.
Fig. 7.
Fig. 8.
Fig. 9.
Fig. 10.
Fig. 11.
Fig. 12.