Sources of non-fossil fuel emissions in carbonaceous aerosols during early winter in Chinese cities

Di Liu¹, Jun Li¹*, Zhineng Cheng¹, Guangcai Zhong¹, Sanyuan Zhu¹, Ping Ding², Chengde Shen², Chongguo Tian³, Yingjun Chen⁴, Guorui Zhi⁵, Gan Zhang¹

¹State Key Laboratory of Organic Geochemistry, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Guangzhou, 510640, China
²State Key Laboratory of Isotope Geochemistry, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Guangzhou, 510640, China
³Key Laboratory of Coastal Environmental Processes and Ecological Remediation, Yantai Institute of Coastal Zone Research, Chinese Academy of Sciences, Yantai 264003, China
⁴State Key Laboratory of Pollution Control and Resources Reuse, Key Laboratory of Cities’ Mitigation and Adaptation to Climate Change, College of Environmental Science and Engineering, Tongji University, Shanghai 200092, China
⁵State Key Laboratory of Environmental Criteria and Risk Assessment, Chinese Research Academy of Environmental Sciences, Beijing 100012, China

*To whom correspondence may be addressed:

Dr. Jun Li; Email: junli@gig.ac.cn; Tel: +86-20-85291508; Fax: +86-20-85290706

Abstract

China experiences frequent and severe haze outbreaks from the beginning of winter. Carbonaceous aerosols are regarded as an essential factor in controlling the formation and evolution of haze episodes. To elucidate the carbon sources of air pollution, source apportionment was conducted using radiocarbon (¹⁴C) and unique molecular organic tracers. Daily 24-hour PM₂.₅ samples were collected continuously from October 2013 to November 2013 in 10 Chinese cities. The ¹⁴C results indicated that non-fossil fuel (NF) emissions were predominant in total carbon (TC; average = 65 ± 7%). Approximately half of the EC was derived primarily from biomass burning (BB) (average = 46 ± 11%), while over half of the OC fraction comprised NF (average = 68 ± 7%). On average, the largest contributor to TC was NF-derived secondary OC (SOCₙf), which accounted for 46 ± 7% of TC, followed by SOC derived from fossil fuels (FF) (SOCₙf; 16 ± 3%),
BB-derived primary OC (POC_{bb}; 13 ± 5%), POC derived from FF (POC_{f}; 12 ± 3%), EC derived from FF (EC_{f}; 7 ± 2%) and EC derived from BB (EC_{bb}; 6 ± 2%). The regional background carbonaceous aerosol composition was characterized by NF sources; POCs played a major role in northern China, while SOCs contributed more in other regions. However, during haze episodes, there were no dramatic changes in the carbon source or composition in the cities under study, but the contribution of POC from both FF and NF increased significantly.

1. Introduction

Recently, a wide range of fine particle (PM\(_{2.5}\)) pollution has affected northern, central and southern China, particularly on haze days, which has had significant effects on air quality, atmospheric visibility and public health, and caused extensive public and scientific concern (Liu et al., 2013b; Wang et al., 2014). Haze events in Chinese urban areas, especially in megacities, have become a common phenomenon, appearing in every season, because of large and intensive pollutant emissions and unfavorable meteorological conditions (He et al., 2014; Liu et al., 2013c).

Generally, heavy and serious haze pollution outbreaks start at the beginning of winter.

Carbonaceous aerosols are the important component of PM\(_{2.5}\) (~20–80%) (Rogge et al., 1993; He et al., 2004; Dan et al., 2004; Kanakidou et al., 2005) and are regarded as essential for controlling the formation and evolution of haze episodes. Relatively high concentrations of carbonaceous aerosols have been observed during typical haze days in northern, southern and central China (Zhao et al., 2013; Deng et al., 2008; Zhang et al., 2014c). Generally, carbonaceous aerosols (total carbon, TC) can be divided into elemental carbon (EC) and organic carbon (OC) according to their different physical and chemical properties (Krivácsy et al., 2001; Kleefeld et al., 2002). EC is formed either from biomass burning (BB; e.g., wood fires, heating) or fossil fuel
combustion (FF; e.g., vehicle or industry emissions such as coal combustion), and can be used as a tracer for primary combustion-generated OC because primary OC and EC are mostly emitted from the same sources (Turpin and Huntzicker, 1995; Strader et al., 1999). OC can be directly derived from primary emissions (primary OC; POC), or formed through oxidation of reactive organic gases followed by gas-to-particle conversion in the atmosphere (secondary OC; SOC) (Choi et al., 2012; Subramanian et al., 2007). Moreover, further subcategories of OC exist, including water-soluble organic carbon (WSOC) and water-insoluble organic carbon (WINSOC), which are distinguished on the basis of water-solubility; these may be essential for assessing the different sources of OC emissions during haze episodes, since WSOC is a proxy for SOC and BB OC, while a large fraction of WINSOC is from POC (Weber et al., 2007b; Docherty et al., 2008; Mayol-Bracero et al., 2002; Weber et al., 2007a; Huang et al., 2014).

Several methods have been introduced to identify and quantify OC emission sources, such as the use of organic molecular tracers (Simoneit et al., 1999), receptor models (PMF, CMB) (Singh et al., 2017; Bove et al., 2014; Marcuzzan et al., 2003), and dispersion models (Colvile et al., 2003); however, their reliability is limited by their low atmospheric lifetimes, in turn due to chemical reactivity and highly variable emission factors (Fine et al., 2001, 2002, 2004; Gao et al., 2003; Hedberg et al., 2006; Robinson et al., 2006). Recently, radiocarbon ($^{14}$C) analysis has been used as a powerful tool for facilitating the direct differentiation of non-fossil fuel (NF) carbon sources from fossil fuel (FF) sources, because $^{14}$C is completely absent from FF carbon (e.g., diesel and gasoline exhaust, coal combustion), whereas NF carbon (e.g., biomass burning, cooking and biogenic emissions) shows a high contemporary $^{14}$C level (Szidat et al., 2009). Hence, $^{14}$C measurements can provide information about the contributions of FF, BB and biogenic emissions
to carbonaceous aerosols. Numerous studies have been performed at urban sites and background sites to assess carbonaceous aerosol sources. For example, contemporary carbon was the dominant pollutant in carbonaceous aerosols at a background sites such as Ningbo and Hainan stations (Liu et al., 2013a; Zhang et al., 2014c). In urban, the relative carbon contributions have shown a significant seasonal difference (Yang et al., 2005; Chen et al., 2013; Liu et al., 2013b; Zhang et al., 2014a; Liu et al., 2014a; Zhang et al., 2017). A combination of $^{14}$C analysis and organic tracer determination allows for more detailed source apportionment of carbonaceous aerosols (Gelencsér et al., 2007; Ding et al., 2008; Lee et al., 2010; Yttri et al., 2011; Zong et al., 2016; Liu et al., 2015; Zhang et al., 2014b).

In this study, sampling was conducted in 10 typical Chinese cities during early winter when heavy haze pollution frequently occurs in this season. Carbonaceous aerosols, including different carbon fractions such as WSOC, WINSOC and EC, along with water-soluble inorganic ions and organic tracers (i.e. anhydrosugars) were analyzed in PM$_{2.5}$ samples. In particular, anhydrosugars such as levoglucosan are used as a molecular marker to indicate biomass-burning emissions. The combination of $^{14}$C analysis and the concentration of levoglucosan has offered new insights into the detailed sources of carbonaceous aerosols. So, source apportionment of carbonaceous aerosols was performed using a source apportionment model based on the $^{14}$C results and measured chemicals.

2. Materials and Methods

2.1 Aerosol sampling

Daily 24-hour PM$_{2.5}$ samples were collected continuously on the rooftops of institutes in 10 Chinese cities (Figure 1) from October 2013 to November 2013. In total, 292 aerosol samples,
including 10 field blanks, were collected on pre-heated (450°C for 5 h) quartz fiber filters (8 × 10 inches; Whatman, UK) using a high volume sampler with a flow rate of 0.3 m³ min⁻¹. The filters were then wrapped in aluminum foil, packed into air-tight plastic bags, and stored at -20°C in a refrigerator until analysis. PM₂.₅ mass concentrations were determined gravimetrically by state regulatory agencies. All samples were analyzed for OC and EC, and 20 samples, including two filters based on the PM₂.₅ concentrations at each site, were selected for further chemical analysis. Details of the sampling information and meteorological parameters used during sampling are shown in the Supporting Information (SI).

2.2 Chemical analysis

OC and EC were obtained with an off-line carbon analyzer (Sunset Laboratory, Inc., USA) using the thermo-optical transmittance method (NIOSH 870). Water-soluble inorganic ions (Na⁺, Cl⁻, Ca²⁺, Mg²⁺, K⁺, NH₄⁺, SO₄²⁻ and NO₃⁻) were analyzed with an ion chromatographer (83 Basic IC Plus, Metrohm, Switzerland). Anhydrosugars (levoglucosan, galactosan and mannosan) were analyzed by gas chromatography-mass spectroscopy (GC-MS) (7890-5975; Agilent) using a capillary column (DB-5MS: 30m, 0.25 mm, 0.25μm). Analysis methods related to OC and EC, water-soluble inorganic ions (Wang et al., 2012) and anhydrosugars (Liu et al., 2014a; Liu et al., 2014b) were presented elsewhere and a detailed analytical procedure and method are available in the SI.

2.3 Separation of carbon species

A punched section of filtrate was cut and sandwiched in a filtration unit, then extracted with 100 mL ultra-pure water (18.2 MΩ). WSOC species were quantified using a total organic carbon (TOC) analyzer (TOC-VCPH; Shimadzu, Japan). The punched filtrate was dried in a desiccator, wrapped
in aluminum foil and then stored in a refrigerator. WINSOC and EC were obtained from the water-filtered sample with an off-line carbon analyzer (Sunset Laboratory, Inc.) using the thermo-optical transmittance method (NIOSH 870).

2.4 Radiocarbon measurements

Isolation procedures for the $^{14}$C measurements of WSOC, WINSOC and EC have been described previously (Liu et al., 2016b; Liu et al., 2013b). Two filters, based on the PM$_{2.5}$ concentrations at each site, were used for $^{14}$C determination of WSOC, WINSOC and EC, to distinguish between FF and NF emissions. To obtain the WSOC, WINSOC and EC fractions from a single punch filter, a circular section of the punch filter was clamped in place between a filter support and a funnel and then 60 ml ultra-pure water was slowly passed through the punch filter without a pump, allowing the WSOC to be extracted delicately. WSOC was quantified as the total dissolved organic carbon in solution using a total organic carbon (TOC) analyzer (Shimadzu TOC_VCPH, Japan) following the nonpurgeable organic carbon protocol. WSOC solution was freeze-dried to dryness at -40 °C. The WSOC residue was re-dissolved with ~500 µl of ultra-pure water and then transferred to a pre-combusted quartz tube, which was then placed in the freeze dryer. After that, the quartz tube was combusted at 850 °C. The remaining carbon on the filter was identified as WINSOC or EC by an OC/EC analyzer (Sunset, U.S.). After WSOC pretreatment and freeze-dried, OC is oxidized to CO$_2$ under a stream of pre-cleaned oxygen pure analytical grade O$_2$ (99.999%, 30 ml min$^{-1}$) during the pre-combustion step at 340°C for 15 min. Before the OC is oxidized, the sample is first positioned in the 650 °C oven for about 45 s flash heating. This flash heating has the advantage of minimizing pre-combustion charring, since it reduces pyrolysis of OC. After the OC separation, the filters were removed from the system, placed into a muffle
furnace at 375°C, and combusted for 4 h. The filters were then quickly introduced back into the
system and oxidized under a stream of pure oxygen at 650°C for 10 min to analyze the EC
fraction. Finally, the corresponding evolved CO₂ (WSOC, WINSOC, and EC) was cryo-trapped,
quantified manometrically, sealed in a quartz tube and reduced to graphite at 600 °C using zinc
with an iron (200 mg, Alfa Aesar, 1.5-3 mm, 99.99%) catalyst for accelerator mass spectrometry
(AMS) target preparation. Approximately 200 µg of carbon was prepared for each carbon fraction.
All ¹⁴C values were reported as the fraction of modern carbon (fₘ) after correcting for
fractionation with δ¹³C. The degree of uncertainty in the ¹⁴C measurements was in the range of
0.2–0.6%. In this study, fₘ was converted to the fraction of contemporary carbon (fₖ), to eliminate
the effects of nuclear bomb tests through application of conversion factors of 1.10 ± 0.05 for EC
and 1.06 ± 0.05 for 2013 OC data. Here, the fₘ values of OC (OC = WSOC + WINSOC) and TC
(TC = WSOC + WINSOC + EC) were calculated by isotopic mass balance. The uncertainties of fₘ
(and fₖ) in WSOC, WINSOC, OC and EC were up to 20% ,20% ,15% and 15%, respectively. The
concentration in the field blank was negligible (0.37 ± 0.05 µg cm⁻²; less than 5% carbon) and no
field blank subtraction was made for ¹⁴C determination. The system blank F¹⁴C was
0.0036(SD=0.0001), which translated to a ¹⁴C age of around 45,000 years BP.

3. Results and Discussion

3.1 PM₂.₅, OC and EC concentrations and spatial distribution

PM₂.₅ levels ranged from 21.9 to 482 µg m⁻³, with an average level of 178 ± 103 µg m⁻³. A total of
98% and 81% of PM₂.₅ exceeded the First Grade National Standard (35 µg m⁻³) and Second Grade
National Standard (75 µg m⁻³) of China, respectively, indicating relatively poor air quality during
sampling days. The OC and EC levels ranged from 0.99 to 75.9 µg m⁻³ (average = 22.8 ± 15.3 µg
m^{-3}) and 0.07 to 19.3 μg m^{-3} (average = 3.66 ± 3.28 μg m^{-3}), respectively; thus, OC and EC were major components of PM$_{2.5}$, accounting for 13 ± 8% and 2 ± 1% of PM$_{2.5}$, respectively. The OC and EC levels in this study were generally higher than those recorded previously in more developed cities (e.g., New York, Los Angeles, Erfurt, Kosan) (Kam et al., 2012; Kim et al., 2000; Gnauk et al., 2005; Rattigan et al., 2010), indicating severe carbonaceous pollution and emphasizing the importance of restricting carbonaceous aerosols in China.

Northern China has high PM$_{2.5}$ concentrations. As shown in Table 1, the average PM$_{2.5}$ concentrations in Beijing (190 ± 79 μg m$^{-3}$), Xinxiang (245 ± 65 μg m$^{-3}$), Taiyuan (285 ± 84 μg m$^{-3}$) and Lanzhou (212 ± 112 μg m$^{-3}$) were significantly higher than those in central and southern China (from 85 μg m$^{-3}$ in Guangzhou to 123 μg m$^{-3}$ in Wuhan). Shanghai, in the eastern coastal region, had the lowest average PM$_{2.5}$ concentration (67 ± 43 μg m$^{-3}$). The ratio of total organic matter (TOM; 1.6 × OC + EC) to total fine particle mass ranged from 17.4% to 32.6%, except in Guiyang. Cities in central and southern China, such as Chengdu, Wuhan, Nanjing, and Guangzhou, had a higher ratio of TOM to PM$_{2.5}$ than other cities. Moreover, the OC/EC ratios in those cities were also higher, with values ranging between 8.1 and 12. The spatial distribution pattern closely reflected energy consumption and regional climate differences. In particular, Guiyang, which is a developing city located on the Western plateau, had a high level of PM$_{2.5}$ (227 ± 77 μg m$^{-3}$), comparable to that in northern China, but also had the lowest levels of OC and EC. Moreover, the TOM to PM$_{2.5}$ ratio was only about 6.0%. This indicates that there are different chemical sources in this developing city compared to megacities in China.

3.2 Radiocarbon results: fraction of modern carbon ($f_m$)

Table 2 shows the proportion (%) of NF sources in various carbon fractions. Overall, NF
emissions represented a more significant proportion of the TC (average = 65 ± 7%; range: 50–79%), at all sites, than FF sources, which underscores the importance of NF sources to carbonaceous aerosols during early winter in China.

EC is only formed by primary emissions, which are inert in ambient air and originate either from BB or FF combustion. In this study, about half of the EC was derived from BB in the 10 urban cities (average 46 ± 11%; range: 24–71%), which represents a slightly higher proportion than that for the same cities in winter and spring, but is similar to previous studies performed in cities in other countries (Szidat et al., 2009; Bernardoni et al., 2013; Liu et al., 2016a). However, this result differs from those obtained in remote regions dominated by BB (Barrett et al., 2015; Zhang et al., 2014a). Compared with other studies in China, the measured biomass burning contributions to EC in Beijing are relatively higher than those in the same city during winter (Zhang et al., 2014b; Zhang et al., 2015b). This is due to the fact that different approach we used for OC/EC separation, and sample selection in this study (selected two filter samples based on relatively lower and higher PM2.5 concentration for each site) because of limitations for $^{14}$C analysis (i.e. the bulk samples required and the high cost for $^{14}$C measurement). However, the result is similar with those using the same approach (Liu et al., 2016c; Zong et al., 2016). Since limitations for A larger contribution of BB to EC was found in central and western China (i.e., Beijing, Lanzhou, Chengdu and Guiyang) (49–63%), where Guiyang had the largest proportion of BB in EC (63 ± 12%), followed by Beijing (50 ± 2.0%), Chengdu (50 ± 1.8%), Wuhan (48 ± 10%) and Nanjing (47 ± 5%); this shows that there are large amounts of BB emissions (e.g., from biofuel burning and outdoor fires) in western and central China during early winter. This phenomenon was also found in central China during the severe haze episode that occurred over
China in January 2013, which suggests that these massive BB emissions were generated indoors (i.e., from domestic heating and cooking) and thus could not be detected by MODIS [Liu et al., 2016b]. Guangzhou had the lowest proportion of BB in EC (32 ± 12%), suggesting that FF emissions (coal combustion and vehicle emissions) dominated in the Pearl Delta region. Similar to Guangzhou, Taiyuan and Xinxiang had lower proportions of BB in EC, of 36 ± 11% and 37 ± 1.7%, respectively. High proportions of BB in EC are due to extremely high levels of BB tracers (levoglucosan). In this study, levoglucosan concentrations were in the range 161 to 672 ng m⁻³ (377 ± 153 ng m⁻³), and were significantly correlated with EC concentrations in BB (r = 0.708, p=0.000).

Over half of the OC fraction was from NF sources at all sites (range: 54–82%), with an average NF source contribution of 68 ± 7%, comparable to previous study reported in four Chinese cities during 2013 winter (Xi’an, Beijing, Shanghai and Guangzhou were 63%, 42%, 51% and 65%, respectively)(Zhang et al. 2015a). Generally, the $f_{nf}$ spatial distribution of OC is similar to that of EC, with NF sources contributing more in central China. Here, OC was divided into WSOC and WINSOC, which has been separated with respect to fossil and NF sources. A large contribution of NF sources to WINSOC (64 ± 7%) was observed in this study, comparable to previous studies performed in urban areas of Europe, e.g., Gothenburg (55 ± 8%) and Zurich (70 ± 7%) (Szidat et al., 2009;Zhang et al., 2013). Moreover, the $f_{nf}$ values for WSOC (70 ± 8%) were slightly higher than those for WINSOC, which showed values comparable to those observed in European and American cities (~70–85%) (Weber et al., 2007a;Szidat et al., 2009;Zhang et al., 2013). A higher $f_{nf}$ value indicated that, for WSOC, the contribution of NF emission sources was greater. WSOC is regarded as a mixture of SOC and BB-derived POC, whereas WINSOC is
mainly composed of POC from FF combustion, BB and biogenic sources. In this study, the ratio of WSOC to OC increased significantly with an increase in the proportion of NF sources in OC \((r = 0.531, p=0.016)\); this implies that POC from BB is more water-soluble, or that more NF-derived VOCs were involved in SOC formation.

### 3.3 Source apportionment of different carbon fractions

A source apportionment model for carbonaceous aerosols, including primary and secondary sources, was applied in this study using measured carbon fractions, anhydrosugars, and \(^{14}\text{C}\) isotopic signals. Detailed information on this model has been provided previously (Liu et al., 2014a; Liu et al., 2016a).

Briefly, EC from FF combustion \((\text{EC}_f)\) and BB-derived EC \((\text{EC}_{bb})\) can be estimated using the following respective equations:

\[
\text{EC}_f = \text{EC} \times (1-f_c) \tag{1}
\]

\[
\text{EC}_{bb} = \text{EC} \times f_c \tag{2}
\]

Similar to EC, OC can be divided into FF OC \((\text{OC}_f)\) and NF OC \((\text{OC}_{nf})\) based on \(^{14}\text{C}\) concentrations. \(\text{OC}_{nf}\) consists of BB-derived primary OC \((\text{POC}_{bb})\), NF-derived SOC \((\text{SOC}_{nf})\) and biological primary carbon (BPC), such as spore and plant debris. BPC particles exist mainly in coarse fractions (> 2.5 μm) and only account for ~1% of OC in PM\(_{2.5}\) [Guo et al., 2012]. Thus, this carbon fraction was ignored in the present study. \(\text{POC}_{bb}\) can be semi-quantitatively estimated from Lev concentrations, due to its unique characteristic of originating from BB, as follows:

\[
\text{POC}_{bb} = \text{Lev} \times (\text{OC}/\text{Lev})_{bb} \tag{3}
\]

According to the levoglucosan/mannosan \((\text{Lev}/\text{Man}; 17.4 \pm 5.9)\) and mannosan/galactosan \((\text{Man}/\text{Gal}; 2.1 \pm 0.3)\) ratios obtained in this study, \(7.76 \pm 1.47\) was adopted as the \((\text{OC}/\text{Lev})_{bb}\)
Thus, the SOC_{nf} fraction can be estimated through subtraction:

\[
\text{SOC}_{nf} = \text{OC}_{nf} - \text{POC}_{bb}
\]  

FF-derived POC and SOC can be estimated by the following respective equations:

\[
\text{POC}_{f} = \text{WINSOC} \times (1-f_c) \]  
\[
\text{SOC}_{f} = \text{WSOC} \times (1-f_c) \]

Figure 2 shows the proportions of different carbon fractions, including EC_{f}, EC_{bb}, POC_{bb}, POC_{f}, SOC_{nf} and SOC_{f}, in total carbon (TC) for the 10 urban cites during the sampling period. On average, the largest contributor to TC was SOC_{nf}, accounting for 46 ± 7% of TC, followed by SOC_{f} (16 ± 3%), POC_{bb} (13 ± 5%), POC_{f} (12 ± 3%), EC_{f} (7 ± 2%) and EC_{bb} (6 ± 2%). The proportion of primary sources (POC_{nf} + POC_{f} + EC_{ad} + EC_{f}) (average = 38 ± 9%; range: 25–56%) was lower than that of secondary sources (SOC_{nf} + SOC_{f}) (average = 62 ± 9%; range: 35–83%), which underlines the importance of SOC in carbonaceous pollution.

It should be noted that the model uncertainties in these contributions depended mainly on correction factors, such as the (POC/Lev)_{bb} emission ratios for wood burning, and on conversion factors used for determining the f_c in ^{14}C analysis. Typical relative uncertainties were recently estimated, using a similar modelling approach, at 20–25 % for SOC_{nf}, SOC_{f}, POC_{bb}, and POC_{f}, and ~13% for EC_{f} and EC_{bb} (Zhang et al., 2015a). A large fraction WINSOC can be from secondary organic aerosol as well. Hence POC_{f} is an upper limit of POC_{f}. SOC_{f} may be overestimated if a small fraction (e.g. <20%) WSOC is not secondary, so SOC_{f} may be an upper limit. Meanwhile, SOC_{nf} may also include other non-fossil sources such as cooking and biogenic emissions, however, they should be limited during wintertime (e.g., <20%). Therefore, our
estimates of SOC many generally represent an upper limit but this will not change our conclusion
towards to the spatial distribution of SOC in China.

POC and EC aerosols are independent from atmospheric gas reaction conditions and thus
directly reflect the characteristics of local emission sources. The total proportions of ECf and POCf
ranged from 10–38%, with an average of 19 ± 9% for all sites. The total proportions of ECf and
POCf in northern and southern China were greater than in western central and eastern coastal
China, indicating a higher impact of FF on local air pollution in both regions. The ratios of POCf
to ECf (0.66–3.32) were within the emission ratios between coal combustion (2.7–6.1) (Zhang et
al., 2008) and traffic exhausts fumes (0.5–1.3) (Zhou et al., 2014; He et al., 2008), indicating that
ccoal combustion and traffic exhaust fumes were the major primary sources at all sites. Beijing (2.6)
and Xinxiang (3.3) were mainly dominated by coal combustion emissions. The total proportions
of ECbb and POCbb ranged from 12–36%, with an average of 19 ± 8%. West central cities, such as
Lanzhou, Chengdu, Guiyang. Nanjing and Wuhan, had large proportions of ECbb and POCbb
(average = 23 ± 7%; range: 14–36%), which confirms the greater impact of BB on local air
pollution in West central China; this should be considered when setting future limits for polluting
corporations.

Total SOC in OC ranged from 42–84% (average = 72 ± 10%) among the sites tested in this study,
which is similar to recent studies, conducted in the haze period in China of January 2013, which
used high-resolution aerosol mass spectrometry; i.e., 41–59% [Sun et al., 2014] and 44–71%
[Huang et al., 2014] obtained from online and offline measurements, respectively. There was no
significant difference in the SOC/OC ratio among the different regions in China studied herein,
except for Guiyang, which had a somewhat lower SOC/OC ratio. Moreover, SOC was comprised
predominantly of NF sources at all sites (67–89%), except at Guiyang with values of 42-53%,
which are similar to areas in developed countries with good air quality, such as Puy de Dôme, France (86–88%) and Schauinsland, Germany (84–93%) [Gelencsér et al., 2007]. However, our values were higher than those of previous studies conducted in China during other winter and spring seasons, indicating the importance of NF to SOC in China during early winter.

3.4 Comparison of chemicals between samples by PM$_{2.5}$ concentration

Two samples, one each with a low and high PM$_{2.5}$ concentration, were obtained from all 10 study sites (Figure S1) for $^{14}$C and inorganic ions analysis, to investigate the composition of carbonaceous aerosols and evaluate the importance of FF and NF carbon in haze formation across China in early winter. During sampling, the air masses generally moved in a northwesterly to northeasterly direction to reach the site. The 5-day back trajectory analysis revealed relatively lower concentrations of PM$_{2.5}$ when the wind speed was higher, and relatively higher PM$_{2.5}$ levels when the wind speed was lower and more stable; synoptic conditions apparently promoted the accumulation of particles (Figure 3).

Theoretically, the aerosol composition at higher wind speeds should reflect regional background aerosol characteristics. Figure 3 shows the PM$_{2.5}$ chemical compositions of the stage for lower PM$_{2.5}$ concentration during sampling period. Here, due to the different conversion factors used to transform WINSOC to WINSOM (1.3), and WSOC to WSOM (2.1), OM calculations were based on the relative contributions of WSOC and WINSOC to OC. TOM is the sum of EC, WINSOM and WSOM. Generally, TOM contributions to PM$_{2.5}$ ranged from 21–38%, except in Guiyang where a value of 8% was observed. Moreover, OM was comprised mainly of NF emissions. In cities in northern China (Beijing, Xinxiang and Taiyuan), the contribution of WINSOM (both FF
and NF) was greater, indicating that POC played a major role in regional air quality during this season. Simultaneously, the lower NO$_3^−$/SO$_4^{2−}$ ratios also implied that POC from FFs might be derived predominantly from coal combustion. The 5-day back trajectory analysis showed that the air mass came from northern China, including regions such as Inner Mongolia and Hebei province, where the ambient temperature is always below 10°C during this season. It is very common for local rural residents to burn coal or biomass fuel to generate heat for their households. Therefore, coal and biomass fuel combustion in northern China might be the major contributor to regional carbonaceous aerosols in northern China during this season. In other cities, WSOM levels in both FF and NF were much higher than those in WINSOM, showing the importance of SOC across China. However, NO$_3^−$/SO$_4^{2−}$ ratios in Shanghai, Nanjing and Wuhan were much higher than in other areas. The back trajectory results showed that the air mass came from northern China or the Yangtze River Delta, implying that traffic exhaust emissions in those regions was more important for carbonaceous aerosol composition.

The chemical compositions of the higher PM$_{2.5}$ samples obtained in each city are shown in Figure 3. There were no dramatic changes in the carbon source or composition in any of the cities; however, the contribution of EC and WINSOM to both fossil and NF fuels increased significantly, along with the NO$_3^−$/SO$_4^{2−}$ ratios, indicating the importance of POC from local regions. The back trajectory results showed that wind speeds were moderate and stable, and that synoptic conditions apparently promoted the accumulation of particles derived either from local or regional sources.

### 4. Conclusion

PM$_{2.5}$ samples were collected continuously from 10 Chinese urban cities during early winter 2013. PM$_{2.5}$, OC and EC levels were highest in northern China, with maximum concentrations of
482 $\mu$g m$^{-3}$ (Taiyuan, n=31), 75.9 $\mu$g m$^{-3}$ (Taiyuan, n=31) and 19.3 $\mu$g m$^{-3}$ (Beijing, n=31), respectively. The $^{14}$C results, for the lower and higher PM$_{2.5}$ concentration sample pairs obtained in each city, indicated that, overall, NF emissions constituted a significant proportion of TC (average = 65 ± 7%) at all sites, i.e., higher than FF sources. Furthermore, about half of the EC was derived primarily from BB (average = 46 ± 11%), and over half of the OC fraction came from NF sources (average = 68 ± 7%). Source apportionment analysis was done using $^{14}$C and unique molecular organic tracers. On average, the largest contributor to TC was SOC$_{nf}$, accounting for 46 ± 7% of TC, followed by SOC$_{f}$ (16 ± 3%), POC$_{bb}$ (13 ± 5%), POC$_{f}$ (12 ± 3%), EC$_{f}$ (7 ± 2%) and EC$_{bb}$ (6 ± 2%). When relatively lower PM$_{2.5}$ concentrations were observed, OM was dominant in carbonaceous aerosols, mainly from NF. POC played a major role in regional air quality in the cities in northern China, while SOC contributed more in cities in other regions of China, such as Nanjing and Wuhan. During haze days, there were no dramatic changes in carbon sources or carbon compositions in the sampled cities, but the contributions of POC were relatively higher than the non-haze days.

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Reference


Fine, P. M., Cass, G. R., and Simoneit, B. R.: Chemical characterization of fine particle emissions from


Table 1 The PM$_{2.5}$, OC and EC data used in this study (average ± standard deviation; μg m$^{-3}$)

<table>
<thead>
<tr>
<th>Sites</th>
<th>N</th>
<th>PM$_{2.5}$</th>
<th>OC</th>
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Note: all fractions are in μg m$^{-3}$, except for levoglucosan (Lev), galactosan (Gal) and mannosan (Man) (all ng m$^{-3}$).
Figure 1. Geographic locations of the 10 Chinese sampling sites. The averages of monitored PM$_{2.5}$ concentrations (daily resolution, n = 31 for each site) during sampling campaign are shown in color plots.

Figure 2. The proportions of different carbon fractions, including elemental carbon derived from fossil fuels (EC$_f$), EC derived from burning biomass (EC$_{bb}$), BB-derived primary organic carbon (POC$_{bb}$), POC derived from FF (POC$_f$), non-FF secondary OC (SOC$_{nf}$) and SOC derived from FF (SOC$_f$) in total carbon (TC) for 10 urban cites during the sampling period.
Figure 3. The chemical compositions of fine particles (PM$_{2.5}$) under non-haze (top) and haze (bottom) conditions during the sampling period.
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