Consistent regional fluxes of CH$_4$ and CO$_2$ inferred from GOSAT proxy XCH$_4$:XCO$_2$ retrievals, 2010-2014

By Feng et al

We thank both reviewers for their support and comments that we believe have helped to improve the revised manuscript. Below are our responses to each individual reviewer comments (denoted in italics). Changes to the text are shown in red.

**Review report 1**

**General comment:**
The paper is of great importance. It provides a useful method to infer regional flux of CH$_4$ and CO$_2$ based on GOSAT XCH$_4$:XCO$_2$ combined with the ground-based observations. The result is convincing that the ratio-based posterior estimation reduced the uncertainties compared to flux only inferred from in-situ. The authors then discuss the regional differences. Although it is certainly not a surprise the largest differences found in the regions with sparse measurements, e.g. tropics and tropical south America, it is very interesting that a clear increase of CH$_4$ emission in the tropical land and the decrease in the south land is shown in the GOSAT version. Overall, this paper is well written. Its tropic and quality is suitable for publication in the journal. I only recommend that the authors to consider the minor comments and

**Specific Comments:**

1) Please comment on the how sensitive the inversion results are to the transport and chemical scheme of the model? I suppose that the results shown here is under the assumption that the transport of model brings fewer uncertainties compared to the uncertainties of input flux. But I think more cautions should be taken for this assumption.

The other reviewer also raised this issue.

We agree of course with the reviewers that model error, including atmospheric transport and model chemistry errors, does affect the resulting flux estimates inferred from top-down flux inversions (for example, Chevallier et al., 2010). We also acknowledge the challenge of quantifying the sources of such errors. In our study, we have assumed a simple (and transparent) description of model error and included it as part of the observation errors for the GOSAT XCH$_4$:XCO$_2$ ratio and for the in-situ CH$_4$ and CO$_2$ observations. To address the reviewers' comment we have included in the revised manuscript an explicit acknowledgement of the challenge associated with quantifying model error.

Lines 213-216:

A robust description of model error remains a major challenge for this and similar studies. We have assumed a simple formulation to describe model error, which will not fully account for impacts of errors from, for example, model atmospheric transport on resulting CO$_2$ and CH$_4$ flux estimates.

2) Please comment on how much the results could change when a different region definition is used, e.g. the comparison of the results based on TransCom regions and higher-resolved region definition in this study. I guess the result could be sensitive to how much ground sites are included in each region. So a coarse definition of the regions perhaps favors the in-situ results more than ratio results, is that true?

We split the TransCom regions and distinguish between sources of CO$_2$ and CH$_4$ to reduce the impact of aggregation error associated with inversions described at coarse spatial resolutions. In the revised manuscript we explicitly make that point (line 181).
We describe the inversion on these smaller geographic regions to help reduce aggregation errors associated with fluxes being estimated on a coarse spatial resolution (Patra et al., 2005).

We agree that increasing the granularity of the inversion grid could affect the relative importance of the in-situ and satellite observations for constraining flux estimates. Investigating this further is outside the scope of this study. However, we have added an acknowledgement of this reviewer point in the summary (line 413)

‘...the sensitivity of our results to model error and to the temporal and spatial resolution of fluxes requires further investigation,’

As a preliminary assessment to answer the reviewer’s question (to satisfy our curiosity but not suitable for publication) we ran another calculation. We artificially enlarged the prescribed errors for in-situ CO2 and CH4 observation by 30% to reduce their relative importance. The results show that for almost every TransCom land regions, the changes in the resulting CO2 and CH4 flux emissions are small. For example, over tropical South American annual net CH4 emissions for 2010 changed by < 10%, where coverage by in-situ data is very sparse.

3) The authors explained and showed in Figure 1 how they defined the regions for the Basis Function. But for the discussion of the results, North lands, Tropical lands and South lands are used. Although they referred the definition to Chevallier et al. 2014, it will be better to show the regions on the map in Figure 1, for example, overlaid on the basis function regions.

Good suggestion. We added another panel to Figure 1 to show the definition of northern, tropical and southern land regions.

4) Line 56 and 400, the name of the region with ‘temperate’ is wrongly typed as ‘temperature’.

Typos corrected.

5) Line 278, the region should be ‘Eurasian temperate’ for more accuracy. Please comment on the large decrease of CH4 in Eurasian temperate region in the GOSAT inversion. From my point of view, two inversions go to opposite direction only in this region for CH4. And there are only 4 sites for CH4 and 3 sites for CO2 (if I count right in Figure 1). How does the GOSAT XCH4:XCO2 data show (large spatial or temporal variation)? Do you see large annual variability or increase tendency in the GOSAT retrieval or GOSAT ratio-based inversion?

The GOSAT XCH4:XCO2 ratios show large spatial variations over Eurasian temperate (see the Figure below). The available in-situ observations are not sensitive to emissions from southeast China, which have strong CH4 sources from wetlands and agriculture. As a result, assimilating in-situ data only will not correct for any prior overestimation of these CH4 sources. In contrast, the GOSAT ratio data provide useful constraints of emissions for this geographical region. To clarify this point in the revised manuscript, we add (Line 285-287):

, which is due to the in situ network having little sensitivity to emissions over a large part of Eurasian temperate, in particular over southeast China where there are large CH4 sources from wetlands and rice paddies.

We find that the GOSAT XCH4:XCO2 ratios show a small but consistent downward trend, which is partially due to the increasing CO2 concentrations. The CH4 emission estimates show no consistent trends from 2010 to 2014. However, due to very limited independent observations over the region, we are unable to evaluate the results.
Figure: Annual mean GOSAT XCH₄:XC₂O₂ ratios for 2013, which are gridded into the 4x5 model boxes.

6) In the caption of Figure 2, color for in-situ experiments should be blue but not green.

Typo corrected.

7) In the first line caption of Figure 7, ‘TAB’ instead of ‘TBA’. The same for line 360 and 368 on page 10.

Typos corrected.

8) On page 9, line 329-330, please add a comma after ‘Because the in situ flux far away’.

Typo corrected.
Review report 2

130 General comments

The paper provides another valuable development in the direction of using XCH₄/XCO₂ ratio in the inversion of both the CO₂ and CH₄ fluxes. This paper is one of the first few papers on the subject, another one was published by Pandey et al., (2016). In this paper a longer analysis period is used, allowing for more extensive validation. Ability of the XCH₄/XCO₂ ratio to constrain fluxes of CO₂ and CH₄ and improve match with independent observations constitute most appealing and encouraging result of this study. The paper is well written and deserves publication with only minor corrections.

Specific comments

140 Larger tropical CH₄ fluxes are inferred with GOSAT ratio as compared to surface data inversion. How to prove that the result is robust with respect to biases in retrieval and even retrieval prior concentration profiles? Another possible suspect could be the transport model bias in the stratosphere for either CO₂ or CH₄ or both. Can authors add more discussion on this issue?

145 Good question. Other GOSAT CH₄ inversions (e.g. Alexe et al., 2015) have also reported elevated CH₄ emissions over tropical regions. While our posterior fluxes, inferred from the ratio inversion, are in better agreement with independent observations over the tropics than our prior fluxes (Figures A2 to A4), we are unable to exclude the impact of observation biases, and particularly model errors. This is partially due to scarcity of independent data over this region.

150 In our response to reviewer 1 (above) we have discussed the challenge surrounding robust quantification of model errors that effectively limits our (and everyone else’s) ability to assess their impact on flux estimates. We have chosen to explicitly acknowledge this issue to ensure the reader is aware of it (Line 213-216):

A robust description of model error remains a major challenge for this and similar studies. We have assumed a simple formulation to describe model error, which will not fully account for impacts of errors from, for example, model atmospheric transport on resulting CO₂ and CH₄ flux estimates.

160 L178 The benefit of dividing Transcom regions into 4 relatively equal ones was extensively explored by Patra et al (2005).

165 Not including this paper was an egregious oversight on our part. In the revision, we add (Line 181):

We describe the inversion on these smaller geographic regions to help reduce aggregation errors associated with fluxes being estimated on a coarse spatial resolution (Patra et al., 2005).

170 See a further discussion of this point in our response to second point from the first reviewer.

Suggestions for technical corrections

L058 Better tell which fluxes are being discussed, suggest to change “fluxes” to “Amazonian fluxes”, the context is ambiguous here.

Done.

L094 Suggest correcting “sufficient” to “sufficiently”

Done.
L100 Houweling et al 2015 is referred to, but not found in references.

Reference added.

L121 Suggest correcting Pandy to Pandey

Typo fixed.

L149 When introducing “prior covariance” need to tell which covariance - fluxes or concentrations?

Changed to ‘prior flux error covariance’.

L396 Text “GOSAT data significantly changed the a priori spatial distribution” should be modified towards saying that posterior changes significantly with respect to prior.

We have changed this to ‘GOSAT data results in significant changes with respect to the a priori spatial distribution’

L462 Wording “XCH4 in : : : lower stratosphere” doesn’t sound right.

Typo fixed.

References

Reference added
Consistent regional fluxes of CH$_4$ and CO$_2$ inferred from GOSAT proxy XCH$_4$-XCO$_2$ retrievals, 2010-2014

Liang Feng$^1$, Paul I. Palmer$^1$, Hartmut Bösch$^2$, Robert J. Parker$^2$, Alex J. Webb$^2$, Caio S. C. Correia$^3$, Nicholas M. Deutscher$^{4,5}$, Lucas G. Domingues$^3$, Dietrich G. Feist$^6$, Luciana V. Gatti$^3$, Emanuel Gloor$^7$, Frank Hase$^8$, Rigel Kivi$^9$, Yi Liu$^{10}$, John B. Miller$^{11,12}$, Isamu Morino$^{13}$, Ralf Sussmann$^{14}$, Kimberly Strong$^{15}$, Osamu Uchino$^{13}$, Jing Wang$^{10}$, Andreas Zahn$^{16}$

1. National Centre for Earth Observation, School of GeoSciences, University of Edinburgh, UK.
2. National Centre for Earth Observation, Department of Physics and Astronomy, University of Leicester, UK.
3 Instituto de Pesquisas Energéticas e Nucleares (IPEN) - Comissao Nacional de Energia Nuclear (CNEN) - Atmospheric Chemistry Laboratory, Cidade Universitaria, Sao Paulo, Brazil.
4. Institute of Environmental Physics, University of Bremen, Germany.
5. Centre for Atmospheric Chemistry, University of Wollongong, Australia.
6. Max Planck Institute for Biogeochemistry, Jena, Germany.
7. School of Geography, University of Leeds, Leeds, UK.
8. Karlsruhe Institute of Technology (KIT), IMK-ASF, 76021 Karlsruhe, Germany.
9. FMI-Arctic Research Center, Sodankylä, Finland.
10. Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029, China.
11. Global Monitoring Division, Earth System Research Laboratory, National Oceanic and Atmospheric Administration, Boulder, Colorado, USA.
12. Cooperative Institute for Research in Environmental Sciences (CIRES), University of Colorado, Boulder, Colorado, USA.
14. Karlsruhe Institute of Technology (KIT), Institute of Meteorology and Climate Research - Atmospheric Environmental Research (IMK-IFU), 82467 Garmisch-Partenkirchen, Germany.
15. Department of Physics, University of Toronto, Toronto, M5S 1A7, Canada.
16. Karlsruhe Institute of Technology (KIT), Institute of Meteorology and Climate Research (IMK), 76344 Eggenstein-leopoldshafen, Germany.
ABSTRACT

We use the GEOS-Chem global 3-D model of atmospheric chemistry and transport and an ensemble Kalman filter to simultaneously infer regional fluxes of methane (CH₄) and carbon dioxide (CO₂) directly from GOSAT retrievals of XCH₄:XCO₂, using sparse ground-based CH₄ and CO₂ mole fraction data to anchor the ratio. This work builds on previously reported theory that takes advantage that: 1) these ratios are less prone to systematic error than either the full physics data products or the proxy CH₄ data products; and 2) the resulting CH₄ and CO₂ fluxes are self-consistent. We show that a posteriori fluxes inferred from the GOSAT data generally outperform the fluxes inferred only from in situ data, as expected. GOSAT CH₄ and CO₂ fluxes are consistent with global growth rates for CO₂ and CH₄ reported by NOAA, and with a range of independent data including in particular new profile measurements (0-7 km) over the Amazon basin that were collected specifically to help validate GOSAT over this geographical region. We find that large-scale multi-year annual a posteriori CO₂ fluxes inferred from GOSAT data are similar to those inferred from the in situ surface data but with smaller uncertainties, particularly over the tropics. GOSAT data are consistent with smaller peak-to-peak seasonal amplitudes of CO₂ than either a priori or the in situ inversion, particularly over the tropics and the southern extra-tropics. Over the northern extra-tropics, GOSAT data show larger uptake than the a priori but less than the in situ inversion, resulting in small net emissions over the year. We also find evidence that the carbon balance of tropical South America was perturbed following the droughts of 2010 and 2012 with net annual fluxes not returning to an approximate annual balance until 2013. In contrast, GOSAT data significantly changed the a priori spatial distribution of CH₄ emission with a 40% increase over tropical South America and tropical Asia and smaller decrease over Eurasia and temperate South America. We find no evidence from GOSAT that tropical South American CH₄ fluxes were dramatically affected by the two large-scale Amazon droughts. However, we find that GOSAT data are consistent with double seasonal peaks in Amazonian fluxes that are reproduced over the five years we studied: a small peak in January to April and a larger peak in June to October, which is likely due to superimposed emissions from different geographical regions.
1. Introduction

Atmospheric growth of the two most abundant non-condensable greenhouse gases (GHGs), carbon dioxide (CO$_2$) and methane (CH$_4$), increases the absorption of Earth’s outgoing infrared radiation (IR) with implications for the radiation budget of Earth’s atmosphere and subsequent manifold changes in climate including an increase in global mean temperatures. The most recent international climate agreement aims to limit the rise in global mean temperature to 2 degrees Celsius, which will be attempted by reducing the emissions of human-driven (anthropogenic) GHGs. This approach necessarily assumes we have good knowledge of emissions from all anthropogenic sectors so that targeted reductions are effective. It also implicitly assumes that the Earth’s biosphere will continue to be a net annual sink for up to 40-60% of anthropogenic CO$_2$ (e.g., Barlow et al., 2015), and the continued stability of natural reservoirs of CH$_4$. Current scientific knowledge, informed by mostly ground-based data and models, does not confidently support either assumption even on a continental scale. Here, we present the first multi-year record of self-consistent regional net fluxes (sources minus sinks) of CO$_2$ and CH$_4$ inferred from the Japanese Greenhouse gases Observing SATellite (GOSAT). We show these fluxes are significantly different to those inferred from ground-based data, particularly over tropical ecosystems, but are generally consistent with independent data throughout the troposphere.

Inferring CO$_2$ and CH$_4$ fluxes directly from atmospheric observations is an ill-posed inverse problem, with a wide range of scenarios that fit these data. Prior information is used to regularize the problem, with care taken to describe data and prior uncertainties to avoid over- or under-fitting the data. There is a growing and progressive literature on estimating GHG fluxes in which an atmospheric chemistry transport model is used to relate observed atmospheric GHG mole fractions to atmospheric surface exchange fluxes. A number of approaches are used to minimize the model minus observation residual to infer spatial and temporal variations in flux. Errors introduced by the incomplete and uneven coverage of current ground-based observation networks are compounded by atmospheric model errors (e.g., transport and chemistry) resulting in significant discrepancies between flux estimates inferred from different models on spatial scales < O(10,000 km) (e.g. Law et al., 2003; Yuen et al., 2005; Stephens et al., 2007; Peylin et al., 2013).

Space-borne observations of short-wave IR (SWIR) that are sufficiently precise to detect small changes in lower tropospheric CO$_2$ and CH$_4$ necessary for flux inference are beginning to improve current understanding of these GHGs. GOSAT (Kuze et al., 2016), launched in 2009, was the first satellite designed purposefully to measure CO$_2$ and CH$_4$ columns using SWIR wavelengths. There is a growing body of literature that has inferred regional CO$_2$ and CH$_4$ fluxes from GOSAT dry-air CO$_2$ (XCO$_2$) and CH$_4$ (XCH$_4$) column mole fractions using the proxy and full-physics data products (Basu et al., 2013; Deng et al., 2013; Houweling et al., 2015; Bergamaschi, et al., 2013; Takagi et al., 2014; Fraser et al., 2014). The resulting flux estimates (particularly for CO$_2$) are often found to be inconsistent with the results based on the surface network, and with each other using different
atmospheric transport models or using different versions of retrievals (Chevallier et al., 2014; Houweling et al., 2015). The reliability of the fluxes inferred from GOSAT XCO₂ retrievals (Reuter et al., 2014; Feng et al., 2016), considering bias in current retrievals (Feng et al., 2016) as well as the variations in temporal and spatial coverage (Liu et al., 2015), is still a subject of ongoing discussions.

We build on previous work that developed a novel approach to estimate simultaneously regional CO₂ and CH₄ flux estimates from the GOSAT XCH₄:XCO₂ ratio measurements, which had been until then used exclusively to develop ‘proxy’ XCH₄ retrievals (Fraser et al., 2014). Previous work has shown that these ratios are less prone to systematic bias that represents a substantial challenge to the full-physics data products. The underlying assumption of the proxy approach is that by taking the ratio of the two retrieved values that have been fitted simultaneously in nearby spectral windows (1.65 µm and 1.61 µm) any interference due to cloud and aerosol scattering will be similar for both retrieved values and will be removed (Frankenberg et al., 2005; 2006). The ratio is then scaled by a model XCO₂ value, under the assumption that atmospheric gradients of XCO₂ are much smaller than XCH₄, to generate XCH₄ proxy retrievals. Data products generated by the proxy approach are more robust against scattering than the full-physics approach so that there are more usable retrievals over geographical regions that are compromised by seasonal aerosol and cloud distributions, e.g. tropical South America. Fraser et al. (2014) used a series of numerical experiments and the Maximum A posteriori (MAP) approach to show that these XCH₄:XCO₂ ratios could be used, in conjunction with in situ observations of CH₄ and CO₂ mole fractions, to simultaneously estimate regional CO₂ and CH₄ fluxes. Pandey et al. (2016) used a similar approach but using a 4-D variational assimilation approach to infer XCO₂ and XCH₄ fluxes for 20 months from April 2009. They found that after correcting biases in the XCH₄:XCO₂ retrievals, the ratio inversion results in similar agreement with independent CO₂ and CH₄ observations, as other inversions based on the in-situ data only or based on individual GOSAT XCH₄ and XCO₂ products. Here, we use an Ensemble Kalman Filter (EnKF) to assimilate the XCH₄:XCO₂ ratio data (UoLV6, Parker et al., 2015) from January 2009 to December 2014, inclusive. A comparison between the UoLV6 data set and the ground-based XCH₄ and XCO₂ data from the Total Carbon Column Observing Network (TCCON) shows a bias of about 0.3%. We use individual in situ and GOSAT observations (instead of monthly means, Fraser et al., 2014) to estimate monthly fluxes at a higher spatial resolution than Fraser et al., 2014.

In the next section we describe the Ensemble Kalman Filter approach, the observations we use to infer the CO₂ and CH₄ fluxes and those we use to evaluate the resulting posteriori flux estimates, and a description of the numerical experiments. In section 3 we describe our result, with a particular focus on tropical South America where we compare our a posteriori model with new aircraft measurements. We conclude the paper in the section 4.

2. Methods and Data

2.1 Ensemble Kalman Filter

We develop an existing EnKF framework that has been used to estimate CO₂ (Feng et al., 2009; 2013; 2016), and CH₄ fluxes from the in-situ or space-based measurements of their atmospheric observations (Fraser et al., 2013). In this study, the state vectors are regional fluxes of CO₂ and CH₄ at location \( x \) and time \( t \) as:
where $g$ denotes CO$_2$ or CH$_4$ tracer gas and $f^g_{p}(x, t)$ describes the a priori estimates of CO$_2$ or CH$_4$ fluxes. Following Fraser et al. (2014), our basis function set $BF_i^g(x, t)$ is defined as the pulse-like (monthly) CO$_2$ or CH$_4$ fluxes from different sectors over pre-defined geographic regions. The coefficients $c_i^g$ for both the CO$_2$ and CH$_4$ fluxes form a joint state vector $c$ to be estimated by optimally fitting the model to the data.

In the Ensemble Kalman Filter framework, the prior flux error covariance $P$ is represented by an ensemble of perturbations of the coefficients $\Delta C$: $P = \Delta C \Delta C^T$, where $T$ represents the matrix transpose. The a posteriori coefficient estimates are given by:

$$c_a = c_f + K (y_{obs} - H(c_f)), \quad (2)$$

where $c_a$, $c_f$ are the prior and posterior estimates, respectively; $y_{obs}$ are the observations; and $H$ is the observation operator that relates surface fluxes (i.e., the coefficients) to the observation data (described below), and includes the atmospheric transport model (Fraser et al., 2014).

The Kalman gain matrix $K$ in Eq. 2 is approximated by (Feng et al., 2009):

$$K \approx \Delta C \Delta Y^T [\Delta Y \Delta Y^T + R]^{-1}, \quad (3)$$

where $R$ is the observation error covariance, and $\Delta Y^T = H(\Delta C)$ projects the flux perturbation (coefficients) ensemble $\Delta C$ to observation space. We use the GEOS-Chem global 3-D chemistry transport model (v9.02) to relate the fluxes to the observation space. For the experiments reported here we run the CTM model at a horizontal resolution of $4^0$ (latitude) $\times 5^0$ (longitude), driven by the GEOS-5 (GEOS-FP for 2013 and 2014) meteorological analyses from the Global Modeling and Assimilation Office Global Circulation Model based at NASA Goddard Space Flight Centre. We use monthly 3-D fields of the hydroxyl radical from the GEOS-Chem HOx-NOx-Ox chemistry simulation to describe the main oxidation sink of CH$_4$ (Fraser et al., 2014). We use a four-month moving lag window to reduce the computational costs related to the projection of the perturbation ensemble into the observation space for longer time periods (Feng et al., 2013; 2016).

Where possible we use consistent emission inventories for CO$_2$ and CH$_4$: monthly biomass burning emission (GFEDv4.0, van der Werf et al., 2010) and monthly fossil fuel emissions (ODIAC, Oda and Maksyutov, 2011). To describe atmospheric CO$_2$ variations, we also use monthly-resolved climatological ocean fluxes (Takahashi et al., 2009), and three-hourly terrestrial biosphere fluxes (CASA, Olsen and Randerson, 2004). To describe atmospheric CH$_4$ variations, following Fraser et al. (2014), we use prescribed annual inventories for emissions from oil and gas production, coal mining, ruminant animals (Olivier et al., 2005), termites, and hydrates (Fung et al., 1991). We use monthly-resolved emissions for rice paddies and wetlands for 2009, 2010 and 2011 (Bloom et al., 2012). From January 2012, we fix the rice paddy and wetland emissions to their monthly means between 2009 and 2011. We also include a simple soil sink of CH$_4$ (Fraser et al., 2014).

We define the pulse-like basis functions (Eq. 1) guided by the TransCom-3 regions (Gurney et al., 2002), with each continental region further divided equally into 4 sub-regions. Figure 1 shows the 44 land regions, and 11 ocean regions that we use in this study; in comparison Fraser et al. (2014) used...
11 land regions and one ocean region. We describe the inversion on these smaller geographic regions to help reduce aggregation errors associated with fluxes being estimated on a coarse spatial resolution (Patra et al., 2005). We distinguish CO₂ fluxes between four categories: 1) ocean fluxes; 2) anthropogenic emissions; 3) biomass burning; and 4) terrestrial biospheric fluxes. For CH₄ fluxes we distinguish between six categories: 1) ocean fluxes; 2) anthropogenic emissions from coal mining; 3) anthropogenic emissions from oil and gas production, fossil fuel combustion, and other; 4) biomass burning; 5) natural fluxes from wetlands, and rice paddies, and; 6) natural fluxes from termites, hydrates and others. In total, we have 143 monthly basis functions for CO₂, and 231 monthly basis functions for CH₄.

We assume an a priori uncertainty of 60% for the coefficients corresponding to the natural CO₂ and CH₄ fluxes, and for CH₄ emissions from coalmines. We assume an a priori uncertainty of 40% for CO₂ anthropogenic emissions, CO₂ and CH₄ ocean fluxes, and anthropogenic emission of CH₄ from the oil and gas industry. We also assume that a priori errors for the same categories are correlated with a spatial correlation length of 800 km, and with a temporal correlation of one month (Feng et al., 2016). We assume that fire emissions of CO₂ and CH₄ are correlated with a correlation coefficient of 0.5, accounting for the variation and uncertainty of the fire emission factors (Parker et al., 2016).

2.2 Observations
We assimilated GOSAT XCH₄:XCO₂ retrievals and in situ surface observations of CO₂ and CH₄ mole fraction. We use version 6 of the proxy GOSAT XCH₄:XCO₂ retrievals from the University of Leicester, UK, including both the nadir observations over lands, and glint observations over oceans. Previous analyses have shown that these retrievals have a bias of 0.3%, with a single sounding precision of about 0.72% (Parker et al., 2015; 2011). In our experiments we globally remove this 0.3% bias from the GOSAT proxy data. We assume that each single GOSAT proxy XCH₄:XCO₂ ratio retrieval has an uncertainty of 1.2% to account for possible model errors, including the errors in atmospheric chemistry and transport.

We also assimilate CO₂ and CH₄ mole fraction observations at surface-based sites, which help anchor the GOSAT ratio observations (Fraser et al, 2014). Figure 1 shows the sites we use, from the NOAA observation network (Dlugokencky et al., 2015). We assume uncertainties of 0.5 ppm and 8 ppb for the in situ observations of CO₂ and CH₄, respectively. We also assume a model error of 1.5 ppm and 15 ppb for CO₂ and CH₄, respectively. We adopt a larger percentage for the CH₄ model error to account for difficulties in modelling chemical sinks of CH₄ in atmosphere (Fraser et al., 2013). A robust description of model error remains a major challenge for this and similar studies. We have assumed a simple formulation to describe model error, which will not fully account for impacts of errors from, for example, model atmospheric transport on resulting CO₂ and CH₄ flux estimates.

To determine the importance of the ratio data we run twin sets of experiments: 1) “ratio” experiment that include the GOSAT data and the in situ data sets, and 2) “in situ” experiment that use only the in situ surface data.

2.3 Independent data to evaluate a posteriori estimates
We use independent observations of atmospheric CO₂ and CH₄ mole fraction to evaluate the atmosphere mole fractions that correspond to the a posteriori fluxes from our inversions. These observations include data collected by TCCON and by four aircraft campaigns. To improve the
readability of the main text, we have placed much of the text and many of the Figures associated with the evaluation of the a posteriori fluxes in Appendix A.

TCCON is a global network of ground-based FTS instruments that measure, among other compounds, the total atmospheric columns of CO$_2$ and CH$_4$ (Wunch et al., 2011). We use the bias-corrected TCCON XCO$_2$ and XCH$_4$ data at all available sites from the recent GGG2014 release of the TCCON dataset (Wunch et al., 2015). For a comprehensive description of the network and the available data from each TCCON site, we refer the reader to the TCCON project page (e.g., Feist et al., 2014; Deutscher et al., 2014; Notholt et al., 2014; Griffith et al., 2014; Iraci et al., 2014; Strong et al., 2014; Dubey et al., 2014; and Sussmann et al., 2014).

We also use aircraft measurements from four projects to evaluate our a posteriori model concentrations: 1) Data collected during experiments one through five from the HIAPER Pole-to-Pole Observations (HIPPO) that provide latitude-altitude cross-sections of tropospheric mole fractions of CO$_2$ and CH$_4$ (and other tracers) covering dates during 2009 to 2011 (Wofsy et al. 2012); 2) Data collected by commercial airliners as part of the Civil Aircraft for the Regular Investigation of the atmosphere Based on an Instrument Container (CARIBIC) experiment, which are mainly at cruise altitudes but also during ascent/descent over airports (Brenninkmeijer et al., 2007; Schuck et al., 2009); 3) Bi-weekly aircraft measurements (surface to 4 km) collected from 2010 to 2012 at four sites over Brazil by IPEN (Instituto de Pesquisas Energeticas e Nucleares) over the Amazon rainforest (AMAZONICA, Gatti et al. (2014)): Rio Branco (RBA), Tabatinga (TAB), Alta Floresta (ALF) and Santaré\'m (SAN); and 4) aircraft measurements conducted by IPEN for the FAPESP/NERC-funded Amazonian Carbon Observatory (ACO, Webb et al., 2016) close to two of the AMAZONICA sites from 2012 to 2014: Salinópolis (SAH) and Rio Branco (RBH). These two sites were chosen to best represent air before and after travelling across the Amazon Basin. The purpose of these flights was to improve validation of GOSAT XCH$_4$ and XCO$_2$ data over the Amazon basin so we flew from the surface to 7 km to capture more of the atmospheric column that GOSAT observes. A detailed description of ACO can be found in Webb et al. (2016) and comparison of these data against GOSAT XCH$_4$:XCO$_2$ data are shown below.

3 Results

3.1 CO$_2$ Fluxes

Figure 2 shows that the in situ only and the ratio inversions result in similar annual net CO$_2$ flux estimates (averaged for 2010 to 2014) over temperate land regions. But compared to the in situ only inversion, the ratio inversion shows a larger net emission over tropical South America, and a smaller net emission from tropical Asia, although the differences are usually within the 1-$\sigma$ uncertainties.

We also find that the a posteriori fluxes for the ratio inversion generally have smaller uncertainties, in particular over tropical lands.

Figure 3 and Table 1 compare the time series of the prior and posterior global net CO$_2$ flux estimates. They show that global annual a posteriori net flux estimates are 40-60% smaller the a priori estimates (Table 1) due to a smaller net emission during boreal winter and a larger net uptake during the boreal summer (Figure 3). The corresponding global annual CO$_2$ growth rate agrees with NOAA
estimates, inferred from in situ observations, typically within 0.15 ppm/a, except for 2013 when the inversions are 0.3 ppm/a lower than the NOAA-reported value.

Figure 3 also shows that the monthly a posteriori flux estimates by the in situ and ratio inversions are similar over the northern landmasses (Figure 1), with the exception of the summer in 2014 when the ratio inversion shows significantly smaller uptake. Over the tropical landmasses, a posteriori fluxes from the ratio inversion show a much smaller seasonal cycle, with exception of boreal summer months in 2014 when these fluxes have larger uptake. In general, uncertainties for the monthly fluxes inferred by the ratio inversion (GOSAT + in situ data) are smaller (up to 30%) than using only the in situ data. This reflects the poor spatial coverage of the current in situ observing network particularly over tropical ecosystems (Figure 1). Over the southern landmasses the a posteriori fluxes for the two inversions are similar and typically within their uncertainties. We find that both inversions show a gradual reduction in the peak-to-trough amplitude, which appears to support a similar downward trend in the a priori estimates from about 9.0 GtC/a between 2010 and 2011 to about 7.5 GtC/a between 2013 and 2014. A posteriori fluxes also consistently show lower net emissions than a priori values during austral winter months.

### 3.2 CH₄ Fluxes

Figure 2 shows that a priori and the a posteriori global annual net CH₄ flux estimates are similar (520 Mt/a for the a priori versus 518 Mt/a for the ratio inversion), but their geographical distributions are significantly different. The ratio and in situ only inversions show much larger emissions than the a priori estimates over tropical lands: by up to 50% larger for tropical South America and for tropical Asia (Figure 2). This increase is partially offset by reduced emissions at mid-latitudes (e.g., temperate South America). Over Eurasian temperate we find that the ratio inversion has 15% smaller emissions than the a priori estimates, but the fluxes inferred from the in situ surface data for the same region are 25% higher than the a priori (Figure 2), which is due to the in situ network having little sensitivity to emissions over a large part of Eurasian temperate, in particular over southeast China where there are large CH₄ sources from wetlands and rice paddies. Figure 2 also shows that the ratio inversion has much smaller (up to 60%) uncertainties than the in situ inversions over almost all TransCom land regions, which is due to better spatial observation coverage of GOSAT proxy data.

Figure 4 shows that at the global scale, the monthly a posteriori fluxes inferred from the ratio and in situ inversions have larger seasonal variations than the a priori: a typical seasonal minimum of about 450 Mt/a and a typical maximum of 680 Mt/a, compared to the a priori that has a minimum of 480 Mt/a and a maximum of 620 Mt/a. The larger a posteriori seasonal variation is largely due to the seasonal cycle over northern landmasses that is driven by varying wetland and fire CH₄ emissions. The ratio inversions also show a muted peak emission of typically 30 Mt/a during January to February, partially due to peak emissions over southern landmasses during the austral summers. Over northern hemisphere landmasses, the in situ inversion is systematically 5-10% higher than the ratio inversion from 2010 to 2014. Over the tropics, we find that a posteriori tropical fluxes from the ratio and in situ inversions are generally larger than the priori estimates. Also the ratio a posteriori fluxes are systematically higher than those inferred from the in situ surface data, and show a small upward annual trend (Table 2). Over this region, we also find that the ratio inversion consistently shows a double-peak structure with a small peak between January and April and a larger peak between June and October (Figure 4). This is not shown by the in situ inversion or by the a priori inventory. A posteriori fluxes for the southern landmasses are generally lower by 30-50 Mt/a than...
the a priori values, which, together with northern landmasses, partially offsets the increase in tropical CH$_4$ emissions (Figure 4). Over southern hemisphere land masses, the seasonal cycles of the ratio and in situ inversions are similar, although the ratio inversion generally has lower seasonal minima, with the exception of 2014 when the phase of the ratio inversion is opposite to the in situ inversion.

### 3.3 Model Evaluation

In general the ratio inversion shows the best agreement with independent CH$_4$ observations, particularly over lower latitudes. A posteriori improvements to the CO$_2$ simulation are relatively small. We find that both the model CO$_2$ and CH$_4$ concentrations reproduce the large-scale spatial (e.g., the North-South gradient) and temporal (seasonal cycle) variations in the HIPPO and CARIBIC data (section 2.3). The a posteriori simulations reproduce the observed TCCON XCH$_4$ and XCO$_2$ variations. Over most TCCON sites, the a posteriori XCO$_2$ model biases are within 0.8 ppm (<0.2%), and the standard deviations are smaller than 1.6 ppm. The typical model biases for model XCH$_4$ data are smaller than 10 ppb (i.e., <0.6%), with a standard deviation smaller than 15 ppb. For more details, we refer the reader to Appendix A where we show pictorially the comparisons between observations and the ratio and in situ a posteriori CO$_2$ and CH$_4$ mole fractions.

Here, we focus on tropical South America (Figure 1) for three reasons. First, in situ surface data are particularly sparse over this geographically region, including two sites (Figure 5) over which we use the observed CO$_2$ and CH$_4$ mole fractions to constrain flux estimates: Arembepe, Bahia, Brazil (ABP, -12.770° latitude, -38.170° longitude) and Ragged Point, Barbados (RBP, 13.165° latitude, -59.432° longitude). Second, they include vulnerable ecosystems that have recently experienced several widespread drought conditions in 2010 and 2012 (see for example, Lewis et al., 2015; Rodrigues and McPhaden, 2014), which have affected their ability of absorbing carbon (Doughty et al., 2015) and increased fire emissions (Gatti et al., 2014; Alden et al., 2016). And third, we report new aircraft profile measurements from the ACO (Webb et al., 2016) that was designed specifically to evaluate GOSAT column observations of CH$_4$ and CO$_2$ (section 3).

Figure 6 shows that the a posteriori monthly CH$_4$ and CO$_2$ flux estimates over tropical South America from the ratio inversion are significantly different from the in situ inversion, as expected given the in situ surface data coverage. However, monthly a posteriori CO$_2$ fluxes from the ratio inversion are not always statistically different from the a priori, reflecting the large a priori uncertainties associated with fluxes over this region. The in situ inversion typically has larger uptake during the dry season (May to September) and smaller emissions during the wet seasons than the ratio inversion. Because the in situ flux estimates over this geographical region rely on observation far away, they are particularly sensitive to a priori uncertainties, as expected. We find that assuming a global a priori uncertainty that is 50% smaller than our control run results in an additional net emission of 0.4 GtC/a over tropical South America in 2010. Including the GOSAT ratio data into that sensitivity inversion leads to a smaller net decrease (of 0.13 GtC/a) in emissions.

Table 3 shows that the a posteriori annual fluxes inferred by the ratio inversion are significantly larger than the in situ inversion in 2010, 2011, and 2012, by about 0.7, 0.4, and 0.5 GtC, respectively. A posteriori fluxes from the ratio inversion shows net emissions are smaller in 2013 and 2014 than in 2010 or 2012, which is due to larger uptake in the dry season and to smaller emissions in the wet seasons (Figure 6). This result reveals the continental-scale impact of the severe droughts in 2010.
and 2012 over Tropical Southern America. Our result for 2010 is consistent with recent studies based on regional-scale AMAZONICA aircraft observations (Gatti et al., 2014; van der Laan-Luijkx et al., 2015; Alden et al., 2016). The in situ inversion fails to reproduce this increase in net emissions during 2010 dry season, instead showing a large uptake (Figure 6).

A posteriori CH₄ fluxes from the ratio inversion are systematically higher than the in situ inversion (Figure 6). This discrepancy is particularly large during October 2013 to March 2014 when the in situ inversion is lower than typical seasonal values observed during previous years. Figure 5 shows that XCH₄:XCO₂ ratio measurements over southwest Amazon increase from 4.55 ppb/ppm to about 4.65 ppb/ppm between October-December 2013 and January-March 2014. This is a small but significant change in the ratio that suggests either enhanced CH₄ emissions and/or lower CO₂ fluxes. The two closest in situ sites to the locus of XCH₄:XCO₂ variability (RPB and ABP) do not reproduce this change. Consequently, the in situ inversion may not accurately describe these CH₄ flux changes over the continental interior.

Figures 7 and 8 show that a posteriori fluxes from the ratio inversion generally decrease the mean model difference against independent AMAZONICA and ACO aircraft observations of CO₂ and CH₄ over the Amazon basin, but with only small improvements to the associated standard deviations. At some sites the fluxes from the ratio inversion significantly mute the rapid variations in atmospheric CO₂ and CH₄ inferred from the in situ data. Figure 7 shows that for CO₂ the greatest improvement is for the central basin sites of RBA+RBH (after 2012), where the bias reduced from -0.62 ppm to 0.01 ppm with accompanying reduction in standard deviation from 3.7 ppm to 2.6 ppm. We find similar but smaller reductions at another AMAZONICA site (TAB). Over other AMAZONICA and ACO sides, the impact of GOSAT XCH₄:XCO₂ ratios are even smaller. The coarse resolution of our model that allows us to exploit efficiently the GOSAT and in situ data is one possible explanation for the large standard deviations (van der Laan-Luijkx et al., 2015; Gatti et al., 2014). Figure 8 shows that overall the ratio inversion better reproduces the AMAZONICA and ACO CH₄ data than the in-situ inversion. The ratio inversion does best at SAN. It also shows a better agreement over RBA as it does for CO₂. After 2012, the ratio inversion shows a positive bias at the two ACO sites SAH and RBH. Assimilating the XCH₄:XCO₂ data reduces the standard deviations (by about 4 ppb to 11 ppb) over ALF, TAB and RBA (RBH after 2012), and slightly (by about 1 ppb) increase the standard deviations at SAN and SAH.

4. Summary

Building on previously reported theory, we simultaneously inferred regional CO₂ and CH₄ fluxes from the proxy GOSAT XCH₄:XCO₂ retrievals 2010-2014, inclusive, anchored by geographically sparse in situ mole fraction data. The main advantage of using these data directly is that the ratio is less compromised by systematic bias on spatial scales greater than typical model grid resolution (<1000 km) and less than large-scale variations captured by ground-based network (<10,000 km), which represents a limiting factor to using full-physics XCO₂ measurements. Inferring CO₂ and CH₄ fluxes together provides a self-consistent methodology.
We showed that a posteriori fluxes inferred from the GOSAT data generally outperformed the fluxes inferred only from in situ data, as expected given their greater measurement coverage. GOSAT CH₄ and CO₂ fluxes are consistent with global growth rates for CO₂ and CH₄ reported by NOAA, and are generally more consistent than the results based on in situ surface data with a range of independent data collected throughout the global troposphere (e.g., aircraft profiles and ground-based total column measurements), and including in particular new profile measurements (0-7 km) over the Amazon basin that were collected specifically to help validate GOSAT over this geographical region.

We found that large-scale multi-year annual a posteriori CO₂ fluxes inferred from GOSAT data are similar to those inferred from the in situ surface data but with smaller uncertainties, particularly over the tropics where in situ surface data are sparse. However, we found that GOSAT data are consistent with smaller peak-to-peak seasonal amplitudes of CO₂ than either a priori or the in situ inversion, particularly over tropical and the southern extra-tropics, where the annual means are similar. Over the northern extra-tropics, GOSAT data infer a larger uptake than supported by the a priori but a smaller uptake than the corresponding in situ data. Using the individual annual means and seasonal variations over 2010-2014, we found evidence from GOSAT that the carbon balance of tropical South America was perturbed following the droughts of 2010 and 2012 when this region was a large annual source of CO₂ (0.5-0.6 PgC/a) to the atmosphere, with net annual fluxes not returning to an approximate annual balance until 2013.

We showed that GOSAT data results in significant changes with respect to a priori spatial distribution of CH₄ emission with a 40% increase over tropical South America and tropical Asia and smaller (partially compensating) decrease over Eurasia and temperate South America. We find no evidence from GOSAT that tropical South American CH₄ fluxes were dramatically affected by the two large-scale Amazon droughts in 2010 and 2012. However, we reported that GOSAT data are consistent with double seasonal peaks in fluxes that are reproduced over the five years we studied: a small peak in January to April and a larger peak in June to October. Currently, we have no explanation for this phenomena but it is likely due to superimposed emissions from different geographical regions.

While the sensitivity of our results to model error and to the temporal and spatial resolution of fluxes requires further investigation, our analysis, in the wider context of other studies, supports the adoption of using space-borne observations of CO₂ and CH₄ to better understand the carbon cycle on the continental scale. Well-known weaknesses of these data (e.g., biases in spatial and temporal coverage) can be partially overcome by integrating them with information from other networks and by judicious use of atmospheric chemistry transport models. The next obvious step is to understand how we can improve source attribution of CO₂ and CH₄ without necessarily resorting to the assumption, as used here and elsewhere, that a priori fossil fuel emission estimates are correct.

Source attribution can be sometimes achieved by exploiting knowledge of spatial distributions of different sources, but techniques that allow more rigorous exploitation of multi-gas correlations must be developed and incorporated into data assimilation systems that will eventually form the backbone to operational systems (e.g., EU Copernicus Atmospheric Monitoring Service to atmospheric CO₂).

Author contributions
L. Feng and P. I. Palmer designed the experiments and wrote the paper; H. Bösch, R. J. Parker, and Alex Webb provided the GOSAT XCO₂ and XCH₄ data; N. M. Deutscher, D. G. Feist, R. Kivi, I. Morino, O. Uchino, F. Hase, R. Sussmann, and K. Strong provided access to TCCON XCO₂ and XCH₄ data; A. Zahn provided access to CARIBIC CO₂ and CH₄ mole fraction data. L. V. Gatti, E. Gloor, C. S. C. Correia, L. G. Domingues, and J. B. Miller provided access to aircraft data (AMAZONICA and ACO) over the Amazon basin. J. Wang and Y. Liu provided a preliminary evaluation of CO₂ and CH₄ fluxes over China. All co-authors provided comments and suggestions on the manuscript.

**Acknowledgements**

Work at the University of Edinburgh was partly funded by the NERC National Centre for Earth Observation (NCEO), and the European Space Agency Climate Change Initiative (ESA-CCI). P. I. Palmer gratefully acknowledges funding from the NCEO and his Royal Society Wolfson Research Merit Award. NCEO and the European Space Agency Climate Change Initiative funded work at the University of Leicester. HB and RJP are supported by the ESA Climate Change Initiative (ESA-CCI). RJP was also funded by an ESA Living Planet Fellowship. We thank NERC and FAPESP for their joint funding of the Amazonian Carbon Observatory Project (NERC Reference: NE/J016284/1). M. Gloor was financially supported by the NERC consortium grant AMAZONICA (NE/F005806/1) which we also thank for providing access to additional aircraft profiles. The TCCON Network is supported by NASA’s Carbon Cycle Science Program through a grant to the California Institute of Technology. The TCCON stations from Bialystok, Orleans and Bremen are supported by the EU projects InGOS, GAIA-CLIM, and ICOS-INWIRE, and by the Senate of Bremen. Nicholas Deutscher is supported by an Australian Research Council - Discovery Early Career Researcher Award: DE140100178. TCCON measurements at Eureka were made by the Canadian Network for Detection of Atmospheric Composition Change (CANDAC) with additional support from the Canadian Space Agency. TCCON Operation at Tsukuba and Rikubetsu sites is supported in part by the GOSAT project. Works by J. Wang and Y. Liu are funded by Helmholtz-CAS Joint Research Groups (HCJRG-307). We also thank the HIPPO team for their observations (http://hippo.ucar.edu/) used in our model evaluation. We thank G. J. Collatz and S. R. Kawa for providing NASA Carbon Monitoring System Land Surface Carbon Flux Products: http://nacp-files.nacarbon.org/nacp-kawa-01/.

The authors would like to thank the two anonymous reviewers for their insightful comments, which help us improve the manuscript significantly.
Appendix A: Wider Geographical Model Evaluation

We use independent observations to evaluate the a posteriori model concentrations that correspond to the flux estimates, acknowledging limitations associated with sparse observation coverage and atmospheric transport model errors (Chevallier et al., 2014). We sample the GEOS-Chem atmospheric chemistry transport at the time and location of each individual observation.

HIPPO
Figures A1 and A2 show that the ratio inversion is marginally more consistent with HIPPO XCO₂ data than the in situ inversion but the spatial error structure are qualitatively similar. The ratio inversion has a positive bias of 0.2 pm and standard deviation of 1.3 ppm compared to the in situ inversion that has a positive bias of 0.3 ppm and standard deviation of 1.3 ppm. The largest standard deviations (up to 0.8%) reflect the ability of models to reproduce small-scale variations, particularly at the lowest (the planet boundary layer) and the highest (the upper troposphere and lower stratosphere) altitudes. We find small differences (generally within 1 ppm) below 4-6 km between 40⁰S and 40⁰N, and much larger differences (up to 2 ppm) in the upper troposphere and in the lower stratosphere north of 45⁰N.

The ratio and in situ inversions show similar spatial structure to HIPPO XCH₄ data. We find a small negative bias (0-15 ppb) in the middle and lower troposphere between 40⁰S and 40⁰N and a larger positive bias (by over 20 ppb) in the extratropical upper troposphere/lower stratosphere. We find the largest discrepancies between model and observed XCH₄ in the higher latitude lower stratosphere, in agreement with previous studies (e.g., Alexe et al., 2015 and Pandey et al., 2016) that is mainly due to difficulties in modelling stratospheric chemical processes. As a result, the ratio inversion and the in situ inversions have similar biases of 0.6 ppb and 0.1 ppb, respectively, as well as similar standard deviations of 27.7 ppb versus 27.5 ppb respectively.

Figure A2 shows that the two a posteriori models reproduce the hemispheric CO₂ gradient, typical for boreal spring months, observed by HIPPO-3 experiment. Compared to the in situ inversion, the ratio inversion has a larger negative bias (-0.8 ppm versus -0.4 ppm) around 20⁰N, in contrast to a slightly larger positive bias over most of the southern hemisphere. We find that the overall model bias and associated standard deviation of the gridded partial CO₂ columns are very small (biases <0.01 ppm and standard deviation < 0.6 ppm). Figure A2 shows that the two a posteriori models also reproduce the hemispheric CH₄ gradient observed by the HIPPO-3 experiment. Compared to the in situ inversion the proxy GOSAT XCH₄:XCO₂ data significantly reduces the negative bias of the CH₄ concentrations (by up to 10 ppb) over the tropical regions. The overall bias for the gridded CH₄ partial columns is reduced from -5.6 ppb for the in-situ inversion to the -1.5 ppb for the ratio inversion.

CARIBIC
Figure A3 shows that the two a posteriori models reproduce the observed annual trend of CO₂ monthly means and the observed seasonal cycle with smaller amplitude. Underestimation of the seasonal cycle of the upper tropospheric CO₂ concentrations is well documented, and believed to be caused by a deficiency in modelling vertical transport (Stephens et al., 2007). Figure A3 also shows that the a posteriori models reproduce the observed trend and seasonal variation of atmospheric
CH$_4$ in the tropical middle/upper troposphere. The ratio inversion has a smaller bias (-0.37 ppb) than the in situ inversion (-8.27 ppb) but has only modestly improved the associated standard deviation by 15% from 7.55 ppb to 6.48 ppb.

**TCCON**

Figure A4 shows that the two a posteriori models have a similar level of agreement with 15 independent TCCON XCO$_2$ retrievals. For most of these sites, the model XCO$_2$ bias is well within 1.0 ppm, and the standard deviation is between 0.6 and 1.5 ppm. The two exceptions are sites around Los Angeles, USA: cj (34.1°N, 118.1°W) and jf (34.2°N, 118.2°W), where the models underestimate atmospheric XCO$_2$ by 1.5-2.0 ppm, which we attribute to our coarse model resolution. Figure A4 also shows that assimilating GOSAT XCH$_4$:XCO$_2$ proxy data significantly reduces the model XCH$_4$ bias by up to 10 ppb over low-latitude TCCON sites. The GOSAT data also helps to reduce the standard deviations over most of the 15 sites.
Tables

Table 1: A priori and a posterior estimates of the annual net CO₂ fluxes for 2010 to 2014 for the global and three contributing regions: 1) north landmasses; 2) tropical landmasses; and 3) south landmasses. One-sigma uncertainties are given in the brackets.

<table>
<thead>
<tr>
<th>Region</th>
<th>Estimate</th>
<th>2010 GtC/a</th>
<th>2011 GtC/a</th>
<th>2012 GtC/a</th>
<th>2013 GtC/a</th>
<th>2014 GtC/a</th>
</tr>
</thead>
<tbody>
<tr>
<td>Global</td>
<td>Prior</td>
<td>8.64(1.64)</td>
<td>7.52(1.76)</td>
<td>8.72(1.57)</td>
<td>7.97(1.63)</td>
<td>8.10(1.64)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>4.83(0.37)</td>
<td>3.54(0.35)</td>
<td>5.10(0.34)</td>
<td>4.61(0.34)</td>
<td>4.14(0.36)</td>
</tr>
<tr>
<td></td>
<td>ratio</td>
<td>4.87(0.25)</td>
<td>3.43(0.25)</td>
<td>5.08(0.24)</td>
<td>4.66(0.24)</td>
<td>4.15(0.26)</td>
</tr>
<tr>
<td>North lands</td>
<td>Prior</td>
<td>6.63(1.47)</td>
<td>6.81(1.60)</td>
<td>7.52(1.44)</td>
<td>7.51(1.48)</td>
<td>7.2(1.53)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>4.60(0.15)</td>
<td>4.47(0.14)</td>
<td>5.07(0.15)</td>
<td>4.89(0.14)</td>
<td>4.90(0.15)</td>
</tr>
<tr>
<td></td>
<td>Ratio</td>
<td>4.68(0.11)</td>
<td>4.81(0.11)</td>
<td>5.38(0.11)</td>
<td>5.05(0.11)</td>
<td>5.30(0.11)</td>
</tr>
<tr>
<td>Tropical lands</td>
<td>Prior</td>
<td>2.57(0.44)</td>
<td>1.55(0.46)</td>
<td>1.95(0.38)</td>
<td>1.53(0.44)</td>
<td>1.76(0.43)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>1.31(0.28)</td>
<td>0.70(0.29)</td>
<td>1.08(0.26)</td>
<td>1.22(0.27)</td>
<td>1.04(0.27)</td>
</tr>
<tr>
<td></td>
<td>Ratio</td>
<td>1.63(0.18)</td>
<td>0.59(0.18)</td>
<td>1.00(0.17)</td>
<td>1.21(0.18)</td>
<td>1.03(0.19)</td>
</tr>
<tr>
<td>South Lands</td>
<td>Prior</td>
<td>0.84(0.57)</td>
<td>0.56(0.57)</td>
<td>0.64(0.49)</td>
<td>0.32(0.56)</td>
<td>0.53(0.45)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>0.03(0.25)</td>
<td>-0.50(0.25)</td>
<td>0.15(0.22)</td>
<td>-0.27(0.23)</td>
<td>-0.38(0.24)</td>
</tr>
<tr>
<td></td>
<td>Ratio</td>
<td>0.09(0.15)</td>
<td>-0.56(0.16)</td>
<td>0.06(0.15)</td>
<td>-0.31(0.16)</td>
<td>-0.52(0.16)</td>
</tr>
</tbody>
</table>

Table 2: The same as Table 1 but for CH₄ fluxes.

<table>
<thead>
<tr>
<th>Region</th>
<th>Estimate</th>
<th>2010 Mt/a</th>
<th>2011 Mt/a</th>
<th>2012 Mt/a</th>
<th>2013 Mt/a</th>
<th>2014 Mt/a</th>
</tr>
</thead>
<tbody>
<tr>
<td>Global</td>
<td>Prior</td>
<td>519.3 (59.9)</td>
<td>517.1 (58.5)</td>
<td>521.1 (58.7)</td>
<td>521.1 (58.7)</td>
<td>521.1 (58.7)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>250.3 (36.4)</td>
<td>253.4 (36.6)</td>
<td>256.2 (36.9)</td>
<td>256.2 (36.9)</td>
<td>256.2 (36.9)</td>
</tr>
<tr>
<td></td>
<td>Ratio</td>
<td>521.2 (6.2)</td>
<td>508.1 (6.5)</td>
<td>508.4 (6.3)</td>
<td>508.4 (6.3)</td>
<td>508.4 (6.3)</td>
</tr>
<tr>
<td>North lands</td>
<td>Prior</td>
<td>262.6 (14.4)</td>
<td>273.2 (16.5)</td>
<td>270.9 (16.4)</td>
<td>269.8 (15.8)</td>
<td>277.0 (14.5)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>230.4 (4.4)</td>
<td>219.2 (4.5)</td>
<td>227.7 (4.5)</td>
<td>226.8 (4.3)</td>
<td>227.8 (4.7)</td>
</tr>
<tr>
<td></td>
<td>Ratio</td>
<td>250.3 (36.4)</td>
<td>253.4 (36.6)</td>
<td>256.2 (36.9)</td>
<td>256.2 (36.9)</td>
<td>256.2 (36.9)</td>
</tr>
<tr>
<td>Tropical lands</td>
<td>Prior</td>
<td>156.4 (15.7)</td>
<td>146.2 (15.3)</td>
<td>147.2 (15.6)</td>
<td>142.4 (15.7)</td>
<td>147.8 (15.2)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>198.0 (5.8)</td>
<td>203.3 (5.8)</td>
<td>200.1 (5.7)</td>
<td>207.1 (5.2)</td>
<td>207.3 (5.9)</td>
</tr>
<tr>
<td></td>
<td>Ratio</td>
<td>201.0 (5.7)</td>
<td>197.1 (5.8)</td>
<td>193.9 (5.8)</td>
<td>199.9 (5.8)</td>
<td>201.1 (5.7)</td>
</tr>
<tr>
<td>South lands</td>
<td>Prior</td>
<td>84.3 (11.6)</td>
<td>70.1 (11.8)</td>
<td>74.5 (10.8)</td>
<td>75.8 (10.8)</td>
<td>83.0 (11.8)</td>
</tr>
<tr>
<td></td>
<td>in-situ</td>
<td>115.4 (26.7)</td>
<td>114.1 (26.2)</td>
<td>114.3 (26.1)</td>
<td>114.3 (26.1)</td>
<td>114.3 (26.1)</td>
</tr>
<tr>
<td></td>
<td>Ratio</td>
<td>115.4 (26.7)</td>
<td>114.1 (26.2)</td>
<td>114.3 (26.1)</td>
<td>114.3 (26.1)</td>
<td>114.3 (26.1)</td>
</tr>
</tbody>
</table>

Table 3: Same as Table 1 but for CH₄ and CO₂ fluxes over tropical South America.

<table>
<thead>
<tr>
<th>CO₂ (GtC/a)</th>
<th>2010</th>
<th>2011</th>
<th>2012</th>
<th>2013</th>
<th>2014</th>
</tr>
</thead>
<tbody>
<tr>
<td>Prior</td>
<td>0.93 (0.36)</td>
<td>0.56 (0.40)</td>
<td>0.53 (0.32)</td>
<td>0.37 (0.34)</td>
<td>0.41 (0.37)</td>
</tr>
<tr>
<td>in-situ</td>
<td>-0.09 (0.23)</td>
<td>-0.05 (0.25)</td>
<td>-0.01 (0.22)</td>
<td>0.18 (0.22)</td>
<td>-0.21 (0.23)</td>
</tr>
<tr>
<td>Ratio</td>
<td>0.63 (0.13)</td>
<td>0.34 (0.14)</td>
<td>0.53 (0.13)</td>
<td>0.05 (0.13)</td>
<td>0.07 (0.14)</td>
</tr>
<tr>
<td>CH₄ (Mt/a)</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Prior</td>
<td>44.1 (18.4)</td>
<td>40.3 (16.4)</td>
<td>40.2 (16.4)</td>
<td>40.2 (16.4)</td>
<td>40.2 (16.4)</td>
</tr>
<tr>
<td>in-situ</td>
<td>67.0 (11.6)</td>
<td>59.5 (11.3)</td>
<td>54.6 (11.6)</td>
<td>52.9 (11.9)</td>
<td>59.5 (11.2)</td>
</tr>
<tr>
<td>Ratio</td>
<td>74.4 (3.6)</td>
<td>78.6 (3.8)</td>
<td>74.0 (3.5)</td>
<td>73.4 (3.2)</td>
<td>73.1 (3.9)</td>
</tr>
</tbody>
</table>
Figures

Figure 1. Top panel: Geographic basis functions used in our CO$_2$ and CH$_4$ flux inversion experiments. There are 44 land and 11 ocean regions. The red dots and the black crosses represent the locations of the NOAA in situ CO$_2$ and CH$_4$ observations we assimilate in both the ratio inversion, and the in situ only inversion. Geographical regions are based on those used by the TransCom experiments (Gurney et al., 2002). Bottom Panel: Definition of the aggregated northern (red), tropical (yellow) and southern (light blue) land regions.
Figure 2: Annual mean (2010-2014, inclusive) regional net fluxes of (top) CO$_2$ and (bottom) CH$_4$ inferred from the (red) ratio experiments and the (blue) in situ experiments. The grey columns represent the a priori estimates and the vertical lines superimposed on the columns denote one-sigma error. Geographical regions are as defined in Figure 1.
Figure 3: The net monthly CO$_2$ fluxes inferred by the in-situ only inversion (blue) and the ratio inversion (red), compared to the prior estimates (black). The vertical lines (envelopes) represent the prior (posterior) uncertainties. In the plots, we aggregate CO$_2$ fluxes of all 4 categories to the net monthly values over 4 predefined global regions (Figure 1): a) global; b) northern lands; c) tropical lands, and; d) south lands.
Figure 4: As Figure 3 but for CH$_4$ fluxes.
Figure 5: GOSAT XCH₄:XCO₂ ratios over tropical South America (Figure 1) described on the GEOS-Chem 4° (latitude) by 5° (longitude) averaged over (left) October to December 2013, inclusive, and (right) January to March 2014, inclusive. Black dots represent two NOAA in situ sites RPB and ABP, and triangles represent independent AMAZONICA sites (RBA, ALF, TAB, SAN) and two ACO sites (RBH, SAH), which are described in the main text.
Figure 6: As Figure 3 but for CO₂ and CH₄ fluxes over tropical South America (Figure 1).
Figure 7: Monthly mean partial CO₂ columns at four sites over the Amazon (RBA, ALF, TAB, and RBH, Figure 1) collected by the AMAZONICA project and two sites (RBH and SAH) after 2012 collected by the ACO project: (left) comparison and (right) differences with the GEOS-Chem model that has been sampled at the time and location of each observation and driven by fluxes inferred from the in situ (blue) and ratio (red) inversions. The mean and standard deviations (ppm) are shown inset of the right hand side panels. In the plot, we have combined the data over the AMAZONICA site RBA (for 2010 to 2012) and the ACO site of RBH (for 2012 to 2014) for a complete time-series from 2010 to 2014 over the same location.
Figure 8: As Figure 7 but for comparison of the monthly mean partial CH$_4$ columns (in ppb) of the model simulations with AMAZONICA and ACO observations. Due to availability, CH$_4$ observations for 2012 have not been included.
Figure A1: Differences between observed and (left) in situ and (right) ratio a posteriori model (top) CO₂ and (bottom) CH₄ mole fractions observed during HIPPO experiments 1-5 (Wofsy et al., 2011) that cover individual periods during 2009, 2010, and 2011. Model and observation are gridded on a latitude interval of 5 degrees and a vertical interval of 500 m.
Figure A2: (Top) HIPPO-3 (Wofsy et al., 2011), May 2010, and a posteriori model partial columns of (left) CO₂ and (right) CH₄ as a function of latitude over the Pacific Ocean, and (bottom) the differences between the observations and the in situ and ratio inversions. The mean biases (standard deviations) between the model and data are shown in set of lower panels. Data and model values are binned into 5° mass-weighted latitude boxes.
Figure A3: Monthly means CARIBIC and a posteriori model (left) CO₂ and (right) CH₄ mole fractions collected in the tropical middle/upper troposphere (<300 hPa) between 30°S and 30°N. The monthly mean biases (standard deviations) of the model minus data differences are shown inset.
Figure A4: Mean multi-year statistics (2010-2014) of the differences between TCCON (top) XCO$_2$ and (bottom) XCH$_4$ measurements and the a posteriori models. Blue and red bars denote the standard deviations between TCCON and the in situ and ratio a posteriori model, respectively. Black circles and green triangles denote the mean deviations TCCON and the in situ and ratio a posteriori models.
References


Oda, T., and Maksyutov, S.: A very high-resolution (1 km×1 km) global fossil fuel CO₂ emission inventory derived using a point source database and satellite observations of nighttime lights, Atmos. Chem. Phys., 11, 543–556, doi:10.5194/acp-11-543-2011, 2011.


