Reviewer #1
The study of Shao et al. Implemented four heterogeneous sulfate formation mechanism into GEOS-Chem model with a focus of the low biases of modeled sulfate production rates in China. The four paths are via H2O2, O3, NO2, and TMI on aerosols. In addition, TMI-catalyzed oxidation in clouds were also considered. To reduce the uncertainties, the oxygen isotopes observations were used for comparison as it’s highly sensitive to the relative importance of different sulfate production mechanism. To investigate the dependence on aerosol PH, sensitivity studies with prescribed values of aerosol PH were also conducted. The design of the experiments are comprehensive and the results are convincing. Overall, the results with the four added heterogeneous reactions significantly reduced the low biases of sulfate and also oxygen isotopes, which show better agreement with the observations. The publication is very well written, clearly structured, and the analyses are comprehensive and convincing. In particular, the authors present a thorough analysis of the heterogeneous oxidation paths, substituting the bulk first-order uptake of SO2 (reaction probability /uptake coefficients) by a more specific calculation approach. The paper has a good chance to become an important reference for future studies on the sulfate heterogeneous reaction mechanism in China. It also has regional/global impacts. I support publication of this manuscript and have only a few small comments that may require minor revision.

Response: We thank the reviewer for the valuable comments. All of them have been implemented in the revised manuscript. Please see our itemized response below.

Comment: 1. P5 L45 Can you add some more details of the methodology of the in-cloud TMI-catalyzed aqueous-phase S(IV) oxidation? As it seemed this in-cloud TMI-relevant path is very important in polluted events.
Response: Thanks for pointing it out. Metal-catalyzed S(IV) oxidation by O2 (TMI) in cloud droplets was considered to be one of the most important sulfate formation pathways during the winter in the North Hemisphere (Huang et al., 2014; Harris et al., 2013; Alexander et al., 2009; Sofen et al., 2011), and we also found that the pathway accounted for over 20% of total sulfate formation during HPP at Beijing (P11L355). We now state in the text (P5L145) “In the model simulation Run_TMI, we implemented the in-cloud TMI-catalyzed aqueous-phase S(IV) oxidation by O2 into the model, which is thought to be one of the most important sulfate formation pathways during the North Hemisphere winter (Huang et al., 2014; Harris et al., 2013; Alexander et al., 2009; Sofen et al., 2011). The parameterization of TMI-catalyzed S(IV) oxidation in cloud for GEOS-Chem follows Alexander et al. (2009).”
We have also added more details on the methodology of the in-cloud TMI-catalyzed aqueous-phase S(IV) oxidation in the Supplement (Text S3; Figure S4; Table S2; Table S3). This also addresses Comment #2 below. We now state in the text (P5L155) “After modification, the average aqueous-phase concentration of Fe(III) in cloud water during our study period is 2.9 µM and Mn(II) is 1.3 µM in the model, which is consistent with previous work (He et al., 2018; Shen et al., 2012; Guo et al., 2012) (see Text S3 in the Supplement for more details).”
Text S3 in the Supplement is as follows:

The parameterization of TMI-catalyzed S(IV) oxidation in clouds in GEOS-Chem is described in Alexander et al., (2009). The natural source of Fe ([Fe]nat) and Mn ([Mn]nat) from mineral dust are scaled to total dust mass. [Fe]nat is 3.5% of total dust mass and [Mn]nat is a factor of 50 lower than [Fe]nat (Alexander et al., 2009). The anthropogenic source of Fe ([Fe]ant) and Mn ([Mn]ant) from coal combustion are scaled to the abundance of primary anthropogenic sulfate due to their common source and atmospheric lifetime. [Mn]ant is 1/300 of primary sulfate concentration and [Fe]ant is 10 times that of [Mn]ant as described in Alexander et al. (2009). Figure S4 shows the distribution of simulated [Fe]nat and [Fe]ant during the entire model simulation period. Table S2 compares autumn-winter mean modeled Mn and Fe with the observations at several sites in east Asia, including 3 sites in the North China Plain (NCP) region (Beijing, Tianjin and Shijiazhuang), 2 sites in the Sichuan Basin (Chengdu and Chongqing), and Lanzhou and Seoul which are the upwind and downwind of Beijing, respectively. The model tends to underestimate both Fe and Mn at the NCP sites by a factor of 0.25-0.72. Model comparison with observations in the Sichuan Basin shows good agreement on Mn concentration (within 20% for both Chongqing and Chengdu), with model calculations overestimating the observations of Fe concentrations. For sites in Lanzhou (northwest of Beijing, upwind) and Seoul (southeast of Beijing, downwind), model comparison with observations indicates good agreement on Fe (within 5%), and underestimates Mn by 60%. The model underestimates Fe and Mn in the NCP region and overestimates in Sichuan Basin. These discrepancies highlight the limitations of our approach applying a global-scale factor of [SO4] primary/[Mn]ant and [dust]/[Fe]nat, as regionally varying emissions control technologies and mass fraction of Fe in dust may impact the relative emission rates.

However, sulfate formation by the TMI-catalyzed oxidation pathway is influenced by soluble Fe^{3+} and Mn^{2+} concentrations, as opposed to total Fe and Mn. Previous studies suggest that the solubility of Fe and Mn ranges from 0.03 to 54% and 1.2 to 97%, respectively. The solubility of metals is influenced by several factors such as the natural versus anthropogenic origin of samples (A.R. Baker et al., 2006; K.V. Desboeufs et al., 2005; K.V. Desboeufs et al., 2001; Spokes et al., 1994; A. Ito1 and Y. Feng, 2010; P.Y. Chuang et al., 2005; Solmon et al., 2009), acidity, and sunlight. In this study, we assume a solubility of 10% for [Fe]nat, 0.45% for [Fe]ant, 50% for [Mn]ant, and 5% for [Mn]nat in cloud water. The modeled soluble Fe^{3+} and Mn^{2+} concentration is shown in Table S3. After modification, the average modeled concentration of soluble Fe(III) in cloud water during our study period is 2.9±1.8 µM and Mn(II) is 1.3±0.7 µM, which is consistent with estimates (He et al., 2018) and observations (Shen et al., 2012; Guo et al., 2012). A sensitivity study is performed based on Run_TMI but with the higher solubility of Fe and Mn as Alexander et al. (2009): the solubility of 1% of [Fe]ant and 50% for [Mn]ant in cloud water. The modeled soluble Fe(III) and Mn(II) concentration reaches 20µM and 10 µM during HPP, up to a factor of 5 higher than the observations. The simulated sulfate concentration is also overestimated the observation by around 100% during HPP in Beijing.
Figure S4. Modeled anthropogenic and natural Fe and Mn (ng m$^{-3}$) at the surface, along with the anthropogenic/total Fe and Mn concentration percentages.

Table S2. Comparison of observed and modeled autumn and winter mean Mn and Fe concentrations (ng m$^{-3}$) from several East Asian locations.

<table>
<thead>
<tr>
<th>Location</th>
<th>Fe (Mn) observation (ng m$^{-3}$)</th>
<th>Fe (Mn) model (ng m$^{-3}$)</th>
<th>Fe (Mn) Model/observation</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>Beijing</td>
<td>1800 (90)</td>
<td>835 (22.7)</td>
<td>0.46 (0.25)</td>
<td>Zhao et al. (2013)</td>
</tr>
<tr>
<td>Tianjin</td>
<td>1980 (120)</td>
<td>1239 (36.2)</td>
<td>0.72 (0.30)</td>
<td></td>
</tr>
<tr>
<td>Shijiazhuang</td>
<td>2250 (150)</td>
<td>1615 (44.3)</td>
<td>0.63 (0.30)</td>
<td></td>
</tr>
<tr>
<td>Chengdu</td>
<td>875 (38)</td>
<td>1245.7 (34.1)</td>
<td>1.42 (0.90)</td>
<td>Wang et al. (2018)</td>
</tr>
<tr>
<td>Chongqing</td>
<td>502.5 (36.5)</td>
<td>1226.0 (39.7)</td>
<td>2.4 (1.09)</td>
<td></td>
</tr>
<tr>
<td>Lanzhou</td>
<td>1534 (59)</td>
<td>1456.9 (33.0)</td>
<td>0.95 (0.56)</td>
<td>Wang et al. (2016)</td>
</tr>
<tr>
<td>Seoul</td>
<td>227 (21)</td>
<td>219.8 (6.1)</td>
<td>0.97 (0.30)</td>
<td>Park et al. (2018)</td>
</tr>
</tbody>
</table>

Table S3. Comparison of observed and modeled soluble Mn(II) and Fe(III) concentrations (µM).

<table>
<thead>
<tr>
<th></th>
<th>This study Range (monthly mean)</th>
<th>Beijing (He et al., 2018)</th>
<th>Mt. Tai$^a$ (Shen et al., 2012)</th>
<th>Mt. Tai (Guo et al., 2012)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Fe(III)</td>
<td>0.4-11.2 (2.9±2.7)</td>
<td>0.6-6.1 (2.6±1.8)</td>
<td>0.8-7.4</td>
<td>2.6</td>
</tr>
<tr>
<td>Mn(II)</td>
<td>0.4-4.8 (1.3±0.9)</td>
<td>1</td>
<td>0.4-1.7</td>
<td>1.2</td>
</tr>
</tbody>
</table>

$^a$Mt. Tai is located in central Shandong province at the eastern edge of the NCP region.

Comment: 2. P5 L45 The TMI-catalyzed oxidation path are important both in-cloud (aqueous) and on-aerosol (heterogeneous). Is there any possibility to verify the assumption of Fe and Mn treatments in the
model, from the natural dust and anthropogenic emissions, solubility, ionic strength of cloud liquid water, etc.? Or if not, is it necessary to add some discussions about the uncertainty of those assumptions and the possibility of the impacts on the analysis?

Response: Thanks for pointing it out. We now state in the text (P5L155): “After modification, the average modeled concentration of Fe(III) in cloud water during our study period is 2.9 µM and that of Mn(II) is 1.3 µM, which are consistent with previous work (He et al., 2018; Shen et al., 2012; Guo et al., 2012) (see Text S3 in the Supplement for more details).” Natural and anthropogenic dust emissions are now shown in Figure S7. The influence of ionic strength on the in-cloud, TMI-catalyzed S(IV) oxidation pathway is negligible, but is significant in aerosol water (heterogeneous pathway). Please refer to the response for Comment #1 and Text S3 in the Supplement for the discussions on Fe and Mn treatments and associated uncertainties.

In addition, we deleted the sentence on P5L155 to avoid misunderstanding. We also rewrite the sentence on P6L165 “In addition, we considered the impacts of acidity and ionic strength on TMI-catalyzed reaction rates following Cheng et al. (2016) (Table S2 in the SI of Cheng et al., 2016), since the ionic strength of aerosol liquid water can reach 20 M during polluted periods (He et al., 2018; Herrmann et al., 2015).”

Comment: 3. P7 L110 The 2010_MEIC emission is used in the study, however the anthropogenic SO2 emissions has been reduced largely since 2009 thus the SO2 overestimation might be expected for Beijing. Is there any impacts/uncertainty induced by this issue?
In the Run_Het results, the TMI-relevant reactions are the most important among the four heterogeneous paths during both clean and polluted events. “In the model, the heterogeneous sulfate production rate from the TMI-catalyzed reaction is calculated as first-order uptake in SO2”. Is that relevant with the SO2 overestimation in the model?
Response: Thanks for pointing this out. Unfortunately, the updated MEIC emission inventory was not released at the time when we ran our model simulations. We now state in the text (P13L388) “Anthropogenic SO2 emissions in China have been reduced sharply since 2009 due to the stringent pollution control measures implemented (Zheng et al., 2018; Van der A et al., 2017; Krotkov et al., 2016). Compared with 2010, anthropogenic SO2 emissions reduced by about 50% in 2015 (Krotkov et al., 2016; Zheng et al., 2018; Van der A et al., 2017). NH3 and non-methane volatile organic compounds (NMVOC) emissions in China remained stable during 2010–2017 due to the absence of effective mitigation measures in current policies (Zheng et al., 2018). The emission changes may affect the abundances of species that influence cloud and aerosol pH, and further influence sulfate production rates and the contribution of each sulfate formation pathway. However, other studies using observations between 2014-2016 (Liu et al., 2017; Song et al., 2018a) found a similar pH range as calculated here, suggesting that a modeled low bias in aerosol pH is not likely to be the source of the modeled discrepancy in △17O(SO42-).”
We believe that the SO2 overestimation in the model will not change the dominance of TMI-relevant pathway in the model because SO2 concentrations influence all four heterogeneous sulfate
formation pathways. As shown in Figure S1, heterogeneous sulfate production rates parameterized as first-order in SO$_2$ or first-order in the oxidant result in similar values when aerosol pH < 6, the only difference is when production rates are limited by the mass transport across the air-water interface at higher pH values. SO$_2$ concentration influence the sulfate formation by gas-phase oxidation by OH and aqueous-phase oxidation by H$_2$O$_2$, O$_3$ and O$_2$ catalyzed oxidation by TMI. Anthropogenic sulfate emission also scales to anthropogenic SO$_2$ emission in the model. Therefore, the SO$_2$ overestimation would not change the main conclusion of this study regarding the fractional contribution of each sulfate formation pathway to total sulfate abundance in Beijing.

Comments: 4. P7 L120 The mass accommodation coefficients for O$_3$ and NO$_2$ are much lower that for SO$_2$. Is this the reason that TMI-relevant reactions are much important in the analysis? How those parameters were determined? Any assumption? Is it necessary to put Text S2 in the manuscript? As the analysis is done for the October-November period, is it possible to apply to other seasons? any sensitivity/uncertainties to regions/seasons etc?

Response: Thanks for suggestions. The mass accommodation coefficients would influence heterogeneous sulfate production rates when aerosol pH > 6, but the aerosol pH seldom reaches these values in the model. Even if aerosol pH reaches 6 in the model, Fe$^{3+}$ and Mn$^{2+}$ concentrations are very low due to low solubility at high pH (Cheng et al., 2016), thus less sulfate would be produced by this pathway under this condition. Therefore, the mass accommodation coefficients are not the reason that TMI-relevant reactions are so important among the four heterogeneous pathways. The mass accommodation coefficients for O$_3$, NO$_2$, H$_2$O$_2$ and SO$_2$ are from laboratory calculations and widely used (Jacob, 2000). Considering the importance of Table S2, we now move it to the manuscript.

The importance of heterogeneous sulfate formation for the global sulfur budget and in different seasons is the focus of our next paper, and is beyond the scope of this manuscript.

Comments: 5. P11 L345 It seemed the sulfate bias is reduced while the PM2.5 bias is still there. I’m just wondering is there any observational data of nitrate or ammonia available for comparison? It might give some indication of the partitioning of sulfate-nitrate. If the nitrate are well simulated or not impacted by the competition from sulfate formation, larger improvements of PM2.5 are expected?

Response:

Thanks for the comment. Both the sulfate and PM$_{2.5}$ low biases still exist, even after implementing the four heterogeneous sulfate production pathways, as discussed in the manuscript. We have discussed in the text the potential influence of HMS in reconciling the remaining low bias in sulfate. This paper focuses on sulfate production mechanisms. Investigation of model biases in nitrate and ammonia are outside the scope of this paper, although we do have a separate paper in preparation focusing on a comparison of modeled and observed nitrate in Beijing. In the figure below, we compare modeled and observed nitrate and ammonia concentrations, which shows that the enhancements in sulfate abundance in Run_Het result in a small increase in ammonia, but have a negligible effect on nitrate concentrations. As noted in previous work (Wang et al., 2013), the model
tends to overestimate nitrate concentrations.

**Figure.** Time series of nitrate and ammonium at the surface in Beijing during the study period of 17 October 2014 – 20 January 2015. 12-hour average nitrate and ammonium observations (black line) are compared with model results from RunStd (orange line) and RunHet (purple line).

**Comments:** 6. P11 L349 The largest sulfate enhancements due to heterogeneous sulfate formation occur in megacities in eastern China and Sichuan Basin. For Sichuan region, is that also due to the high cloud liquid water path/RH?

**Response:** Thanks for pointing this out. We now state in the text (P10L320): “The largest enhancement in sulfate abundance after adding the in-cloud TMI pathway occurs in Sichuan basin (around 6.5 µg m⁻³), where simulated anthropogenic Fe and Mn from coal fly ash (Figure S4) and SO₂ are high (Zhang et al., 2009) in part due to high SO₂ emissions combined with stagnant air and high relative humidity all year (Huang et al., 2014).” We have also compared the observed and modeled Fe and Mn concentrations in the Sichuan region as shown in Table S2.

**Comments:** 7. In Run Het, the simulated oxygen isotopes are improved in average but the median are largely underestimated. In addition to the assumption of O₃ oxidation underestimation, is there any possibility of other missing paths that not taken into account? Any discussions necessary?

**Response:** Ozone is the only oxidant with a Δ¹⁷O value high enough to explain the majority of the discrepancy. The ozone oxidation pathway is mainly limited by the fraction of S(IV) that exists as SO₃²⁻, which increases with increasing aerosol pH. In the manuscript, we have suggested that sulfate formed via oxidation of SO₂ by ozone on alkaline dust could explain at least part of the discrepancy, as this reaction is not included in the model (P14L440).

An underestimate of H₂O₂ could also explain part of the discrepancy. We now state in the text “The average modeled H₂O₂ concentrations during HPP and CP (Figure S3) underestimates the observations (Ye et al., 2018) by up to an order of magnitude (Figure S3). Mao et al. (2013) proposed a HO₂−Cu−Fe catalytic mechanism for H₂O₂ production in the aerosol phase. In their mechanism, the uptake of HO₂ and subsequent heterogeneous reactions with Cu and Fe will lead to production of H₂O₂ when the molar ratio of dissolved Cu to Fe was >0.1. Ye et al. (2018) found that the molar ratio of dissolved Cu to Fe could be >0.1 during moderately polluted days, suggesting the uptake
of HO₂ radicals on particles might be an important source of H₂O₂ during the winter in Beijing. We note however that an underestimate of modeled H₂O₂ cannot explain all of the discrepancy in Δ¹⁷O(SO₄²⁻), as sulfate formed from H₂O₂ oxidation (0.7 ‰) is lower than the observed mean Δ¹⁷O(SO₄²⁻) of 0.9 ‰. All other oxidation pathways yield Δ¹⁷O(SO₄²⁻) = 0 ‰ and cannot explain the model's low bias in Δ¹⁷O(SO₄²⁻).