



CCN measurements at the Princess Elisabeth Antarctica Research Station during three austral summers

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Abstract.

For three austral summer seasons (2013-2016, each from December to February) aerosol particles arriving at the Belgian Antarctic research station Princess Elisabeth (PE), in Dronning Maud Land in East Antarctica were characterized in terms of number concentrations of total aerosol particles (N_{CN}) and cloud condensation nuclei (N_{CCN}), the particle number size distribution (PNSD), the aerosol particle hygroscopicity and the influence of the air mass origin on N_{CN} and N_{CCN} . In general N_{CN} was found to range from 40 to 6700 cm^{-3} with a median of 333 cm^{-3} , while N_{CCN} was found to cover a range between less than 10 and 1300 cm^{-3} for supersaturations (SS) between 0.1 and 0.7 %. It is shown that the aerosol is Aitken mode dominated and is characterized by a significant amount of freshly, secondarily formed aerosol particles, with 94 % and 36 % of the aerosol particles are smaller than 90 nm and ≈ 35 nm, respectively. Measurements of the basic meteorological parameters as well as the history of the air masses arriving at the measurement station indicate that the station is influenced by both, continental air masses originating from the Antarctic inland ice sheet (continental events - CE) and marine air masses originating from the Southern Ocean (marine events- ME). CEs came along with rather constant N_{CN} and N_{CCN} values, which we denote to be Antarctic continental background concentrations. MEs however cause large fluctuations in N_{CN} and N_{CCN} caused by scavenging due to precipitation or new particle formation based on marine precursors. The application of HYSPLIT back trajectories in form of the potential source contribution function (PSCF) analysis indicate, that the region of the Southern Ocean is a potential source of Aitken mode particles. For particles larger than ≈ 110 nm (CCN measured at SS of 0.1 %) the Antarctic ice shelf regions were found to be a potential source region, most likely due to the emission of sea salt aerosol particles, released from snow particles from surface snow layers by sublimation, e.g., during periods of high wind speed, leading to drifting or blowing snow. On the basis of the PNSDs and N_{CCN} , the critical diameter for cloud droplet activation and the aerosol particle hygroscopicity parameter κ were determined to be 110 nm and 1, respectively, for a SS of 0.1 %. The region of the Antarctic inland plateau however was not found to feature a significant source region for CN and CCN measured at the PE station in austral summer.



1 Introduction

Aerosol particles can be emitted into the atmosphere either directly e.g., by mechanical processes or combustion, or indirectly, due to nucleation from the gas-phase. Under specific conditions, aerosol particles can act as cloud condensation nuclei (CCN) and form cloud droplets. Whether a particle forms a cloud droplet depends on the chemical composition and the size of the particle as well as the surrounding supersaturation (SS). Aerosol particles can influence the climate either directly, by scattering or absorption of solar radiation, or indirectly due to their impact on cloud formation and cloud properties as e.g., the cloud albedo (Twomey, 1974) or on the lifetime of clouds (Albrecht, 1989; Rosenfeld et al., 2008). The direct effect of aerosol particles is well understood. In contrast to this, the manifold indirect aerosol effects are less understood. The influence of indirect aerosol effects on the global climate and the radiative forcing still features a low confidence level and large uncertainties (IPCC, 2013). The impact of aerosol particles on climate and global radiative forcing is mainly determined by their physical and chemical properties. Investigations of these properties of aerosol particles in general and of CCN in particular by means of in situ measurements at various different sites and conditions are a necessary to lower these uncertainties. The Antarctic region is particularly interesting for aerosol particle and CCN in situ studies for two reasons. Firstly, Antarctica is located far from anthropogenic activities and is one of the most pristine areas on the globe (Hamilton et al., 2014). Thus, it is a favorable environment for studying natural aerosol particle background conditions and processes that prevailed in a preindustrial atmosphere. A more accurate knowledge about preindustrial aerosol processes, conditions and properties is important for a reduction of uncertainties of model estimates concerning radiative forcing (Hamilton et al., 2014; Carslaw et al., 2013). Secondly, similar to the Arctic, the Antarctic region is extremely sensitive to climate change. Jacka and Budd (1998) analyzed surface temperature data of 16 stations on the Antarctic continent and 22 stations on Southern Ocean island and found warming rates of 0.9-1.2 °C and 0.7-1 °C, respectively. In particular in the West Antarctic and the Antarctic Peninsula the warming is several times higher than in other regions (Jacka and Budd, 1998; Vaughan et al., 2003; Kravchenko et al., 2011; IPCC, 2013). The Antarctic sea ice as well as the inland ice sheet are potentially subject to change in such a changing environment. However, at the moment, the southern hemisphere has not shown a decrease in sea ice extent. Cavalieri and Parkinson (2008) and Parkinson and Cavalieri (2012) even found a positive trend in the annual maximum Southern Ocean sea ice extend. Both, the sea ice area and the open water area have the potential to emit aerosol particles into the atmosphere. Sea ice is a potential source for sea salt aerosol particles (Huang and Jaeglé, 2017; Yang et al., 2008; Wagenbach et al., 1998) and nitrogen (Dall'Osto et al., 2017) and open sea water may emit sea spray aerosol and precursors for new particle formation (Liss and Lovelock, 2008; Modini et al., 2015). Therefore, variations in sea ice coverage will likely lead to changes in the nature of aerosol particle sources. The mass balance of the Antarctic ice sheet also shows quite unexpected trends. Velicogna and Wahr (2006) and Shepherd et al. (2012) found that the ice sheets of West Antarctica and the Antarctic Peninsula had lost mass, whereas the East Antarctic ice sheet had gained mass. The gain of ice mass in East Antarctica is also confirmed by Martin-Español et al. (2017), however they found it to be smaller than losses in West Antarctica. As precipitation, which is linked to the abundance of CCN, is the only source of mass gain to the Antarctic ice sheet, it is necessary to study the properties of these aerosol particles as well as their impact on cloud formation and precipitation, their sources, sinks and pathways in the changing environment of Antarctica.



Although Antarctica is a harsh environment where access for field work is difficult, various aerosol particle studies have been conducted at different Antarctic research stations during the last decades. A wide range of topics has already been investigated, including new particle formation, seasonal cycles of number and mass concentrations as well as size distributions, chemical composition, hygroscopicity and optical properties of aerosol particles. Furthermore, some studies have reported on Antarctic CCN properties, however the locations they cover are limited to the Antarctic Peninsula (DeFelice, 1996; DeFelice et al., 1997; Kim et al., 2017) or the area of the Weddell Sea on the Brunt Ice Shelf (O’Shea et al., 2017). Both locations are part of West Antarctica and especially the Antarctic Peninsula is mainly influenced by marine air masses that directly originate from the Southern Ocean. To achieve a more detailed picture of Antarctic CCN, further measurements that can be used to characterize CCN in the eastern and especially in the central part (Antarctic inland plateau) of Antarctica, are needed.

To gain further knowledge about aerosol particle and particularly CCN properties in East Antarctica, we conducted measurements at the Belgian Antarctic research station Princess Elisabeth (PE), in Dronning Maud Land. For three austral summer seasons (2013-2016, always from December to February) a Condensation Particle Counter (CPC), a Cloud Condensation Nucleus counter (CCNc) and a Laser Aerosol Spectrometer (LAS) were used to measure simultaneously aerosol particle and CCN properties inside the East Antarctic boundary layer. In addition, the present study introduces meteorological data, collected by an automatic weather station and precipitation rates derived from a vertically pointing precipitation radar, as well as the history of the measured air masses, calculated by means of the Numerical Atmospheric-dispersion Modelling Environment (NAME) and the NOAA HYSPLIT trajectory model. This data set has enabled the study of the variability of the Condensation nuclei (CN) and CCN number concentrations, to identify their sources, sinks and transport pathways and to analyze the particle hygroscopicity during austral summer in East Antarctica. The special location of the PE station in the escarpment zone with katabatic winds coming from the Antarctic inland ice sheet, further allows an insight into the state of aerosol particle and CCN properties of continental Antarctica.

2 Experimental procedure and methods

2.1 Measuring site and meteorology

The measurements presented in this study were all performed at the Belgian Antarctic research station Princess Elisabeth (PE, Figure 1a), in Dronning Maud Land, East Antarctica (71.95° S, 23.35° E, 1390 m asl., around 200 km inland from the Antarctic coast). The PE station is located upon the granite ridge of Utsteinen Nunatak in the Dronning Maud Land region of East Antarctica and lies north of the Sør Rondane Mountain Range, that has peaks up to an altitude of 3300 m asl. This area is located in the escarpment zone between the Antarctic inland plateau and the coast which can be seen in the topographic map of Antarctica in Figure 1b. A more detailed description of the conditions at the measurement station and its near surroundings is given by Pattyn et al. (2010) and Gorodetskaya et al. (2013). The PE station is designed as a zero emission station with power production mainly based on wind and solar energy (see www.antarcticstation.org). This reduces local emissions which makes the PE station an excellent base for conducting in situ aerosol particle measurements. Nevertheless, general station activities, traffic by skidoos or bulldozers, and irregular diesel generator operation times cause contamination which is however discarded



in the final data (see section 2.2). The station is inhabited from November through the end of February. During the other month the station and most of its scientific instruments are operated under remote control. As the Cloud Condensation Nuclei counter used for this study needs an operator on site, we mainly present data collected from December to February during three subsequent austral summers (2013-2016).

5 The basic meteorological parameters (temperature, pressure, relative humidity, wind speed and wind direction) were measured by means of an automatic weather station (AWS, Gorodetskaya et al., 2013; Souverijns et al., 2017b) which was located 300 m east of PE. Precipitation rates were derived from Metek's micro-rain radar (MRR-2) calibrated with an optical disdrometer (Precipitation Imaging Package) (Souverijns et al., 2017). The radar is a vertically pointing 24-GHz precipitation radar operating at PE since January 2010 (Gorodetskaya et al 2015). Effective reflectivities (Z) were estimated from the radar's raw
10 Doppler spectra using the algorithm of Maahn and Kollias (2012) specifically designed for MRR snowfall applications. Snowfall rates (S) at 300m agl level were estimated using a mean Z - S relationship derived by Souverijns et al (2017) based on snowfall microphysical measurements at PE. More details about the AWS, the precipitation radar and the estimation of precipitation can be found in Gorodetskaya et al. (2013), Gorodetskaya et al. (2015) and Souverijns et al. (2017), respectively
15 by either synoptic regimes, which usually correspond to strong easterly winds with sometimes a slight northerly component, or a katabatic regime, that is mostly associated with relatively weak south-southeasterly winds (Gorodetskaya et al., 2013; Souverijns et al., 2017b).

The two different meteorological situations can be identified based on wind speed and wind direction, which are both depicted in the form of a wind rose in Figure 1c (exemplarily based on measurements between December 18, 2015 to February
20 20, 2016 in the third season). The more frequent easterly winds clearly correspond to higher wind speeds, mainly over 5 m/s, whereas the less frequently occurring southerly winds are usually below 5 m/s. Additional meteorological parameters for each year are shown in Figure 2 as a time series of hourly (gray lines) and daily averaged (red lines) values. Respective seasonal mean values together with the standard deviation, minimum and maximum for the period from December 1 to February 20 are shown in Table 1. The mean values as well as the fluctuation in the meteorological parameters show no large differences
25 between the three measurement periods. Due to the shielding effect of the Sør Rondane Mountains to the South, the climate at PE is relatively mild for an Antarctic site (Gorodetskaya et al., 2013).

2.2 Instrumentation and data processing

The total particle number concentration (N_{CN}) was measured by a Condensation Particle Counter (CPC, TSI model 3776), which has a lower cut off at 3 nm and was operated at a total flow rate of 1.5 l/min. The CPC was first installed for continuous
30 operation in November 2012. Due to several power outages in austral winter, data coverage of the winter months was not equal between the years and the CPC was restarted in the respective austral summers. The last data for this study was measured in May 2016. The inlet tubing for the CPC consisted of a 1 m long vertical 0.5 in stainless steel tubing (not heated) installed tightly through the roof of the measurement container. Inside, 0.7 m of a 3/8 in (0.19 in inner diameter) conductive flexible tubing made the connection to the CPC with only a smooth bend just before the inlet at the front of the CPC. On the roof of



the container, 0.15 m of the same flexible tubing was connected to the stainless steel tubing in order to serve as inlet without a size cut-off. With this kind of inlet, there were never issues with inlets clogged by snow during storms, in particular important during the non-inhabited winter periods. Clogging of inlets caused by riming never happened due to the extreme dryness at the measurement site. The CPC was operated with a 4 l butanol reservoir bottle. Consumption of butanol was between 3 to 3.5 l for a complete year of measurements. Each austral summer the CPC was checked on leaks and the butanol was exchanged. The procedure to assure non-contaminated data is described further below.

In parallel to N_{CN} the particle number size distribution ($PNSD$) was measured by means of a Laser Aerosol Spectrometer (LAS, TSI model 3340) in the size range from 90 nm up to 6.8 μm (99 log-distributed channels). The inlet setup and tubing for the LAS is similar to the one for the CPC and installed directly next to that one. However, inside, first 0.5 m of a 1/8 in (inner diameter) and then 0.2 m of a 1/16 in (inner diameter) conductive flexible tubing made the connection (no bend) to the measurement chamber of the LAS. The LAS was operated with a sample flow rate of 0.07 l/min with a sheath flow of 0.6 l/min. The LAS was maintained and re-calibrated in spring 2015 by TSI Inc. In October 2015, before shipment to Antarctica, the LAS was compared to an SMPS system (DMA type Medium Hauke; CPC, TSI model 3010) at the cloud laboratory of the Leibniz Institute for Tropospheric Research (TROPOS). 18 selected sizes (80, 90, 100, 110, 120, 130, 140, 150, 175, 200, 250, 300, 350, 400, 500, 600, 700, 800 nm) of ammonium sulphate particles and 4 sizes (100, 500, 700, 800 nm) of PSLTM standard solutions were used to validate the counts of the LAS. Necessary corrections were high in the two LAS channels below 100 nm (around +70 %), distinct in the two channels around 100 nm (+ 10 %) and low in the other size ranges up to 800 nm (between 1 to 5 %, negative and positive corrections). These corrections were applied to the LAS data set used for this study. In this study, we continuously use hourly averaged N_{CN} values and $PNSDs$.

The number concentration of cloud condensation nuclei (N_{CCN}) was measured using a Cloud Condensation Nuclei counter (CCNc, Droplet Measurement Technologies (DMT), Boulder, USA). The CCNc is a continuous-flow thermal-gradient diffusion chamber which is described in detail by Roberts and Nenes (2005). The inlet tubing for the CCNc consisted of a 2.2 m long vertical conductive flexible tubing (similar to the one used for the CPC and LAS) with only a smooth bend just before the inlet of the CCNc. The inlet outside was directly next to the inlets of the CPC and LAS. The CCNc was operated as recommended by Gysel and Stratmann (2013) for polydisperse CCN measurements. The CCNc was operated at a constant total flow rate of 0.5 l/min and at 5 different supersaturations (SS ; 0.1, 0.2, 0.3, 0.5 and 0.7 %), each for 12 minutes per hour. To ensure stable column temperatures, the first 5 min and the last 30 sec at each SS setting were excluded from the data analysis. The remaining data points were averaged, so that the result is one N_{CCN} value per SS per hour. For consistency checks between N_{CN} and N_{CCN} also 1 % SS was adjusted at times (but data at that SS were not included in the analysis presented in here). Prior to each of the three measurement periods in Antarctica a SS calibration of the CCNc was done at the cloud laboratory of the Leibniz Institute for Tropospheric Research (TROPOS). These calibrations were performed with size selected ammonium sulfate particles and for pressure conditions relevant for the PE station (approximately 820 hPa), based on recommendations given by Gysel and Stratmann (2013) and Rose et al. (2008). All values (N_{CN} , N_{CCN} and $PNSD$) are presented with respect to standard conditions, i.e., a pressure of 1013.25 hPa and a temperature of 293.15 K.

The container for the aerosol measurements is located 60 m south of the PE main station (Figure 1a). Given that the PE station



is designed as a zero-emission station, the daily activities are concentrated in the W-NW sector, and the distribution of the wind direction (Figure 1c), the container was most often exposed to non-contaminated air. The container is well-insulated and equipped with a small heater. Heating was hardly necessary in austral summer (due to 24 hour sun light). However, there is no air conditioning system (due to energy demand; remote control during austral winter; necessary filter systems and no exchange possible of them during austral winter). Therefore, in austral summer, the temperature inside the container varied between \approx 10 and 40 °C. This range exceeds the recommended operating temperature range of the CPC and the LAS of 10 to 35 °C and 10 to 30 °C, respectively, as well as the temperature range for which the CCNc was calibrated, which is 20 to 30 °C. Therefore, N_{CN} , N_{CCN} and $PNSDs$ measured during time periods in which the temperature inside the measurement container was outside of the operating temperature ranges were excluded from the analysis presented herein.

Further, as mentioned in section 2.1, the data still contained values caused by emissions from the activities at the station. In order to identify hourly intervals with contamination, the following data sets were examined: i) the minute-CPC data; ii) simultaneously measured hourly data for the mass concentration of light-absorbing aerosol (Magee Sci. AE31, 7-wavelength aethalometer; set up in the same measurement container); iii) wind speed and wind direction measured by the AWS. As indicators for contamination abrupt peaks, outliers, and strong variations between higher and lower minute-CPC values and/or distinctly higher mass concentrations of light-absorbing aerosol ($>50 \text{ ng/m}^3$) were used. Because the PE station was designed as zero-emission station, there was no relationship between wind speed or wind direction with elevated values for N_{CN} or light-absorbing aerosol. However, each hourly interval with wind speed $< 3 \text{ m/s}$ and/or wind direction outside the sector 20° to 225° was examined again for conspicuous signals in its variation in time.

The hygroscopicity of the aerosol particles was determined by applying the κ -Köhler-theory (Petters and Kreidenweis, 2007). The inferred hygroscopicity parameter κ represents the average particle chemistry. To infer κ , first the critical diameter (d_{crit}) needs to be determined, based on the measured N_{CCN} and $PNSD$. d_{crit} is the diameter at which particles are just large enough to be activated to a droplet when exposed to a certain supersaturation (SS). For a pair of simultaneously measured $PNSD$ and N_{CCN} , d_{crit} is obtained by calculating the cumulative particle number concentration from that $PNSD$, from the largest diameter on downward, and it is the diameter at which this cumulative concentrations is equal to N_{CCN} . Using the assumption that the surface tension is equal to that of pure water, d_{crit} and the SS are used to derive κ values of the Antarctic aerosol particles. A detailed description of this method, including the application of a Monte Carlo simulation to precisely determine uncertainties in d_{crit} and κ , is presented by Herenz et al. (2017). This procedure of inferring κ values could only be done for N_{CCN} measurements at a SS of 0.1 %, as d_{crit} for larger SS is below the lower size limit of the $PNSDs$ of 90 nm.

2.3 Identification of air mass origins and potential source regions

To analyze the influence of the air mass origin on N_{CN} and N_{CCN} measured at the PE station, we applied two different models to obtain information on the air mass history. The first one is the Numerical Atmospheric-dispersion Modelling Environment (NAME), which was used to perform a simple residence time analysis (Fleming et al., 2012). The second one is the Potential Source Contribution Function (PSCF), a more advanced type of residence time analysis that results in a probability field (Fleming et al., 2012).



The NAME dispersion model

The NAME atmospheric dispersion model (Jones et al., 2007) is a Lagrangian particle-trajectory model, that is operated by the UK Meteorological Office. For this study 10000 abstract particles per hour were released at an altitude of 10 m above the location of the PE station. On the basis of the Meteorological Office Unified Model (UM) meteorological field data, 10-day back trajectories for these particles were calculated. Summing up the concentration of these particles at time steps of 15 minutes back in time (in total 960 time steps) results in a footprint that shows the history of the air masses during the last 10 days. For this procedure only particles that are located within the surface layer (0-100 m) are taken into account. An example footprint of the first of December 2013 (midnight) is shown in Figure 3. Footprints were derived every three hours, resulting in a total number of 2019 NAME footprints used in this study. To further analyze the impact of different surface properties on the measured aerosol particle properties, the area around Antarctica was divided into the following 5 different regions (see also Figure 4): the Antarctic escarpment zone and inland plateau (continental area at or above 200 m above sea surface level (asl)), Southern Ocean, South America, Africa and Reactive Zone. The Reactive Zone includes the following regions that are known to have the potential to emit either primary particles or precursors for secondarily formed particles:

- The Antarctic continental area below 200 m asl and 8 islands in the Southern Ocean (South Georgia, South Sandwich, Falkland, South Orkney, Prince Edward, Crozet, Kerguelen, Heard and McDonald Island). These regions are included because they are habitats for numerous different types of penguins. Bird guano (Croft et al., 2016) and in this special case penguin guano (Legrand et al., 1998), acts as a source of ammonia and may contribute to the formation of new particles in coastal Antarctic areas.
- The permanent and seasonally covered sea ice areas. These are known to have the potential to act as source of organic nitrogen that contributes to secondarily formed aerosol particles (Dall’Osto et al., 2017) or to emit primary sea salt particles (Huang and Jaeglé, 2017; Yang et al., 2008; Wagenbach et al., 1998).
- The marine area up to 200 km from the coasts of the islands and continents (for Antarctic, continent plus ice shelves). These areas are included due to an enhanced chlorophyll concentration in the coastal areas of the Southern Ocean. Chlorophyll can be used as a proxy for dimethyl sulfide (DMS) (Vallina et al., 2006), which in turn plays a role in new particle formation (Liss and Lovelock, 2008).

The proportional residence time that the air masses spent over the 5 different regions during the last 10 days was determined in order to assess to what extent these regions influence the aerosol particle and CCN properties. This type of a residence time analysis was already used for an Antarctic site by O’Shea et al. (2017).

Potential Source Contribution Function

The Potential Source Contribution Function (PSCF) is a receptor modeling method that originally was developed by Ashbaugh et al. (1985) and was applied in a number of high latitude studies before e.g., in Dall’Osto et al. (2017) for the Antarctic and



in Yli-Tuomi et al. (2003) for the Arctic. The PSCF model is based on air mass back trajectories, and it is commonly used to identify regions that have the potential to contribute to high values of measured concentrations at a receptor site. In this study we apply the PSCF on N_{CN} and N_{CCN} .

The NOAA HYSPLIT trajectory model (Stein et al., 2015) was used to calculate hourly resolved 10-day back trajectories based on $1 \times 1^\circ$ GDAS (Global Data Assimilation System) meteorological data. To account for uncertainties in back trajectory analysis, every hour a set of 15 back trajectories was calculated, which is composed of 5 different plane locations (one exactly at the measurement station and 4 in close proximity around it) at three altitudes (100 m, 200 m and 300 m above the surface level). In total this results in a set of 88152 back trajectories that were used for the PSCF analysis. Each back trajectory consists of trajectory segment endpoints, which represent the central geographical position of the air parcel at a particular time. To calculate the PSCF the whole region that is covered by these trajectory segment endpoints is divided into an array of $5 \times 5^\circ$ grid cells (i,j). The assumption is that aerosol particles that are emitted in such a cell are incorporated into the air parcel and transported to the receptor cite. The PSCF can be calculated as follows:

$$PSCF_{i,j} = \frac{m_{i,j}}{n_{i,j}}, \quad (1)$$

where $n_{i,j}$ is the total number of trajectory segment endpoints that fall into a cell and $m_{i,j}$ is the number of trajectory segment endpoints that fall into that cell and fulfill a given criterion, where this criterion typically is the exceedance of a certain threshold. In this study we used the 75 % percentile of either N_{CN} or N_{CCN} as that threshold. According to Hopke (2016): “Cells containing emission sources would be identified with conditional probabilities close to 1 if trajectories that have crossed the cells effectively transport the emitted contaminant to the receptor site. The PSCF model thus provides a means to map the source potentials of geographical areas. It does not apportion the contribution of the identified source area to the measured receptor data.” As it is probable that small values of $n_{i,j}$ would lead to uncertain and high PSCF values it is necessary to apply a weighting function. For this study a discrete weighting function based on $\log(n+1)$, which is a measure of the back trajectory density, was applied (Waked et al., 2014):

$$W = \begin{cases} 1.00 & \text{for } n_{i,j} \geq 0.85 \cdot \max(\log(n+1)) \\ 0.725 & \text{for } 0.6 \cdot \max(\log(n+1)) > n_{i,j} \geq 0.85 \cdot \max(\log(n+1)) \\ 0.35 & \text{for } 0.35 \cdot \max(\log(n+1)) > n_{i,j} \geq 0.6 \cdot \max(\log(n+1)) \\ 0.1 & \text{for } 0.35 \cdot \max(\log(n+1)) > n_{i,j} \end{cases} \quad (2)$$

The measured concentration of total particles and CCN is also affected by losses that occur along the path of the air parcel between the source and the receptor site. As precipitation, which is known to be one of the major sinks for aerosol particles, in particular for CCN, is an output parameter of the calculated NOAA HYSPLIT back trajectories, it can be taken into account. Hence, we run the PSCF model with a precipitation filter. Back trajectories were cut off, and not considered for the PSCF analysis, as soon as a trajectory segment endpoint shows a precipitation of 0.1 mm/h and the total precipitation (sum of



precipitation of 240 segment endpoints) of this back trajectory exceeds a value of 5 mm/240h. The second criterion was added as it seemed not to be reasonable to discard a trajectory only because of showing a low precipitation of some mm/h at some segment endpoints. Note, the filter criteria as well as the criteria of the weighting function are empirical.

3 Results and discussion

5 3.1 Total Particle and CCN number concentrations and regional analysis of the NAME model footprints

This section presents the measured N_{CN} , N_{CCN} and $PNSD$ as well as the proportional residence time of the air masses over the regions introduced in Section 2.3. Time series are given for the three austral summer seasons of 2013/2014, 2014/2015 and 2015/2016 in Figure 6, 7 and 8, respectively.

Measurements of N_{CN} throughout the whole year were performed between 2012 and 2016. Their visualization in Figure 5 shows a clear seasonal cycle with the lowest monthly median values during the austral winter and a maximum during late austral summer. The monthly 10 % and 90 % percentiles also indicate the highest variability of N_{CN} during February. Several studies at different Antarctic sites found that the physical and chemical aerosol particle properties are subject to a similar seasonality e.g., Hara et al. (2011), Weller et al. (2011), Virkkula et al. (2009) and Kim et al. (2017). Just like N_{CN} , also N_{CCN} follows a seasonal cycle with a minimum in austral winter and a maximum in austral summer (Kim et al., 2017). Hence, our measurements during austral summer capture the season in which the aerosol production in Antarctica and the surrounding source regions is most active.

We found N_{CN} (black dots in panel C Figure 6, 7 and 8) to cover a range between 40 and 6700 cm^{-3} (on the base of hourly averaged values) with a median value of 333 cm^{-3} . Our measured N_{CCN} (bluish dots in panel C Figure 6, 7 and 8) cover a range between less than 10 cm^{-3} at $SS=0.1\%$ to 1300 cm^{-3} for the largest SS of 0.7 %. The integration of the $PNSD$ over the whole size range ($N_{CN>90\text{nm}}$, red dots in panel C Figure 6, 7 and 8) shows the aerosol particle number concentration in the size range between 90 nm and 6.8 μm . $N_{CN>90\text{nm}}$ has a median value of 20 cm^{-3} . The median, 10 and 90 % percentile values for N_{CN} , $N_{CN>90\text{nm}}$ and N_{CCN} at all measured SS are summarized in the first column of Table 2. The third column of Table 2 shows the fraction of $N_{CN>90\text{nm}}$ and N_{CCN} at different SS to N_{CN} (based on the median values of the first column in Table 2). The values of $N_{CN>90\text{nm}}/N_{CN}$ and $N_{CCN,0.7\%}/N_{CN}$ are 0.06 and 0.64, respectively. This indicates, that the aerosol particles feature an Aitken mode dominance, as 94 % of the aerosol particles are smaller than 90 nm. Assuming a hygroscopicity parameter κ of 0.8 for the coastal area of East Antarctica, taken from Pringle et al. (2010), the critical diameter d_{crit} for $SS = 0.7\%$ was determined by means of the κ -Köhler theory to be ≈ 35 nm. On the basis of this assumption 36 % the aerosol particles are smaller than roughly 35 nm. That is indicative for a high amount of freshly, secondarily formed aerosol particles, which form from precursor gases emitted from the Reactive Zone as e.g., ammonia and DMS (see Section 2.3). Primary emitted natural aerosol particles that are known to occur in Antarctica from e.g., mineral dust (Wegner et al., 2015) or sea salt (Huang and Jaeglé, 2017; Yang et al., 2008; Wagenbach et al., 1998), are known to clearly exceed this size (Lamb and Verlinde, 2011). Unfortunately, we can not examine the Aitken mode particles in much more detail, as our $PNSD$ data is in the size range between 90 nm and 6.8 μm and hence only shows the accumulation and coarse mode particles. However, several



other studies at coastal Antarctic sites report *PNSD* measurements that show pronounced and dominant Aitken modes during austral summer (e.g., Asmi et al., 2010; O'Shea et al., 2017; Kim et al., 2017).

Panel D of Figure 6, 7 and 8 shows the regional analysis of the NAME footprints, as described in Section 2.3. It can give insights on the influence of the air mass origin on N_{CN} and N_{CCN} . The regional analysis shows, that during the 10 days prior to the measurements, air masses only have been influenced by the Antarctic continent, the Southern Ocean and the Reactive Zone region but not by South America or Africa. Thus, we can be confident, that we mainly measured pristine air masses and aerosol particles of a natural origin without much anthropogenic influence.

The contributions from Antarctica, the Southern Ocean and the Reactive Zone region show a large variability. During 61 % of the measurement times, the air masses spent ≥ 90 % of the 10 days prior to their arrival at the measurement site over the continental region. These times are called Continental Events (CE) from now on. During CEs, we record only a low variability in the measured N_{CN} and N_{CCN} . To illustrate this, Figure 9 shows a scatter plot of N_{CN} and the fraction of time that the respective air masses spent over the Antarctic region (continental fraction). All data from the three seasons are included. During CEs, N_{CN} rarely exceeds 475 cm^{-3} , with maximum values of 990 cm^{-3} , while 90 % of N_{CN} (i.e., the 90 % percentile) cover a range from 170 to 475 cm^{-3} . The concentration ranges during CEs for N_{CN} , $N_{CN>90nm}$ and N_{CCN} at all *SS* are shown in the second column of Table 2. These concentration ranges can be assumed to be pristine Antarctic continental background concentrations during austral summer.

Vice versa, during 39 % of the time the proportion of the Reactive Zone plus the Southern Ocean region was larger than 10 %, which we from now on call Marine Events (ME). During MEs we record an enhanced variability in N_{CN} and N_{CCN} . Also the precipitation, depicted in panel A of Figure 6, 7 and 8, shows a connection to MEs. Especially strong precipitation events only occur during certain most intense MEs affecting PE, e.g., on December 2 in 2013, January 18 in 2015 and January 30 in 2016, in line with the findings of Gorodetskaya et al. (2014) and Souverijns et al. (2017b). These precipitation events significantly lower N_{CN} and N_{CCN} on a time scale of some hours to one day, due to scavenging and wet deposition. The minimum values that we report for N_{CN} and N_{CCN} were measured during these strong precipitation events. As the Antarctic region does not act as a significant source for water vapor (see katabatic meteorological regime in Section 2.1) it is self-explanatory that strong precipitation events only occur during MEs. But also the largest values for N_{CN} and N_{CCN} are only observed during MEs, as the Reactive Zone and the Southern Ocean region potentially represent source regions for primary and secondary formed aerosol particles. As already mentioned in Section 2.3 the region of the Reactive Zone can contribute to the Antarctic aerosol particle loading due to sea bird and penguin guano and the release of ammonia that potentially contributes to the formation of new particles (Croft et al., 2016; Legrand et al., 1998). The Reactive Zone and the Southern Ocean region also emit precursors for secondary aerosol particle formation as e.g., DMS, whose oxidation products sulfuric and methane sulfonic acid have the ability to form aerosol particles that grow to CCN sizes (Liss and Lovelock, 2008). Also, these regions have the potential to contribute to the aerosol particle loading by primary emissions of sea salt particles due to blowing snow on sea ice surfaces (Huang and Jaeglé, 2017; Yang et al., 2008; Wagenbach et al., 1998) or bubble bursting from wave action (Lamb and Verlinde, 2011).

The time series of N_{CN} in Figure 6, 7 and 8 often show a spontaneous increase during MEs of several thousand particles



per cm^3 . Figure 10 exemplarily shows such an event, that took place on December 6 in 2014. Between 7 and 10 a.m. N_{CN} increased from ≈ 200 to $\approx 6000\text{ cm}^{-3}$. This was accompanied by an increase of N_{CCN} . In total we detected 12 comparable events with an increase of N_{CN} up to several thousand cm^{-3} that all took place in a time frame between several hours and ≈ 1 day. In the vast majority these events of increased N_{CN} were followed by an increase of N_{CCN} by a factor of roughly two, in a few cases even up to 15. This only holds for N_{CCN} measured at SS between 0.2 %-0.7 %. N_{CN} measured at a SS of 0.1 % usually shows a different trend and seems to be decoupled from the other measurements. Other studies at Antarctic sites report events of new particle formation (NPF) during austral summer, e.g., Asmi et al. (2010) and Weller et al. (2015) at the Finnish research station Aboa and the German Neumayer station, respectively, which are both coastal sites or Järvinen et al. (2013) at Dome C, a site in Central Antarctica. Growth rates of particles from NPF reported in these studies were $\approx 2.5\text{ nm/h}$ at Dome C and $\approx 1\text{ nm/h}$ for the two coastal sites. However, it was also described that particles rarely grow to sizes larger than $\approx 25\text{ nm}$, i.e., that they do not reach sizes at which they can readily act as CCN.

Our measured *PNSDs* do not cover the size range of the nucleation and Aitken mode, however, in combination with measurements of N_{CN} and N_{CCN} we can state that comparably freshly formed particles originating from new particle formation events were observed during our measurements. Our measurements at the PE station show, that these freshly formed aerosol particles seem to reach size ranges relevant for CCN activation and thus are climatically relevant.

3.2 Air mass origins and potential source regions

Figure 11 shows the spatial distribution of the PSCF calculated for N_{CN} , $N_{CN}-N_{CCN,0.7\%}$, $N_{CCN,0.7\%}$ and $N_{CCN,0.1\%}$. These four parameters represent concentrations of all particles (with a lower size cut at 3 nm), particles in the size range up to $\approx 35\text{ nm}$, particles with sizes above $\approx 35\text{ nm}$ and the largest particles above $\approx 110\text{ nm}$, respectively. The analysis was done using the data of all three austral summer periods, which is a data set of approximately 230 days and a corresponding set of 88152 back trajectories. The 75 % percentile values of N_{CN} , $N_{CN}-N_{CCN,0.7\%}$, $N_{CCN,0.7\%}$ and $N_{CCN,0.1\%}$, on the basis of which the PSCF analysis was done, are 466, 184, 268 and 13, respectively. High values in the maps in Figure 11 indicate, which regions have a high potential to contribute to the 25 % of the highest number concentrations measured at the receptor site. The PSCF of N_{CN} shows enhanced values over the region of the Southern Ocean, mostly between 60° S and 40° S , but not over the Antarctic continental region. Hence, the marine region of the Southern Ocean is likely to be the dominant source region leading to an enhancement in N_{CN} measured at PE, while the Antarctic continent itself is not likely to act as a particle source. This is in accordance with a result discussed in Section 3.1, i.e., the low variability of measured number concentrations during CEs.

$N_{CN}-N_{CCN,0.7\%}$ and $N_{CCN,0.7\%}$ are two complementary parameters, adding up to N_{CN} . The PSCF maps of $N_{CCN,0.7\%}$ and $N_{CN}-N_{CCN,0.7\%}$ show clearly distinct patterns, indicating that different source regions are likely to contribute to high concentrations of particles with sizes below and above $\approx 35\text{ nm}$. However, both share that their highest signals are again in the Southern Ocean between 60° S and 40° S , however, at different longitudes. The PSCF of $N_{CN}-N_{CCN,0.7\%}$ shows a large area of high signals between 40° W and 60° E . When calculating transport times based on air mass back trajectories, an average



transport time of 5.1 days from this area to PE station is obtained. The PSCF of $N_{CCN,0.7\%}$ shows the largest area of high signals in a region between 140° W and 80° W, for which the average transport time to the PE station is 8.8 days. These air masses usually travel either along the west wind drift through the Drake Passage and circumnavigate Antarctica before making landfall close to PE station, or they travel along the easterly winds over coastal East Antarctica till they reach the PE station. As already discussed in Section 3.1, the aerosol observed at the PE station features a dominant Aitken mode. This can be brought in line with the results discussed here. The aerosol particles that originate from the marine areas that show up dominantly in the PSCF likely are mainly secondary aerosol particles that grow during the transport to the PE station. The size of the measured aerosol particles can be assumed to be a function of average transport time, corresponding to source regions for larger particles that are further away (when considering air mass traveling times).

The PSCF map for $N_{CCN,0.1\%}$ differs from the others. Overall, values are lower, pointing towards a more uniformly distributed origin of particles with sizes above ≈ 110 nm. But it should also be stressed that values for $N_{CCN,0.1\%}$ are generally low (see Table 2). The PSCF map shows almost no areas of enhanced values over the Southern Ocean, but several spots of comparably enhanced values show up along the coast of Antarctica, i.e., over the Reactive Zone region. The congruence between these spots and the different shelf ice areas that are shown in Panel d of Figure 1 is striking. The PSCF shows significantly increased values at the locations of the Ross, Ronne-Filchner and Amery shelf ice (1, 2 and 6 in Panel d of Figure 1), and slightly increased values at the location of the Fimbul, West and Shackleton shelf ice (5, 7 and 8). Hence, the Antarctic shelf ice regions seem to be potential source regions for enhanced values of $N_{CCN,0.1\%}$. We will elaborate on that further in the next section.

3.3 Hygroscopicity

For the data set presented here, the hygroscopicity parameter κ can only be inferred for $SS=0.1\%$, for which the median d_{crit} was determined to be 110 nm. For higher SS , N_{CCN} is above $N_{CN>90nm}$, i.e., d_{crit} is below the lower size limit of the measured $PNSD$. Therefore, the hygroscopicity derived here is only valid for the low number of comparably large particles that are activated at 0.1% (see Table 2). All κ values from the three seasons have a median value of 1 and are shown in a histogram in Figure 12. These are generally high atmospheric κ values covering a broad range between 0.5 and 1.6. Large κ values as those observed here typically are only found for particles consisting of inorganic substances (Petters and Kreidenweis, 2007). Particularly values of roughly 1 or above are only known to occur for sea salt. 0.95 was reported in Wex et al. (2010) as the mean value for the sea spray signal in marine air masses, derived from a collection of ambient hygroscopic growth measurements. Zieger et al. (2017) give a value of 1.1 for inorganic sea salt particles at 90% relative humidity, and Petters and Kreidenweis (2007) give mean values of 1.12 and 1.28 for NaCl, based on hygroscopic growth and CCN measurements, respectively. (It may be worth noting that κ derived from hygroscopic growth typically is below that derived from CCN measurements, see Petters and Kreidenweis (2007).) The lower values we derived for κ are too low to originate from pure sea salt particles. But besides for inorganic compounds, marine aerosol particles may also contain internally mixed organic substances which reduces their hygroscopicity (Swietlicki et al., 2008). Secondarily formed aerosol particles of marine origin are a result of DMS oxidation and further reactions. They can be expected to contain sulphates, and Petters and Kreidenweis (2007) give a κ value of e.g., $\kappa = 0.61$ for ammonium sulfate, derived from CCN measurements. Overall, the range of κ values



we derived for particles with sizes of ≈ 110 nm indicates that they are mostly composed of inorganic substances. The median κ of 1 might even point towards a dominance of sea salt.

Before we compare our results to literature, we want to mention that the uncertainty of the κ values was inferred with a method based on Monte Carlo simulations as described in Herenz et al. (2017) and Kristensen et al. (2016). This analysis shows that the uncertainties in our κ values are in the same order than the variability of the values itself. This allows no interpretation of the variability in κ with respect to different air mass origins.

A few other studies already examined the hygroscopicity of Antarctic aerosol particles, as well as the impact of sea ice regions on it. During the PEGASO ship cruise that took place in the austral summer in 2015 in the proximity of the Antarctic Peninsula and the Filchner-Ronne shelf ice, Dall'Osto et al. (2017) found an increased N_{CN} (aerosol particles larger than 3 nm) within air masses with an origin over sea ice regions in comparison to air masses that originated over open water. Other studies further suggest, that sea ice regions efficiently emit sea salt aerosol particles, e.g., Huang and Jaeglé (2017); Yang et al. (2008); Wagenbach et al. (1998). O'Shea et al. (2017) measured CCN at the Halley research station, ≈ 30 km from the Weddell Sea on the Brunt Ice Shelf. They report a median κ value of 0.66 during measurements in December for five different SS (0.08, 0.2, 0.32, 0.41, 0.53 %). Also, they had an event of a median κ value of 1.13 during two days, during which back trajectories indicate that air masses had passed over sea ice regions of the Weddell Sea. This is indicative for ice surfaces being able to emit aerosol particles with a high hygroscopicity and is in line with our findings. Pringle et al. (2010) applied the ECHAM-MESSy Atmospheric Chemistry (EMAC) model to simulate the global distribution of κ at the surface. That study results in values between 0.6 and 0.9 for Antarctic coastal areas and >0.9 for the Southern Ocean region. Asmi et al. (2010) measured the hygroscopicity of Antarctic aerosol particles at the Aboa station using a Hygroscopicity-Tandem Differential Mobility Analyser. They also found the Antarctic aerosol particles to be very hygroscopic with an average hygroscopic growth factor of 1.75 for 90 nm particles at 90 % RH, which is larger than the hygroscopic growth factor of ammonium sulfate. Unlike these studies and our findings, Kim et al. (2017) report a lower particle hygroscopicity. Their results are based on CCN and $PNSD$ measurements that were conducted at the King Sejong Station in the Antarctic Peninsula between 2009 and 2015. For CCN measurements at a SS of 0.4 % they found an annual mean κ value of 0.15 ± 0.05 , which, however, is the only time such low κ values were reported for Antarctica.

Summarizing, we conclude that the few large aerosol particles we observe for sizes of and above ≈ 110 nm may partially originate from NPF and subsequent growth. However the majority of these aerosol particles likely consist of sea spray particles or even more so of sea salt particles emitted over sea ice regions. This fits to the results presented for $N_{CCN,0.1\%}$ in Section 3.2, showing the marine areas in coastal proximity and especially the shelf ice regions as potential source regions.

4 Summary and conclusions

The data set presented here contains in situ ground-based aerosol particle data sampled at the Belgian Antarctic research station Princess Elisabeth (PE), in Dronning Maud Land in East Antarctica. During three austral summer seasons (2013-2016, each from December to February) we measured total aerosol particle number concentration and size distribution as well as the total



CCN number concentration at 5 different supersaturations. An automatic weather station, located in the vicinity of the PE station, and a precipitation radar were used to gain further information about the meteorological conditions. The history of the air masses arriving at the PE station was modeled by using the NAME dispersion model and the PSCF model based on HYSPLIT back trajectories.

5 N_{CN} was found to range between 40 and 6700 cm^{-3} with a median of 333 cm^{-3} . For particles being larger than 90 nm ($N_{\text{CN}>90\text{nm}}$) we found a median concentration of 20 cm^{-3} . N_{CCN} covers a range between less than 10 cm^{-3} at $SS=0.1\%$ and 1300 cm^{-3} for the largest SS of 0.7%. The median values of N_{CCN} for supersaturations of 0.1, 0.2, 0.3, 0.5 and 0.7% are 14, 81, 121, 177 and 212 cm^{-3} , respectively. All of the previous values are calculated on the basis of the entire measurement period of three austral summers. The fractions of $N_{\text{CN}>90\text{nm}}/N_{\text{CN}}$ and $N_{\text{CCN},0.7\%}/N_{\text{CN}}$ indicate that 94% and 36% of the
10 particles are smaller than 90 nm and ≈ 35 nm, respectively. From this we conclude, that an Aitken mode dominated aerosol prevailed, that includes a significant amount of freshly, secondarily formed aerosol particles.

The fluctuations in N_{CN} and N_{CCN} can be associated with the history of the air masses and the precipitation measured at the PE station. Both methods, the regional analysis on the basis of the NAME dispersion model as well as the PSCF analysis show, that high N_{CN} values are directly linked to the advection of marine air masses, which we call marine events (ME), having their
15 origin in the region of the Southern Ocean. The occurrence of precipitation is also directly linked to the occurrence of MEs, as marine air masses are the only significant source for water vapor in Antarctica. Strong precipitation events caused the lowest N_{CN} and N_{CCN} values presented in this study, due to particle scavenging and wet deposition. Therefore, MEs showed the lowest but also the highest particle concentrations measured. In contrast, when air masses had spent more than 90% of the 10 days prior to arrival over the Antarctic continent, which are times we called continental events (CE), measured N_{CN} and N_{CCN}
20 values were comparably constant, and we assume these to be continental background concentrations during austral summer. The Antarctic continent itself was found to not act as a significant source for aerosol particles and CCN measured at the PE station during these times. MEs and CEs occur 39% and 61% of the time, respectively.

The hygroscopicity of the CCNs could only be determined for measurements at $SS=0.1\%$, as the PNSDs could only be measured in a size range between 90 nm and 6.8 μm . The median d_{crit} and κ of the entire measurement period were determined
25 to be 110 nm and 1, respectively. This high hygroscopicity, which is in agreement with most of the other studies dealing with Antarctic hygroscopicity, can mainly be attributed to the presence of sea salt and sulphate aerosol particles. This is in agreement with the PSCF analysis, for which the Antarctic ice shelf areas were found to cause elevated values for particles with sizes above ≈ 110 nm, again pointing to sea salt aerosol particles. These particles could have been released and formed from snow particles from surface snow layers by sublimation, e.g., during periods of high wind speed when fresh snow is available and
30 winds are high enough to cause drifting or blowing snow (Gossart et al., 2017).

Although this is to our knowledge the most comprehensive set of CCN data in the region of East Antarctica, it is limited on the austral summer seasons. To get a complete picture of CCN properties in East Antarctica, full time measurements throughout the whole year are needed, together with PNSD measurements covering diameters down to some nm. This would enable more in-depth investigations of new particle formation and particle hygroscopicity in different size ranges. However, the here presented
35 data increases our knowledge of aerosol particle and in particular CCN properties in Antarctica.



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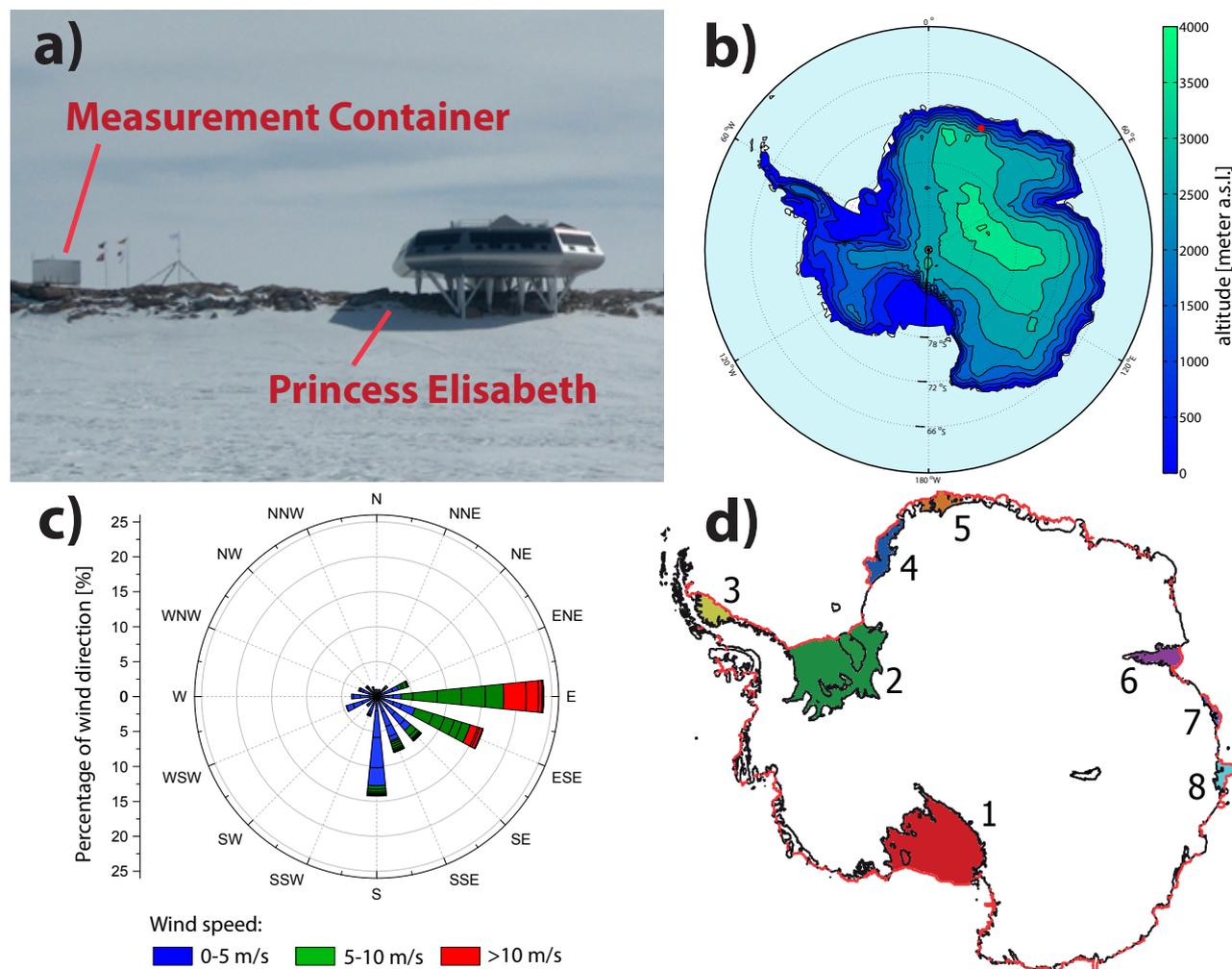


Figure 1. a) Picture of the Princess Elisabeth research station and the measurement container, where the aerosol measurements were performed, both located on the Utsteinen Nunatak ridge (view from ESE). b) Topographic map of Antarctica, the red dot shows the location of PE. This map was done using the Matlab mapping package M_Map. c) Wind direction and wind speed depicted as a wind rose for the third measurement period (December 18, 2015 to February 20, 2016). d) Map showing the location of the largest shelf ice regions in Antarctica: 1) Ross 2) Ronne-Filchner 3) Larsen C 4) Riiser-Larsen 5) Fimbul 6) Amery 7) West and 8) Shackleton shelf ice. The black line represents the coast line and the red line represents the ice edge. This map was created using Matlab and Antarctic Mapping Tools (Schaffer et al., 2016; Greene et al., 2017).

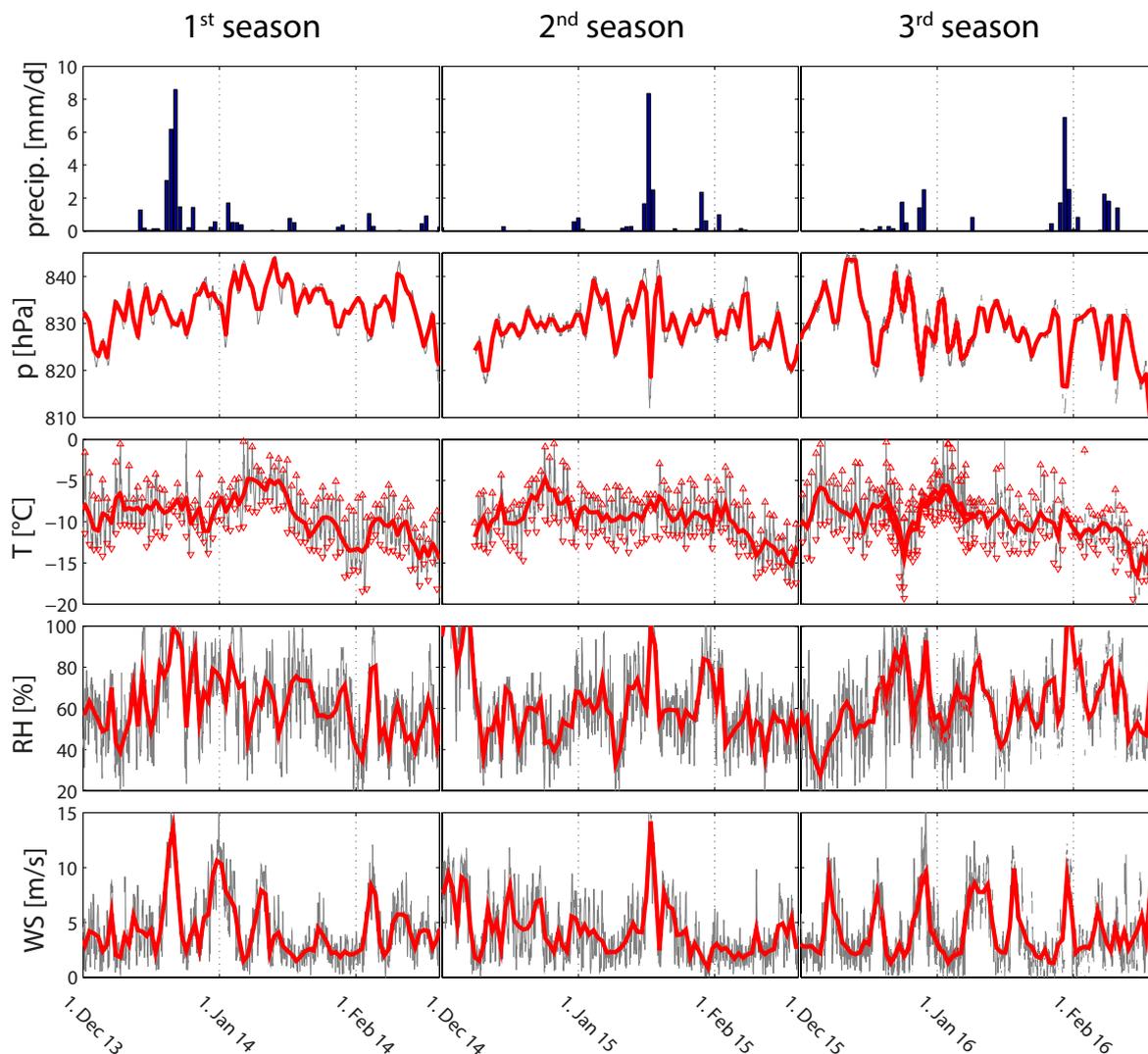


Figure 2. Time series of hourly (gray) and daily (red) mean values for temperature, pressure, relative humidity with respect to ice (RH) and wind speed (WS) measured by the AWS. Maximum and minimum temperature values are shown as triangles. The daily precipitation measured by the precipitation radar is shown as bars.

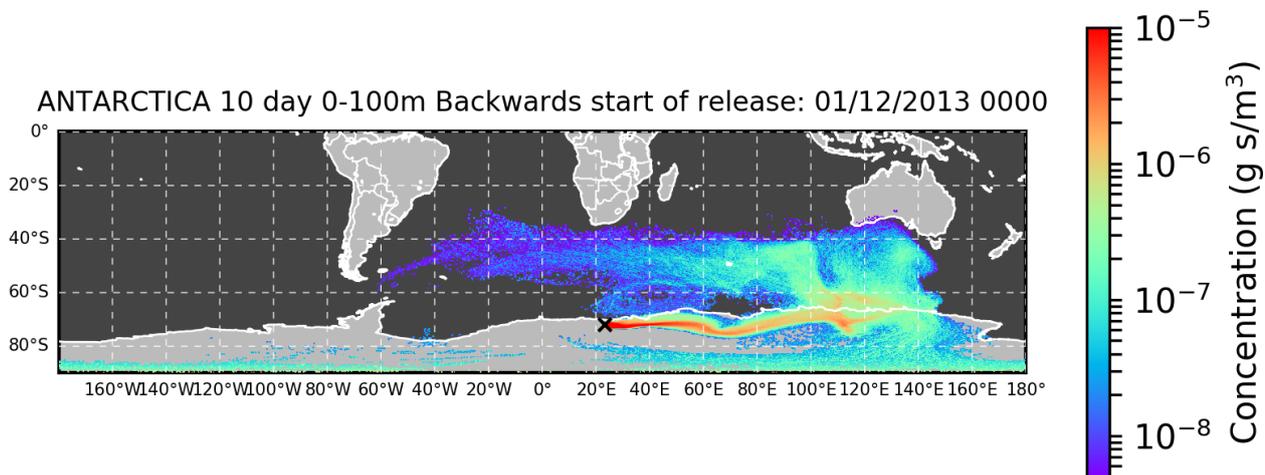


Figure 3. NAME dispersion model 10 day backwards footprint.

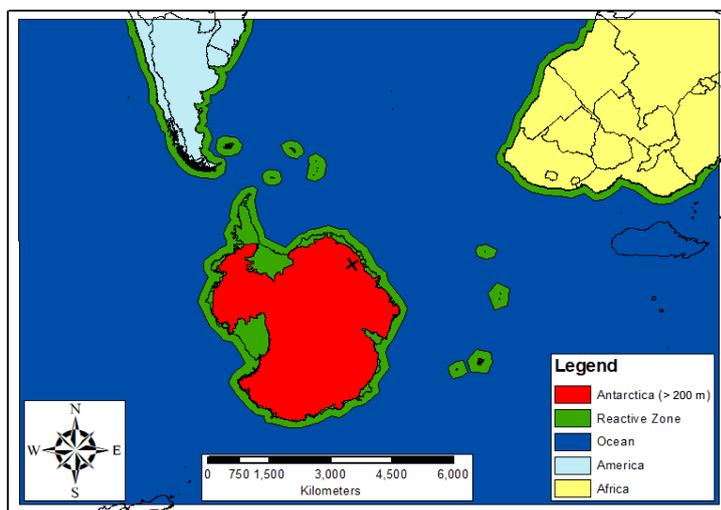


Figure 4. 5 different regions (Southern Ocean, Antarctic inland plateau, South America, Africa and Reactive Zone) that are used to track the percentage residence time in each region before arriving at PE from the NAME footprints.

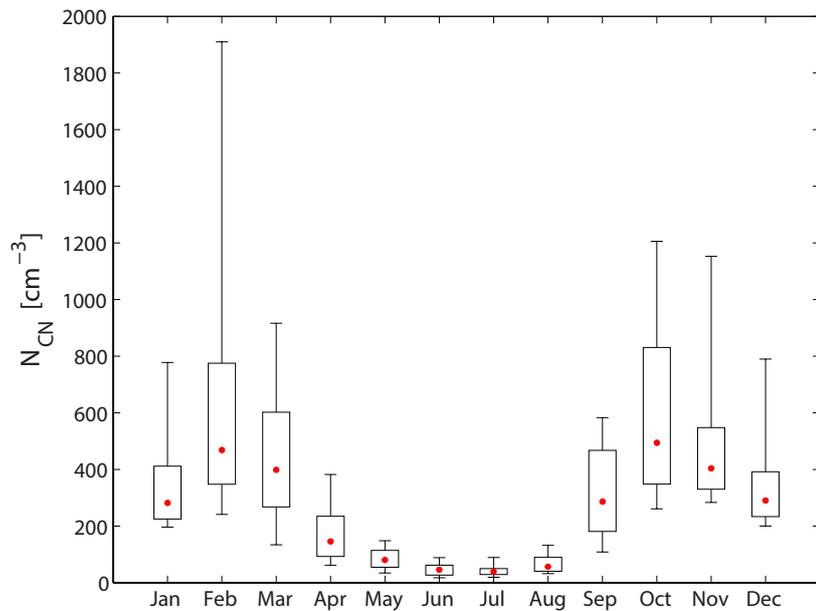


Figure 5. Box plot of monthly median values (red dots), interquartile range (black box) and 10 % and 90 % percentile (black bars) of N_{CN} measured at the PE station between 2012 and 2016.

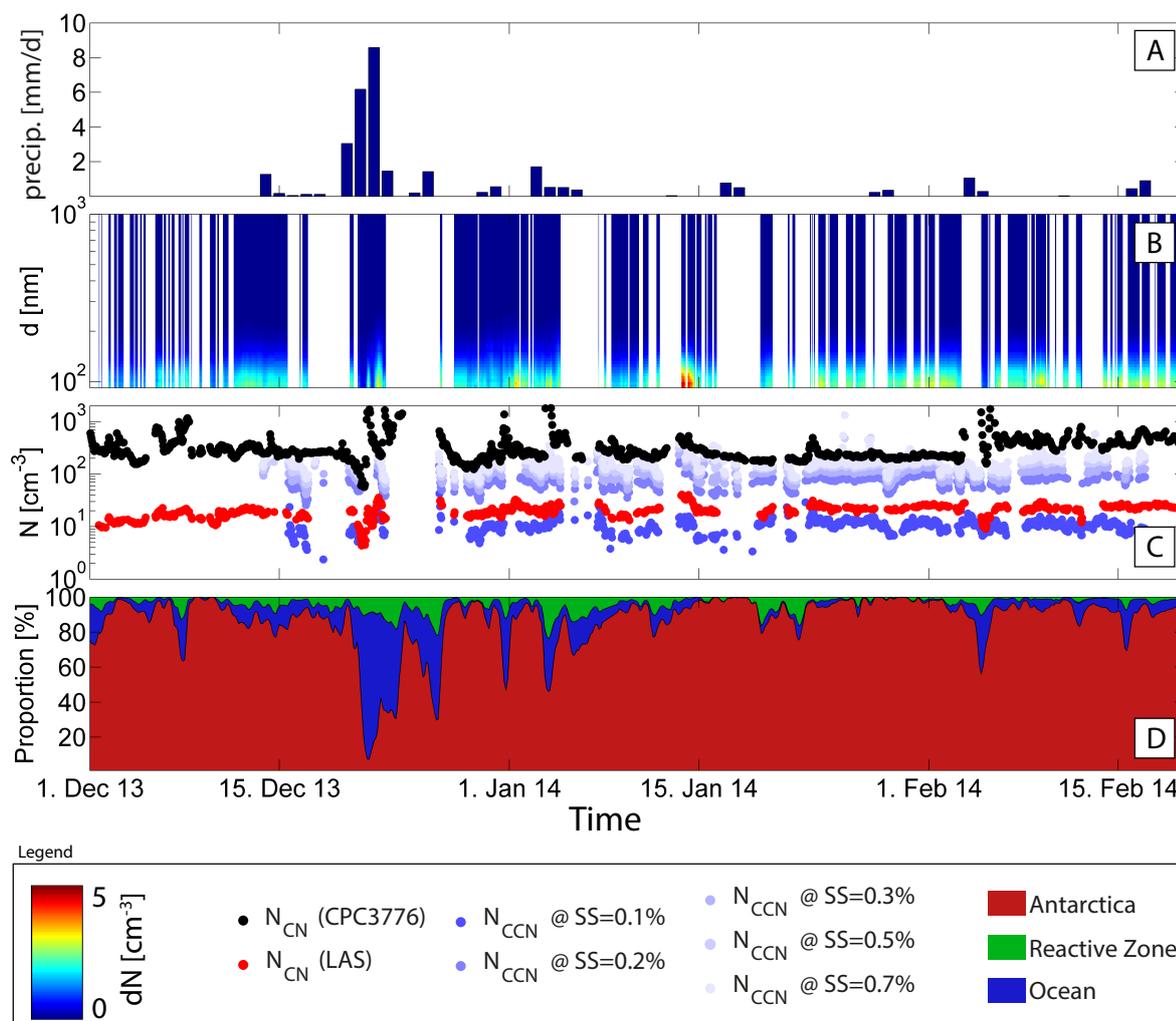


Figure 6. Time series of the first season (December 2013 to February 2014) of A) daily precipitation shown as bars (same data as in Figure 2), B) *PNSD*, depicted between 90 nm and 1 μm, C) N_{CN} measured by the CPC in black, N_{CN} measured by the LAS (integrated concentration between 90 nm and 6.8 μm) in red and N_{CCN} measured by the CCNc at SS between 0.1 and 0.7 % in different blue colors, D) proportion of residence of the air masses over the Antarctic continent (red area), the Reactive Zone (green area) and the Southern ocean (blue area) areas during the past ten days (based on the NAME model footprints).

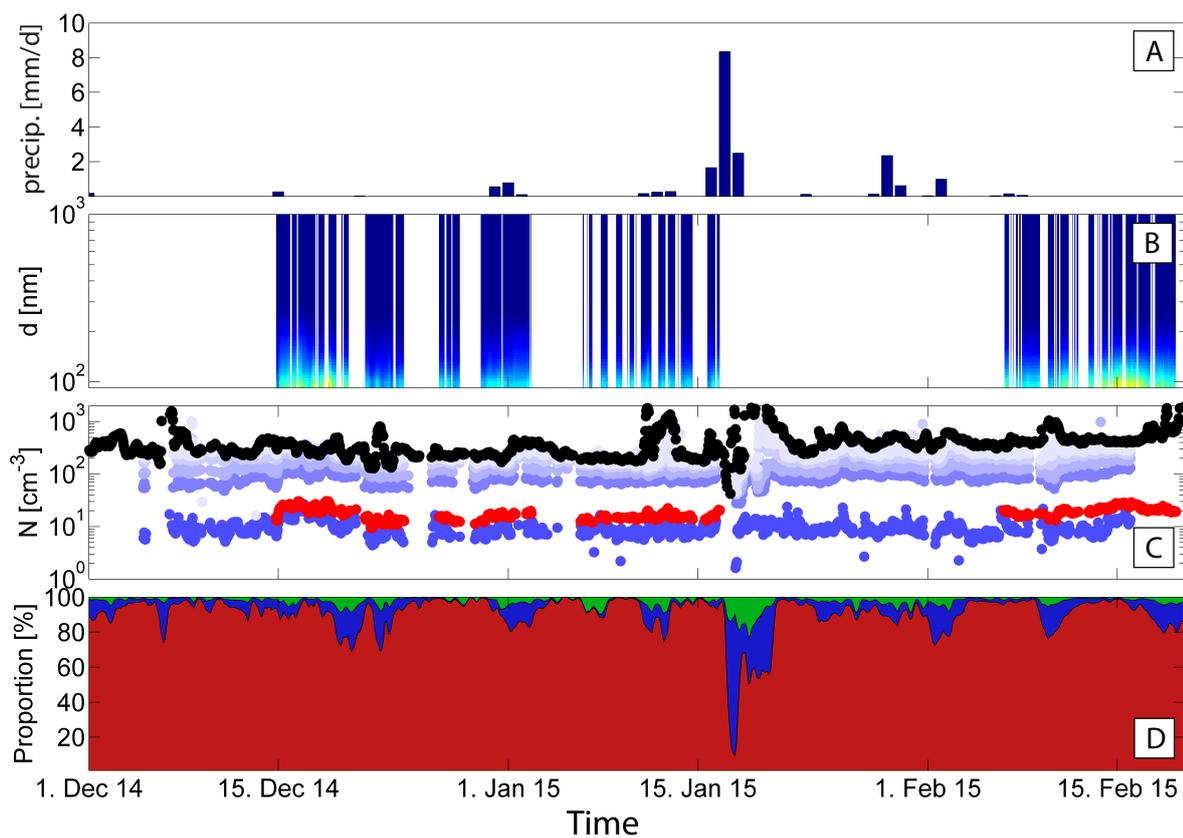


Figure 7. Time series of the second season (December 2014 to February 2015). For further details see caption of Figure 6.

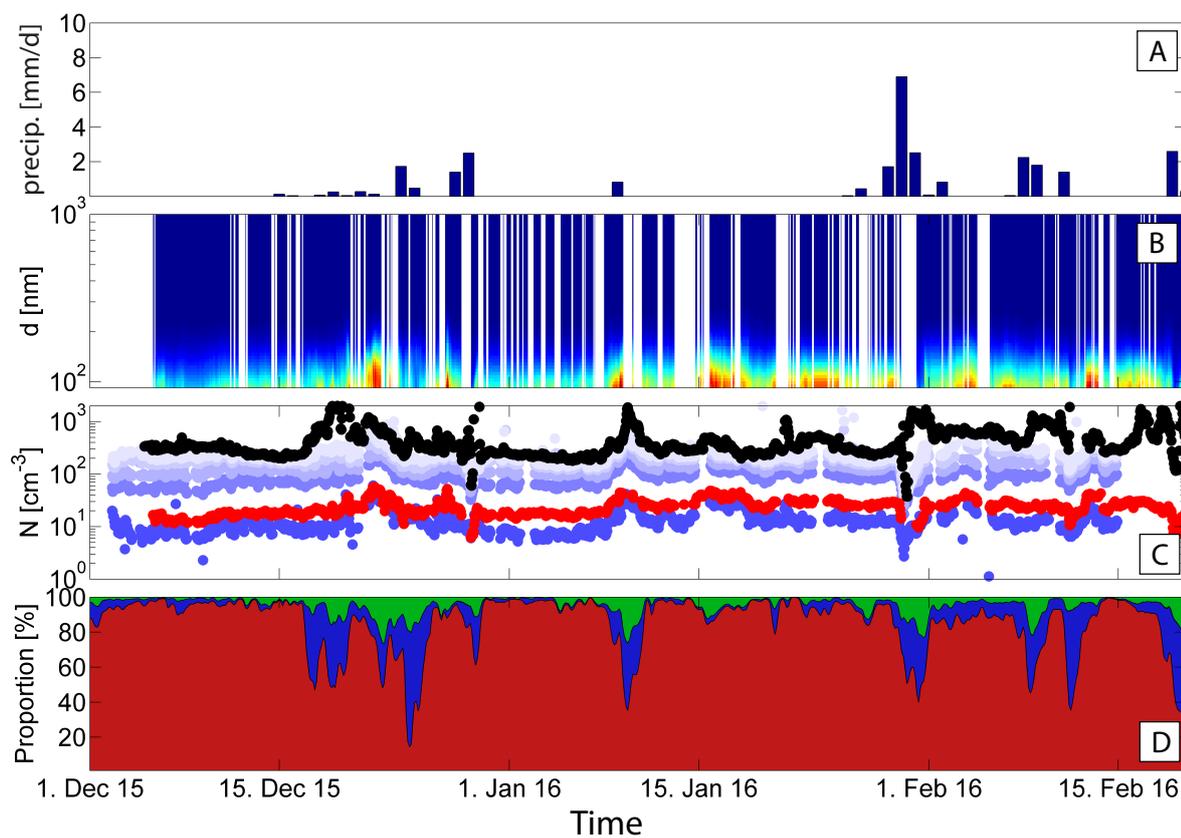


Figure 8. Time series of the third season (December 2015 to February 2016). For further details see caption of Figure 6.

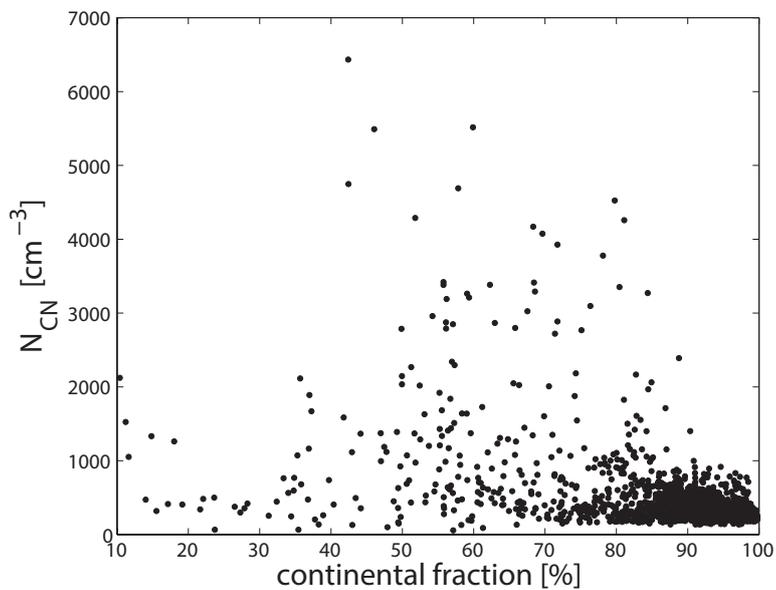


Figure 9. Connection between N_{CN} and the occurrence of continental air masses.

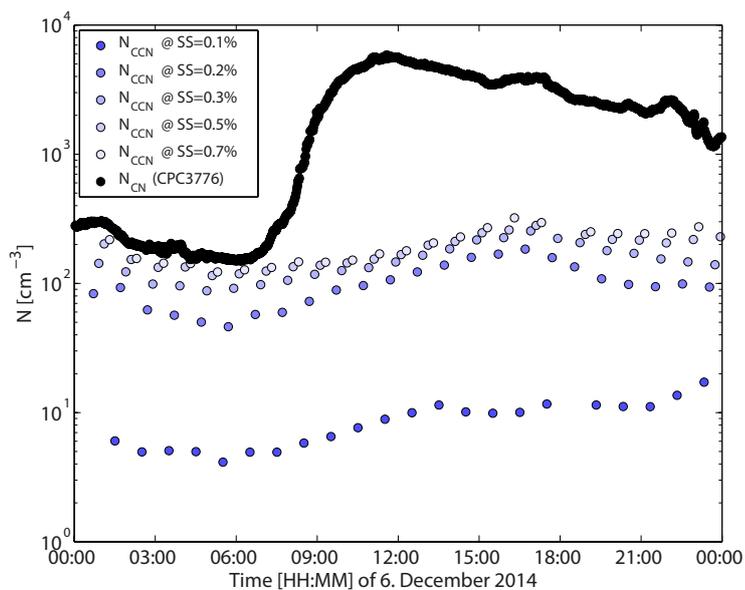


Figure 10. N_{CN} and N_{CCN} during an event of new particle formation at the PE station.

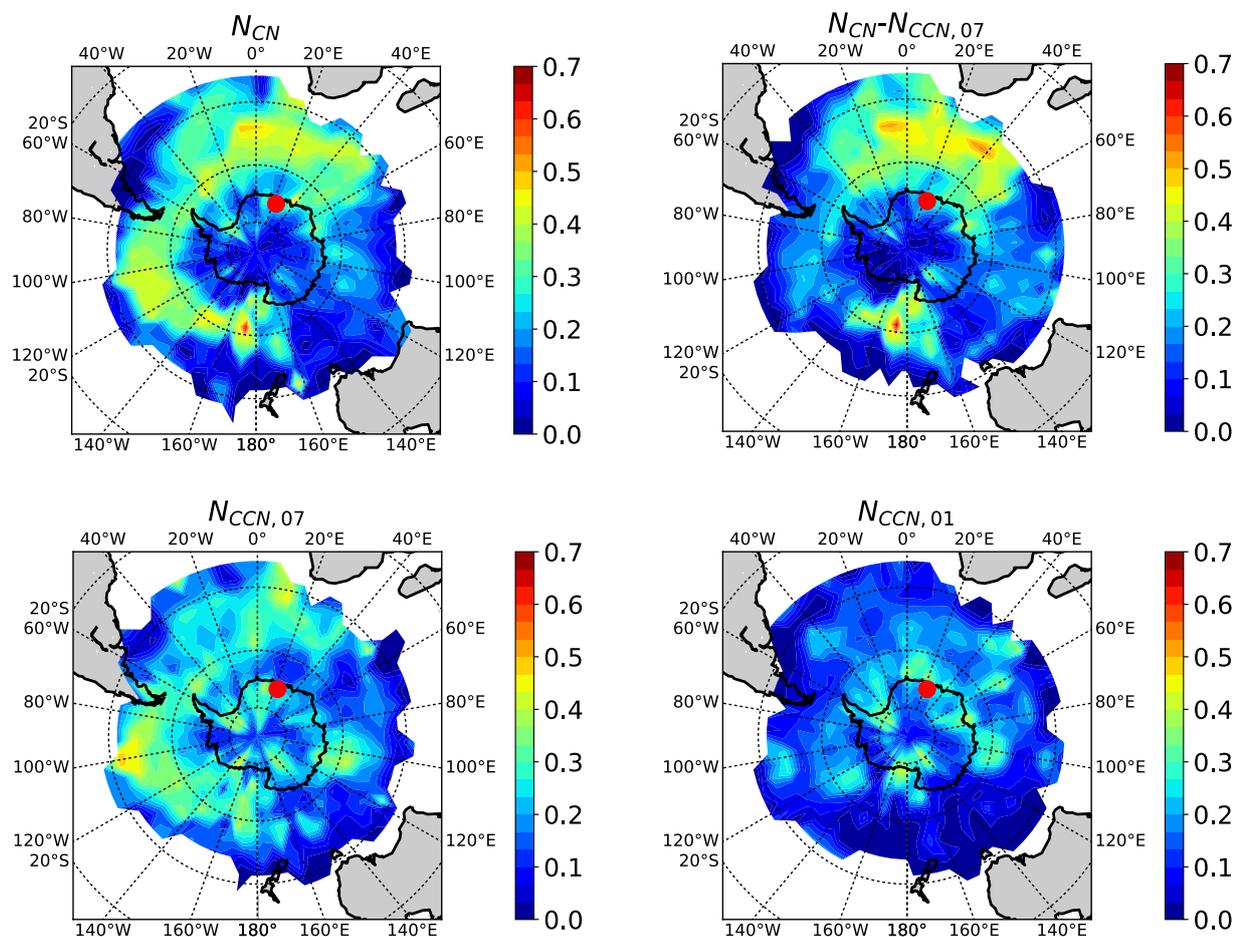


Figure 11. PSCF results that are plotted over a map of Antarctica for N_{CN} , $N_{CN}-N_{CCN,0.7\%}$, $N_{CCN,0.7\%}$ and $N_{CCN,0.1\%}$. The colorbar indicates the value of the PSCF.

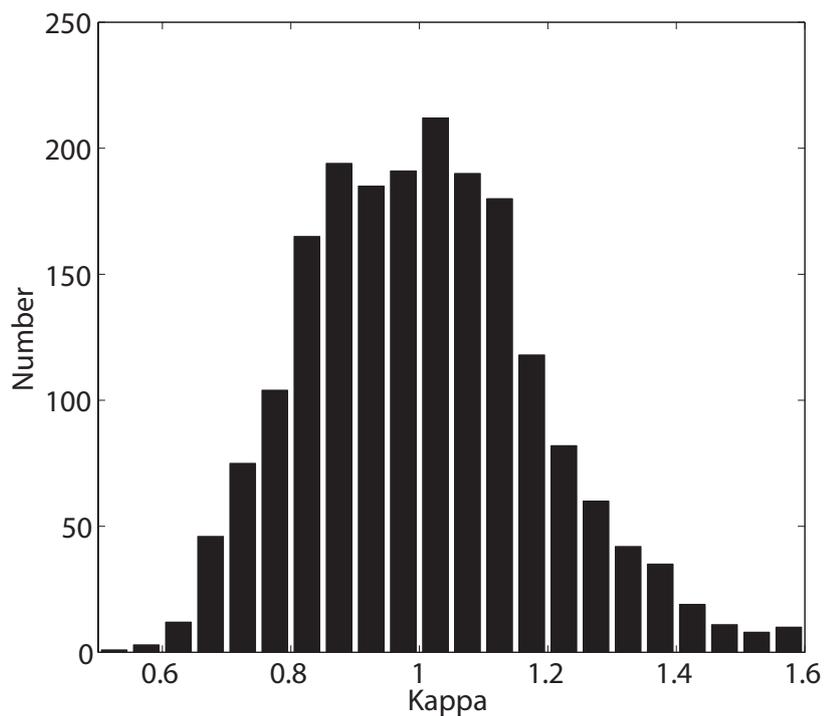


Figure 12. Histogram showing the 2171 κ values of all three seasons.



Table 1. Basic meteorological parameters (temperature, pressure, relative humidity with respect to ice, wind speed and precipitation) measured by the AWS and the precipitation radar. Shown are the mean, standard deviation, minimum and maximum values based on hourly mean values, in case of precipitation daily mean values, for the three measurement periods (each from 1 December to 20 February).

variable	2013/2014	2014/2015	2015/2016
	Mean (std), Min, Max	Mean (std), Min, Max	Mean (std), Min, Max
Air temperature [°C]	-9.4 (3.1), -19.3, 0.5	-9.7 (2.7), -17.6, -0.5	-10.1 (2.7), -20.5, 2.3
Pressure [hPa]	833.4 (5), 817.3, 844.3	829.7 (4.7), 812, 843.5	828.6 (6.6), 807, 845.3
RH [%]	61.9 (18.6), 13.1, 100	58.9 (17.2), 14, 100	64.3 (18.2), 14.7, 100
Wind speed [m/s]	4.35 (2.87), 0.13, 16.21	4.26 (2.89), 0.03, 22.59	4.21 (2.99), 0, 16.6
Precipitation [mm/d]	0.38(-), -, 8.6	0.24 (-), -, 8.3	0.35 (-), -, 6.9

Table 2. Overview showing N_{CN} , $N_{CN>90nm}$ and N_{CCN} at different supersaturations, given as median (and 10% and 90% percentiles in brackets) in column 1 for all data, in column 2 for CEs (continental events, based on the regional analysis of the NAME model output). Column 3 shows the fraction of $N_{CN>90nm}$ and N_{CCN} to N_{CN} (based on the median values of column 1).

Parameter	Median concentration	Median concentration during CEs	N_{CN} (LAS)/ N_{CN} (CPC)
	(10%, 90% percentile) [cm^{-3}]	(10%, 90% percentile) [cm^{-3}]	or N_{CCN}/N_{CN} (CPC)
N_{CN} (CPC)	333 (206, 893)	292 (205, 474)	-
$N_{CN>90nm}$ (LAS)	20 (14, 29)	20 (14, 29)	0.06
$N_{CCN,0.1\%}$	14 (10, 23)	14 (10, 21)	0.04
$N_{CCN,0.2\%}$	81 (56, 110)	79 (58, 105)	0.24
$N_{CCN,0.3\%}$	121 (90, 168)	120 (95, 161)	0.36
$N_{CCN,0.5\%}$	177 (125, 260)	177 (133, 232)	0.53
$N_{CCN,0.7\%}$	212 (138, 326)	210 (150, 292)	0.64