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Supplement for

2 **The single-particle mixing state and cloud scavenging of black carbon at**
3 **a high-altitude mountain site in southern China**

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23 **SPAMS**

24 Individual particles are introduced into SPAMS through a critical orifice. They are focused
25 and accelerated to specific velocities, determined by two continuous diode Nd:YAG laser beams
26 (532 nm), which are used to trigger a pulsed laser (266 nm) to desorp/ionize the particles. The
27 produced positive and negative molecular fragments are recorded. In summary, a velocity, a
28 detection moment, and an ion mass spectrum are recorded for each ionized particle, while there is
29 no mass spectrum for not ionized particles. The velocity could be converted to d_{va} based on a
30 calibration using polystyrene latex spheres (PSL, Duke Scientific Corp., Palo Alto) with
31 predefined sizes.

32

33 **Aethalometer data analysis**

34 The absorption coefficient is defined by Beer–Lambert’s law (Arnott et al., 2005). A
35 variable attenuation (ATN), is defined to represent the filter attenuation through the sample spot
36 on a filter (Arnott et al., 2005; Weingartner et al., 2003; Backman et al., 2016). Aerosol light
37 absorption coefficient and BC mass concentration can be calculated directly based on the
38 measured ATN. It is well known that the measured ATN may differ from the true aerosol
39 absorption due to ‘filter loading effect’, a phenomenon which appears as a gradual decrease of
40 instrumental response as the aerosol loading on the filter increases (Arnott et al., 2005).
41 Therefore, two calibration factors are introduced to convert aethalometer attenuation
42 measurements to “real” absorption coefficient (Weingartner et al., 2003).

43 The AE-31 used in the present study may suffer from the effects described above. Differently,
44 the AE-33 has been improved by the incorporation of a filter loading correction part, based on a
45 two parallel spot measurement of optical absorption. It could provide a real-time output of the
46 “loading compensation” parameter to compensate for the “loading effect”. The details of the
47 principle of operation, data deduction, and error budget of the AE-33, the inherent uncertainties in
48 its technique and the corrections are extensively available in the literature (Drinovec et al., 2015).

49 Therefore, we reported EBC concentration from the results of AE-33. The noise level of the AE-
50 33 at the time-base of 1 minute is $< 0.016 \text{ Mm}^{-1}$ for the b_{abs} , corresponding to 1 ng m^{-3} for the mass
51 concentration of BC. As noted in the manuscript and Fig. S10, the EBC measured by AE-31 is
52 significantly correlated ($R^2 = 0.9$, $p < 0.001$) with that measured by AE-33. Therefore, EBC
53 concentrations derived from AE-31 were not corrected for the calculation of $\text{Mf}_{\text{scav,EBC}}$.

54 As shown in Fig. S10, AE-31 might underestimate $\sim 15\%$ of EBC for cloud INT particles in the
55 calculation of $\text{Mf}_{\text{scav,EBC}}$. It is also noted that a threshold of $8 \mu\text{m}$ might underestimate the mass
56 concentration of cloud RES EBC, since the size of droplets might extend to as low as $3 \mu\text{m}$.
57 Unfortunately, the size distribution of cloud droplets was not available for our study. Therefore, we
58 assumed that the largest underestimate of cloud RES particles is 30% to assess the uncertainties for
59 $\text{Mf}_{\text{scav,EBC}}$ calculation. The mean $\text{Mf}_{\text{scav,EBC}}$ was $30\text{-}36\%$ when they were taken into account. Overall,
60 the uncertainties for the calculation of $\text{Mf}_{\text{scav,EBC}}$ is with 10% .

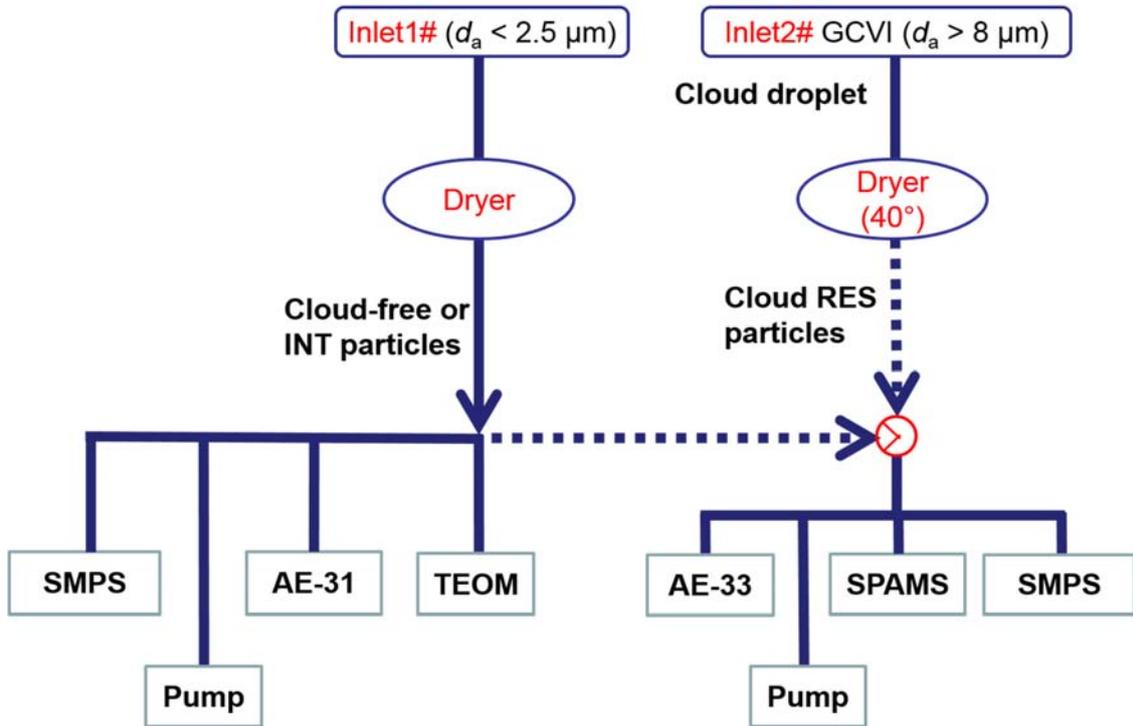
61 Table S1. Average mass concentrations, mass fractions relative to fine particles and scavenged fractions of BC from the literatures.

Site	site type	season (year)	ave (\pm std) ($\mu\text{g m}^{-3}$)	mass fraction	Mf _{scav,EBC} (%)	References
Shenzhen, Southern China	urban	Summer (2011)	4.0 \pm 3.1	~11%	- ^a	(Lan, et al., 2013)
Guangzhou, Southern China	urban	Summer (2008)	8.86	-	-	(Wu, et al., 2013)
Guangzhou, Southern China	urban	Fall(2010)	4.3	~4% ^b	-	(Zhang, et al., 2013)
Shenzhen, Southern China	urban	Fall(2009)	6.0 \pm 6.3	-	-	(Huang, et al., 2012)
Guangzhou, Southern China	Rural	Summer (2008)	2.62	-	-	(Wu, et al., 2013)
Ba Guang village, southern China	Rural	Fall(2009)	2.6 \pm 1.0	-	-	(Huang, et al., 2012)
Mt. Soledad (251 m m.s.l.)	marine	Summer (2012)	0.07	-	-	(Schroder, et al., 2015)
Yongxing Island, Southern China	marine	Summer (2008)	0.54	-	-	(Wu, et al., 2013)
A coastal Chilean hill, (Valparaíso), 450 m a.s.l.	low-altitude	Winter (2013)	0.34 - 0.95	-	13 - 50	(Hitzenberger et al., 2016)
Puy de Dome (France), 1465 m a.s.l.	mid-altitude	Winter-spring (2001)	-	-	33 - 74	(Sellegri et al., 2003)
Nova Scotia, Canada (Below 1 km)	mid-altitude	Summer (1993)	0.06 \pm 0.01	-	2 - 32	(Chylek et al., 1996)

62 ^a not available.

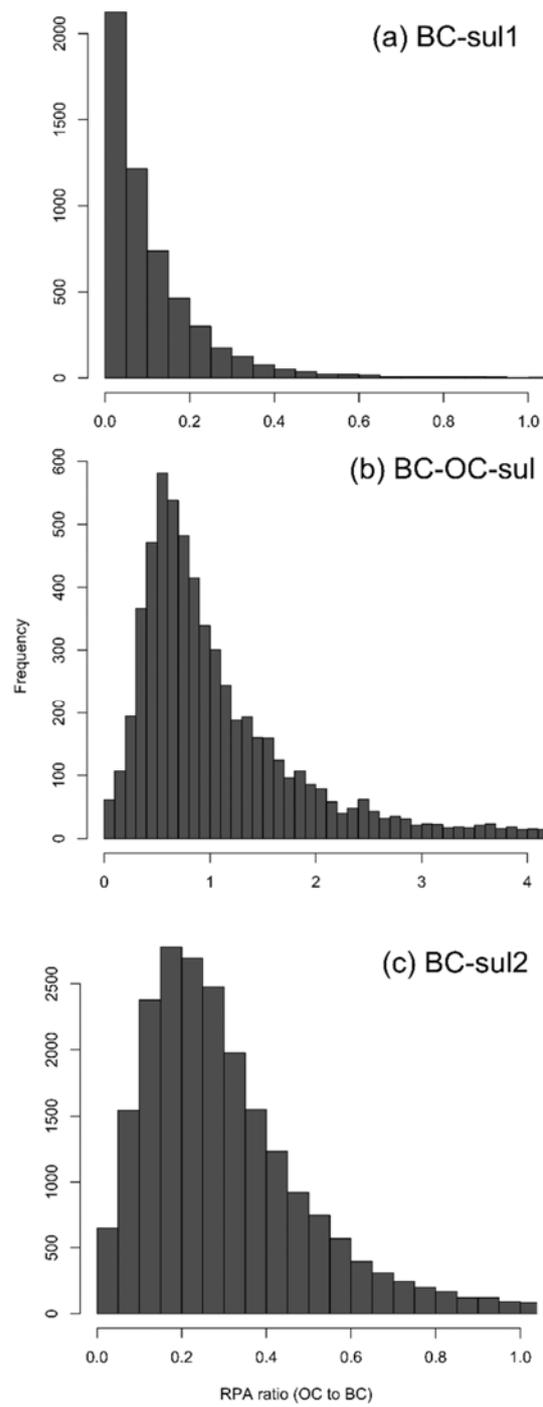
Nova Scotia, Canada (1-3 km)	mid-high-altitude	Summer (1993)	0.22 ± 0.03	-	-	(Chylek et al., 1996)
Mt. Rax (1644 m a.s.l.)	high-altitude	Spring (1999)	0.43	-	-	(Hitzenberger et al., 2001)
Mt. Rax (1644 m a.s.l.)	high-altitude	Spring (2000)	0.72	-	54 ± 25	(Hitzenberger et al., 2001)
Alpine Jungfrauoch (Switzerland), 3850 m a.s.l.	high-altitude	Summer (2004)	0.06	-	61	(Cozic et al., 2007)
Alpine Jungfrauoch (Switzerland), 3850 m a.s.l.	high-altitude	Winter (2004)	0.05	-	-	(Cozic et al., 2007)

63 ^b mass fraction relative to PM₃.



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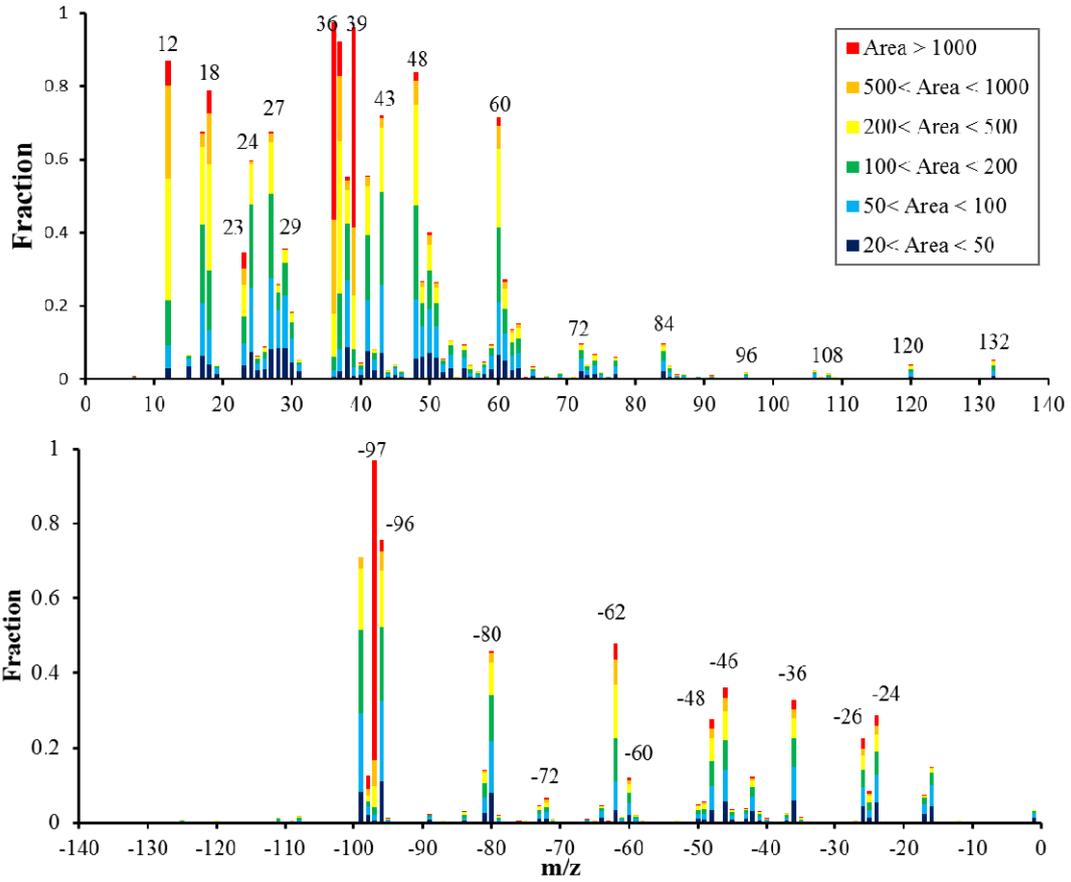
65 **Figure S1.** A scheme of the instrumentation setup in this study. The dash line illustrates
 66 that the sampling pipe was either connected to Inlet 1# or Inlet 2#. As described in
 67 section 2.1, the cloud INT and RES particles were intermittently measured by these
 68 instruments during Cloud III, through manually connect the sampling pipe to either Inlet
 69 1# or Inlet 2# at approximately one-hour intervals.



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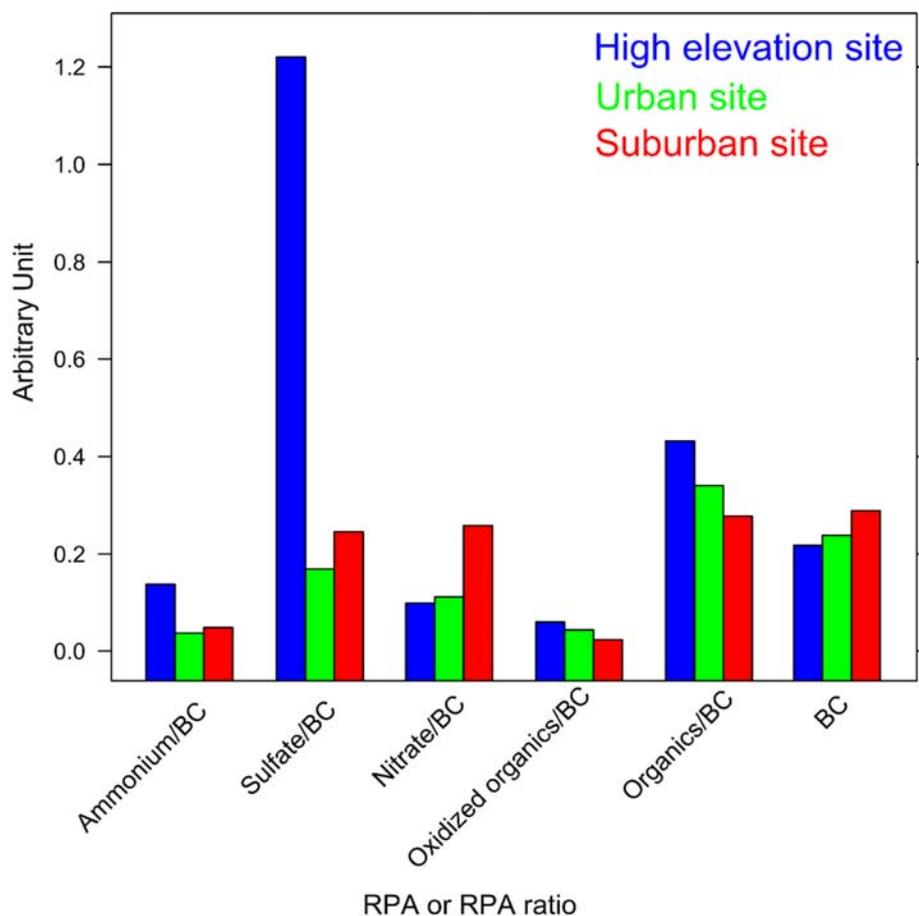
72 **Figure S2.** Statistic analysis on the RPA ratio of OC to BC in BC types. Markers were
73 selected as m/z 27, 43, 50, 51, 61, 63, -26 for OC, and carbon ion clusters ($C_n^{+/-}$, $n \leq 5$)
74 for BC, the same as those in Fig. 3.



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76 Figure S3. The number-based digitized mass spectrum of cloud-free BC-containing

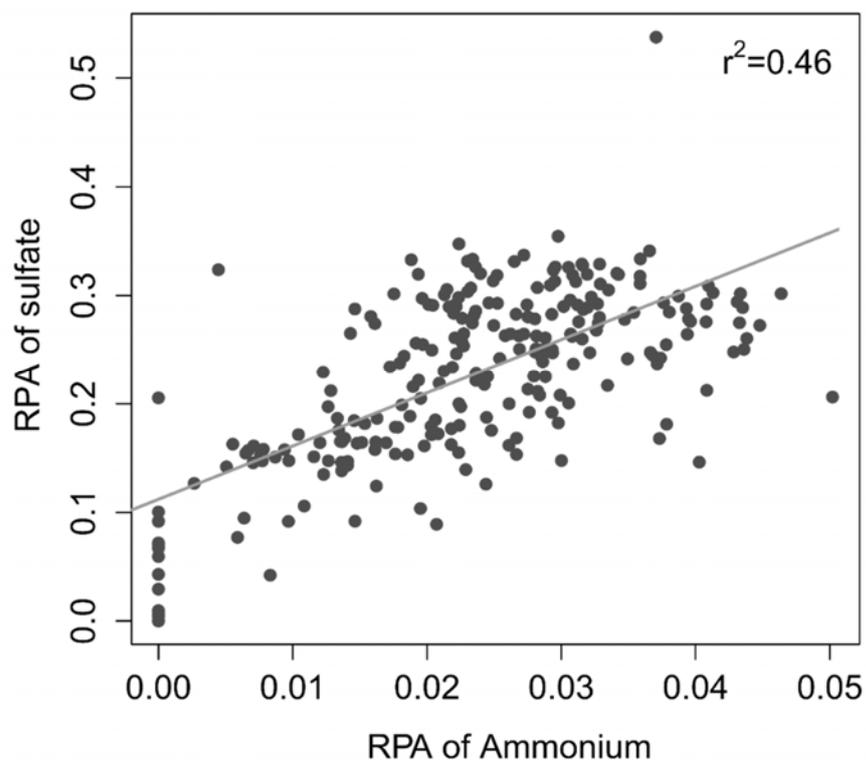
77 particles at the remote high-altitude site.



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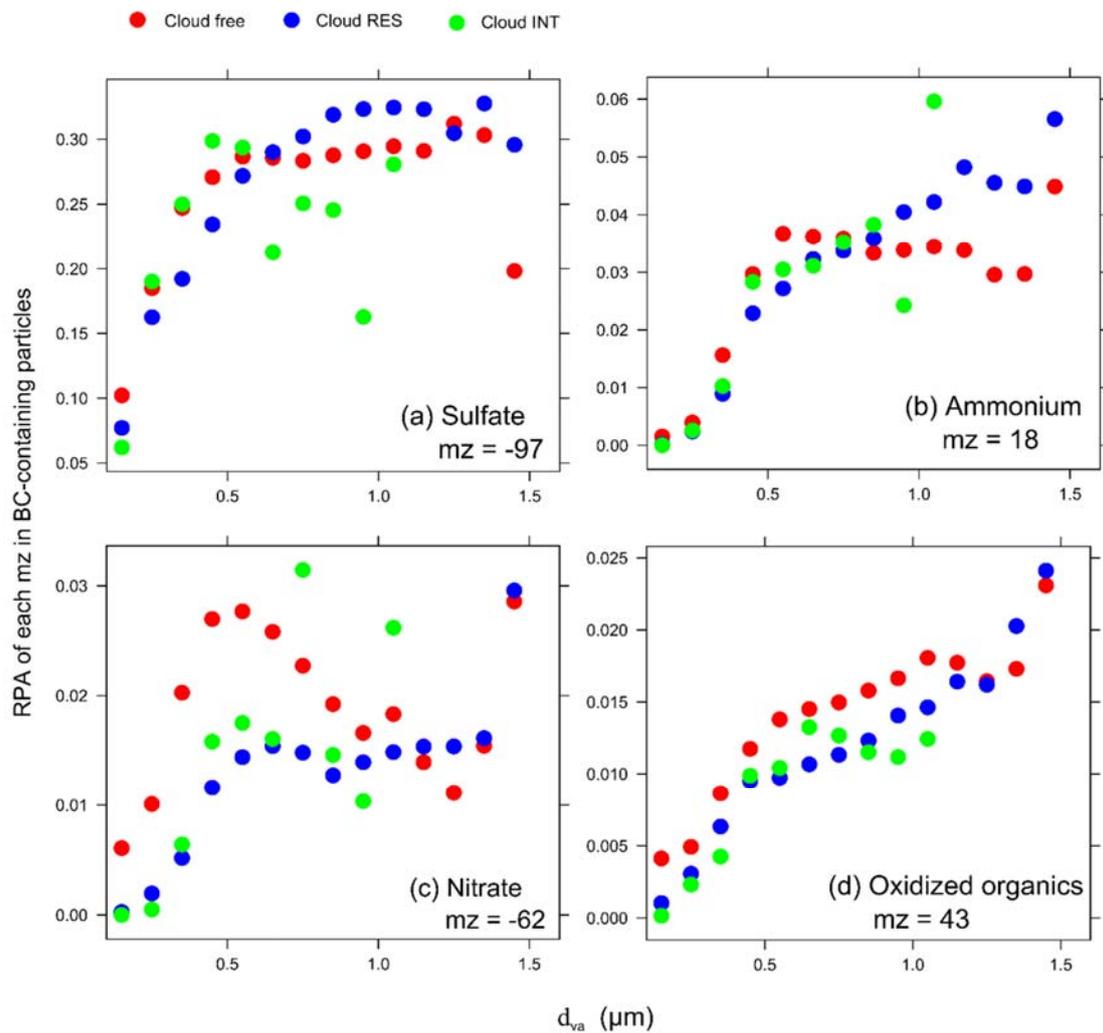
79 Figure S4. RPA ratio of ammonium (m/z 18), sulfate (m/z -97), nitrate (m/z -62),
 80 oxidized organics (m/z 43), and other organics (m/z 27, 50, 51, 61, 63, -26) to BC, and
 81 RPA of BC (carbon ion clusters ($C_n^{+/-}$, $n \leq 5$)) at the high elevation site, urban
 82 (Guangzhou), and suburban sites (Heshan) during winter in southern China. The particles
 83 in Guangzhou and Heshan were similarly measured by SPAMS during winter. Despite of
 84 matrix effects due to the laser desorption/ionization for SPAMS, advances have been
 85 made in semi-quantifying individual chemical species, either through multivariate
 86 analysis or by applying peak intensities for specific ions (e.g., Xing et al., 2011; Jeong et
 87 al., 2011; Healy et al., 2013). RPA, defined as the peak area of each m/z divided by the

88 total dual ion mass spectral peak area, is related to the relative amount of a species on a
89 particle. Compared to absolute peak area, RPA was commonly applied because it is less
90 sensitive to the variability in ion intensities associated with particle-laser interactions. It
91 is also noted that matrix effects might be lower when calculation was performed for
92 similar particle type, i.e., BC-containing particles.



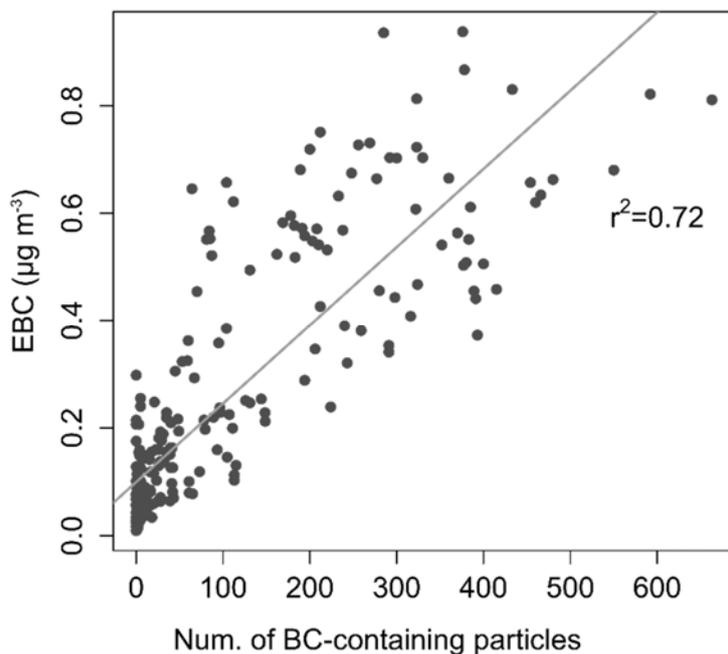
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94 Figure S5. Correlation analysis of hourly average RPA for ammonium and sulfate
95 associated with BC-containing particles. The correlation coefficient is a bit lower than
96 expected might partly due to matrix effect in single particle mass spectrometry (e.g., Xing
97 et al., 2011; Jeong et al., 2011; Healy et al., 2013).



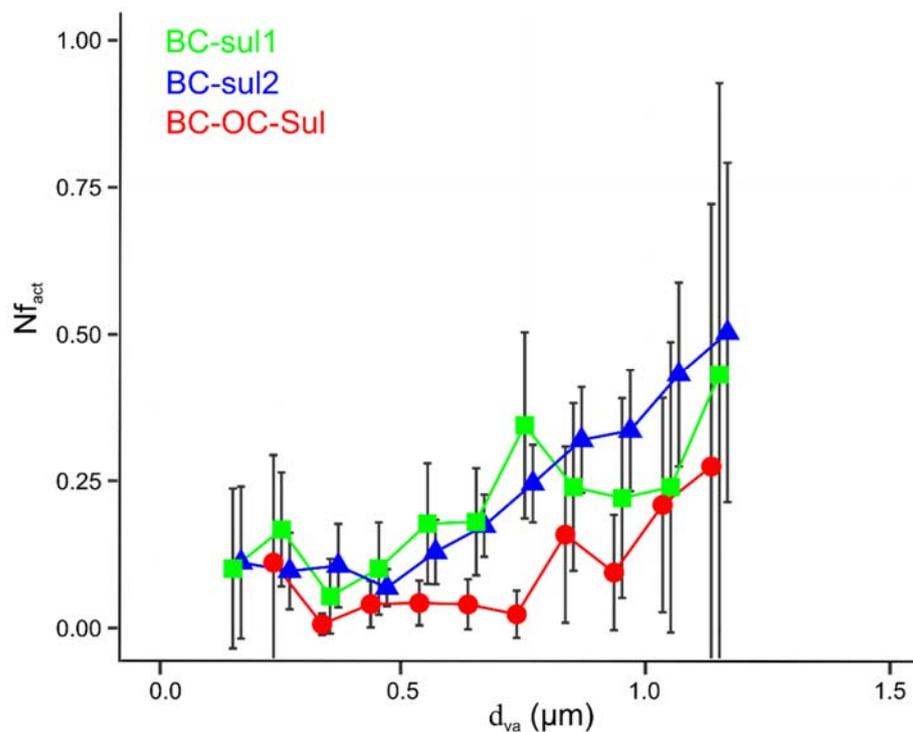
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99 Figure S6. RPA of each secondary species associated with BC-containing particles in
 100 cloud-free, INT, and RES particles as a function of particle sizes.



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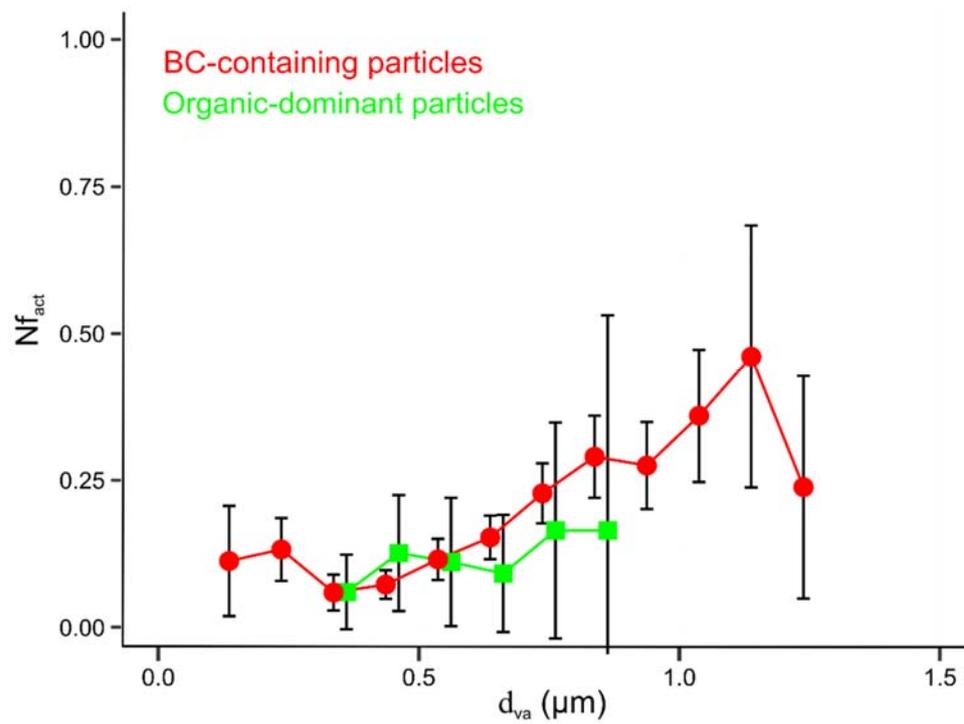
102 Figure S7. Correlation between time series of Num. of BC-containing particles and
 103 concentration of EBC. The volume equivalent diameter of BC particles cores measured in
 104 southern China was typically around 200 nm (Huang et al., 2012; Huang et al., 2011).
 105 Huang et al. (2011) showed that a large fraction (> 60%) of BC particles are internally
 106 mixed with a significant amount of non-refractory materials (coating thickness > 70 nm)
 107 at a rural site in southern China. Furthermore, Yu et al. (2010) showed that over 50% of
 108 BC are above 500 nm, also indicating internally mixed of BC, with regard that majority of
 109 BC particles cores have volume equivalent diameter less than 500 nm (Huang et al., 2012;
 110 Huang et al., 2011). As also discussed in section 3.1, BC-containing particles were already
 111 heavily mixed with secondary species arriving at our site, and therefore they should be
 112 larger enough for the detection by SPAMS.



113

114

115 Figure S8. Size-resolved Nf_{act} estimated for three particle types of BC-containing
 116 particles. Note that this data only collected during Cloud III event when both cloud RES
 117 and INT particles were collected, however, not simultaneously but intermittently. It is
 118 noted that although the Nf_{act} for BC-OC-sul type is lower than BC-sul types, the Nf_{act} for
 119 all the BC-containing particles is similar to that of all the detected particles. We attributed
 120 it to two reasons: (1) BC-OC-sul particles only accounted for ~20% of BC-containing
 121 particles, and (2) the other particles also contained OC particles (~10%).

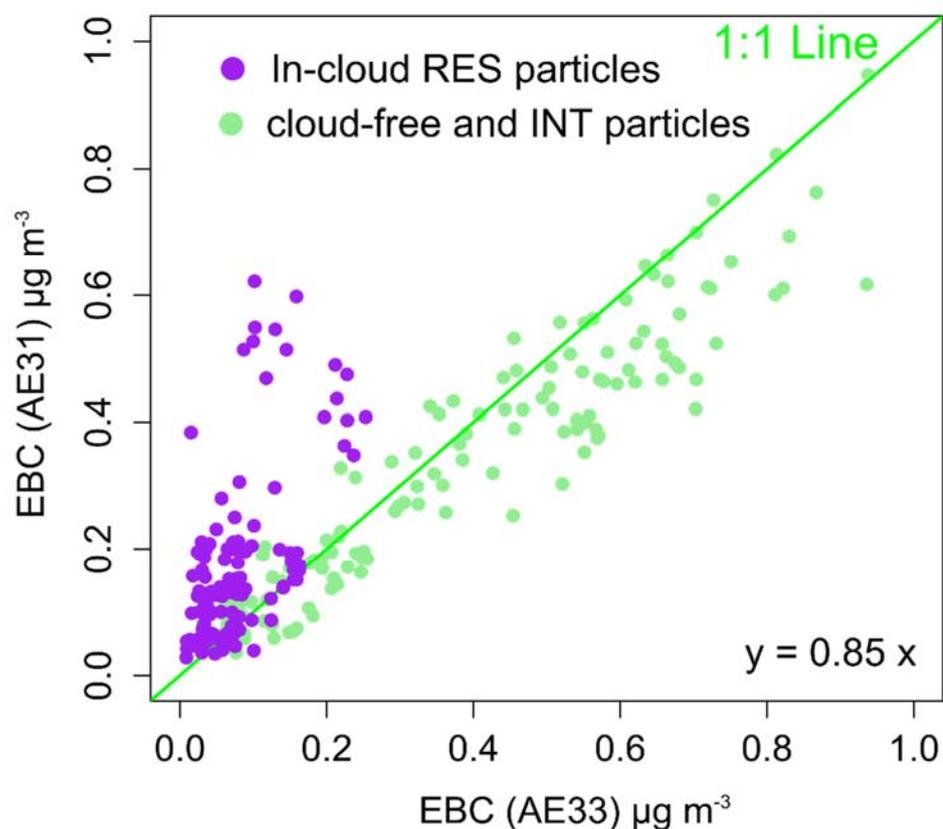


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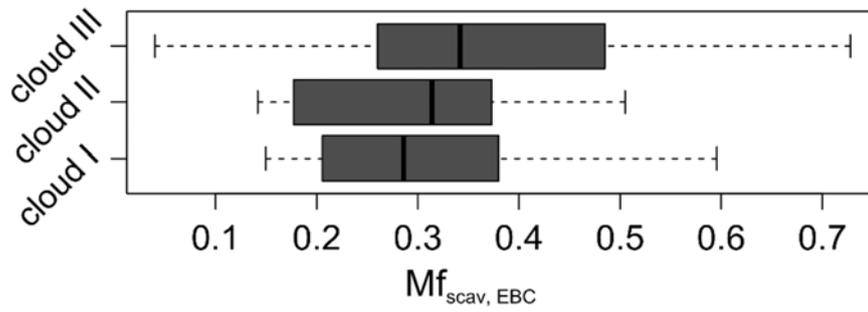
124 Figure S9. Size-resolved Nf_{act} estimated for BC-containing particles and organic-

125 dominant particles, similar to Fig. S6.



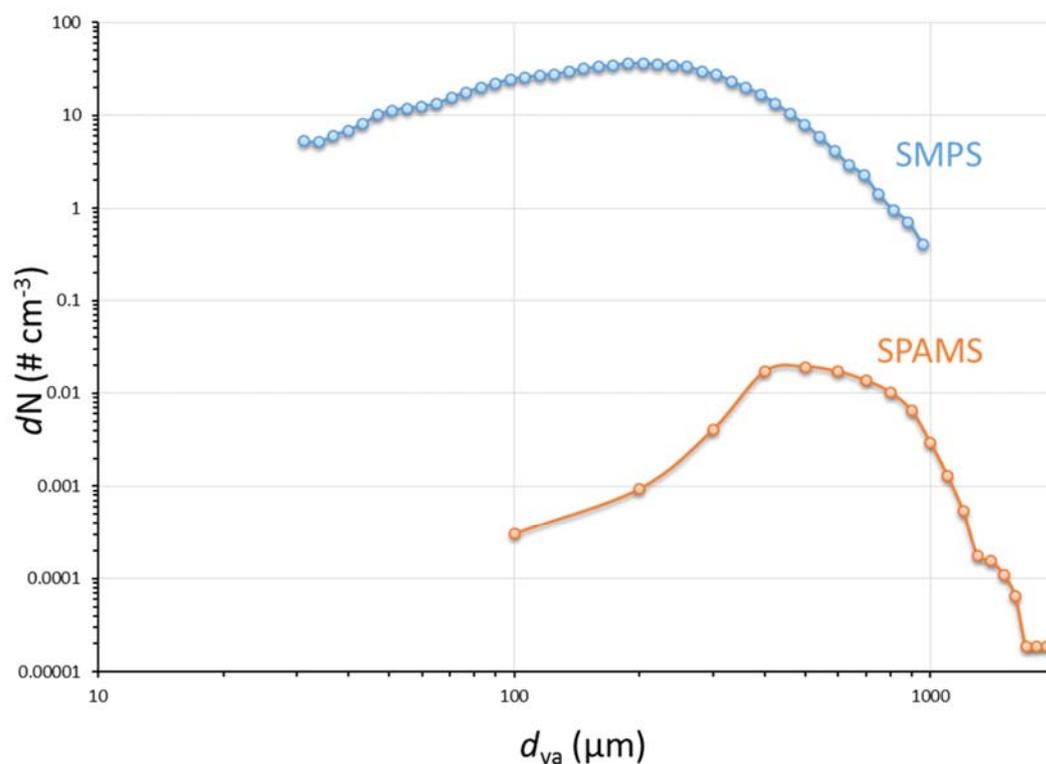
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127 Figure S10. Correlation analysis of EBC measured by AE31 and AE33. They measured
 128 the same aerosol for out-of-cloud (including cloud INT and cloud-free) particles. However,
 129 during cloud events, AE33 measured cloud RES particles or cloud INT particles for some
 130 periods, while AE31 measured cloud INT particles. Therefore, the EBC were compared
 131 when the same aerosol were measured, as shown in green dots. The result indicates that
 132 they are highly correlated, with EBC measured by AE31 only slightly lower than those by
 133 AE33.



134

135 Figure S11. Box and whisker plot of $Mf_{scav,EBC}$ for each cloud event. In a box and whisker
 136 plot, the lower, median and upper lines of the box denote the 25th, 50th, and 75th
 137 percentiles, respectively, and the lower and upper edges of the whisker denote the 10th
 138 and 90th percentiles, respectively.



139

140 Figure S12. A representative comparison between the size distributions measured by the
 141 SPAMS and the SMPS within 12 hours measurements. It should be noted that the
 142 diameter is represented as d_{va} by SPAMS, while the diameter measured by the SMPS is
 143 represented as electrical mobility diameter (d_m). Herein, the d_m was first converted to the
 144 d_{va} for the comparison. The conversion could be simplified to $d_m = d_{va} \cdot \rho_{eff} / \rho_0$ (DeCarlo et
 145 al., 2004), where ρ_{eff} refers to the effective density, ρ_0 is the unit density 1.0 g cm^{-3} . The
 146 ρ_{eff} is assumed to be 1.5 g cm^{-3} for the calculation (Hu et al., 2012).