

Dear Dr. Frank Dentener,

Thank you very much for dealing our manuscript, titled “A 15-year record (2001–2015) of the ratio of nitrate to non-seasalt sulfate in precipitation over East Asia,” (#acp-2017-848) which we would like to resubmit for publication in *Atmospheric Chemistry and Physics*.

We would also like to thank the anonymous referee and Dr. Wenche Aas for reading our manuscript and providing helpful comments. We have revised our manuscript according to the reviewers' comments and suggestions. We believe that these revisions address all points raised by the reviewers. Our point-by-point responses are provided below, and revisions are indicated in blue in the revised manuscript.

Sincerely,  
Syuichi Itahashi

The long-term trend of NO<sub>3</sub><sup>-</sup> and SO<sub>4</sub><sup>2-</sup> concentration in precipitation in China, Korea, and Japan was analyzed based on EANET monitoring data. The results showed certain correlation between China's emission and the wet deposition in East Asia. The manuscript was quite well written. However, the method for trend analysis was not clear introduced. Since there are a lot of anthropogenic sources and other natural sources of Na<sup>+</sup> than sea-salt in China, the necessity of sea-salt correction for SO<sub>4</sub><sup>2-</sup> concentration in precipitation need be reconsidered. Some detailed comments are:

In the revised manuscript, the methodology to calculate linear trend are explicitly introduced. Please refer the reply to the specific comment of 5). Regarding Na sources, we performed and checked the analysis of the correspondence between SO<sub>4</sub><sup>2-</sup> and nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation around Beijing in China. Please refer the reply to specific comments of 3) and 6).

1) Page 1, Line 15: The ratio is an index only for relative contribution of SO<sub>2</sub> and NO<sub>x</sub>.

We have added a short note to clarify this point as follows on P1, L17:

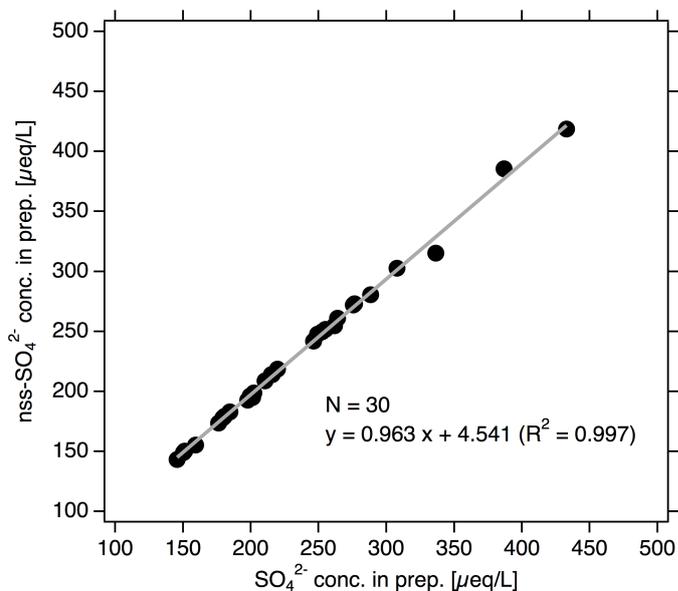
“...to investigate the relative contributions of these acidifying species.”

2) Page 3, Line 21: As I know, a long-term (2001-2013) monitoring of precipitation chemistry at one of the IMPACTS site (TSP) was recently reported by Yu et al. (ES&T, 2017).

Thank you for bringing this to our attention. We have added the suggested reference and a brief explanation of the study (e.g., P3, L22; P5, L28-29, P13, L21-22).

3) Page 4, Line 21: I suggest not to do sea-salt correction for SO<sub>4</sub><sup>2-</sup>, at least in China.

We agree that there are Na sources other than seasalt, particularly in China. Na might be affected by wind-blown dust in spring, biomass burning in autumn, and coal burning in winter. The impact of Na can be considered as minor on annual mean base because rainfall event mostly occurred during summer. The following figure shows the relation between annual mean (2008-2010) of SO<sub>4</sub><sup>2-</sup> and nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation analyzed on ten sites around Beijing in China. From this analysis, we can confirm that there are no significant differences between SO<sub>4</sub><sup>2-</sup> and nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation in China. On the other hand, in Korea and Japan, where surrounded with oceans, Na source from sea-salt would have impact. In this study, to ensure consistent data types for China, Korea, and Japan, we would like to show results for non-seasalt SO<sub>4</sub><sup>2-</sup>.



Supplemental figure: Scatter-plot of annual mean (from 2008 to 2010)  $\text{SO}_4^{2-}$  and nss- $\text{SO}_4^{2-}$  concentration in precipitation at ten sites around Beijing in China.

4) Page 4, Line 24: Is the monthly mean concentration volume-weighted?

Yes, we have explicitly indicated that volume-weighted monthly means were used in the revised manuscript on P4, L29-30.

5) Page 5, Line 4: I can not see the introduction on the method for trend analysis.

Linear trends were determined based on a linear regression analysis. We have described these methods as follows on P 8, L 24-25:

“Linear trends during each of the three phases were analyzed using linear regression, and significance levels were calculated using the Student’s *t*-test.”

6) Page 5, Line 32:  $\text{Na}^+$  can not be used as sea-salt tracer, at least in China.

In this study, to facilitate comparisons among values for China, Korea, and Japan, we would like to use the same data types and thereby would like to present non-seasalt  $\text{SO}_4^{2-}$ . Please also refer the reply 3).

7) Page 11, Line 7: Double ‘Xiang Zhou’.

This was an error and should read “Zhuxiandong.” In the revised manuscript, the correct site classification has been included in Table 1.

8) Page 26, Line 4: Better use equivalence as the unit for the sum.

We have revised the calculation of the sum as suggested. Figure 7 and the relevant explanation in the main manuscript have been revised.

9) Page 28 and 29, Table 2 and 3: How to calculate the trend?

The trend was estimated using linear regression. We have clarified this point in the revised manuscript. Please refer the reply 5).

W. Aas (waa@nilu.no)

This study compares observed trends in in-situ observations of sulfate and nitrate in precipitation, and satellite data of NO<sub>2</sub> and SO<sub>2</sub> with emission estimates of SO<sub>x</sub> and NO<sub>x</sub> in East Asia for the last 15 years. This is an important topic since there has been large changes in the emissions in Asia the last decades. The comparison between satellite data with emission, I think is good, but the comparison with in situ data I find a bit shallow. The spatial variability of the emissions and the wet deposition are large, and the site representativity is a critical question. Are the sites in question able to show the overall trends for the different regions? For South Korea and Japan it is probably OK, but in China it is too few sites. The authors are aware of the problem and add campaign data in China, but these are only snapshots in time and cannot be used for assessing trends. I do miss more discussion of this. As I understand it, the authors claim that the ratio of NO<sub>3</sub>/SO<sub>4</sub> is a good parameter for assessing the regional trends (better than the individual concentrations), but this is mainly showing that the chemical regime is changing, which is probably correct, maybe not all over the continent though, and does it help us to understand whether the individual trends of NO<sub>3</sub> and SO<sub>4</sub> are correct?

A critical point is that the paper is merging all type of sites. In the EANET network, there are many urban sites and (some of) these may not show representative trends for the region. Even though the authors at page 6 line 8-14 discuss this a bit, and claim they (the Chinese urban sites) show similar trends, it is very few sites and only trends for the ratio is presented. I will suggest to separate urban and regional sites in the study, and even better calculate the trends at the individual sites rather than averaging all the sites, which have totally different siting criteria.

In this study, we aimed to present a comprehensive view of precipitation chemistry at the country scale for China, Korea, and Japan. As pointed out by the reviewer, EANET observation sites over China provide insight into long-term trends but are limited with respect to spatial coverage. To overcome this limitation, we combined these data with the IMPACTS data over southern China and coordinated observation campaign data around Beijing by LAPC/IAP/CAS. These data are limited with respect to the timescale and only provide snapshots; however, we believe that the comparison between these NO<sub>3</sub><sup>-</sup> and nss-SO<sub>4</sub><sup>2-</sup> concentrations and *Ratio* can provide important insights. By comparing EANET with these data (Figs. 2 (c) and (e); Fig. 5 (b)), we can confirm that the annual means are highly consistent among data types. Although the individual site and precipitation event might be affected by regional (or smaller scale) characteristics, precipitation chemistry will be triggered by large-scale characteristics of the annual mean. To reinforce this analysis, the separation of urban and regional (rural and remote) sites in China is important. In the revised manuscript, we have included the classification analysis in Figs. 2 (c) and (e) and Fig. 5 (b) for the NO<sub>3</sub><sup>-</sup> concentration in precipitation,

nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation, and *Ratio*. The correlations between mean values averaged over urban sites and rural and remote sites were 0.21, 0.78, and 0.65 for the NO<sub>3</sub><sup>-</sup> concentration in precipitation, nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation, and *Ratio*, respectively. The relatively low correlation for the NO<sub>3</sub><sup>-</sup> concentration in precipitation reflects the regional characteristics, and particularly differences in nitrogen species. A strong correlation was found between the nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation and *Ratio*; hence, we concluded that *Ratio* observed in the EANET network is a representative dataset for the precipitation chemistry in China. Please also see our responses to specific comments.

I also find the trend calculations a bit too optimistic since you calculate trends from only five years periods, usually one needs at least 7 years (preferably 10) for calculating trends. You have not described how you calculate trends and the significance, except that it is linear.

Linear trends during the three phases were estimated using linear regression, and significance levels were calculated using the Student's *t*-test. We have described this analysis in the revised manuscript on page 8, lines 17–18. Linear trends over five years were used to evaluate emission variation in each phase (Figs. 2 (b) and (d)). This time period is short, and we have established criteria to exclude outliers by carefully checking the original data. The analyses of individual sites will be interesting; however, due to the strict screening of outliers, data for some years are lacking for some sites. Therefore, our aim was to show average values for characteristics of precipitation chemistry over China, Korea, and Japan.

More specific comments to the text:

1) Page 1. The references Endo et al and Ban et al are mainly studying EANET data and not directly with US and Europe. You should rather use US and European references when comparison.

Some examples:

EMEP: Tørseth, K. et al. Introduction to the European Monitoring and Evaluation Programme (EMEP) and observed atmospheric composition change during 1972-2009. *Atmospheric Chemistry and Physics* 12, 5447-5481, doi:10.5194/acp-12-5447-2012 (2012).

CASTNET: Sickles Ii, J. E. & Shadwick, D. S. Air quality and atmospheric deposition in the eastern US: 20 years of change. *Atmos. Chem. Phys.* 15, 173-197, doi:10.5194/acp-15-173-2015 (2015).

IMPROVE: Hand, J. L., Schichtel, B. A., Malm, W. C. & Pitchford, M. L. Particulate sulfate ion concentration and SO<sub>2</sub>; emission trends in the United States from the early 1990s through 2010. *Atmos. Chem. Phys.* 12, 10353-10365, doi:10.5194/acp-12-10353-2012 (2012).

NADP: Lehmann, C. M. B., Bowersox, V. C., Larson, R. S. & Larson, S. M. Monitoring Long-term Trends in Sulfate and Ammonium in US Precipitation: Results from the National Atmospheric

Deposition Program/National Trends Network. Water, Air, & Soil Pollution: Focus 7, 59-66, doi:10.1007/s11267-006-9100-z (2007).

UNECE also has relatively new assessments of trends from Europe and North America, which might be relevant as well:

<http://www.unece.org/index.php?id=42861> and <http://www.unece.org/index.php?id=42947>

We appreciate the summary of references related to the US and European observation network. Based on these additional references, we have revised the sentence as follows on P2, L10-13:

“According to the Clean Air Status and Trends Networks (CASTNET) in the US (Sickles and Shadwick, 2015) and the European Monitoring and Evaluation Programme (EMEP) in Europe (Torseth et al., 2012), in Japan, which is located in the downwind region of the Asian continent, the total wet and dry deposition amounts have surpassed those of both the US and Europe (Endo et al., 2011; Ban et al., 2016).”

2) Page 3 line 13-16. I don't understand what you mean here that “the deposition were centred”. with a reference to Pan 2015, who looks at deposition of trace elements in Northern China. I assume the authors are discussing the point I address in the beginning with representativity of the EANET sites, but I don't understand how you can state that “The approach taken here will further promote our understanding of precipitation chemistry for all of China” since the additional sites only cover a short period and cannot be used for trends. If you had looked at one specific year to assess the spatial deposition it would have been different.

We have revised the reference to Pan et al. (2012, 2013), which discusses sulfur and nitrogen deposition based on ten coordinated observations around Beijing. We agree that the additional datasets for southern China and northern China were limited with respect to the time period; however, in combination with the EANET observations over four regions, we believe that it is possible to characterize the precipitation chemistry over China. We have explained this point in the revised text as follows on P3, L15-17:

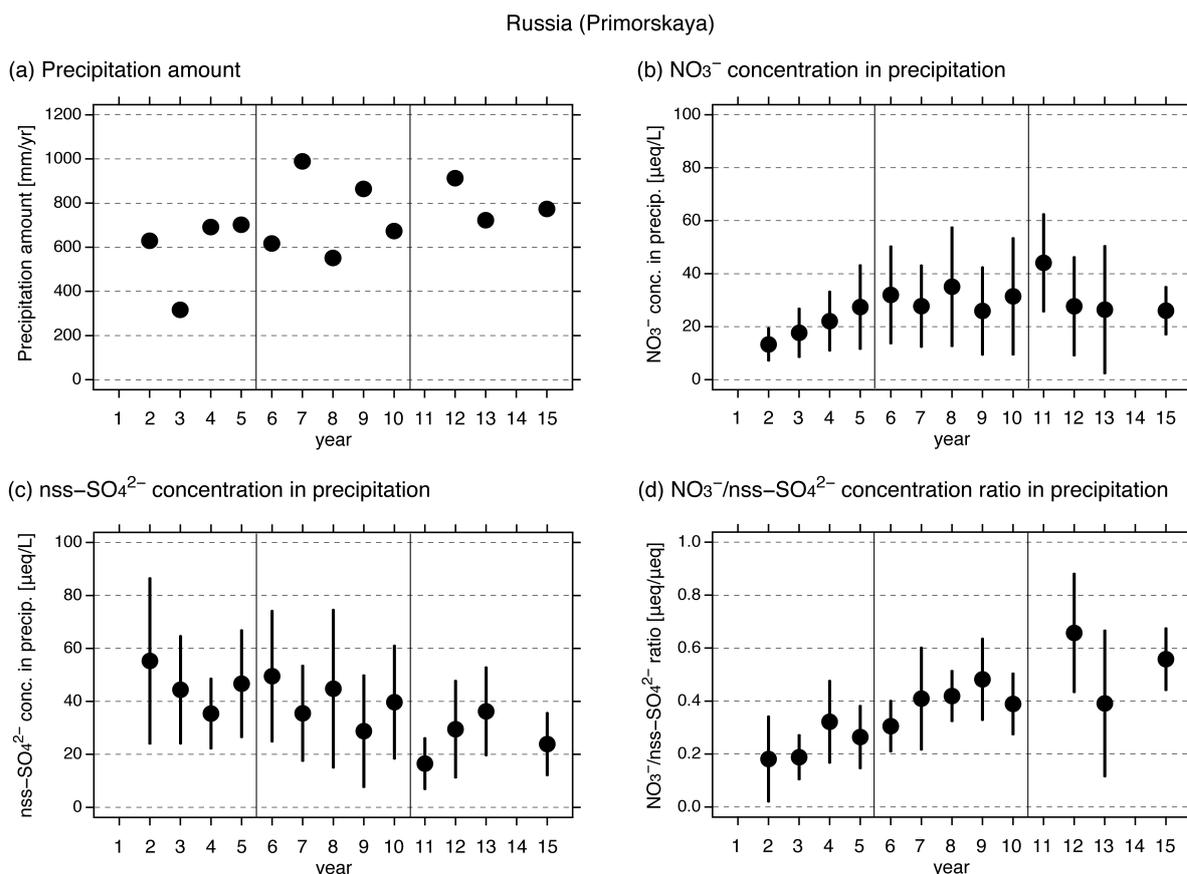
“Although these additional data over southern and northern China are limited to one year or a few years, this approach of combining and comparing observational datasets will further improve our understanding of precipitation chemistry over China, broadly.”

Considering the representativeness of EANET sites and the reviewer's comment, we have added analyses in which ten EANET sites over China, of which we classified as urban sites (n = 6) and rural and remote sites (n = 4).

3) Page 4 line 6. The Ogasawara site was excluded with quite strict criteria. 25% is more appropriate. Further you could have included the Russian site close to the Korean border, Primorskaya, which is also downwind from the large emission sources in China.

Because the Ogasawara islands are located on the Pacific Ocean, values below the detection limit were found, namely, 11.8% for  $\text{NO}_3^-$  and 7.7% for  $\text{nss-SO}_4^{2-}$ . For other observation sites in Japan, values are typically over the detection limit. Hence, we have excluded the Ogasawara site from this study.

Thank you for your comment regarding the site in Russia. In fact, Primorskaya showed similar trends to those observed in China, Korea, and Japan as is illustrated in the figure below. Analyses of the Russian site are beyond the scope of this study; hence, we prefer not to include these results in the manuscript.



Supplemental figure: Long-term temporal variation in (a) precipitation, (b)  $\text{NO}_3^-$  concentration in precipitation, (c)  $\text{nss-SO}_4^{2-}$  concentration in precipitation, and (d) *Ratio* are illustrated in the figure below.

4) Page 8 line 13-14: “For the treatment of precipitation amount, months where data were insufficient were the same as when applying the Smirnov-Grubbs test for Ratio calculation.” I don’t understand this sentence

We have revised the sentence as follows and moved to section 2.1 on P5, L11-14):

“For the treatment of the monthly precipitation amount, the Smirnov–Grubbs test for *Ratio* was used to discard outliers. If at least 9 months were available, the annual accumulated precipitation amount was calculated.”

5) Page 8 line 16-17 The sentence “Statistical analysis revealed that, except for the increasing and decreasing trend over China and Korea during Phase III .. there was no clear change in precipitation amount...” is somewhat in contrast to the conclusion later on page 13 line 21: “In spite of the increasing trends of precipitation amount, decreasing trends for nss-SO<sub>4</sub><sup>2-</sup> wet deposition amounts over China, Korea, and Japan were seen after 2005–2006”. The increase in precipitation amount in China is after 2010 (though the variability is very high) and South Korea has decreasing amount.

We apologize for the apparently conflicting statements. We have carefully revised the conclusion as follows on P14, L9-12:

“Despite the increasing trends in the precipitation amount, decreasing trends for nss-SO<sub>4</sub><sup>2-</sup> wet deposition over China were observed after 2005–2006 and matched reductions in SO<sub>2</sub> emissions from China. This decrease in China triggered the decreases in nss-SO<sub>4</sub><sup>2-</sup> wet deposition over Korea and Japan.”

6) Page 8-9 “The temporal variation found in the NO<sub>3</sub> concentration in precipitation did not correspond to the NO<sub>x</sub> emissions variation”. This non-linearity can be due to several factors. E.g.: The non representativity of the sites, the change in atmospheric composition and chemical regimes (changes in base cations and ammonium), oxidation capacity of the atmosphere, all may change the lifetime of NO<sub>x</sub> and NO<sub>3</sub> (course or fine).

We have added potential explanations for the nonlinearity as follows on P9, L20-22:

“This may be explained by changes in atmospheric composition and chemical regimes with changes in base cations and the oxidation capacity of the atmosphere, leading to changes in the lifetime of nitrogen species.”

7) Page 9 line 18. “Contamination”. It’s a bit misleading word. The satellite measurements are not contaminated they are influence by sources outside Korea (like China)

We have replaced “contamination” with “impacted” on P10, L7.

8) Page 9 line 23. Have the volcanic activity changed during the period to influence the trend of the SO<sub>2</sub> column data? Should maybe also have been included also in the emission inventories?

According to the volcanic activity report by Japan Meteorological Agency

(<http://www.data.jma.go.jp/svd/vois/data/tokyo/volcano.html>), the activity above south Kyushu (a western island of Japan) increased from Phase II to Phase III. This can be detected in satellite observations (Fig. 4), and the activity of Miyakejima, located south of Tokyo, gradually decreased after the eruption in 2000. As we have stated on P10, L32-34, relatively

high nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation in Japan in 2001–2002 were partially related to the activity of Miyakejima.

9) Page 9-10. IMPACTS sites have lower levels compared to EANET since these are mainly regional/rural sites while the BNU site is urban and naturally higher than the average EANET site.

We have clarified these points as follows on P9, L25-28:

“When comparing EANET observations with additional datasets, IMPACTS sites around south China showed lower NO<sub>3</sub><sup>-</sup> concentrations in precipitation, reflecting the characteristics of remote sites, while BNU and LAPC/IAP/CAS sites around Beijing showed relatively high concentrations, reflecting the emission intensity over northern China.”

10) Page 10: There are some contradicting statements:

1) Over China, NO<sub>2</sub>/SO<sub>2</sub> column ratio were flat during Phase I, sharply increasing during Phase II, and almost flat during Phase III.

2) NO<sub>x</sub>/SO<sub>2</sub> emission ratio were well correlated with the variation in Ratio over China

3) Ratio was almost 0.3 during Phase I and subsequently increased.. during Phase II, with a trend of  $+14.8 \pm 1.9\%/year$  and around 0.4–0.6 during Phase III with a trend of  $+10.1 \pm 3.8\%/year$  ( $p < 0.05$ ).

This leads to the conclusion on page 11. “Ratio observed in EANET network can be a representative dataset of China for the precipitation chemistry” This is not obvious for me

By separating the EANET observation sites, we clearly observed that the mean values for urban sites and for rural and remote sites were highly correlated (correlation coefficient, 0.65). The manuscript has been revised for clarity as follows on P11, L20-32:

“*Ratio* was almost 0.3 during Phase I and subsequently increased to 0.5 during Phase II, with a trend of  $+14.8 \pm 1.9 \%/year$  ( $p < 0.001$ ), and to approximately 0.4–0.6 during Phase III, with a trend of  $+10.1 \pm 3.8 \%/year$  ( $p < 0.05$ ). The analyses of different site types for EANET observations over China (Fig. 5 (b), see Table 1 for classification) revealed that the mean values for urban sites and rural and remote sites were highly correlated, with a correlation coefficient of 0.65. In addition to the variation of *Ratio* based on EANET observations for urban sites and rural and remote sites, other observations through IMPACTS, BNU, and LAPC/IAP/CAS also indicated similar variation in *Ratio*. Based on the IMPACTS dataset, which covered southern China with extensive data in 2001–2003, *Ratio* in China during Phase I was approximately 0.2, which was much lower than the current (Phase III) status. Although NO<sub>3</sub><sup>-</sup> and nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation were higher than those in the EANET observation dataset, observations at BNU revealed a *Ratio* of 0.31, which was consistent with EANET observation results. A coordinated observation network by LAPC/IAP/CAS, which

operated for three years around Beijing, indicated an increasing trend during Phase II from 0.39 in 2008 to 0.59 in 2010. These results obtained from multiple observations reinforce the idea that *Ratio* observed in the EANET network is representative of the precipitation chemistry in China.”

11) Page 12 line 31. Nr. Is this reactive nitrogen? It is not defined. Can also be interpret as reduced nitrogen. I assume reactive nitrogen since you at page 13 line 1 states that Nr can cause eutrophication and this means both NO<sub>3</sub> and NH<sub>4</sub>

Nr refers to reactive nitrogen; it has been defined on P13, L1.

12) Figure 7c. It is not clear how sum NO<sub>3</sub> and SO<sub>4</sub> is calculated Figure C do not add up from fig a and b, even if changing units. Maybe better to look at equivalent and not mass.

We have added a brief explanation of the calculation used in Figure 7 (c). Considering the suggestion, the revised manuscript shows the equivalent basis (keq/ha) for the sum of NO<sub>3</sub><sup>-</sup> and SO<sub>4</sub><sup>2-</sup>, and we have recalculated the NO<sub>3</sub><sup>-</sup> fraction on an equivalent basis.

13) Table 2. It would have been nice to indicate how many sites are included in these calculations, and the time period for the different phases should be included in the table so the reader don't need to search in the text. Is it the same number sites in all the phases? In the table there is no unit (kg/ha pr year?)

We have indicated the total numbers used to calculate the means and trends in Tables 2 and 3. The time periods are also included in Tables 2, 3, and 4. The methods used to calculate the linear trends and the units are included in the note.

# A 15-year record (2001–2015) of the ratio of nitrate to non-seasalt sulfate in precipitation over East Asia

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**Abstract.** Acidifying species in precipitation can have severe impacts on ecosystems. The chemical composition of precipitation is directly related to the amount of precipitation; accordingly, it is difficult to identify long-term variation in chemical concentrations. The ratio of the nitrate ( $\text{NO}_3^-$ ) to non-seasalt sulfate ( $\text{nss-SO}_4^{2-}$ ) concentration in precipitation on an equivalent basis (hereinafter, *Ratio*) is a useful index to investigate the relative contributions of these acidifying species. To identify the long-term record of acidifying species in precipitation over East Asia, the region with the highest emissions worldwide, we compiled ground-based observations of the chemical composition of precipitation over China, Korea, and Japan from 2001 to 2015 based on the Acid Deposition Monitoring Network in East Asia (EANET). The spatial coverage was limited, but additional monitoring data for Japan, southern China, and northern China around Beijing were utilized. The period of analysis was divided into three phases: Phase I (2001–2005), Phase II (2006–2010), and Phase III (2011–2015). The behaviors of  $\text{NO}_3^-$  and  $\text{nss-SO}_4^{2-}$  concentrations and hence the *Ratio* in precipitation were related to these precursors. The anthropogenic  $\text{NO}_x$  and  $\text{SO}_2$  emissions and the  $\text{NO}_x/\text{SO}_2$  emission ratio were analyzed. Further, satellite observations of the  $\text{NO}_2$  and  $\text{SO}_2$  column density to capture the variation in emissions were applied. We found that the long-term trend in the  $\text{NO}_3^-$  concentration in precipitation was not related to the variation in  $\text{NO}_x$  emission and the  $\text{NO}_2$  column. In comparison, the  $\text{nss-SO}_4^{2-}$  concentration in precipitation over China, Korea, and Japan was partially connected to the changes in  $\text{SO}_2$  emissions from China, but the trends were not significant. The long-term trends of *Ratio* over China, Korea, and Japan were nearly flat during Phase I, increased significantly during Phase II, and were essentially flat again during Phase III. This variation in *Ratio* in East Asia clearly corresponded to the  $\text{NO}_x/\text{SO}_2$  emission ratio and the  $\text{NO}_2/\text{SO}_2$  column ratio in China. The initial flat trend during Phase I was due to increases in both  $\text{NO}_x$  and  $\text{SO}_2$  emissions in China, the significantly increasing trend during Phase II was triggered by the increase in  $\text{NO}_x$  emissions and decrease in  $\text{SO}_2$  emissions in China, and the return to a flat trend during Phase III was caused by declines in both  $\text{NO}_x$  and  $\text{SO}_2$  emissions in China. These results

suggest that emissions in China had a significant impact not only on China, but also on downwind precipitation chemistry during the 15-year period of 2001–2015. In terms of wet deposition, the  $\text{NO}_3^-$  wet deposition over China, Korea, and Japan did not change dramatically, but the  $\text{nss-SO}_4^{2-}$  wet deposition declined over China, Korea, and Japan from Phase II to III. These declines were caused by a strong decrease in the  $\text{nss-SO}_4^{2-}$  concentration in precipitation accompanied by a reduction in  $\text{SO}_2$  emission from China, which counteracted the increase in precipitation. These findings indicated that the acidity of precipitation shifted from sulfur to nitrogen.

## 1 Introduction

Accompanying the recent acceleration of anthropogenic emissions in Asia, atmospheric deposition in Asia has been a major focus worldwide (Vet et al., 2014). According to the Clean Air Status and Trends Networks (CASTNET) in the US (Sickles and Shadwick, 2015) and the European Monitoring and Evaluation Programme (EMEP) in Europe (Torseth et al., 2012), in Japan, which is located in the downwind region of the Asian continent, the total wet and dry deposition have surpassed those of both the US and Europe (Endo et al., 2011; Ban et al., 2016). In terms of the wet deposition of nitrogen and sulfur compounds over Japan, the influence of anthropogenic emissions from China has been determined via chemical transport model simulations (Kuribayashi et al., 2012; Kajino et al., 2011; 2013; Morino et al., 2011).

The amount of wet deposition is affected by the amount of precipitation. For this reason, the ratio of the nitrate ( $\text{NO}_3^-$ ) concentration to non-seasalt sulfate ( $\text{nss-SO}_4^{2-}$ ) concentration in precipitation on an equivalence basis, hereafter referred to as *Ratio*, is useful for evaluating the relative contributions of nitrogen and sulfur to the acidity of precipitation. *Ratio* in Japan was about 0.41 between 1987 and 1990 (Fujita et al., 2003). Over western Japan, *Ratio* increased substantially by 0.09–0.17 between 1987 and 1996, reaching about 0.5 in the late 1990s. This corresponded to a large increase in  $\text{NO}_x$  emissions compared with  $\text{SO}_2$  emissions across East Asia (Takahashi and Fujita, 2000). In our previous studies (Itahashi et al., 2014a, 2015), we showed that the *Ratio* trend in precipitation over Japan remained flat at around 0.5–0.6 between 2000 and 2005, with a subsequent increase to 0.6–0.7 between 2006 and 2011. During this period, the  $\text{NO}_x/\text{SO}_2$  emission ratio in Japan constantly decreased, and changes in *Ratio* closely followed the changes in the  $\text{NO}_x/\text{SO}_2$  emission ratio in China. A significant increase in *Ratio* between 2006 and 2011 was also found over both China and Korea. A correlation between *Ratio* in China, Korea, and Japan and the  $\text{NO}_x/\text{SO}_2$  emission ratio in China was detected. A sensitivity simulation via a regional chemical transport model of Chinese anthropogenic emissions indicated the important contribution of  $\text{NO}_x$  emissions near the source region and the high impact of  $\text{SO}_2$  emissions over downwind regions. Identifying the importance of acidity via nitrogen and sulfur is critical to mitigate ecosystem effects, such as soil acidification (Zhao et al., 2009) and surface water acidification (Yamashita et al., 2016). Similar studies of *Ratio* (or the inverse of *Ratio*, defined as the ratio of the  $\text{nss-SO}_4^{2-}$  to  $\text{NO}_3^-$  concentration in precipitation) have been conducted in northern China (Wang et al., 2012), Nanjing (Tu et al., 2005), southeastern China (Cui et al., 2014), and the Pearl River Delta region (Lu et al., 2015).

We next consider recent changes in *Ratio* in East Asia according to recent changes in anthropogenic emissions. In our previous studies (Itahashi et al., 2014a, 2015), which considered data obtained in 2000–2011, emissions from China exhibited a transition from a continuously increasing trend. NO<sub>x</sub> emissions in China have been rising steadily (Kurokawa et al., 2013; Itahashi et al., 2014b), whereas SO<sub>2</sub> emissions in China peaked in 2005–2006 and subsequently declined due to the introduction of flue-gas desulfurization systems in China’s 11th Five-Year Plan (2006–2010) (Kurokawa et al., 2013; Itahashi et al., 2012a). Recently, it has been reported that NO<sub>x</sub> emissions in China declined after 2011–2012 (Irie et al., 2016; Xia et al., 2016; Krotkov et al., 2016; van der A et al., 2017). To determine the cause of this change, we revisited and updated an analysis of precipitation data in relation to emission variation over East Asia. We compiled precipitation observations over the 15-year period between 2001 and 2015. Considering the drastic change in emissions from China, the analysis period was divided into three parts: Phase I (2001–2005), Phase II (2006–2010), and Phase III (2011–2015). This study extends our previous studies (Itahashi et al., 2014a; 2015) in the following respects. First, this study incorporates additional data obtained over southern China and northern China around Beijing. The ground-based observation network in East Asia does not cover northern China, which is characterized by a large urban population and therefore contributes substantial anthropogenic emissions (e.g., Pan et al., 2012; 2013). **Although these additional data over southern and northern China are limited to one year or a few years, this approach of combining and comparing datasets will further improve our understanding of precipitation chemistry over country-wide China.** Additionally, this study uses satellite observations of NO<sub>2</sub> and SO<sub>2</sub> column density as a proxy to estimate recent emission changes. This can further enhance our knowledge of the current emission status. Analyses of long-term precipitation chemistry are still limited. For example, precipitation chemistry over periods of longer than 10 years has been evaluated for Guangzhou in 1983–2010 (Fang et al., 2013), Shenzhen in 1986–2006 (Huang et al., 2008), Lijiang in 1989–2006 (Zhang et al., 2012), Nanjing in 1992–2003 (Tu et al., 2005), and Tieshanping in 2001–2013 (Yu et al., 2017). This long-term, 15-year analysis provides a comprehensive overview of the precipitation chemistry over East Asia.

This paper is structured as follows. Section 2 introduces the dataset used in this study, which includes ground-based observations, emission inventories, and satellite observations that were applied as a proxy for emission levels. Section 3 is dedicated to the results and discussion. First, Section 3.1 presents the long-term trends for NO<sub>3</sub><sup>-</sup> and nss-SO<sub>4</sub><sup>2-</sup> in China, Korea, and Japan. Section 3.2 describes the intensive analysis of *Ratio* and its relation to emissions. Section 3.3 explores the analysis of wet deposition. Finally, Section 4 summarizes this research and outlines future perspectives.

## 2 Dataset

### 2.1 Ground-based observations

The ground-based observations used in this study are spatially mapped in Fig. 1. Detailed information for the observation period, including latitude ( $^{\circ}$ N), longitude ( $^{\circ}$ E), elevation (meters above sea level [a.s.l.]), and the classification of each site, is listed in Table 1. The observation dataset for the chemical composition of precipitation, which was compiled by the Acid Deposition Monitoring Network in East Asia (EANET) program, was mainly used in this study. In China, EANET observations were obtained over three areas in southern China (Zhuhai, Xiamen, and Chongqing) and one area in central China (Xi'an), for a total of 10 sites. In Korea, EANET observations from three sites (Cheju, Imsil, and Kanghwa) were available. In Japan, data for a total of 11 sites were available on EANET. Among these 11 sites, the data obtained at Ogasawara, which is located in the northwest Pacific Ocean ( $27.09^{\circ}$ N,  $142.22^{\circ}$ E, 230 m a.s.l.), includes estimates of the chemical composition of precipitation that are below the detection limit, that is, 11.8% for  $\text{NO}_3^-$  and 7.7% for  $\text{nss-SO}_4^{2-}$ . Therefore, the data from Ogasawara were excluded. Samples were collected using a wet-only sampler at a daily interval, except for Banryu (sampled weekly during 2001–2015) and Ijira (sampled weekly in 2001). Concentrations of  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  in precipitation were determined by ion chromatography and qualified via ion balance and conductivity agreement. The completeness of the data was determined from the precipitation coverage duration and total precipitation amount (EANET, 2000; 2010). In Japan, two sites (Ryori and Komae) were included in the study. Under the Global Atmosphere Watch (GAW) program of the World Meteorological Organization, the Japan Meteorological Agency (JMA) has been conducting observations of the atmospheric concentration and deposition at the northern, remote Ryori site since 1976; observations of deposition at Ryori ended in 2011. Samples were collected using a wet-only sampler at a daily interval at Ryori. The Central Research Institute of Electric Power Industry (CRIEPI) conducted continuous monitoring of precipitation chemistry. The CRIEPI data obtained at Komae, which is located near Tokyo, span 1987 to the present (Fujita et al., 2000). As we have more fully described in our previous studies (Itahashi et al., 2014a; 2015), CRIEPI monitoring was also conducted at Goto Island, which is located at the western edge of Japan. However, the period of coverage was limited to February 2000 to April 2003, so we excluded the Goto data from this study. Samples were collected using a wet-only sampler at 10-day intervals in Komae. For the EANET, JMA, and CRIEPI monthly mean datasets, the  $\text{nss-SO}_4^{2-}$  concentration (in mol/L) was calculated from the conservative assumption that sodium ( $\text{Na}^+$ ) is a seasalt tracer, using the following equation:

$$\text{nss-SO}_4^{2-} = \text{SO}_4^{2-} - 0.06028 \times \text{Na}^+. \quad (1)$$

To analyze the long-term behavior of the dataset over the 15-year period of 2001–2015, outliers in the observation of EANET, JMA, and CRIEPI were carefully examined according to method used in Itahashi et al. (2015). The volume-weighted monthly mean concentrations of  $\text{NO}_3^-$ ,  $\text{nss-SO}_4^{2-}$ , and *Ratio* in precipitation were analyzed using the Smirnov–Grubbs outlier test at each site. In this study, the outliers for *Ratio* were directly checked. In this method, outliers were

detected one at a time, assuming that the data likely followed an approximately normal distribution. The hypothesis of no outliers is rejected when

$$G = \frac{\max_{i=1,\dots,N}|C_i - \bar{C}|}{s} > \frac{N-1}{\sqrt{N}} \sqrt{\frac{t_{(\alpha/2N, N-2)}^2}{N-2+t_{(\alpha/2N, N-2)}^2}}, \quad (2)$$

where  $N$ ,  $\bar{C}$ , and  $s$  are the number, mean, and standard deviation, respectively, of  $\text{NO}_3^-$  or  $\text{nss-SO}_4^{2-}$  concentrations in precipitation or *Ratio* in precipitation ( $C_i$ ).  $t_{(\alpha/2N, N-2)}^2$  denotes the critical value of the  $t$ -distribution with  $(N - 2)$  degrees of freedom and a significance level of  $(\alpha/2N)$ . Outlier detection and removal were iterated until the dataset satisfied the specified significance level of 0.05. Following the above criteria, 4.7%, 3.8%, and 4.4% of data points for  $\text{NO}_3^-$ ,  $\text{nss-SO}_4^{2-}$ , and *Ratio* in precipitation, respectively, were discarded from the China dataset; 5.0%, 3.0%, and 3.0% of these were discarded from the Korea dataset; and 2.4%, 1.5%, and 1.6% of these were discarded from the Japan dataset. Finally, the annual mean concentrations of  $\text{NO}_3^-$ ,  $\text{nss-SO}_4^{2-}$ , and the annual mean *Ratio* in precipitation were calculated from the monthly mean data when at least 9 months of data were available for a given year at the site. [For the treatment of the monthly accumulated precipitation amount, the Smirnov–Grubbs test for \*Ratio\* was applied to discard outliers. If at least 9 months were available, the annual accumulated precipitation was calculated.](#)

These observation datasets taken from EANET, JMA, and CRIEPI were essentially the same as those used in our previous studies (Itahashi et al., 2014a; 2015). A limitation of our previous studies was a lack of spatial coverage over northern China, especially around the capital of Beijing, because EANET covered only the area from southern to central China (Fig. 1). It has been recognized that anthropogenic emissions centered over this region (Kurokawa et al., 2013; Li et al., 2017), the related atmospheric concentration, and depositions are severe in China. To overcome this limitation and advance our knowledge of precipitation chemistry over the whole of China, we evaluated additional sources of data for the chemical concentration of precipitation over China.

The integrated monitoring program on acidification of the Chinese terrestrial system (IMPACTS) was established through a Chinese–Norwegian cooperative project (Larssen et al., 2004; 2006) from 2001 to 2003. Under the IMPACTS program, atmospheric concentration, precipitation composition, and soil, water, and vegetative effects were studied at five forested sites (LXH, LGC, LCG, CJT, and TSP; refer to Fig. 1 and Table 1) over southern China. In terms of deposition, four measurements (wet-only, bulk, throughfall collected below the tree canopy, and below-ground vegetation) were obtained. Wet-only sampling data were used in this study. Observations of the chemical composition of precipitation by a wet-only sampler were reported for four sites (LGS, LCG, CJT, and TSP) from 2001 to 2003 and for one site (LXH) from 2002 to 2003. Data for LGS in 2001 were not used, owing to insufficient coverage. [At TSP, continuous monitoring results up to 2013 have recently been reported \(Yu et al., 2017\).](#)

The precipitation chemistry over northern China has important implications. The precipitation sample at Beijing Normal University (BNU) in 2003 (Sun et al., 2004; Tang et al., 2005) and the recent work by the State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry/Institute of Atmospheric Physics/Chinese Academy of

Sciences (LAPC/IAP/CAS) from December 2007 to November 2010 (Pan et al., 2012; 2013; Pan and Wang, 2015; Wang et al., 2012) were included in this study. Using the observation framework from BNU, data for a total of 53 rain events were collected at Beijing. To prevent contamination from dry deposition, the collector surface was covered with a plastic lid. A detailed description of this collection method is provided by Tang et al. (2005). In the recent work conducted by LAPC/IAP/CAS, a three-year observation from December 2007 to November 2010 was conducted at 10 sites around Beijing. Daily rainwater samples were collected using a customized wet/dry automatic collector. The precipitation sensor opened the collection funnel of the cover device when rainfall began. In this study, December 2007 to November 2008 was regarded as the year 2008, December 2008 to November 2009 was regarded as the year 2009, and December 2009 to November 2010 was regarded as the year 2010. For the IMPACTS and LAPC/IAP/CAS datasets, the nss-SO<sub>4</sub><sup>2-</sup> concentration was calculated using Na<sup>+</sup> as a seasalt tracer based on Eq. (1). Owing to the lack of information on Na<sup>+</sup> within the samples at BNU, SO<sub>4</sub><sup>2-</sup> was used; however, it was reported that Na<sup>+</sup> is not a major ion component of these samples.

## 2.2 Emission inventories

It is predicted that variation in NO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, and hence *Ratio* is directly related to the emission of NO<sub>x</sub> and SO<sub>2</sub>. We used the following emission inventories in this study. The Regional Emission inventory in ASia (REAS) version 2.1 (Kurokawa et al., 2013), which covers Asia from 2000 to 2008, was the main dataset used in this study to obtain NO<sub>x</sub> and SO<sub>2</sub> emissions over China, Korea, and Japan. The recent status of Asian anthropogenic emissions was assessed by harmonizing different local emission inventories with a mosaic approach named MIX (Li et al., 2017). MIX covers the years 2008 and 2010. The data from MIX were also used to acquire data for NO<sub>x</sub> and SO<sub>2</sub> emissions over China, Korea, and Japan.

The latest country-level status of emissions in China are described by Xia et al. (2016). Their emissions data cover the period from 2000 to 2014. The primary case, which analyzed the contributions of advanced combustors with improved energy efficiency and air pollutant control devices with improved pollutant removal efficiency, was used in this study. In Korea, the National Institute of Environmental Research (NIER) provided estimates of the national emissions via the National Air Pollutants Emission Service, and the latest data reported cover 1999 to 2013 (NIER, 2017). In Japan, the Japan Auto-Oil Program (JATOP) provided five-year interval emission datasets beginning in 1995 (JATOP, 2012a; 2012b). The datasets for 2005 and 2010 were used in this study.

Using the total amounts of NO<sub>x</sub> and SO<sub>2</sub> emissions estimated via inventories over China, Korea, and Japan, the NO<sub>x</sub>/SO<sub>2</sub> emission ratio on a molar basis was calculated by calculating NO<sub>x</sub> as NO<sub>2</sub>. The behavior of the NO<sub>x</sub>/SO<sub>2</sub> emissions ratio is correlated to *Ratio* in precipitation.

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### 2.3 Satellite observations

The emission inventories had some time lags. Accordingly, satellite observations of the NO<sub>2</sub> and SO<sub>2</sub> vertical column density were combined to capture the recent status of NO<sub>x</sub> and SO<sub>2</sub>. The NO<sub>x</sub>/SO<sub>2</sub> column ratio reflects the NO<sub>x</sub>/SO<sub>2</sub> emission ratio and was effective for characterizing the correspondence with *Ratio* in precipitation. Recently, satellite observations have  
5 been widely used as a proxy for emissions data. The NO<sub>2</sub> column has been used to capture NO<sub>x</sub> emissions (e.g., Miyazaki et al., 2012; Mijling et al., 2013; Itahashi et al., 2014b; Han et al., 2015; Irie et al., 2016) and SO<sub>2</sub> column has been used for SO<sub>2</sub> emissions (e.g., Lee et al., 2011; Li et al., 2010) and/or volcanic eruptions (e.g., Brenot et al., 2014). Several studies have indicated the importance of different technologies to control emissions (e.g., Li et al., 2010; Wang et al., 2015; Krotkov et al., 2016; van der A et al., 2017). For example, the ratio of ozone monitoring instrument (OMI)-derived SO<sub>2</sub>/NO<sub>2</sub> was  
10 used to determine the effectiveness of the flue-gas desulfurization devices for power plants in China (Li et al., 2010; Wang et al., 2015).

The NO<sub>2</sub> and SO<sub>2</sub> column dataset, which was observed by OMI onboard the National Aeronautics and Space Administration (NASA) Earth Observing System Aura satellite, was used in this study (NASA, 2017). The Aura satellite was launched on  
15 15 July 2004 in a sun-synchronous ascending polar orbit with a local equator crossing time of 13:45 ± 0:15. During the data period, it measured sunlight backscattered from the Earth over a wide range of ultraviolet and visible wavelengths to derive abundances of ozone and other trace gases important for air quality and climate. Science-quality data operations began on 1 October 2004; hence, the data from 2005 to 2015 were used to cover our analysis period of 2001–2015. Retrieval algorithms were based on the products provided by NASA.

In terms of the NO<sub>2</sub> column, we used the most recent version of level 3 daily global nitrogen dioxide product (OMNO2d, version 3.0), which was released in August 2016 and is gridded at a resolution of 0.25° × 0.25° (Krotkov et al., 2013). This  
20 product contains the total and tropospheric column for all atmospheric conditions and for sky conditions where the cloud fraction is less than 30%. We analyzed the tropospheric column with clouds screened on the condition of a cloud fraction of less than 30%.

In terms of the SO<sub>2</sub> column, we used a level 3 daily global sulfur dioxide product (OMSO2e) of the latest version (3.0),  
25 which was released in February 2015 and is gridded at a resolution of 0.25° × 0.25° (Krotkov et al., 2015). The dataset contains the total column of SO<sub>2</sub> in the planetary boundary layer. The algorithm was based on a principal component analysis, as introduced by Li et al. (2013). Cloud fraction, scene number, solar and satellite viewing angles, and row anomaly flags were provided as ancillary parameters. The data filtering of this level 3 dataset included the removal of rows with any  
30 of the following: anomaly flags, radiative cloud fraction greater than 20%, solar zenith angle greater than 70.0°, or scene number greater than 58 or less than 3. In addition, we adopted the smoothed method to average out the noise levels of the SO<sub>2</sub> column, following the research of Koukouli et al. (2016), who provided the anthropogenic loading of SO<sub>2</sub> over China as obtained from different satellite sensors. This method smoothed the SO<sub>2</sub> column assigned to each of the 0.25° × 0.25° grid cells, which were weighted by the SO<sub>2</sub> column of the surrounding eight cells. In this process, the negative values were

regarded as zero values because our aim was to construct the NO<sub>x</sub>/SO<sub>2</sub> column ratio from satellite observations. Two data periods associated with volcanic activities were excluded from our analysis as follows. The Sarychev Volcano in the Kuril Islands (48.09°N, 153.20°E) had an explosive eruption that emitted a huge amount of ash and SO<sub>2</sub> at altitudes of 10–16 km (Brenot et al., 2014). The data from 14–22 June 2009 included this large amount of SO<sub>2</sub> (over 10 D.U.) in and around the analyzed domain; accordingly, they were discarded from the calculations of monthly and annual means. An eruption of the Nabro Volcano in Eritrea (13.37°N, 41.07°E) from 12 June to 7 July 2011 was also reported. During the night of 12 June 2011, this volcano started to erupt, and on 14 June 2011, it spewed a volcanic plume across the route of many flights over east Africa and the Middle East (Brenot et al., 2014). The data from 15 June to 9 July 2011 were excluded from the calculations of monthly and annual means according to the approach detailed in van der A et al. (2017). In Japan, where many active volcanoes are located, SO<sub>2</sub> was continuously emitted at a level that surpassed anthropogenic emissions (e.g., Itahashi et al., 2017a). Owing to the difficulties of attempting to separate the effect of volcanic activity, the SO<sub>2</sub> column data for Japan, excluding the two data periods mentioned above, were used.

Based on the daily gridded data of NO<sub>2</sub> and SO<sub>2</sub> columns, monthly averages were calculated first and then the annual averages were calculated. In the calculation of annual averages, cells with monthly averaged data not available for at least 9 months were regarded as deficient cells for consistency with the criteria adopted for the ground-based observations of NO<sub>3</sub><sup>-</sup>, nss-SO<sub>4</sub><sup>2-</sup>, and *Ratio* in precipitation. The NO<sub>x</sub>/SO<sub>2</sub> column ratio was obtained from the annual averaged gridded data for NO<sub>2</sub> and SO<sub>2</sub> columns.

### 3 Results and Discussion

#### 3.1 Long-term trend in the chemical composition of precipitation with respect to emissions

The long-term trends in the precipitation amount, NO<sub>x</sub> emission from NO<sub>2</sub> columns, NO<sub>3</sub><sup>-</sup> concentration in precipitation, SO<sub>2</sub> emission from SO<sub>2</sub> columns, and nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation in 2001–2015 over China, Korea, and Japan are shown in Fig. 2. The averaged values and statistical analyses to estimate linear trends during Phase I (2001–2005), II (2006–2010), and III (2011–2015) are summarized in Table 2. Linear trends during each of the three phases were analyzed using linear regression, and significance levels were calculated using the Student's *t*-test. To support our finding of variation in the spatial distribution, satellite observations from 2005 to 2015 were used. NO<sub>2</sub> and SO<sub>2</sub> columns were mapped for 2006–2007 (the first half of Phase II), 2010–2011 (the transition from Phase II to Phase III), and 2014–2015 (the latter half of Phase III), as shown in Fig. 3. The annual variation in NO<sub>2</sub> and SO<sub>2</sub> columns, based on linear regression analyses during Phases II and III, are also mapped in Fig. 4.

As shown in Fig. 2 (a), year-to-year variation in precipitation was found and the average values for annual precipitation accumulation were around 1300, 1100, and 1500 mm/year over China, Korea, and Japan, respectively. A statistical analysis

revealed that, except for the increasing and decreasing trends over China and Korea during Phase III ( $p < 0.05$ ), there were no clear changes in precipitation levels during the 15-year period.

The NO<sub>x</sub> emissions and satellite observations of NO<sub>2</sub> columns are shown in Fig. 2 (b). Over China, emissions increased from 2001 to 2010, peaked in 2011–2012, and decreased after 2012, consistent with previous findings (e.g., Irie et al., 2016; Krotkov et al., 2016; van der A et al., 2017). Based on the spatial distribution, the NO<sub>2</sub> column above China peaked during the transition from Phase II and III (left panel of Fig. 3 (b)) and contrasting trends, that is, increasing and decreasing trends, were revealed during Phases II and III (left panel of Fig. 4 (a) and (b)). Over China, the NO<sub>x</sub> emissions of REAS and Xia et al. (2016) as well as satellite observations of NO<sub>2</sub> columns were highly consistent over long-term observations. Over Korea, NO<sub>x</sub> emissions obtained from REAS and NIER showed a slight increase in 2003–2004 and a flat trend after 2008. There were some inconsistencies between NO<sub>x</sub> emissions and the NO<sub>2</sub> column in 2005–2007. On the one hand, NO<sub>x</sub> emissions showed a decreasing trend; on the other hand, the NO<sub>2</sub> column showed a flat trend. Over Japan, both NO<sub>x</sub> emissions and the NO<sub>2</sub> column exhibited a slight decrease during the 15-year period. Overall, the correlation between NO<sub>x</sub> emissions and NO<sub>2</sub> columns suggests that satellite observations of the NO<sub>2</sub> column can serve as a proxy for NO<sub>x</sub> emissions.

Such variation in NO<sub>x</sub> emissions should be related to the change in the NO<sub>3</sub><sup>-</sup> concentration in precipitation. The long-term trend in the NO<sub>3</sub><sup>-</sup> concentration in precipitation is shown in Fig. 2 (c). The NO<sub>3</sub><sup>-</sup> concentration in precipitation decreased in the order of China, Korea, Japan, with concentrations of approximately 50 µeq/L in China and 40 µeq/L in Korea. These levels were around twofold greater than that in Japan, 15–20 µeq/L. The temporal variation found in the NO<sub>3</sub><sup>-</sup> concentration in precipitation did not correspond to the variation in NO<sub>x</sub> emissions (Fig. 2 (b)). This was particularly notable for China; although NO<sub>x</sub> emissions from 2001 to 2010 doubled, the NO<sub>3</sub><sup>-</sup> concentration in precipitation did not increase substantially. The NO<sub>3</sub><sup>-</sup> concentration in precipitation over Korea and Japan also did not show dramatic variation. [This may be explained by changes in atmospheric composition and chemical regimes with changes in base cations and the oxidation capacity of the atmosphere, leading to changes in the lifetime of nitrogen species.](#) Statistical analyses of the NO<sub>3</sub><sup>-</sup> concentration in precipitation (Table 2) did not indicate a significant overall trend for Phases I–III, with values remaining within -5 to +3 %/year over China, Korea, and Japan, with the exception of a negative trend of  $-3.4 \pm 1.3$  %/year ( $p < 0.05$ ) in Japan during Phase II. When comparing EANET observations with additional datasets, IMPACTS sites around south China showed lower NO<sub>3</sub><sup>-</sup> concentrations in precipitation, [reflecting the characteristics of remote sites](#), while BNU and LAPC/IAP/CAS sites around Beijing showed relatively high concentrations, [reflecting the emission intensity over northern China](#). These levels reflected the emission intensity of the particular regions and were within nearly one standard deviation of values obtained from the EANET observation network. [For China, 10 EANET sites were classified into an urban category \(six sites\) and rural and remote sites \(four sites\) \(Fig. 2 \(c\), see Table 1 for classification\). Generally, concentrations were higher for urban sites than for rural and remote sites, and the correlation coefficient for mean values over urban sites and rural and remote sites was 0.21; these results can be attributed to differences in regional characteristics.](#)

SO<sub>2</sub> emissions and satellite observations of the SO<sub>2</sub> column are shown in Fig. 2 (d). Over China, a steady increase was detected for 2001–2004, with a peak in 2005–2006 and a decrease after 2006; there thus seems to be a 5-year time-lag prior

to the decrease in NO<sub>x</sub> emissions and NO<sub>2</sub> columns. The SO<sub>2</sub> column, which was available from 2005, showed a continuous decline over all of China (center of Fig. 3) and a decreasing trend was estimated by linear regression (center of Fig. 4 (a) and (b)). In China, the SO<sub>2</sub> emissions of REAS and Xia et al. (2016) and SO<sub>2</sub> columns were highly consistent with the observed variation. Over Korea, SO<sub>2</sub> emissions obtained from REAS and NIER showed a slight decrease, which was consistent with the SO<sub>2</sub> column reduction, but some discrepancies were found for 2007–2008. The temporal variation in the SO<sub>2</sub> column in Korea was similar to that observed in China. The spatial distribution data (Fig. 3) suggested that the SO<sub>2</sub> column over Korea is impacted by upwind regions, e.g., China. The SO<sub>2</sub> column obtained from SO<sub>2</sub> PBL products have a center of mass altitude at about 900 m, which is expected to lead to longer transport distances and longer lifetimes of SO<sub>2</sub>. Therefore, compared with the correlation between NO<sub>x</sub> emission and NO<sub>2</sub> columns, it was difficult to clearly capture the relation between SO<sub>2</sub> emissions and SO<sub>2</sub> columns. Over Japan, SO<sub>2</sub> emissions exhibited a slight decrease during the 15-year period, whereas a nearly flat trend was found for the SO<sub>2</sub> column. This is partially related to the fact that the SO<sub>2</sub> column includes SO<sub>2</sub> emissions from volcanoes. Hot spots of the SO<sub>2</sub> column above south Kyushu (a western island of Japan) and central Honshu (near Tokyo), shown in Fig. 3, were related to the location of volcanic activity during this period.

The Nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation also decreased in the order of China, Korea, and Japan, similar to the NO<sub>3</sub><sup>-</sup> concentration in precipitation. The Nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation was 200–400 µeq/L above China, around 60 µeq/L above Korea, and 30 µeq/L above Japan. The levels over China were about tenfold greater than those of Japan, and the levels over Korea were almost twice those over Japan (Fig. 2 (e)). SO<sub>2</sub> emissions from China reached a peak in 2005–2006 (Fig. 2 (d)) and the nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation above China reached approximately 393.0 µeq/L in 2006, with a decline in 2015. The nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation above China were around 230 µeq/L during Phase I and II, but were 171.5 µeq/L during Phase III (Table 2). Trends calculated by linear regression were  $12.7 \pm 8.3$  %/year (not significant) during Phase I,  $-20.3 \pm 8.8$  %/year (not significant) during Phase II, and  $-13.6 \pm 5.2$  %/year ( $p < 0.05$ ) during Phase III. It seems that the variation in the nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation was partially related to SO<sub>2</sub> emission changes. In comparing EANET observations with other datasets, IMPACTS sites around south China showed lower nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation, BNU sites showed higher concentrations, and LAPC/IAP/CAS sites showed similar values. These levels were within one standard deviation of the means obtained from the EANET observation network, similar to the NO<sub>3</sub><sup>-</sup> concentration in precipitation. By classifying 10 EANET sites over China as urban sites or rural and remote sites (Fig. 2 (e), see Table 1 for classification), the mean values over urban sites and rural and remote sites were highly correlated, with a correlation coefficient of 0.78. This result implied that the nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation was largely determined by large-scale characteristics related to SO<sub>2</sub> emissions. Variation in SO<sub>2</sub> emission from Korea and Japan were constant or declined between 2000 and 2008, and the nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation above Korea and Japan did not exhibit a clear relation with SO<sub>2</sub> emissions in these countries. In addition, the increasing trend of  $+10.0 \pm 0.8$  %/year ( $p < 0.001$ ) over Korea during Phase I was not related to the variation in SO<sub>2</sub> emissions in Korea. Relatively high nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation in Japan in 2001–2002 were partially explained by the SO<sub>2</sub> emissions from Miyakejima volcano in 2000 (Itahashi et al., 2012, 2014a, 2015). With the exception of an increase in Korea during Phase I and the peak in Japan in

2001–2002, nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation above Korea and Japan were high in 2005–2007 and subsequently decreased. These patterns were similar to those over China and appeared to be connected to the variation in SO<sub>2</sub> emissions from China, though these trends were not significant.

### 5 3.2 Long-term trend in *Ratio* in precipitation and its relationship to changes in emissions in China

The long-term variation in precipitation and NO<sub>3</sub><sup>-</sup> and nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation did not indicate clear or significant trends during the 15-year period above China, Korea, and Japan, with a few exceptions. To further clarify the correlations between anthropogenic emission changes and the precipitation chemistry over East Asia, we focus on *Ratio*. The long-term trends in *Ratio* were analyzed for 2001–2015 and the results are shown in Fig. 5. The NO<sub>x</sub>/SO<sub>2</sub> emission ratio and NO<sub>2</sub>/SO<sub>2</sub> column ratio were calculated using average values over each country. Over China, both the NO<sub>x</sub>/SO<sub>2</sub> emission ratio and NO<sub>2</sub>/SO<sub>2</sub> column ratio were flat during Phase I, sharply increased during Phase II, and were nearly flat during Phase III. The substantial increase during Phase II was caused by an increase in NO<sub>x</sub> emissions and a decrease in SO<sub>2</sub> emissions (Fig. 2 (b) and (d)), as discussed in Itahashi et al. (2015). It was interesting to find that the trends in the NO<sub>x</sub>/SO<sub>2</sub> emission ratio and the NO<sub>2</sub>/SO<sub>2</sub> column ratio were again flat during Phase III. This can be explained by the declines in both NO<sub>x</sub> and SO<sub>2</sub> emissions. Based on the spatial mapping of the NO<sub>2</sub>/SO<sub>2</sub> column ratio (right part of Fig. 3), the NO<sub>2</sub>/SO<sub>2</sub> column ratio was substantially lower than 1.0 (red in the right panels of Fig. 3) above China during the first half of Phase II, and subsequently increased above 1.0 (yellow to green in the right panels of Fig. 3) during the latter half of Phase II and into Phase III. An increase was found over all of China during Phase II, but only a slight increase over central China and decreases over north and south China were observed during Phase III (right panels of Fig. 4 (a) and (b)). Such changes in the NO<sub>x</sub>/SO<sub>2</sub> emission ratio were highly correlated with the variation in *Ratio* over China. *Ratio* was almost 0.3 during Phase I and subsequently increased to 0.5 during Phase II, with a trend of  $+14.8 \pm 1.9$  %/year ( $p < 0.001$ ), and to approximately 0.4–0.6 during Phase III, with a trend of  $+10.1 \pm 3.8$  %/year ( $p < 0.05$ ). The analyses of different site types for EANET observations over China (Fig. 5 (b), see Table 1 for classification) revealed that the mean values for urban sites and rural and remote sites were highly correlated, with a correlation coefficient of 0.65. In addition to the variation of *Ratio* based on EANET observations for urban sites and rural and remote sites, other observations through IMPACTS, BNU, and LAPC/IAP/CAS also indicated similar variation in *Ratio*. Based on the IMPACTS dataset, which covered southern China with extensive data in 2001–2003, *Ratio* in China during Phase I was approximately 0.2, which was much lower than the current (Phase III) status. Although NO<sub>3</sub><sup>-</sup> and nss-SO<sub>4</sub><sup>2-</sup> concentrations in precipitation were higher than those in the EANET observation dataset, observations at BNU revealed a *Ratio* of 0.31, which was consistent with EANET observation results. A coordinated observation network by LAPC/IAP/CAS, which operated for three years around Beijing, indicated an increasing trend during Phase II from 0.39 in 2008 to 0.59 in 2010. These results obtained from multiple observations reinforce the idea that *Ratio* observed in the EANET network is representative of the precipitation chemistry in China.

In our previous studies (Itahashi et al., 2014a; 2015), we have highlighted the impact of the NO<sub>x</sub>/SO<sub>2</sub> emission ratio in China on the *Ratio* over downwind countries, including Korea and Japan. NO<sub>x</sub>/SO<sub>2</sub> emissions and the NO<sub>2</sub>/SO<sub>2</sub> column ratio in Korea showed complex variation. These parameters increased during Phase I, stayed essentially flat during Phase II, and increased slightly or remained flat during Phase III. It should be noted that the NO<sub>x</sub>/SO<sub>2</sub> emission ratio and NO<sub>2</sub>/SO<sub>2</sub> column ratio were not correlated in some parts of Korea. In particular, a flat trend was observed for the NO<sub>x</sub>/SO<sub>2</sub> emission ratio from 2006 to 2007, with a decline in the NO<sub>2</sub>/SO<sub>2</sub> column ratio; later, there was a flat trend for the NO<sub>x</sub>/SO<sub>2</sub> emission ratio from 2010 to 2011, with an increase of the NO<sub>2</sub>/SO<sub>2</sub> column ratio. These discrepancies are attributed to complications from using the SO<sub>2</sub> column as a proxy for SO<sub>2</sub> emissions over Korea. *Ratio* over Korea was almost 0.6 during Phase I, with a trend of  $-4.9 \pm 1.9$  %/year ( $p < 0.05$ ). This subsequently increased to 0.8 during Phase II, with a trend of  $+13.6 \pm 4.7$  %/year ( $p < 0.05$ ), and around 0.8–1.0 during Phase III, with a trend of  $+3.9 \pm 2.6$  %/year (not significant). Considering the NO<sub>x</sub>/SO<sub>2</sub> emission ratio during Phase II, the variation in *Ratio* in Korea may be connected to the variation in *Ratio* in China.

In Japan, both the NO<sub>x</sub>/SO<sub>2</sub> emission ratio and NO<sub>2</sub>/SO<sub>2</sub> column ratio exhibited declining trends during the 15-year period. However, the long-term variation in *Ratio* did not show a decreasing trend. The *Ratio* over Japan was almost 0.6 during Phase I, with a trend of  $-0.3 \pm 1.9$  %/year (not significant), subsequent increased to 0.7 during Phase II, with a trend of  $+3.6 \pm 1.3$  %/year ( $p < 0.05$ ), and remained level at approximately 0.7 during Phase III, with a trend of  $+2.5 \pm 1.3$  %/year (not significant). Considering the continuous decline in the NO<sub>x</sub>/SO<sub>2</sub> emission ratio during the 15-year period, the variation in *Ratio* in Japan appears to be related to the variation in *Ratio* in China.

The relationships between *Ratio* over China and *Ratio* over Korea or Japan are displayed in a scatter-plot (Fig. 6). These data clearly illustrate that the *Ratio* over China and Korea and the *Ratio* over China and Japan dramatically changed over time. Along with the variation in the NO<sub>x</sub>/SO<sub>2</sub> emission ratio and NO<sub>2</sub>/SO<sub>2</sub> column ratio in China, *Ratio* in precipitation over China was flat, then increased, and returned to a flat trend during the 15-year period. The variation over China was related to patterns of variation over Korea and Japan, with correlation coefficients of 0.84 and 0.81, respectively ( $p < 0.001$ ). These results revealed the impact of emission changes in China and the variation in precipitation chemistry at a regional scale in East Asia.

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### 3.3 Long-term trend in wet deposition

Finally, we focused on wet deposition, which was calculated by multiplying the chemical concentration in precipitation with the precipitation amount. Monthly data were checked via an outlier test for the chemical concentration, and the annual accumulated wet deposition was computed for years in which there was at least 9 months of coverage. The long-term temporal variation in NO<sub>3</sub><sup>-</sup> and nss-SO<sub>4</sub><sup>2-</sup> wet deposition over China, Korea, and Japan are shown in Fig. 7. Statistical analyses of averages values and trends during Phases I, II, and III are summarized in Table 4. NO<sub>3</sub><sup>-</sup> wet deposition amounts were approximately 6, 4, and 3 kg-N/ha over China, Korea, and Japan, respectively, and the year-to-year variation was high. In a statistical analysis of the NO<sub>3</sub><sup>-</sup> wet deposition amount, there were no significant trends for China, Korea, and Japan

during each phase (Fig. 7 (a)). With respect to reactive nitrogen (Nr) deposition, the threshold value of 10 kg-N/ha has been established (e.g., Bleeker et al., 2011). In the 15-year long-term analysis, wet deposition of  $\text{NO}_3^-$  accounted for more than half of this threshold Nr deposition in China and about one-third in Korea and Japan.

The  $\text{nss-SO}_4^{2-}$  wet deposition exhibited decreasing trends during Phase II and III. The amount over China was around 30 kg-S/ha during Phase I and below 20 kg-S/ha during III, and the amounts over Korea and Japan decreased to about 2 and 1 kg-S/ha, respectively. Based on a statistical analysis,  $\text{nss-SO}_4^{2-}$  wet deposition showed increasing (non-significant) trends during Phase I and decreasing trends during Phases II and III. Decreasing trends over China and Japan during Phase II and over Korea during Phase III were significant ( $p < 0.05$ ) (Fig. 7 (b)). It should be emphasized that precipitation over China, Korea, and Japan during Phase II and over China during Phase III increased. Our results indicate that the decreasing trends in the  $\text{nss-SO}_4^{2-}$  wet deposition were caused by a strong decline in the  $\text{nss-SO}_4^{2-}$  concentration in precipitation, which counteracted the increase in overall precipitation. The reduction in  $\text{SO}_2$  emissions over Korea and Japan might partially contribute to the decline in the  $\text{nss-SO}_4^{2-}$  concentration in precipitation. However, taking into account the correlation between precipitation chemistry over East Asia and emission changes in China, the  $\text{SO}_2$  emission reduction in China after 2005–2006 strongly impacted both local deposition and downwind deposition.

To consider the impact of wet deposition from  $\text{NO}_3^-$  and  $\text{nss-SO}_4^{2-}$ , the sum was calculated and is shown in Fig. 7 (c) as an equivalent basis; additionally, the fraction of this sum attributed to  $\text{NO}_3^-$  wet deposition was also analyzed (expressed as a percentage). For the 15-year analysis period, it has been suggested that  $\text{nss-SO}_4^{2-}$  wet deposition is reduced in East Asia. The results indicated that the fraction of  $\text{NO}_3^-$  wet deposition increased gradually during the 15-year period over East Asia, indicating that further attention should be paid to the deposition of  $\text{NO}_3^-$  and related Nr species. These results are consistent with those of previous analyses of several cities in China, including Nanjing (Tu et al., 2005), Shenzhen (Huang et al., 2008), Guangzhou (Fang et al., 2013), the Pearl River Delta region (Lu et al., 2015), Tieshanping, which is one of the IMPACTS sites (Yu et al., 2017), and Beijing (Wang et al., 2012). Our study reinforced these inferences at a larger scale over East Asia. Excess Nr deposition will result in eutrophication, which normally reduces biodiversity. Transboundary nitrogen air pollution and deposition have been found over East Asia (Itahashi et al., 2016, 2017b). To reduce their impact, further understanding of  $\text{NH}_3$  emissions will be important to develop a multipollutant control strategy (Zhao et al., 2009). In China, the pH of precipitation has not been lower than those of Korea and Japan due to substantial buffering species (Duan et al., 2016); therefore, base cations should be included in forthcoming analyses.

#### 4 Summary and Future Perspectives

In this study, the chemistry of precipitation in East Asia in 2001–2015 was evaluated, with a focus on the behavior of *Ratio* (defined as the  $\text{NO}_3^-/\text{nss-SO}_4^{2-}$  concentration in precipitation). The monitoring networks over East Asia were initially used, and the study was augmented by intense observation campaigns over southern and northern China.  $\text{NO}_2$  and  $\text{SO}_2$  column density satellite observations were used as proxies for  $\text{NO}_x$  and  $\text{SO}_2$  emissions, and this was a key update from our previous

studies (Itahashi et al., 2014a; 2015). The results for the  $\text{NO}_3^-$  concentration in precipitation suggested that there is no clear relation between  $\text{NO}_x$  emissions and corresponding trends in China, Korea, and Japan. In comparison, the  $\text{nss-SO}_4^{2-}$  concentration in precipitation was partially correlated to the  $\text{SO}_2$  emissions in China, and a corresponding (but non-significant) decline in Korea and Japan was also found. The analysis of *Ratio* clarified the trend (i.e., flat, followed by an increase and subsequent flattening) during the 15-year period in China, Korea, and Japan, and this variation was correlated to the  $\text{NO}_x/\text{SO}_2$  emission ratio and the  $\text{NO}_2/\text{SO}_2$  column ratio in China. The initial flat trend was due to the increases in both  $\text{NO}_x$  and  $\text{SO}_2$  emission in China, the subsequent increase was caused by the increase in  $\text{NO}_x$  emissions and the decrease of  $\text{SO}_2$  emission in China, and the eventual flat trend was due decreases in both  $\text{NO}_x$  and  $\text{SO}_2$  emission in China. Due to the impact of upwind regions, it is difficult to use the  $\text{SO}_2$  column as an accurate proxy of  $\text{SO}_2$  emissions in Korea. **Despite the increasing trends in the precipitation amount, decreasing trends for  $\text{nss-SO}_4^{2-}$  wet deposition over China were observed after 2005–2006 and matched reductions in  $\text{SO}_2$  emissions from China. This decrease in China triggered the decreases in  $\text{nss-SO}_4^{2-}$  wet deposition over Korea and Japan.** Both nitrogen and sulfur compounds pose acidification risks in ecosystems via atmospheric deposition processes. During the 15-year period analyzed in this study, sulfur wet deposition in East Asia decreased. This reduction seems to be important for future patterns, and additional studies focused on nitrogen are needed. Although this study was focused on wet deposition, a synergetic analysis of dry deposition processes is also needed to understand the impact of total wet and dry deposition on East Asian ecosystems.

#### **Author contributions**

Syuichi Itahashi performed the analysis of observation, emission inventory, and satellite data, and prepared the manuscript with contributions from all co-authors. Keiya Yumimoto and Itsushi Uno conducted the satellite data analysis and discussed the results. Hiroshi Hayami and Shin-ichi Fujita measured precipitation at the Komae site and performed the analysis. Yuepeng Pan and Yuesi Wang measured precipitation at ten observation sites around Beijing and performed the analysis.

#### **Competing interests**

The authors declare that they have no conflict of interest.

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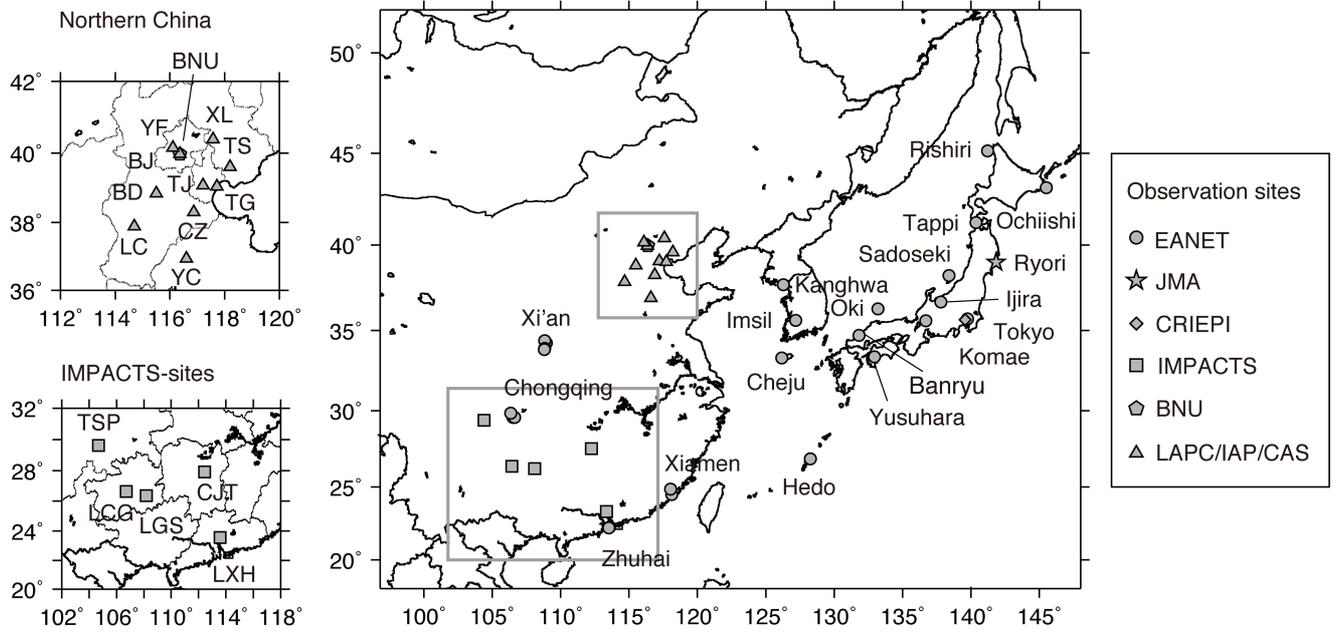
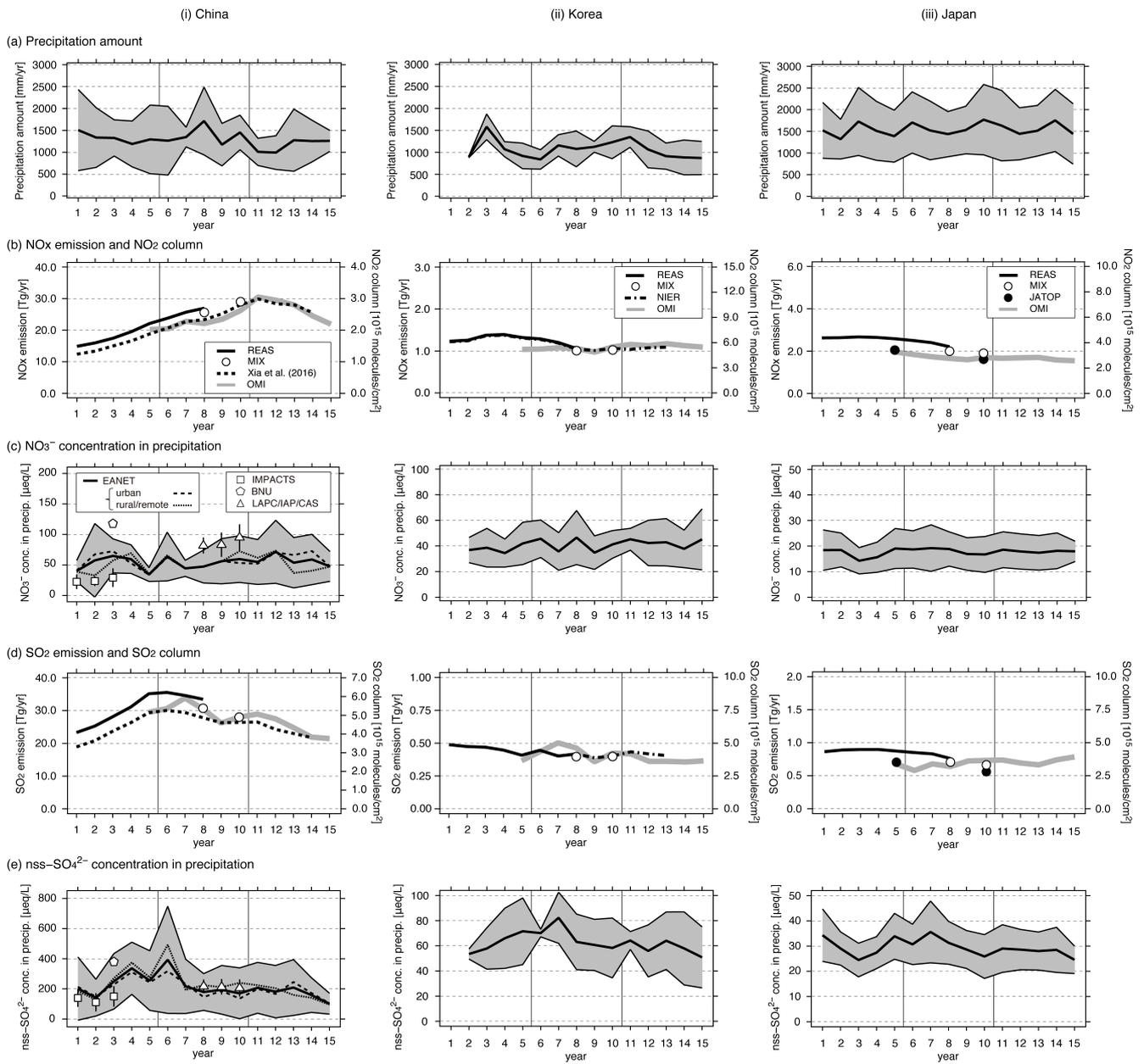
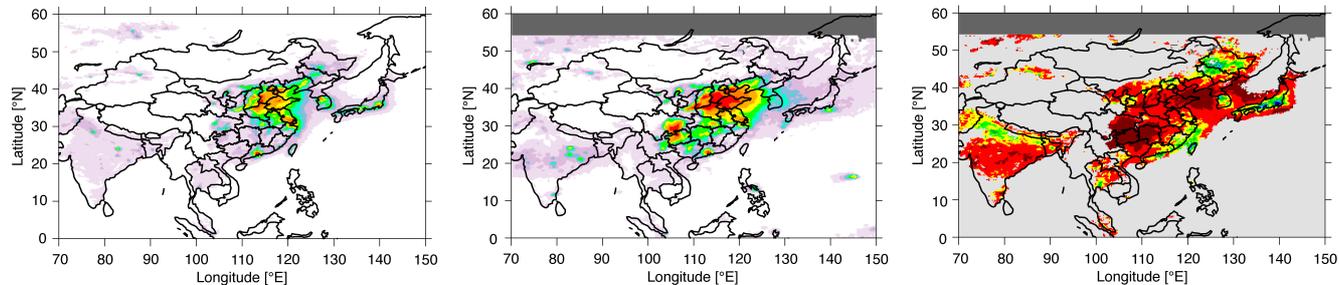


Figure 1: Geographical mapping of observation sites included in this study. Detailed information about each observation site is provided in Table 1.

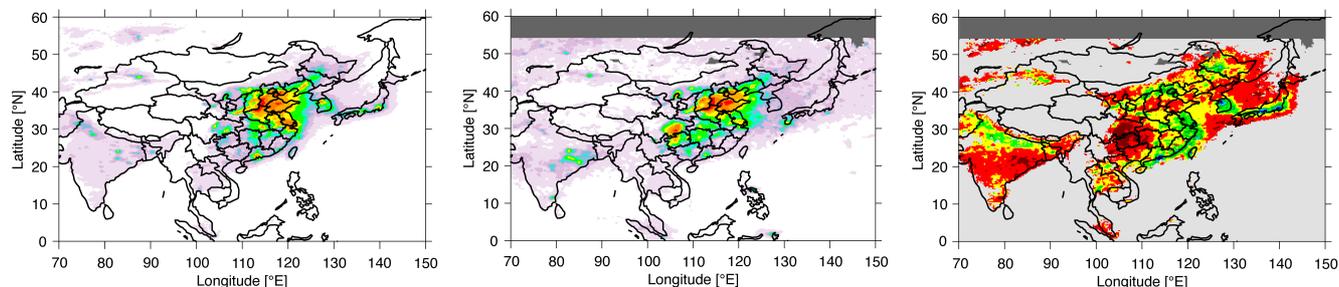


**Figure 2: Long-term temporal variation of (a) precipitation, (b) NO<sub>x</sub> emission with the NO<sub>2</sub> column, (c) NO<sub>3</sub><sup>-</sup> concentration in precipitation, (d) SO<sub>2</sub> emission with the SO<sub>2</sub> column, and (e) nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation during 2001–2015 over (i) China, (ii) Korea, and (iii) Japan. Mean and one standard deviation across EANET observation sites (JMA and CRIEPI sites are also included in Japan) are indicated by the thick lines and shaded areas, respectively, in (a), (c), and (e). For China, EANET observation sites are denoted with dashed lines for urban sites and dotted lines for rural and remote sites, and observation campaign data are indicated by marks corresponding to marks on Fig. 1. The numbers on the bottom indicate the number of years after 2000 (e.g., 5 indicates 2005 and 10 indicates 2010).**

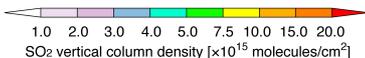
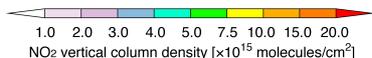
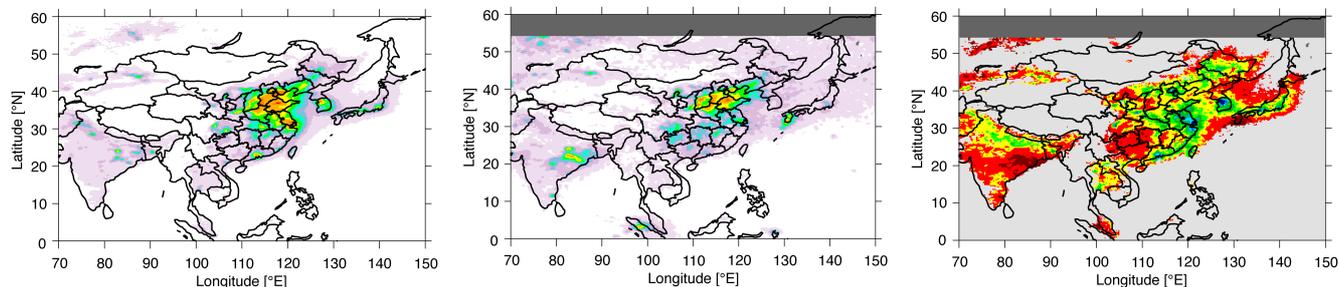
(a) Averaged between 2006 and 2007 (the first half of Phase II)



(b) Averaged between 2010 and 2011 (transition between Phase II and Phase III)

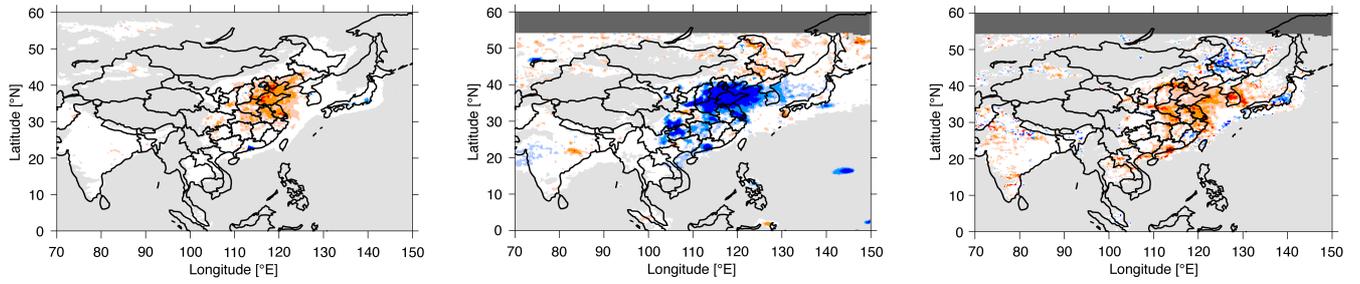


(c) Averaged between 2014 and 2015 (the latter half of Phase III)

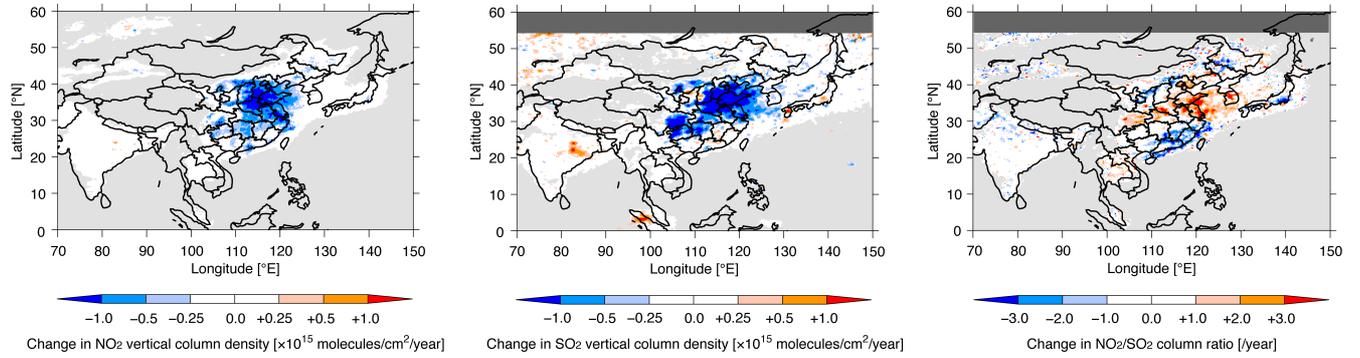


5 **Figure 3: Satellite observations of (left) the NO<sub>2</sub> column, (center) SO<sub>2</sub> column, and (right) NO<sub>2</sub>/SO<sub>2</sub> column ratio averaged over (a) 2006–2007 (the first half of Phase II), (b) 2010–2011 (transition between Phase II and Phase III), and (c) 2014–2015 (the latter half of Phase III). Dark gray cells indicate places where annual mean could not be calculated. If either the NO<sub>2</sub> or SO<sub>2</sub> column value is less than 1.0 × 10<sup>15</sup> molecules/m<sup>2</sup>, this is indicated by a light gray cell to clarify the NO<sub>2</sub>/SO<sub>2</sub> column ratio.**

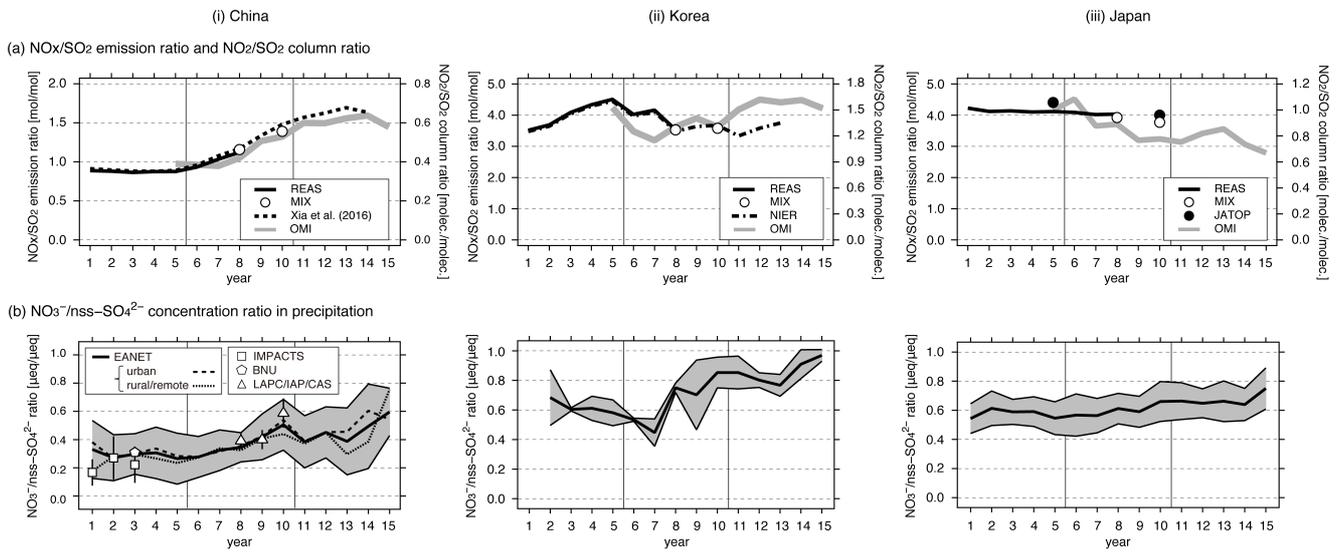
(a) Annual changes between 2006 and 2010 (Phase II)



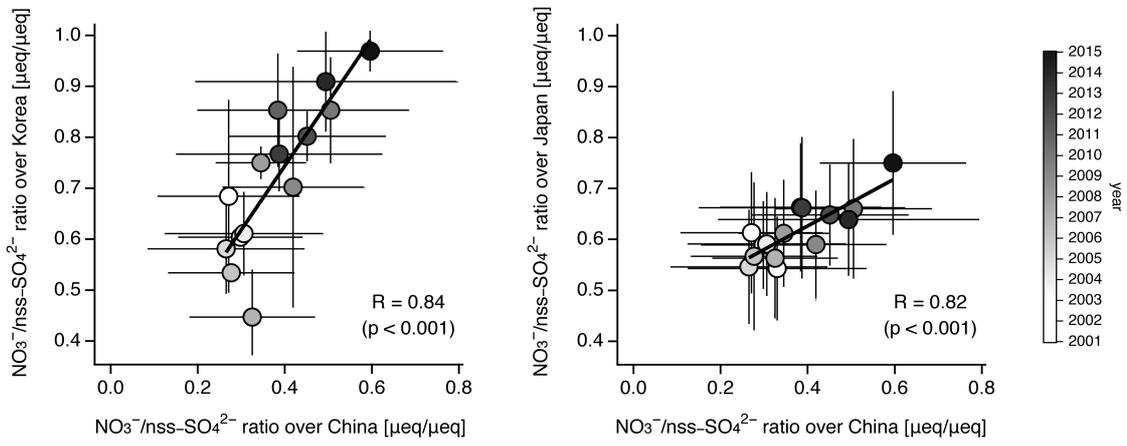
(b) Annual changes between 2011 and 2015 (Phase III)



5 **Figure 4: Annual changes based on the linear regression results for the (left) NO<sub>2</sub> column, (center) SO<sub>2</sub> column, and (right) NO<sub>2</sub>/SO<sub>2</sub> column ratio during (a) Phase II and (b) Phase III. Dark gray cells indicate areas where annual mean calculation could not be performed. If either the NO<sub>2</sub> or SO<sub>2</sub> column value is less than  $1.0 \times 10^{15}$  molecules/m<sup>2</sup>, this is indicated by a light gray cell for clarity.**

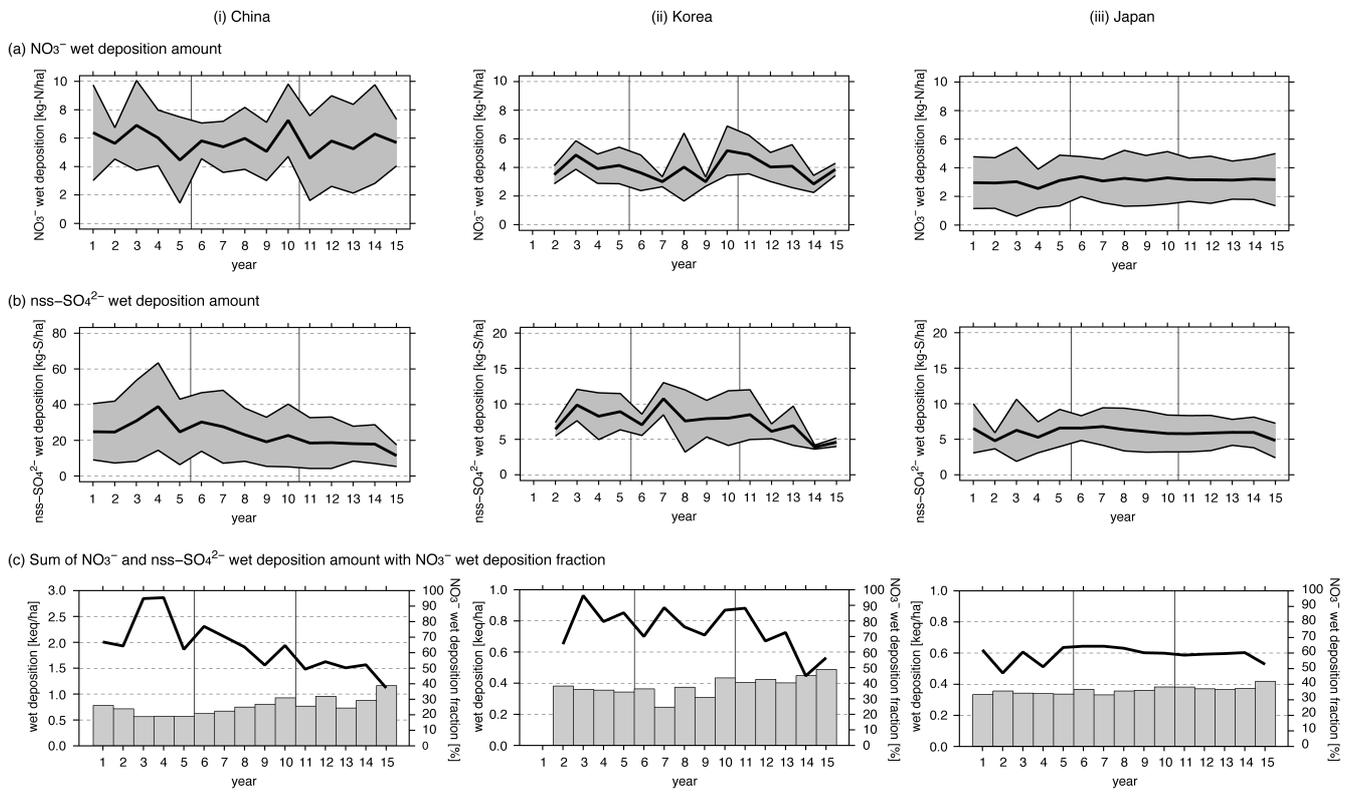


5 **Figure 5: Long-term temporal variation in the (a) NO<sub>x</sub>/SO<sub>2</sub> emission ratio and NO<sub>2</sub>/SO<sub>2</sub> column ratio, and (b) NO<sub>3</sub><sup>-</sup>/nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation (*Ratio*) during 2001–2015 over (i) China, (ii) Korea, and (iii) Japan. Mean and one standard deviation across EANET observation sites (JMA and CRIEPI sites are also included in Japan) are indicated by the thick lines and shaded areas, respectively, in (a), (c), and (e). For China, EANET observation sites are denoted with dashed lines for urban sites and dotted lines for rural and remote sites, and observation campaign data are indicated by marks corresponding to Fig. 1. The numbers on the bottom indicate the number of years after 2000 (e.g., 5 indicates 2005 and 10 indicates 2010).**



**Figure 6:** Scatter-plots of (left) the  $\text{NO}_3^-/\text{nss-SO}_4^{2-}$  concentration in precipitation (*Ratio*) in China and Korea, and (right) those in China and Japan. Each circle indicates the annual mean *Ratio*, and the color indicates the year.

5



5 **Figure 7: Long-term temporal variation of (a)  $\text{NO}_3^-$  wet deposition (kg-N/ha), (b)  $\text{nss-SO}_4^{2-}$  wet deposition (kg-S/ha), and (c) the sum of  $\text{NO}_3^-$  and  $\text{nss-SO}_4^{2-}$  wet deposition (kgq/ha) with the  $\text{NO}_3^-$  fraction during 2001–2015 over (i) China, (ii) Korea, and (iii) Japan. One standard deviation across observation sites is indicated by the shaded area in (a) and (b). The numbers on the bottom indicate the number of years after 2000 (e.g., 5 indicates 2005 and 10 indicates 2010).**

**Table 1. Locations of observation sites included in this study.**

	Country	Site	Period	Latitude (°N)	Longitude (°E)	Elevation (m a.s.l.)	Classification
EANET	China	Zhuhai					
		- Xiang Zhou	2001–2015	22.27	113.57	40	urban
		- Zhuxiandong	2008–2015	22.20	113.52	45	urban
		Xiamen					
		- Hongwen	2001–2015	24.47	118.13	50	urban
		- Xiaoping	2001–2015	24.85	118.03	686	remote
		Chongqing					
		- Guanyinqiao	2001–2007	29.57	106.52	262	urban
		- Haifu	2008–2015	29.62	106.50	317	urban
		- Jinyunshan	2001–2015	29.82	106.37	800	rural
		Xi'an					
		- Shizhan	2001–2015	34.23	108.95	400	urban
		- Weishuiyuan	2001–2006	34.37	108.85	366	rural
		- Jiwozi	2001–2015	33.83	108.80	1800	remote
	Korea	Cheju	2001–2011	33.30	126.17	72	remote
		Imsil	2001–2011	35.60	127.18	-	rural
		Kanghwa	2001–2011	37.70	126.28	150	urban
	Japan	Hedo	2001–2015	26.87	128.25	60	remote
		Yusuhara	2001–2015	33.38	132.94	790	remote
		Banryu	2001–2015	34.68	131.80	53	urban
		Ijira	2001–2015	35.57	136.69	140	rural
Tokyo		2007–2015	35.69	139.76	26	urban	
Oki		2001–2015	36.29	133.19	90	remote	
Sadoseki		2001–2015	38.25	138.40	136	remote	
Tappi		2001–2015	41.25	140.35	106	remote	
Ochiishi		2003–2015	43.16	145.50	49	remote	
Rishiri		2001–2015	45.12	141.21	40	remote	
JMA	Japan	Ryori	2001–2015	39.03	141.82	260	
CRIEPI	Japan	Komae	2001–2015	35.64	139.58	27	
IMPACTS	China	Guangdong					
		- Liu Xi He (LXH)	2002-2003	23.55	113.58	500	
		Guizhou					
		- Lei Gong Shan (LGS)	2002-2003	26.37	108.18	1630-1735	
		- Liu Chong Guan (LCG)	2001-2003	26.63	106.72	1320-1400	
		Hunan					
- Cai Jia Tang (CJT)	2001-2003	27.92	112.43	450-500			
Chongqing							
- Tie Shan Ping (TSP)	2001-2003	29.63	104.68	450-500			
BNU	China	Beijing					
		- Beijing Normal University	2005	39.96	116.37	40	
LAPC/IAP/CAS	China	Shandong					
		- Yucheng (YC)	2008-2010	36.85	116.55	23	
		Hebei					
		- Luancheng (LC)	2008-2010	37.89	114.69	57	
		- Cangzhou (CZ)	2008-2010	38.30	116.87	10	
		- Baoding (BD)	2008-2010	38.85	115.50	21	
		- Tangshan (TS)	2008-2010	39.60	118.20	24	
		- Xinglong (XL)	2008-2010	40.38	117.57	872	
		Tianjin					
		- Tanggu (TG)	2008-2010	39.04	117.72	0	
		- Tianjin (TJ)	2008-2010	39.08	117.21	6	
Beijing							
- Beijing (BJ)	2008-2010	39.96	116.36	57			
- Yangfang (YF)	2008-2010	40.15	116.10	73			

**Table 2. Statistical analysis of average values and linear trends for precipitation, NO<sub>3</sub><sup>-</sup> concentration in precipitation, and nss-SO<sub>4</sub><sup>2-</sup> concentration in precipitation over China, Korea, and Japan during Phases I, II, and III.**

	Phase I (2001-2005)		Phase II (2006-2010)		Phase III (2011-2015)	
	Mean	Trend	Mean	Trend	Mean	Trend
Precipitation [mm/year]						
China	1330.2 (N = 35)	-4.3±1.9	1390.2 (N = 31)	+1.5±5.4	1158.3 (N = 28)	+6.6±2.5*
Korea	1114.9 (N = 11)	-3.9±15.4	1084.3 (N = 15)	+6.9±3.0	1015.1 (N = 15)	-11.2±3.3*
Japan	1493.5 (N = 44)	-0.5±3.8	1591.7 (N = 56)	+0.9±3.2	1555.7 (N = 56)	-0.5±3.1
NO <sub>3</sub> <sup>-</sup> concentration in precipitation [μmol/L]						
China	51.3 (N = 31)	-1.9±9.4	54.3 (N = 31)	+0.6±5.5	57.3 (N = 34)	-4.7±5.0
Korea	38.0 (N = 11)	+3.0±4.1	40.8 (N = 15)	-2.4±4.7	42.7 (N = 15)	-1.1±2.5
Japan	17.2 (N = 44)	-0.9±4.4	18.1 (N = 55)	-3.4±1.3*	18.0 (N = 55)	-0.6±0.8
nss-SO <sub>4</sub> <sup>2-</sup> concentration in precipitation [μmol/L]						
China	237.0 (N = 32)	+12.7±8.3	230.6 (N = 32)	-20.3±8.8	171.5 (N = 33)	-13.6±5.2*
Korea	62.2 (N = 11)	+10.0±0.8***	66.9 (N = 15)	-6.8±3.6	58.6 (N = 15)	-4.2±2.6
Japan	29.8 (N = 44)	-0.8±5.2	30.4 (N = 56)	-5.4±2.9	27.8 (N = 55)	-3.3±1.5

Note: Total numbers of observation sites are shown in parenthesis. Linear trends were estimated using linear regression and are shown as rates (/year) with corresponding units. \*p < 0.05, \*\*p < 0.01, and \*\*\*p < 0.001, and lack of a mark indicates a

5 lack of significance. Analysis for Korea during Phase I was for the 2002–2005 period.

**Table 3. Statistical analysis of average values and linear trends for *Ratio* in precipitation over China, Korea, and Japan during Phases I, II, and III.**

	Phase I (2001–2005)		Phase II (2006–2010)		Phase III (2011–2015)	
	Mean	Trend	Mean	Trend	Mean	Trend
<i>Ratio</i> [mol/mol]						
China	0.29 (N = 35)	-3.3±2.7%	0.37 (N = 31)	+14.8±1.9%***	0.46 (N = 28)	+10.1±3.8%*
Korea	0.62 (N = 11)	-4.9±1.9%*	0.66 (N = 15)	+13.6±4.7%*	0.86 (N = 15)	+3.9±2.6%
Japan	0.58 (N = 44)	-0.3±1.9%	0.60 (N = 56)	+3.6±1.3%*	0.67 (N = 56)	+2.5±1.3%

Note: Total numbers of observation sites are shown in parenthesis. Linear trends were estimated using linear regression and results are shown in (mol/mol)/year. \* $p < 0.05$ , \*\* $p < 0.01$ , and \*\*\* $p < 0.001$ , and a lack of mark indicates a lack of

5 significance. Analysis for Korea during Phase I was for the 2002–2005 period.

**Table 4. Statistical analysis of average values and linear trends for NO<sub>3</sub><sup>-</sup> and nss-SO<sub>4</sub><sup>2-</sup> wet deposition over China, Korea, and Japan during Phases I, II, and III.**

	Phase I (2001-2005)		Phase II (2006-2010)		Phase III (2011-2015)	
	Mean	Trend	Mean	Trend	Mean	Trend
NO <sub>3</sub> <sup>-</sup> wet deposition [kg-N/ha]						
China	5.88	-5.9±4.6	5.90	+4.4±4.5	5.53	+4.9±3.2
Korea	4.10	+2.4±7.5	3.76	+8.2±7.3	3.95	-8.3±4.8
Japan	2.92	-0.3±2.7	3.23	-0.4±1.5	3.17	+0.2±0.3
nss-SO <sub>4</sub> <sup>2-</sup> wet deposition [kg-S/ha]						
China	28.80	+4.9±7.4	24.57	-9.6±3.4*	16.90	-8.8±4.3
Korea	8.36	+7.0±8.0	8.26	-1.1±6.3	6.01	-16.7±5.6*
Japan	5.88	+1.0±5.0	6.32	-3.5±0.9*	5.68	-3.2±2.5

Note: Linear trends were estimated using linear regression and results are shown as rates (/year) with corresponding units. \*p

< 0.05, \*\* p < 0.01, and \*\*\* p < 0.001, and a lack of a mark indicates a lack of significance. Analysis for Korea during Phase I

5 was for the 2002–2005 period.