

Response to reviewers comments for:
“A new diagnostic for tropospheric ozone production”
(acp-2017-378)

The reviewers comments have been addressed. Below are the authors responses (red) and changes implemented (blue). The reviewers comments (black) have been numbered to avoid confusion.

Anonymous Referee #1

This paper presents a novel analysis of ozone production in terms on the spin states of the bonds in the precursor species. This is an interesting and original concept, and is a commendable attempt to generate a diagnostic of ozone production that has a sound physico-chemical basis, and one that provides more process insight than the standard methods based on NO_x cycling. The paper is worthy of publication, but needs revision to address a number of weaknesses and to enhance its value to the scientific community.

General Comments

1. The background theory behind the diagnostic could be presented more clearly. While the concept of electron spin is well understood in the physical chemistry community, it is necessary to provide a brief introduction for a wider audience, along with references to literature where readers can learn more.

The paragraph in the introduction that introduces spin has been expanded and a reference to an explanation of the fundamental principles included. This paragraph now reads:

The inefficiency of ground state O₂ as an atmospheric oxidant is due to its electronic structure. In quantum mechanics, all atomic particles have an intrinsic angular momentum known as spin [Atkins and De Paula, 2014]. The spin of an electron is described by the spin quantum number, *s*, and can have values of either +½ or -½ for a single electron. The Pauli exclusion principle states that if two electrons occupy the same orbital then their spins must be paired, and thus cancel. With two unpaired electrons ground state O₂ is a spin-triplet with a total spin quantum number $S=½+½=1$ (giving a term symbol of ${}^3\Sigma_g^-$). In contrast, virtually all trace chemicals emitted into the atmosphere contain only paired electrons and are thus spin-singlets ($S=0$). The quantum mechanical spin selection rule $\Delta S=0$ means that allowed electronic transitions must not result in a change in electron spin. From a simplistic perspective (i.e. ignoring nuclear spin interactions, inter-system crossings, nuclear dipole effects etc.) this spin selection rule means that the reaction of ground state O₂ with most emitted compounds is effectively spin forbidden. Electronically excited O₂ (${}^1\Delta_g$ or ${}^1\Sigma_g^+$) is a spin singlet and is more reactive in the atmosphere but low concentrations limit its role [Larson and Marley, 1999]. Instead, atmospheric oxidation proceeds predominantly via reactions with spin-doublet oxygen-derived species ($S=½$), notably the hydroxyl (OH) and peroxy radicals (RO₂ = HO₂, CH₃O₂, C₂H₅O₂, etc.), or spin-singlet species (e.g. ozone (O₃)).

2. The paper addresses the rate of ozone production, but discussion focuses solely on long-term integrated ozone production on an annual global scale. It is not clear how applicable the new diagnostic is to smaller regions and shorter timescales where the assumption of steady state (line 141) may be less appropriate, and where emissions may be less important than transport. What is needed to extend the diagnostic to these smaller spatial and temporal scales? The potential for analysis of regional

budgets is alluded to on line 307, but no detail is provided.

The strengths and limitations of the approach should be set out more clearly. What additional insight does the new metric provide and how might this be applied to real problems (e.g., to the sensitivity of ozone production to assumptions of VOC speciation, to simplification of isoprene chemistry, or to treatments of deposition processes?) How does the approach compare with previous attempts to generate diagnostics, e.g., though the concept of photochemical ozone creation potentials (POCPs) for individual VOCs? There is little reference to earlier approaches in the field.

The aim of this paper is to describe a new approach for the study of ozone production in chemical transport models, and to illustrate this through a global budget analysis and comparison with the most commonly used diagnostic for this. The application of the new diagnostic to other scales and problems, as well as comparison to other available metrics is of interest but unfortunately outside the scope of this work. The following paragraph has been added to the conclusion section to discuss the strengths of the approach in relation to other possible applications, and also identify things that would need to be considered for this to be successful.

Future work is necessary to identify the usefulness of this approach on smaller spatial and temporal scales. For regional modelling scale, the transport flux of bonds into the domain would need to be considered alongside the emissions of bonds. However, this might help to disentangle O₃ production due to local VOC emissions from that due to VOC emissions outside of the domain. This bond focussed approach may also have usefulness on shorter timescales. For example, when considering vertical fluxes in and out of the boundary layer, a bond centred approach could help. What fraction of the bonds emitted at the surface are exported to the free troposphere. If a measurement of reactivity flux could be made this could be tested experimentally.

Specific Comments

3. Figure 4: please explain how the contributions of the R and F terms presented in the figure are derived. It is easy to see for the standard diagnostic, where the terms sum linearly to the total PO₃, but it is not as clear for the new diagnostic as the terms are no longer independent of each other (as defined in Eq. 1).

The following text has been added to the figure caption.

The P_sO₃ diagnostic parameters are derived for each model simulation using the diagnostic implementation described in Sect. 3, and the fractional change in each parameter from the base simulation calculated.

4. The meaning of the horizontal dashed lines in Figure 4 is not clear.

These are gridlines to aid comparison between plots. We have not changed this as we feel it helps interpretation of the figure, but are happy to take the editors guidance.

5. I.270: How many simulations were performed for these sensitivity studies? Please state this in the text.

Text now includes following sentence.

A set of 5 simulations was performed for each model sensitivity investigated (NO_x , isoprene and CH_4), with a common base simulation, resulting in 13 simulations in total.

- Figure 7 is not well conceived. It is not clear why a log- NO_x scale is used, given that the relationships expected are not exponential (neither line drawn here is expected to be straight, as would quickly become evident at larger or smaller NO_x levels). Perhaps plot OH vs CH_4 bond emission directly, and label the points with the NO_x level?

We thank the reviewer for this suggestion and have updated the figure to that shown below.

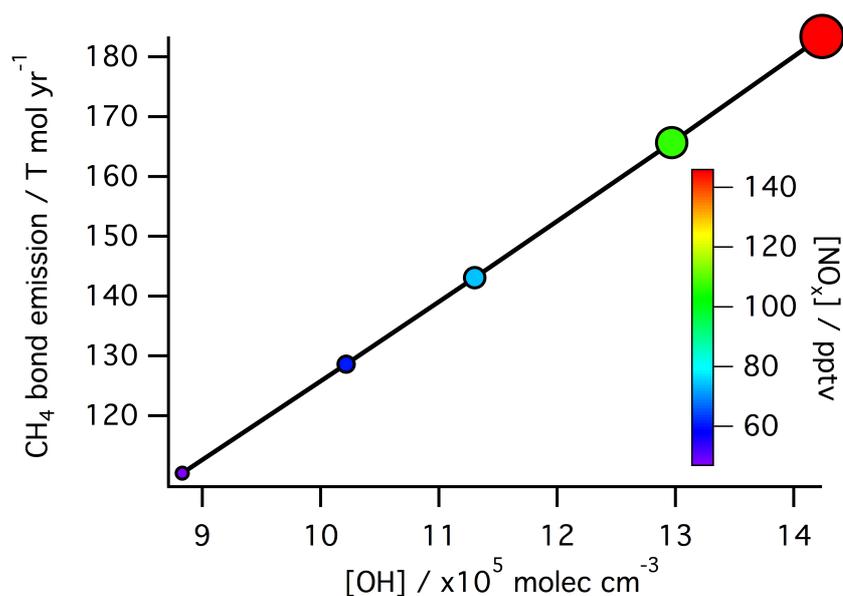


Figure 7. Effective CH_4 emissions as a function of global mean OH concentration, for simulations where NO_x emissions were changed. Marker size and colour indicate global NO_x concentration.

- Figures 8 and 10 would be more effectively presented through the use of a bar chart, so that the relative changes can be seen more clearly.

We again thank the reviewer for this suggestion and have remade the figures as bar charts (see below).

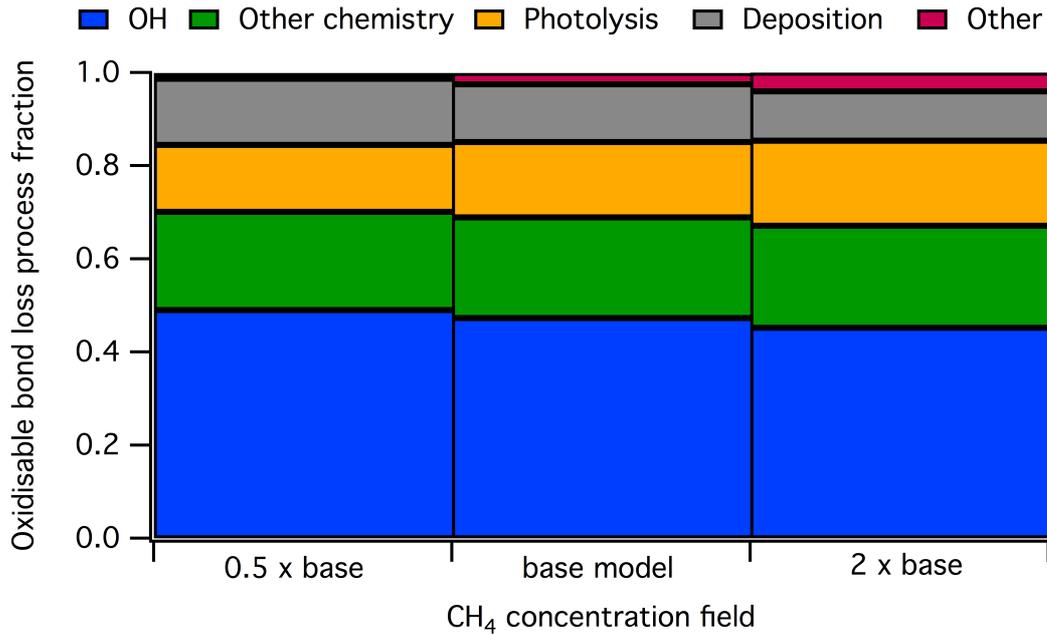


Figure 8. Oxidisable bond loss mechanism fractions under changing effective CH₄ emissions (0.5 x CH₄ concentration field, base simulation and 2 x CH₄ concentration field).

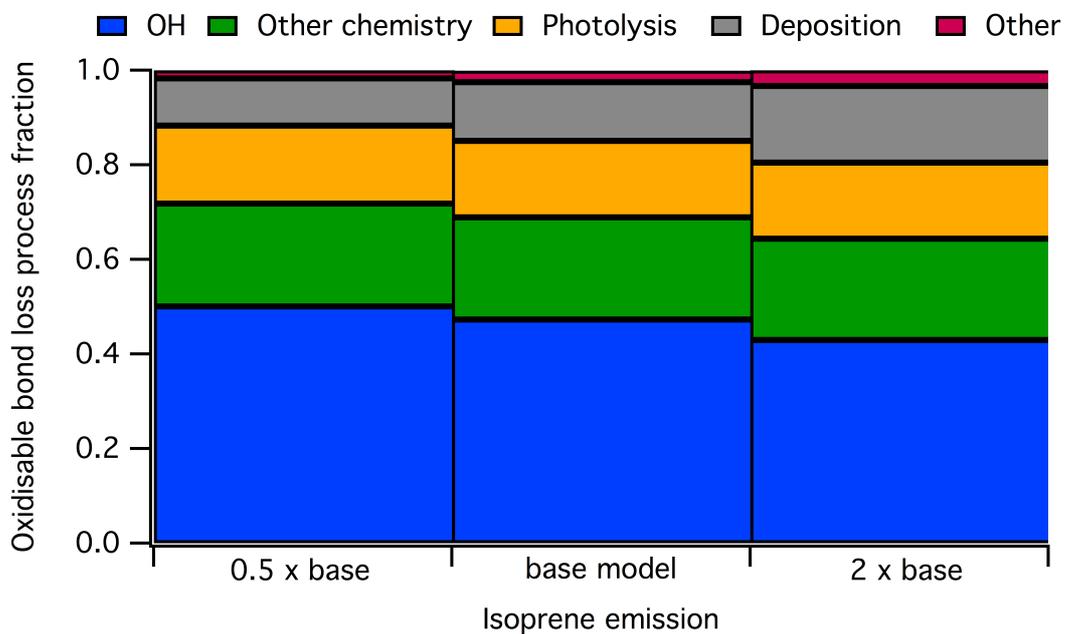


Figure 10. Oxidisable bond loss mechanism fractions under changing isoprene emissions.

- Supplement: The "errors in chemistry scheme" need some explanation, and these entries should be at the bottom of the table, as it doesn't aid the reader's comprehension to put them at the top.

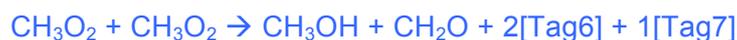
We have added the following text to the supplement and moved the entries to the bottom of table S1.

Inconsistencies within the chemistry scheme, where the lumped nature of some reactions result in a non-physical production or loss of oxidisable bonds, are also tracked as errors in the chemistry scheme.

9. The supplement needs more detail on the implementation of the approach. It would be difficult for anyone to replicate in a different model without more information about the reaction classification. It would be helpful to provide a worked example of how the multiple in the Table is arrived at, and this could be included in the supplement.

The following has been added to the supplement.

Reaction tags were added to all reactions in the chemistry scheme, and the GEOS-Chem diagnostic was used to provide a direct measure of their production. An example of how this was implemented is shown below for a select few steps of the methane oxidation scheme illustrated in Fig. 1.



Reaction tags used in example reactions: Tag1 = Oxidisable bond lost via OH chemical reaction; Tag2 = Oxidisable bond + OH → 1 radical (RO₂); Tag3 = OH + CH₄ reaction (special tag used to calculate effective CH₄ emission); Tag4 = RO₂ to peroxide; Tag5 = HO₂ to peroxide; Tag6 = RO₂ lost to carbonyl forming peroxy radical self reaction; Tag7 = Bond lost to RO₂ + RO₂ → 0 radicals.

Typos and minor issues

10. The English grammar needs a little work in places, particularly where the subject of a verb is inappropriate (e.g., "GEOS-Chem fixes CH₄ concentrations..." on line 318 would be clearer as "CH₄ concentrations are fixed in GEOS-Chem...")

This has been addressed.

11. I.133: add the before top

This has been addressed.

12. I.251: grammar in first sentence needs correcting.

This has been addressed.

13. I.358: MO₂ should be written as CH₃O₂ for consistency with line 356, and it would be helpful to do this throughout the text, e.g., line 427/8.

This has been addressed.

14. Numbers less than 10 without units are better presented as text than numerals.

This stylistic change has not been implemented, but the authors are happy to do so if the editor wishes.

Anonymous Referee #2

General Comments

This paper presents a new method for diagnosing ozone production based on the processing of chemical bonds. The authors show that this new diagnostic changes our view of the relative importance of different hydrocarbon emissions, which is an improvement over previous methods using a simple total carbon-based approach. The authors also quantify the ozone-producing efficiency of the emitted bonds. The ability of this diagnostic to separate the difference between shifting the NO/NO₂ ratio and its impact on ozone production vs. the increase in the fraction of RO₂ reacting with NO is valuable. Overall, the discussion of the diagnostic and model sensitivities is quite lengthy and could be shortened by spending less time on the discussion of methane, per the comment below. This paper should be published after addressing the comments below, in particular, how this diagnostic could be relevant to our understanding of the differences in ozone production across models without actually implementing the diagnostic in every single chemical transport model.

Specific Comments

15. The discussion of methane and isoprene is confusing due to the model implementation of methane as a fixed concentration. It might be better to focus the discussion on evaluating perturbations to isoprene emissions, and contrast that to methane, as opposed to the way it is presented now, with the caveat about model treatment of methane. Then the discussion of the dependence of methane 'emission' on OH would not be needed (i.e. Figure 7) which is difficult to follow.

We accept the reviewers comment that the discussion of methane and isoprene could be confusing. However, the fundamental differences in both their chemistries and treatment in the majority of chemical transport models mean we strongly feel that they warrant individual treatment. We have significantly shortened and simplified the discussion of methane, and have simplified Fig. 7 (see comment 6 above). As the 1st reviewer did not have an issue with the individual discussions of methane and isoprene we respectfully leave it to the editor to decide if our response to this comment is adequate.

16. This analysis would also be strengthened by presenting the types of information that global model comparisons of ozone production should include to take advantage of this type of diagnostic. For example, it seems that if all models presented their total methane, isoprene, CO, and NO_x budgets, this diagnostic would help interpret the resulting impact on ozone production without actually implementing the diagnostic in each model. This might increase the scientific contribution of this paper.

We have added the following paragraph to the conclusions section of the paper.

Another potentially important application is in model-model comparisons. Increases in our understanding of why different models calculate different O₃ production and burdens has been slow [Stevenson *et al.*, 2006; Wu *et al.*, 2007; Young *et al.*, 2013]. Although a complete tagging like that described here is unlikely to occur for all of the models involved in the comparison, a small number of additional diagnostics is likely to produce a significantly better understanding of the models. Diagnosing (1) the total bond flux (direct emissions plus the flux for those species kept constant), (2) the rate of production of RO₂ and (3) the rate of production of O₃, could help differentiate why certain models produce more or less O₃ than others. The ratios between these fluxes would help identify what aspect of the emissions of chemistry differs between the models.

17. The paragraph starting on line 341 needs clarification. What do you mean by “the final 20% due to the increased OH competing for the available oxidisable bonds.” Doesn’t this just mean that with higher NO_x, you get higher OH concentrations and thus you increase the concentration of RO₂ as well and NO?

This sentence has been reworded to avoid confusion.

“the final 20% due to the increased OH concentration competing for the available oxidisable bonds and resulting in increased RO₂ production.”

Technical Corrections

18. Is discussing SO₂ oxidation relevant to ozone in any way? If not, it is confusing and should be removed.

Although SO₂ oxidation has minimal direct impact on O₃ production it is still a source of peroxy radicals that are central to this diagnostic approach ($\text{SO}_2 + \text{OH} + \text{O}_2 \rightarrow \text{SO}_3 + \text{HO}_2$). We therefore would prefer to keep the sentence on SO₂ for completeness and also to aid others in reproducing the diagnostic approach.

19. You say that over a long enough timescale, the global atmosphere can be considered to be in steady-state, and thus equation (1) applies. Please clarify the conditions where this diagnostic is useful/applicable. For example, could it be used for a daily analysis of ozone production.

See response to comment 2.

20. Please be consistent with the use of CH₃O₂ or MO₂.

This has been addressed.

21. On line 438, the sentence that starts with “With the majority” is not a full sentence.

This has been addressed.

22. On line 440, remove the comma after OH.

This has been addressed.

A new diagnostic for tropospheric ozone production

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Abstract

Tropospheric ozone is important for the Earth's climate and air quality. It is produced during the oxidation of organics in the presence of nitrogen oxides. Due to the range of organic species emitted and the chain like nature of their oxidation, this chemistry is complex and understanding the role of different processes (emission, deposition, chemistry) is difficult. We demonstrate a new methodology for diagnosing ozone production based on the processing of bonds contained within emitted molecules, the fate of which is determined by the conservation of spin of the bonding electrons. Using this methodology to diagnose ozone production in the GEOS-Chem chemical transport model, we demonstrate its advantages over the standard diagnostic. We show that the number of bonds emitted, their chemistry and lifetime, and feedbacks on OH are all important in determining the ozone production within the model and its sensitivity to changes. This insight may allow future model-model comparisons to better identify the root causes of model differences.

1. Introduction

The chemistry of the troposphere is one of oxidation [Levy, 1973; Kroll *et al.*, 2011]. Organic compounds together with nitrogen and sulfur containing molecules are emitted into the troposphere where they are oxidised into compounds which can either be: absorbed by the biosphere; are involatile enough to form aerosols; can deposit to the surface; or be taken up by clouds and rained out. The oxidation of these compounds is significantly slower than might be expected based on the atmospheric composition of 20% molecular oxygen (O₂).

31 The inefficiency of ground state O₂ as an atmospheric oxidant is due to its electronic
32 structure. In quantum mechanics, all atomic particles have an intrinsic angular
33 momentum known as spin [Atkins and De Paula, 2014]. The spin of an electron is
34 described by the spin quantum number, s, and can have values of either +½ or -½ for a
35 single electron. The Pauli exclusion principle states that if two electrons occupy the
36 same orbital then their spins must be paired, and thus cancel. With two unpaired
37 electrons, ground state O₂ is a spin-triplet with a total spin quantum number S=½+½=1,
38 (giving a term symbol of ³Σ_g⁻). In contrast, virtually all trace chemicals emitted into
39 the atmosphere contain only paired electrons and are thus spin-singlets (S=0). The
40 quantum mechanical spin selection rule ΔS=0 means that allowed electronic
41 transitions must not result in a change in electron spin. From a simplistic perspective
42 (i.e. ignoring nuclear spin interactions, inter-system crossings, nuclear dipole effects
43 etc.) this spin selection rule, means that the reaction of ground state O₂ with most
44 emitted compounds is effectively spin forbidden. Electronically excited O₂ (¹Δ_g or
45 ¹Σ_g⁺) is a spin singlet and is more reactive in the atmosphere but low concentrations
46 limit its role [Larson and Marley, 1999]. Instead, atmospheric oxidation proceeds
47 predominantly via reactions with spin-doublet oxygen-derived species (S=½), notably
48 the hydroxyl (OH) and peroxy radicals (RO₂ = HO₂, CH₃O₂, C₂H₅O₂, etc.), or spin-
49 singlet species (e.g. ozone (O₃)).

50 One of the few spin-triplet species in the atmosphere other than O₂ is the ground state
51 of atomic oxygen (O(³P)), which readily undergoes a spin allowed reaction with O₂ to
52 produce the spin-singlet O₃ molecule. This spin allowed reaction is responsible for the
53 creation of O₃ in both the stratosphere, where it forms the protective O₃ layer, and the
54 troposphere. The ability of O₃ to oxidise other spin-singlet species makes it a powerful
55 oxidant, and it is thus considered a pollutant with negative health effects. Sources of
56 O(³P) within the troposphere are limited because solar photons at sufficiently short
57 wavelengths to directly photolyse O₂ to O(³P) are essentially unavailable.

58 Aside from the photolysis of O₃ itself, the only other significant source of
59 tropospheric O(³P) is the photolysis of nitrogen dioxide (NO₂) [Crutzen, 1971].
60 Nitrogen oxides are emitted into the troposphere as nitrogen oxide (NO), which can be
61 oxidised to NO₂ by O₃ and other oxidants. A large thermodynamic energy barrier
62 prevents oxidation of NO to NO₂ by the OH radical [Nguyen et al., 1998], and
63 therefore NO oxidation occurs through reaction with either O₃ or RO₂. In terms of O₃

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71 production, the oxidation of NO by O₃ forms a null cycle. Thus only the reaction of
72 NO with RO₂ leads to a net production of O₃.

73 Exploring the distribution, source and sinks of tropospheric O₃ is a central theme of
74 atmospheric science. Chemical transport models (online and offline) are essential
75 tools enabling this understanding but their validity needs to be continually assessed.
76 Model-model comparison exercises are commonly performed to assess performance,
77 and comparisons of modelled O₃ budgets traditionally form part of this assessment
78 [Stevenson *et al.*, 2006; Wu *et al.*, 2007; Wild, 2007; Young *et al.*, 2013]. Ozone
79 production is diagnosed from the flux of NO to NO₂ via reaction with each of the
80 speciated RO₂ in the model's chemical schemes. This approach provides information
81 on the relative importance of the different RO₂ in the fast NO + RO₂ reactions within
82 the model, but gives very little detail on how the longer time scale model processes
83 (emissions, chemistry, deposition) influence O₃ production. Thus exploring the
84 reasons that models differ in their O₃ production is difficult and progress has been
85 slow.

86 A new diagnostic framework that links large scale model drivers such as emission,
87 chemistry, and deposition to O₃ production would allow an improved assessment of
88 why model ozone budgets differ. We attempt to provide such a framework here.

89 **2. A new diagnostic framework.**

90 The rate of production of tropospheric O₃ is limited by the rate of oxidation of NO to
91 NO₂, which is in turn limited by the rate of production of peroxy radicals (RO₂).
92 Peroxy radicals form through association reactions of hydrogen (H) atoms or alkyl
93 radicals (both spin-doublets, S=1/2) with O₂, forming a highly reactive spin-doublet
94 radical on an oxygen atom. This spin allowed reaction converts spin-triplet O₂ that
95 cannot react with spin-singlet pollutants into a spin-doublet O₂ containing species that
96 can. As such the formation of RO₂ is central to the atmosphere's oxidation capacity,
97 and its production is limited by the rate of production of H atoms or alkyl radicals.
98 Thus the maximum potential rate of tropospheric O₃ production is equal to the rate at
99 which H atoms and alkyl radicals are produced.

100 Hydrogen atoms and alkyl radicals are predominantly produced via the spin allowed
101 breaking of the spin-pairing between the two electrons in a C or H containing covalent
102 bond (S=0), such as those in hydrocarbons. These spin-pairings can be broken in the

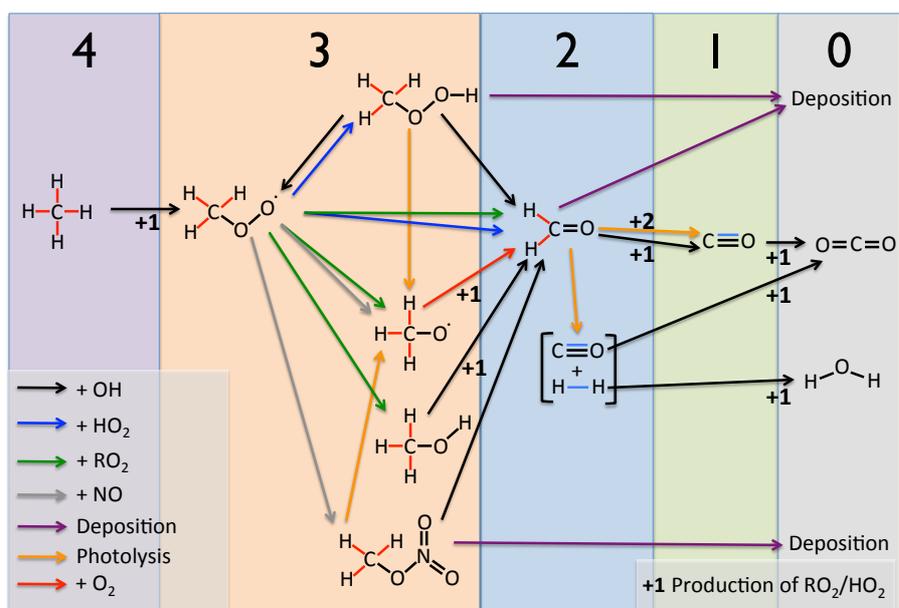
103 atmosphere either chemically or photolytically, with the products necessarily
104 conserving spin. The breaking of a covalent bond by a photon ($s=1$) can result in two
105 products with $S=1/2$ or two products with $S=0$. Likewise, oxidation by a radical ($S = 1/2$)
106 will result in one product with $S=0$ and one with $S=1/2$, because the unpaired electron
107 on the radical reactant pairs with one of the covalent bond electrons to produce a spin-
108 singlet.

109 Although the majority of RO_2 is formed from emitted C or H containing covalent
110 bonds, there are a few notable exceptions. Hydrogen atoms can also be produced
111 through the oxidation of CO to CO_2 by OH. During this reaction the coordinate bond
112 between the C and O atom is broken and the H atom is produced via the breaking of
113 the O-H bond. The other notable exception is the oxidation of an SO_2 lone pair of
114 electrons to SO_3 by OH, where again the H atom produced comes from the OH. In
115 both of these exceptions a spin-singlet electron pairing (CO coordinate bond or SO_2
116 lone pair) is broken during the production of the H atom, and we can therefore
117 consider these reactions as similar to the breaking of C or H containing covalent bond.
118 For simplicity these spin-singlet electron pairings that can be broken in the
119 troposphere to produce either a H atom or alkyl radical will be referred to as
120 “oxidisable bonds” (C-C, C-H, C=C, CO coordinate bond, S:).

121 Tropospheric O_3 production occurs through the oxidation of NO by RO_2 . Following
122 the above rationale, these RO_2 are produced during the spin allowed breaking of
123 oxidisable bonds predominantly contained within emitted VOCs. This perspective
124 allows us to build a new metric for the production of tropospheric O_3 based around the
125 spin conserving properties of oxidisable bond breaking. In the extreme case, all
126 oxidisable bonds are photolysed to produce two spin-doublet RO_2 products, which
127 then react exclusively with NO to generate O_3 . Thus at steady state, the maximum rate
128 of O_3 production is equal to the rate of production of RO_2 , which is equal to twice the
129 rate of destruction of the number of oxidisable bonds. This in turn is equal to twice the
130 rate of emission of oxidisable bonds. Deviation from this maximum is determined by:

- 131
- The relative importance of processes that produce spin-singlet vs. spin-
132 doublet products during oxidisable bond breaking;
 - The fraction of spin-doublet products from oxidisable bond breaking which
133 form RO_2 ;
 - The fraction of RO_2 that go on to oxidize NO to NO_2 .
- 135

136 To illustrate this Fig. 1 shows the tropospheric oxidation of a methane (CH_4) molecule
 137 through various steps to either a carbon dioxide (CO_2) molecule or a species that is
 138 deposited (CH_3OOH , CH_2O , CH_3NO_3). Methane contains 4 x C-H oxidisable bonds
 139 (8 paired bonding-electrons) and as the oxidation proceeds, the number of oxidisable
 140 bonds decays to zero. Figure 1 highlights the steps in the tropospheric CH_4 oxidation
 141 mechanism that form spin-doublet products, with between 1 and 5 RO_2 produced
 142 depending on the oxidation pathway. This compares with the theoretical maximum of
 143 8 if all the original C-H bonds were photolysed to yield 2 spin-doublet products.



144
 145 **Figure 1. Peroxy radical production during the tropospheric oxidation of CH_4 .**
 146 **Moving from left to right, the oxidisable bonds (emitted = red, produced = blue)**
 147 **present in CH_4 are removed via a range of tropospheric processes, indicated by**
 148 **the coloured arrows. The large numbers across the top of the figure indicate the**
 149 **number of oxidisable bonds at each stage of this oxidation. The production of**
 150 **RO_2 is indicated by the +1/+2 numbers with the associated process arrows for**
 151 **producing 1 or 2 RO_2 respectively.**

152 The principal atmospheric source of oxidisable bonds is the emission of C-H, C-C and
 153 C=C bonds in hydrocarbons, with the only other significant sources being the
 154 emission of CO and the chemical production of CO and H_2 during hydrocarbon
 155 oxidation. Over a long enough timescale, the global atmosphere can be considered to

156 be in a chemical steady state, where the rate of loss of oxidisable bonds is balanced by
157 the rate of production or emission. Thus the O₃ production rate can be described by
158 equation (1), where the O₃ production metric P_sO_3 is equal to the number of spin-
159 paired electrons in oxidisable bonds (i.e. twice the sum of the number of oxidisable
160 bonds emitted (E_{bonds}) and chemically produced (P_{bonds})), multiplied by the number of
161 spin-doublet radicals produced per oxidisable bond break divided by the maximum of
162 2 ($F_{Radicals}$), multiplied by the fraction of the radicals produced which are RO₂ (F_{RO2}),
163 multiplied by the fraction of RO₂ that goes on to react with an NO to produce an O₃
164 molecule (F_{NO}). A small correction (I) for the production of RO₂ via reactions of spin-
165 doublet radicals other than those that result in the breaking of oxidisable spin-pairings
166 (e.g. $O_3 + OH \rightarrow HO_2 + O_2$) is included.

$$167 \quad P_sO_3 = \left((2 \times (E_{bonds} + P_{bonds}) \times F_{radicals} \times F_{RO_2}) + I \right) \times F_{NO} \quad (1)$$

168 **3. Implementation**

169 We use the GEOS-Chem model to evaluate this new O₃ production diagnostic. GEOS-
170 Chem is a global chemical transport model of tropospheric chemistry, aerosol and
171 transport (www.geos-chem.org version 9-02). The model is forced by assimilated
172 meteorological and surface fields (GEOS-5) from NASA's Global Modelling and
173 Assimilation Office, and was run at 4°x5° spatial resolution. The model chemistry
174 scheme includes O_x, HO_x, NO_x, BrO_x and VOC chemistry as described in Mao et al.
175 [2013] as are the emissions. The new P_sO_3 diagnostic has been implemented via the
176 tracking of reactions by type in the GEOS-Chem chemical mechanism file (further
177 details given in the SI). This tracking of reactions enables the fate of all oxidisable
178 bonds as well as the production and loss of all RO₂ within the model to be determined
179 using the standard GEOS-Chem production and loss diagnostic tools. Model
180 simulations were run for 2 years (July 1st 2005 – July 1st 2007) with the first year used
181 as a spin up and the diagnostics performed on the second year.

182 The standard GEOS-Chem diagnostic for O₃ production (PO_3) is shown on the left
183 side of Table 1. This emphasizes the very fast cycling between NO and NO₂, but
184 provides little in terms of higher process level information. The right side of Table 1
185 shows the new budget for P_sO_3 , which tracks the processing of oxidisable bonds
186 within the model. Both diagnostic methods give the same final answer but our new
187 methodology provides more process level detail. Figure 2 illustrates this new process

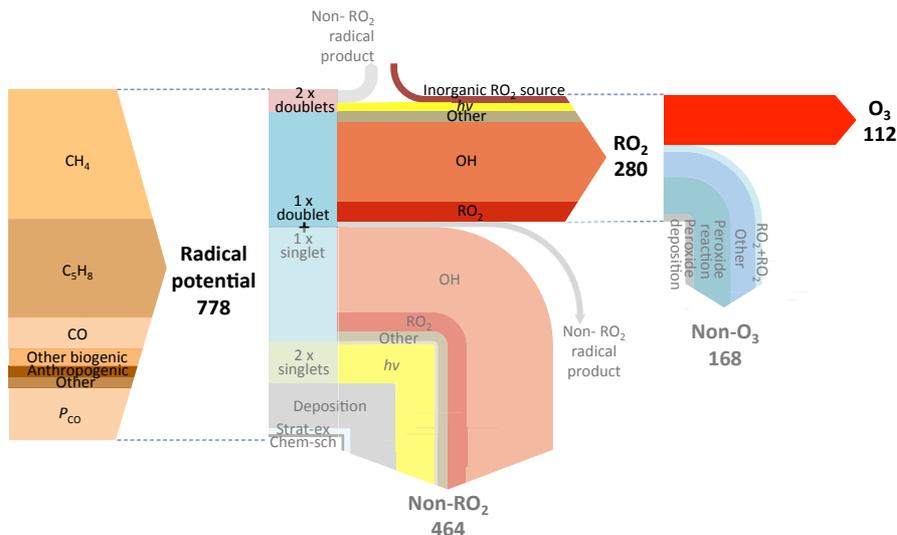
188 based approach, showing the flow of emitted oxidisable spin-paired electrons (bonds)
 189 to O₃ and the magnitude of the various mechanisms that contribute to and compete
 190 with O₃ production. The annual oxidisable bond emission of 389 T mol yr⁻¹ has the
 191 potential to create 778 T mol yr⁻¹ of radicals. If all oxidisable bonds were broken by
 192 photons to produce two radical products the RO₂ production would be 778 T mol yr⁻¹.
 193 If the oxidisable bonds were instead broken via radical reaction (e.g. OH) then RO₂
 194 production would be 389 T mol yr⁻¹. The various oxidisable bond breaking / removal
 195 pathways within the model result in the production of 280 T mol yr⁻¹ of RO₂, with the
 196 remainder largely producing stable spin singlet products.

197 Of the 280 T mol yr⁻¹ RO₂ produced, 112 T mol yr⁻¹ reacts with NO to produce O₃.
 198 The remainder is lost through the reaction or deposition of RO₂ reservoir species
 199 (RO_{2y}= RO₂ + peroxides + peroxy-acetyl nitrates). For example the production of
 200 methylperoxide (CH₃O₂ + HO₂ = CH₃OOH) results in the loss of 2 RO₂'s. However,
 201 the reaction of methylperoxide with OH can re-release CH₃O₂ (CH₃OOH + OH =
 202 CH₃O₂ + H₂O). Thus, the production of methylperoxide represents the loss of a HO₂
 203 and the movement of a CH₃O₂ into a peroxide RO_{2y} reservoir species. The deposition
 204 of a peroxide molecule is thus the loss of a RO_{2y} reservoir species. Notable in Fig. 2 is
 205 that the role of PAN and nitrate removal of global RO_{2y} is negligible, instead being
 206 dominated by peroxide production and loss and the reaction of RO₂ with O₃.

<i>PO</i> ₃ / T mol Yr ⁻¹		<i>PO</i> ₃ / T mol Yr ⁻¹ (except <i>F</i> _{Radicals} , <i>F</i> _{RO2} , and <i>F</i> _{NO} which are all unitless)	
NO + HO ₂ → NO ₂	74	<i>E</i> _{bonds}	330
NO + CH ₃ O ₂ → NO ₂	27	<i>P</i> _{bonds}	58
Other RO ₂ + NO → NO ₂	10	<i>F</i> _{radicals}	0.40
Other	1	<i>F</i> _{RO2}	0.86
		Inorganic RO ₂ source	15
		<i>F</i> _{NO}	0.40
<i>PO</i>₃	112	<i>P</i>_sO₃	112

207 **Table 1. Comparison of ozone production diagnostics for GEOS-Chem base**
 208 **simulation. Standard model *PO*₃ diagnostics (left column) show reactions**
 209 **responsible for NO to NO₂ conversions but provide little process level**

210 information. The new P_sO₃ (right) provides increased information on the
 211 processes controlling O₃ production within the model.

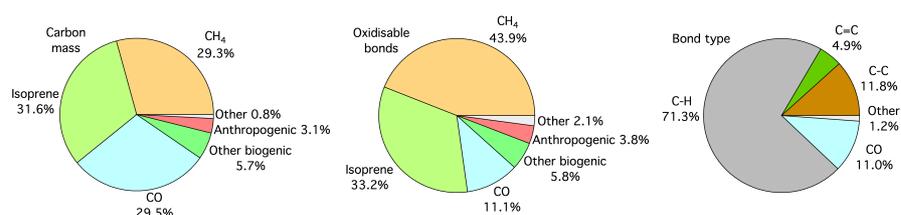


212
 213 **Figure 2. Flow of oxidisable bonds to O₃ production in the GEOS-Chem base**
 214 **simulation. Arrows are coloured according to process and the arrow thickness is**
 215 **proportional to the flux through that channel. Spin-paired electrons are input as**
 216 **oxidisable bonds into the model (left arrow), with the potential to create 778 T**
 217 **mol yr⁻¹ of radicals. The actual fate of these bonds is shown in the central arrow,**
 218 **producing 280 T mol yr⁻¹ of RO₂, of which 112 T mol yr⁻¹ reacts with NO to**
 219 **produce O₃ (right arrow).**

220 3.1 Emitted oxidisable bonds

221 The fuel for tropospheric oxidation chemistry is the emission of oxidisable bonds,
 222 predominantly in the form of hydrocarbons. The production of tropospheric O₃ from
 223 the spin-paired bonding electrons emitted into the standard GEOS-Chem model
 224 occurs with an efficiency of 14% (112 T mol yr⁻¹ molecules of O₃ produced / 778 T
 225 mol yr⁻¹ spin-paired electrons emitted as oxidisable bonds, Fig.2). These spin-paired
 226 bonding electrons are predominantly emitted in the form of CH₄, isoprene (C₅H₈) and
 227 CO (37%, 28%, and 9% respectively). Oxidisable bonds produced during chemical
 228 reactions (P_{bonds}), account for 15% of the net source. Figure 3 shows emissions of CO
 229 and hydrocarbons in the standard GEOS-Chem simulation in terms of mass of carbon
 230 per compound, number of oxidisable bonds per compound and as number of bonds in

231 different oxidisable bond types. The commonly used carbon mass approach splits
 232 emissions approximately equally between each of the major sources (CH₄ (29%),
 233 Isoprene (32%) and CO (30%)). In contrast, the oxidisable bonds accounting approach
 234 apportions hydrocarbon emissions 44%, 33% and 11% for CH₄, isoprene and CO
 235 respectively. This highlights the high number of oxidisable bonds per carbon atom in
 236 CH₄ (4) compared to isoprene (2.8) and CO (1). Thus efforts to consider emissions on
 237 a per-bond basis may provide more insight into chemical processes, as it is these
 238 bonds that ultimately determine the chain-like chemistry rather than the mass of
 239 carbon atoms. This helps to emphasise the relative importance of CH₄ emissions on
 240 global tropospheric chemistry compared with other emissions such as isoprene or CO.
 241 The type of oxidisable bond emitted is overwhelmingly C-H (71%).



242
 243 **Figure 3. Pie charts showing hydrocarbon emissions in the base GEOS-Chem**
 244 **simulation. Emissions split by carbon mass (left), number of oxidisable bonds**
 245 **(centre) and bond type (right).**

246 The total emission and production of oxidisable bonds has the potential to create 778
 247 T mol yr⁻¹ of radicals. However, only 6% of the oxidisable spin-pairings are broken to
 248 give the maximum 2 spin-doublet products (e.g. radical channel of CH₂O photolysis).
 249 The majority (68%) are oxidized via reaction with a spin-doublet species (OH) to
 250 produce 1 spin-singlet and 1 spin-doublet product (e.g. OH + VOC). The remaining
 251 26% of spin-paired electrons are removed to form two spin-singlets (e.g. the non-
 252 radical channel of CH₂O photolysis). Thus, of the 778 Tmol yr⁻¹ spin-paired electrons
 253 emitted or produced only 265 T mol yr⁻¹ (34%) are converted into RO₂, with an
 254 additional 15 T mol yr⁻¹ produced from reactions such as O₃ + OH → HO₂ + O₂ (I).
 255 The efficiency of O₃ production from the available oxidisable bonds is further reduced
 256 as only 40% of the 280 T mol yr⁻¹ of RO₂ produced react with NO to produce NO₂.
 257 The remainder is lost either through the self-reaction of RO₂ or via loss through
 258 deposition or reaction of RO_{2y} reservoir species (e.g. peroxides). Thus overall 14% of
 259 the emitted bonding electrons go on to make O₃.

260 The new O₃ production diagnostic presented here (*P_sO₃*) shows the impact of
261 processes such as emission, deposition and chemical mechanism, and provides
262 significantly more detail than the standard *PO₃* diagnostic approach (Table 1). We
263 now explore the sensitivity of model O₃ production to changing emissions of NO_x and
264 VOC from the perspective of the two diagnostic methods.

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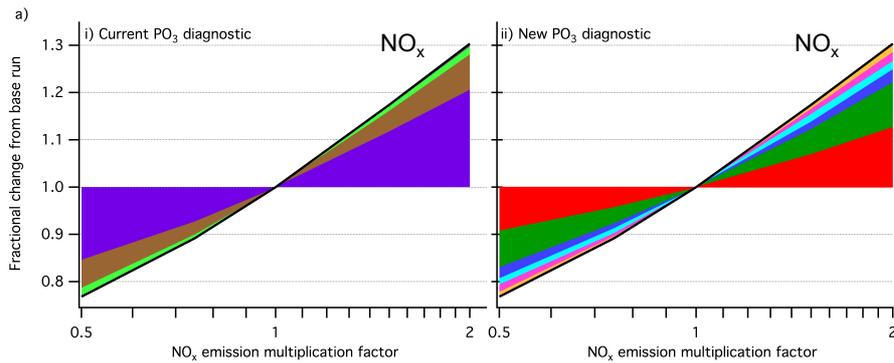
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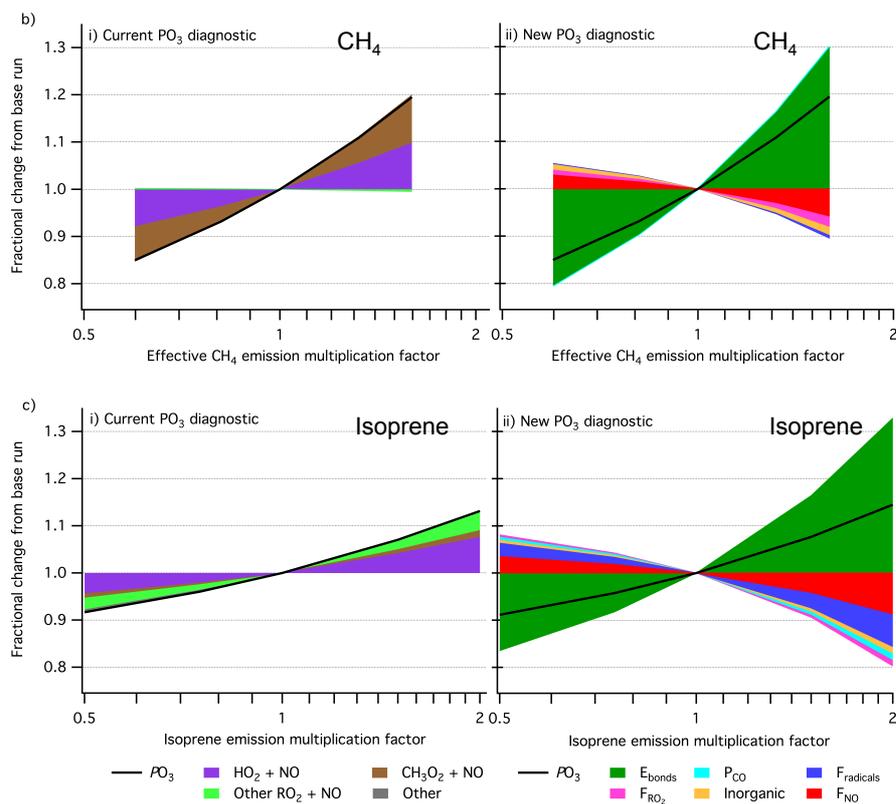
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265 4 Model sensitivities

266 Understanding model response to changing emissions is an important tool for
267 considering policy interventions. The major controls on O₃ production are emissions
268 of NO_x and VOCs. We show in Fig. 2 that from the perspective of global O₃
269 production, oxidisable bond emissions are dominated by CH₄ and isoprene. Figure 4
270 shows the impact of changing emissions of NO_x, isoprene and CH₄ on O₃ production
271 from both the perspective of this new methodology and the conventional NO+RO₂
272 diagnostic approach. A set of 5 simulations was performed for each model sensitivity
273 investigated (NO_x, isoprene and CH₄), with a common base simulation, resulting in 13
274 simulations in total. The following sections investigate these model responses and use
275 the new diagnostic to provide insight into the processes driving the observed response
276 in O₃ production.



277



280

281

282 **Figure 4. Understanding the effect of NO_x and VOC emissions on ozone**
 283 **production at the process level. Stack plots showing fractional change in model**
 284 **PO_3 compared to base simulation and associated contributions from the current**
 285 **PO_3 (i) and new P_sO_3 (ii) diagnostic parameters under changing NO_x emissions**
 286 **(a), effective CH_4 emission (b) and isoprene emission (c). The P_sO_3 diagnostic**
 287 **parameters are derived for each model simulation using the diagnostic**
 288 **implementation described in Sect. 3, and the fractional change in each parameter**
 289 **from the base simulation calculated.**

290 4.1 NO_x emissions

291 Figure 4a diagnoses the relative response of GEOS-Chem O_3 production to changing
 292 NO_x emissions, using simulations where NO_x emissions from anthropogenic, biomass
 293 burning, biofuels, soil and lighting sources were multiplied by factors of 0.5 - 2.
 294 Increasing NO_x emissions increases O_3 production. The standard RO_2+NO diagnostic
 295 (Fig.4a(i)) shows that fractional contributions to the total change in PO_3 from HO_2

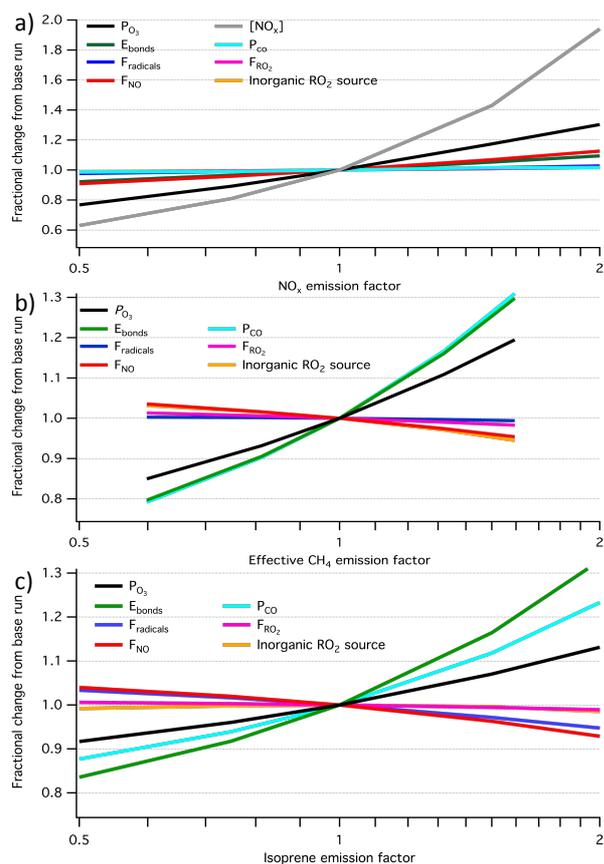
296 | (67%), methyl-peroxy (CH_3O_2) (25%), and other RO_2 (8%) remain approximately
297 constant across the NO_x emission range investigated. This diagnostic provides little
298 detail on the processes driving the change in O_3 production under changing NO_x
299 emissions. In contrast, Fig. 4a(ii) is based on the new $P_s\text{O}_3$ diagnostic and shows a
300 range of process level changes occurring as NO_x emissions change.

301 4.1.1 Impact of changing NO_x emission on F_{NO}

302 Unsurprisingly, as NO_x emissions increase the fraction of RO_2 reacting with NO to
303 produce NO_2 (F_{NO}) increases (red section in Fig. 4a(ii)). However, this impact only
304 accounts for around 40% of the increase in $P_s\text{O}_3$. Figure 5a shows the fractional
305 change in all the $P_s\text{O}_3$ efficiency parameters and the global mean NO_x concentration
306 as a function of the changing NO_x emission. As NO_x emissions increase the increase
307 in NO_x concentration in the model is somewhat dampened. Halving the NO_x emission
308 leads to NO_x burdens dropping by ~35%, and doubling leads to an increase of 95%.
309 This dampening is due to the impact of NO_x emissions on OH (see section 4.1.2),
310 which is the dominant sink for NO_x . Increasing NO_x increases OH concentrations,
311 which in turn shortens the NO_x lifetime thus dampening the response of concentration
312 to emission.

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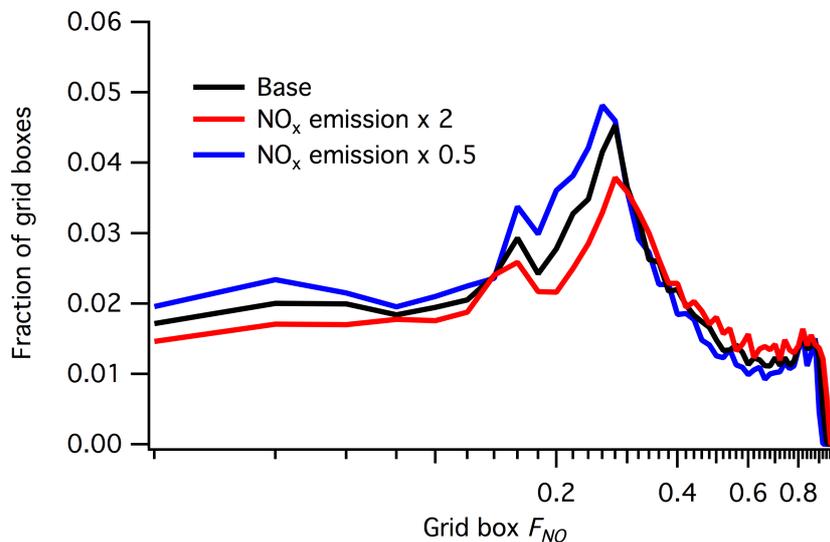


314

315 **Figure 5. Fractional change in new P_5O_3 diagnostic parameters from base run**
 316 **against changing NO_x emission (a); effective CH_4 emission (b); and isoprene**
 317 **emission (c).**

318 The response of F_{NO} to changes in NO_x emissions is also dampened relative to the
 319 change in NO_x emissions. This is due to spatial variability in F_{NO} , which is not
 320 affected uniformly by changing NO_x emissions. Figure 6 shows the probability
 321 distribution of F_{NO} values across all model grid boxes for the base simulation and the
 322 half and doubled NO_x emission simulations (black, blue and red lines respectively).
 323 For example, in a grid-box in the continental boundary layer where RO_2 reacts
 324 overwhelmingly with NO , doubling the NO_x emission may move F_{NO} from 0.90 to
 325 0.95 but it can't double it. Similarly, in the remote boundary layer where RO_2 reacts
 326 overwhelmingly with other RO_2 doubling NO_x emissions may move F_{NO} from 0.3 to

327 0.4 but again it doesn't double. Thus the geographical spread of NO_x chemistry limits
 328 the change in F_{NO} caused by changing NO_x emissions. The spatial variability in the
 329 new P_sO_3 diagnostic parameters shows that this approach has significant potential in
 330 the analysis of regional O_3 budgets as well as global.



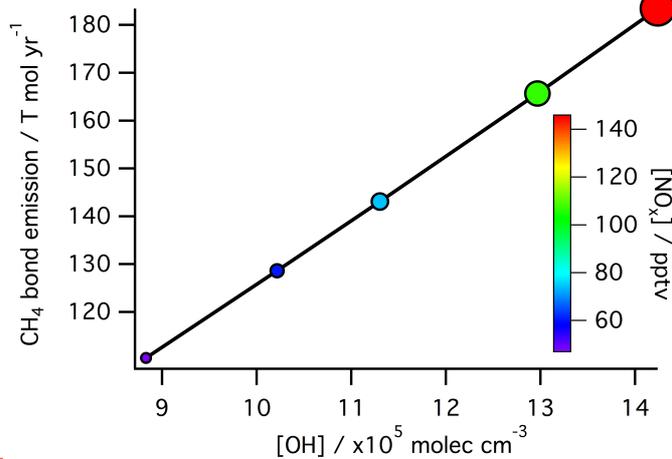
331
 332 **Figure 6. Effect of NO_x emission on distribution of F_{NO} values (log scale).** F_{NO}
 333 values for each model grid box in the base and NO_x emission x 0.5 and x 2
 334 simulations, split into 50 x 0.02 width bins.

335 **4.1.2 Impact of changing NO_x emission on E_{bonds}**

336 Figure 4a(ii) shows that 60% of the response in P_sO_3 to changing NO_x emission is due
 337 to factors other than F_{NO} , with 40% of the increase due to changes in the emissions
 338 (E_{bonds} : 32%) and chemical production (P_{bonds} : 8%) of oxidizable bonds. This increase
 339 in E_{bonds} is surprising given VOC emissions are unchanged in these simulations.
 340 However, increasing NO_x emissions results in an increased OH concentration in the
 341 model, which then leads to an increase in CH_4 oxidation. Methane (CH_4)
 342 concentrations are fixed in GEOS-Chem, resulting in an increase in the effective CH_4
 343 emission, as OH concentrations increase, causing an increase in the total bond
 344 emission (E_{bonds}). Figure 7 shows the response of effective CH_4 bond emission to
 345 global mean OH concentration as it changes with global mean NO_x concentration,

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354 More CH₄ oxidation also leads to more CH₂O production and in turn more CO
 355 production (P_{CO}), accounting for a significant fraction of the increase in this term.

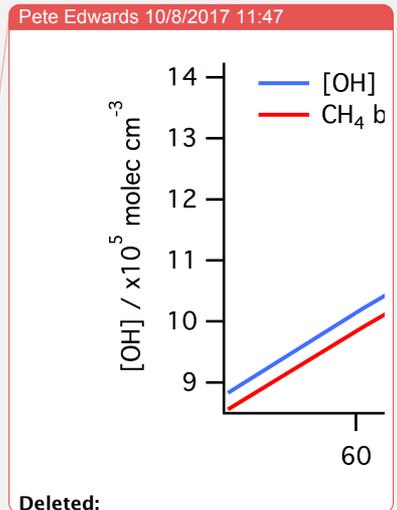


356
 357 **Figure 7. Effective CH₄ emissions as a function of global mean OH concentration,**
 358 **for simulations where NO_x emissions were changed. Marker size and colour**
 359 **indicate global NO_x concentration.**

360 4.1.3 Impact of changing NO_x emission on F_{radicals}, F_{RO2} and I

361 The fraction of radicals produced from bond oxidation (F_{radicals}) and the fraction of
 362 those radicals which are RO₂ (F_{RO2}) show slight positive increase with NO_x emission,
 363 accounting for 9% and 6% of the change in P_sO₃ respectively. This reflects changes in
 364 the partitioning of the fate of the oxidisable bonds, and is largely due to the changes in
 365 OH. As OH increases with NO_x emission, the rate of chemical oxidation of bonds
 366 increases at the expense of other losses, in particular deposition. The inorganic RO₂
 367 source term (I) also correlates with NO_x emission, as it is largely determined by the
 368 concentrations of OH and O₃. This change accounts for 5% of the observed change in
 369 P_sO₃.

370 Thus, with this new diagnostic methodology it is evident that only 40% of the model
 371 O₃ production response to changing NO_x emission is due to the direct effect of
 372 increasing NO concentration on the rate of RO₂ + NO reactions. Another 40% is due
 373 to fixing the concentration of CH₄ within the model, with the final 20% due to the
 374 increased OH concentration competing for the available oxidisable bonds and
 375 resulting in increased RO₂ production.



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384 4.2 Changing effective CH₄ emissions

385 As Fig. 2 shows CH₄ to be the largest single source of oxidisable bonds, this section
386 investigates the response of the O₃ production diagnostics to changing CH₄ emissions.

387 Figure 4b shows the O₃ production diagnostics response to varying the CH₄ emission
388 rate within the model. As the model uses prescribed CH₄ concentrations, these were
389 varied by factors of between 0.5 and 2 from the base simulation, and the CH₄ emission
390 diagnosed from the loss rate of CH₄ to reaction with OH, the only CH₄ loss in the
391 model. We describe this as the effective CH₄ emission.

392 As effective CH₄ emission increases, O₃ production also increases. The standard
393 diagnostic (Fig.4b(i)) shows that this increase occurs through an increased rate of
394 reaction of HO₂ and CH₃O₂ with NO, as would be expected as these are the RO₂
395 produced during CH₄ oxidation. The rate of other RO₂ + NO reactions actually
396 decreases slightly as CH₄ emissions increase, due to lower OH concentrations and
397 increased competition for NO from HO₂ and CH₃O₂. The new diagnostic (Fig.4b(ii)),
398 however, shows the increase in O₃ production with increasing effective CH₄ emission
399 is not simply a result of more HO₂ and CH₃O₂.

400 4.2.1 Impact of changing effective CH₄ emission on F_{NO}

401 The observed change in P_sO₃ is around one third smaller than would be expected from
402 the increase in the oxidisable bond emission (E_{bonds}) and bond production (P_{bonds})
403 terms alone. This is due to a countering decrease in the other efficiency parameters
404 with increasing effective CH₄ emission. Figure 5b shows the fractional change in all
405 the efficiency parameters as a function of the changing effective CH₄ emission. The
406 decrease in the fraction of RO₂ reacting with NO to produce NO₂ (F_{NO}) is driven by
407 increasing O₃ concentrations, which push the NO/NO₂ ratio towards NO₂. This
408 reduces the availability of NO to react with RO₂ thereby reducing O₃ production. This
409 shift in the NO/NO₂ ratio also increases NO_x loss within the model with increasing
410 CH₄ emission, as the increased CH₄ oxidation increases RO₂ concentrations resulting
411 in larger losses of NO₂ via compounds such as peroxyacetyl nitrate (PAN) and
412 peroxyntic acid (PNA).

413 4.2.2 Impact of changing effective CH₄ emission on E_{bonds}

414 Increasing the effective CH₄ emission results in an increase in E_{bonds}. Changing the
415 fraction of total emitted oxidisable bonds from CH₄ does however have significant

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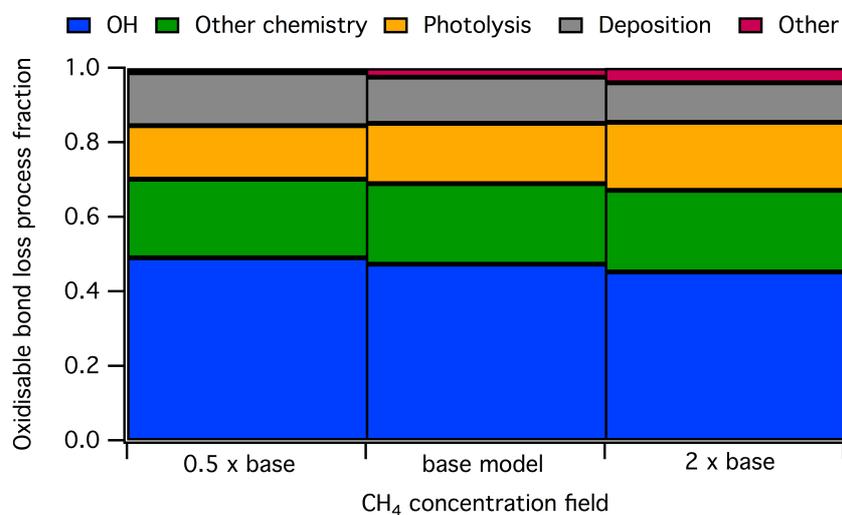
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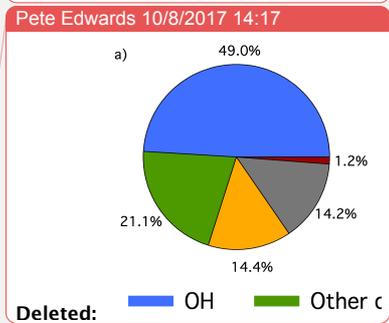
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422 consequences on the loss mechanisms of these bonds, which influences the other
 423 efficiency parameters. Figure 8 show the split of oxidisable bond loss mechanisms in
 424 the base simulation and those with the CH₄ concentration fields multiplied by 0.5 and
 425 2. As the effective CH₄ emission increases the fraction of bonds lost via OH
 426 decreases, despite the actual number of oxidisable bonds lost to OH increasing. A
 427 larger fraction of bonds are therefore lost via the other mechanisms shown in Fig. 8
 428 rather than reaction with OH. As CH₄ removal occurs predominantly in the free
 429 troposphere, increasing the effective CH₄ emission also results in a reduction in the
 430 fraction of oxidisable bonds lost via deposition. The largest fractional increase in bond
 431 loss mechanism with increasing effective CH₄ emission is for photolysis, with the
 432 increase in the “other” fraction due to increased loss of bonds to the stratosphere with
 433 increasing CH₄.



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434
 435 **Figure 8. Oxidisable bond loss mechanism fractions under changing effective**
 436 **CH₄ emissions (0.5 x CH₄ concentration field, base simulation, and 2 x CH₄**
 437 **concentration field).**

438 **4.2.3 Impact of changing effective CH₄ emission on F_{radicals}, F_{RO2} and I**

439 The fraction of oxidisable bonds that goes on to produce radicals (F_{radicals}) and the
 440 fraction of these that are RO₂ (F_{RO2}) also decrease with increasing effective CH₄
 441 emissions. This is due to decreasing global OH concentration resulting from increased
 442 loss by reaction with CH₄ and a decreasing NO concentration. This favours bond loss

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460 via pathways that produce less RO₂ (e.g. CH₂O photolysis). The long lifetime of CH₄
461 compared with the majority of other sources of oxidisable bonds, also results in a
462 decrease in the fraction of bonds lost to deposition as total bond oxidation increases
463 fractionally in the free troposphere where deposition is a less significant loss
464 mechanism than in the boundary layer.

465 4.3 Changing isoprene emission

466 The species through which the oxidisable bonds are emitted has a significant impact
467 on O₃ production, due to their subsequent removal mechanisms. For example, in a
468 simulation where the only emission of oxidisable bonds is CO, F_{radicals} is 0.5 and F_{RO₂}
469 is 1 as the only CO sink is reaction with OH to produce one HO₂ (OH + CO → HO₂ +
470 CO₂). The CO coordinate bond, which in theory has the potential to produce 2
471 radicals, only produces 1 radical, which is an RO₂.

472 Isoprene has the most complex chemistry in the model and is the second largest
473 source of bonds into the atmosphere after CH₄ (Fig. 3). Figure 4c shows the response
474 of the two O₃ production diagnostics to varying the isoprene emission within the
475 model. The standard diagnostic (Fig.4c(i)) shows that the most significant increase in
476 P₃O₃ from increasing isoprene emissions is from NO + HO₂ and non-CH₃O₂ peroxy
477 radicals, with a smaller increase from CH₃O₂. The new P₃O₃ diagnostic (Fig.4c(ii))
478 again provides more insight, showing significant offsetting of around a half between
479 the terms.

480 4.3.1 Impact of changing isoprene emission on F_{NO}

481 The increased isoprene emission leads to a similar change in the magnitude of the
482 total number of oxidisable bonds emitted (E_{bonds}) as the simulations in which effective
483 CH₄ emission were varied. However, the countering decrease in all of the efficiency
484 parameters is much larger for isoprene than for CH₄. Figure 5c shows the fractional
485 change in the new P₃O₃ ozone production diagnostic parameters as a function of
486 isoprene emissions compared to the base simulation. The change in F_{NO} is due to both
487 a decrease in global mean NO_x concentrations with increasing isoprene and the spatial
488 distribution of isoprene emissions. The majority of global isoprene emissions are in
489 regions with low NO_x emissions, and thus low values of F_{NO}. Figure 9 shows a
490 decrease in global mean NO_x, and global mean OH_v concentrations with increasing
491 isoprene emissions, however, the effect is less than that seen when CH₄ is responsible

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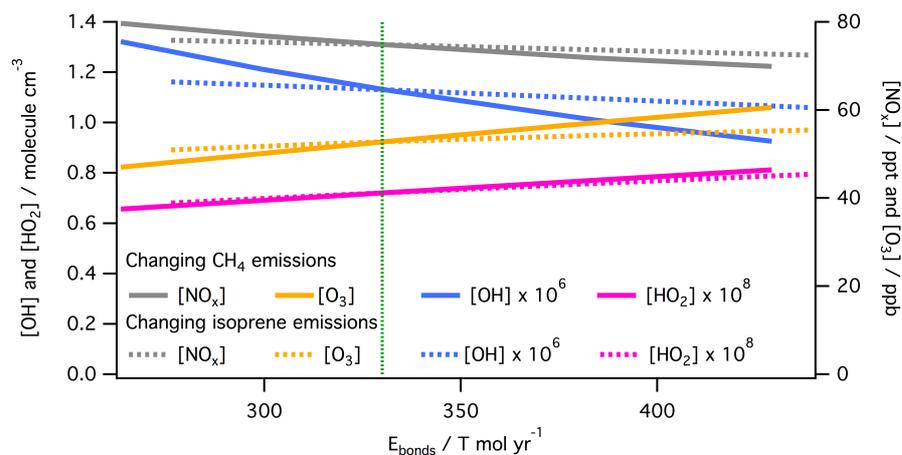
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525 for the same increase in oxidisable bond emission. This is due in a large part to the
 526 spatial scales over which the two compounds impact.



527

528 **Figure 9. The effect of oxidisable bond parent species on OH, HO₂, O₃ and NO_x**
 529 **concentrations. Global mean [OH], [HO₂], [O₃] and [NO_x] for simulations where**
 530 **the effective CH₄ emission (solid lines) and isoprene emission (dashed lines) were**
 531 **changed, against model E_{bonds}. The dashed vertical green line indicates E_{bonds} in**
 532 **the base simulation (330 T mol yr⁻¹).**

533 **4.2.2 Impact of changing isoprene emission on E_{bonds}**

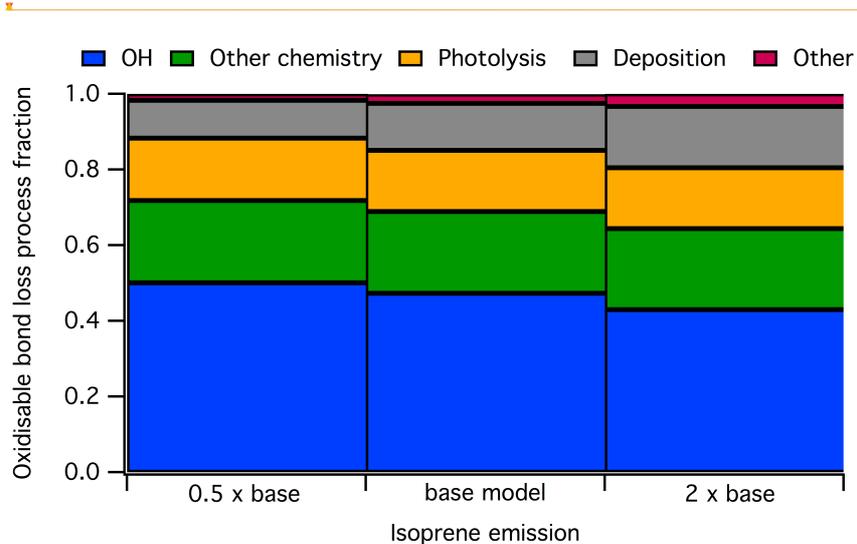
534 As isoprene is the second largest source of oxidisable bonds (Fig. 3), increasing the
 535 isoprene emission results in a significant increase in E_{bonds}. Differences in both the
 536 spatial distribution of emissions and the oxidation chemistry of isoprene and CH₄,
 537 however, means that the impact of the increases in E_{bonds} on O₃ production are
 538 significantly different for the two compounds. This is predominantly because the
 539 fraction of oxidisable bonds that are physically deposited for isoprene is high
 540 compared to those emitted as CH₄. This increase is due to i) the higher solubility of
 541 isoprene oxidation products compared to those of CH₄, and ii) the higher reactivity of
 542 isoprene means its oxidation occurs in the boundary layer where both dry and wet
 543 deposition is most effective.

544 | Figure 10 shows the fate of oxidisable bonds in the base simulation and those with the
 545 isoprene emissions multiplied by 0.5 and 2. The complex myriad of products formed
 546 during the isoprene oxidation mechanism also results in the production of many

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549 highly oxygenated multifunctional compounds with high Henry's law solubility
550 constants, meaning they are more readily lost to deposition.

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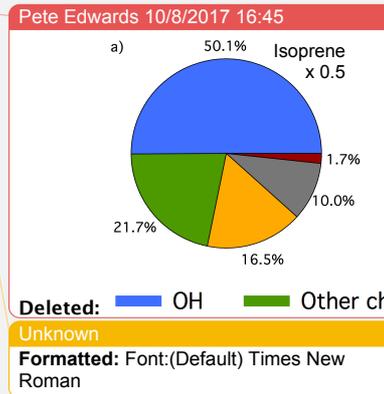


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553 **Figure 10. Oxidisable bond loss mechanism fractions, under changing isoprene**
554 **emissions.**

555 Increasing the isoprene emission also has a slight offsetting impact on the effective
556 CH₄ emission, as increased isoprene concentrations decrease OH concentrations, and
557 thus decrease the effective CH₄ emission. A doubling in isoprene emission causes a
558 6% reduction in the effective emission of CH₄.

559



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567 4.3.3 Impact of changing isoprene emission on F_{radicals} , F_{RO_2} and I

568 As shown in Fig. 3c(ii), increasing the isoprene emission results in a reduction in all
569 P_sO_3 efficiency parameters. The reductions in F_{radicals} is due to the higher fraction of
570 oxidisable bonds that are lost via non-radical forming pathways (e.g. deposition) for
571 isoprene relative to the other main oxidisable bond emission sources CH_4 and CO .
572 The slight decreases of F_{RO_2} and I with increasing isoprene emission are
573 predominantly due to changes in OH and NO_x (Fig. 9).

574 The complex chemistry of isoprene oxidation combined with the spatial distribution of
575 isoprene emissions means the increase in O_3 production due to increases in isoprene
576 emissions is roughly half what might be expected from the increase in oxidisable bond
577 emission alone (i.e. if the increase was *via* CO instead of isoprene).

578 5. Conclusions

579 We have shown that this bond-focussed approach to O_3 production provides a
580 significantly more detailed understanding of the processes involved. The role of
581 modelled VOC emissions and O_3 burden has been reported previously [Wild, 2007;
582 Young *et al.*, 2013]. However previous efforts extending this to a general process led
583 approach has not been successful. This new approach provides a tool with which the
584 processes controlling O_3 production can be investigated, and a metric by which
585 different emissions can be compared. For example, the differing chemistry of isoprene
586 and CH_4 shows that even though their emissions of carbon mass are comparable, the
587 atmosphere responds in different ways, with the isoprene bonds being less effective in
588 producing O_3 than CH_4 bonds. By quantifying multiple steps in the O_3 production
589 process, competing changes in the system become apparent (as shown in Fig. 4b(ii)
590 and c(ii)) and are thus testable. This enables the effect of model approximations on O_3
591 production to be quantified (e.g. the effect of NO_x on CH_4 emissions when using CH_4
592 concentration fields).

593 This new diagnostic also points towards the importance of observational datasets for
594 assessing our understanding of tropospheric chemistry. Although the budget presented
595 in Fig. 2 provides an annually integrated global estimate it points towards local
596 comparisons that can be made to assess model fidelity. Comparisons, both their
597 magnitude and their ratios, between observed and modelled bond concentration, bond
598 emission and loss fluxes (e.g. OH reactivity [Yang *et al.*, 2016] or depositional fluxes

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599 [\[Wesely and Hicks, 2000\]](#)), and O₃ production [[Cazorla and Brune, 2010](#)] would all
600 provide comparisons for outputs from the P_sO₃ diagnostic and help assess model
601 performance.

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602 Future work is necessary to identify the usefulness of this approach on smaller spatial
603 and temporal scales. For regional modelling scale, the transport flux of bonds into the
604 domain would need to be considered alongside the emissions of bonds. However, this
605 might help to disentangle O₃ production due to local VOC emissions from that due to
606 VOC emissions outside of the domain. This bond focussed approach may also have
607 usefulness on shorter timescales. For example, when considering vertical fluxes in and
608 out of the boundary layer, a bond centred approach could help. What fraction of the
609 bonds emitted at the surface are exported to the free troposphere. If a measurement of
610 reactivity flux could be made this could be tested experimentally.

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612 Another potentially important application is in model-model comparisons. Increases
613 in our understanding of why different models calculate different O₃ production and
614 burdens has been slow [[Stevenson et al., 2006](#); [Wu et al., 2007](#); [Young et al., 2013](#)].
615 Although a complete tagging like that described here is unlikely to occur for all of the
616 models involved in the comparison, a small number of additional diagnostics is likely
617 to produce a significantly better understanding of the models. Diagnosing (1) the total
618 bond flux (direct emissions plus the flux for those species kept constant), (2) the rate
619 of production of RO₂ and (3) the rate of production of O₃ could help differentiate why
620 certain models produce more or less O₃ than others. The ratios between these fluxes
621 would help identify what aspect of the emissions of chemistry differs between the
622 models.

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Moved up [1]: This new diagnostic also points towards the importance of observational datasets for assessing our understanding of tropospheric chemistry. Although the budget presented in Fig. 2 provides an annually integrated global estimate it points towards local comparisons that can be made to assess model fidelity. Comparisons, both their magnitude and their ratios, between observed and modelled bond concentration, bond emission and loss fluxes (e.g. OH reactivity [[Yang et al., 2016](#)] or depositional fluxes [[Wesely and Hicks, 2000](#)]), and O₃ production [[Cazorla and Brune, 2010](#)] would all provide comparisons for outputs from the P_sO₃ diagnostic and help assess model performance.

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Deleted: A comparison between models based on this methodology may well help identify at a process level why models differ in their O₃ production. The application of this diagnostic to regional O₃ production should also increase insight into the processes controlling model O₃.

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690 **Author contributions**

691 All work presented here was conceived by P.M.E. and M.J.E. The implementation,
692 model simulations and analysis were carried out by P.M.E., and the manuscript was
693 written by P.M.E. with substantial input from M.J.E..

694 **Additional information**

695 The authors declare no competing financial interests.

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