

Impacts of the July 2012 Siberian Fire Plume on Air Quality in the Pacific Northwest

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Abstract. Biomass burning emissions emit a significant amount of trace gases and aerosols and can affect atmospheric
15 chemistry and radiative forcing for hundreds or thousands of kilometers downwind. They can also contribute to exceedances
of air quality standards and have negative impacts on human health. We present a case study of an intense wildfire plume
from Siberia that affected the air quality across the Pacific Northwest on July 6-10, 2012. Using satellite measurements
(MODIS True Colour RGB imagery and MODIS AOD), we track the wildfire smoke plume from its origin in Siberia to the
Pacific Northwest where subsidence ahead of a subtropical Pacific High made the plume settle over the region. The
20 normalized enhancement ratios of O₃ and PM₁ relative to CO of 0.26 and 0.08 are consistent with a plume aged 6-10 days.
The aerosol mass in the plume was mainly submicron in diameter (PM₁/PM_{2.5} = 0.96) and the part of the plume sampled at
the Whistler High Elevation Monitoring Site (2182 m ASL) was 88% organic material. Stable atmospheric conditions along
the coast limited the initial entrainment of the plume and caused local anthropogenic emissions to buildup. A synthesis of air
quality from the regional surface monitoring networks describes changes in ambient O₃ and PM_{2.5} during the event and
25 contrasts them to baseline air quality estimates from the AURAMS chemical transport model without wildfire emissions.
Overall, the smoke plume contributed significantly to the exceedances in O₃ and PM_{2.5} air quality standards and objectives
that occurred at several communities in the region during the event. Peak enhancements in 8-hr O₃ of 34-44 ppbv and 24-hr
PM_{2.5} of 10-32 µg/m³ were attributed to the effects of the smoke plume across the Interior of British Columbia and at the
Whistler Peak High Elevation Site. Lesser enhancements of 10-12 ppbv for 8-hr O₃ and of 4-9 µg/m³ for 24-hr PM_{2.5}
30 occurred across coastal British Columbia and Washington State. The findings suggest that the large air quality impacts seen
during this event were a combination of the efficient transport of the plume across the Pacific, favorable entrainment
conditions across the BC interior and the large scale of the Siberian wildfire emissions. A warming climate increases the
risk of increased wildfire activity and events of this scale re-occurring under appropriate meteorological conditions.

Keywords: long-range atmospheric transport, smoke, enhancement ratio, threshold exceedance, ozone, particulate matter

1 Introduction

Wildfires emit significant amounts of primary pollutants including fine particulate matter ($PM_{2.5}$), carbon monoxide (CO), nitrogen oxides (NO_x), and non-methane hydrocarbons (NMHCs). These pollutants can react to form ozone (O_3) and secondary organic aerosols (SOA), which can affect downwind air quality and climate over a wide range of scales from local (Akagi et al., 2011, 2012, 2013; Smolyakov et al., 2014), to regional (Wigder et al., 2013), to inter-continental (Jaffe et al. 2001, 2004; Bertschi et al., 2004; Bertschi and Jaffe, 2005; Lapina et al., 2006; Pfister et al., 2006; Weiss-Penzias et al., 2007; Bein et al., 2008; Strada et al., 2012; Jaffe and Wigder, 2012). Short-term exposures to air pollution from biomass burning have been found to be associated with a range of health impacts, including respiratory symptoms, increased hospital admissions and emergency room visits and premature mortality (Henderson and Johnston, 2012).

The rate of O_3 and SOA formation in wildfire plumes, and hence the potential for adverse air quality and health impacts, can vary considerably due to several factors, including biomass fuel type, combustion characteristics, atmospheric conditions, and downwind distance (Akagi et al., 2011; Jaffe and Wigder, 2012). The O_3 production in biomass burning is driven by photochemical reactions of NO_x and NMOCs and typically occurs in a NO_x limited environment. O_3 enhancements relative to CO, also known as the normalized enhancement ratio (NER; $\Delta O_3/\Delta CO$), for temperate and boreal wildfires generally increase as the plume ages and is typically in the range of 0.1 to 0.7 (Jaffe and Wigder, 2012). A similar trend is present for the NER of PM_1 ($\Delta PM_1/\Delta CO$) for plumes aged <2 days; yet for older plumes, SOA formation can be overshadowed by PM loss due to wet and dry deposition, cloud processing and/or evaporation (Wigder et al., 2013). Further study is still required to better understand the interactions of the above factors on downwind air quality for long-range transport (LRT) events.

The Russian boreal forests represent the largest forested region on Earth with an aerial coverage of approximately 8 million km^2 (Stocks, 2004). About 1% of this region is damaged from 10,000 to 35,000 forest fires annually (Isaev et al., 2002). Jaffe et al. (2004) estimated that the active Siberian wildfire season in 2003 caused summertime O_3 enhancements of 9-17 ppbv, which led to episodic O_3 exceedances in western North America. Other studies have also demonstrated the LRT of wildfire smoke from Siberia can affect air quality over Asia (Jeong et al., 2008; Jung et al., 2016) and contribute to arctic haze (Brock et al., 2011).

This study examines an episodic LRT event of Siberian wildfire smoke to the Pacific Northwest that had significant and widespread air quality impacts on July 6-10, 2012. Record temperatures and dry weather across Siberia in 2012 led to increased wildfire activity. Over 17,000 wildfires were detected in July and August 2012 alone. The Fire Inventory from NCAR model (FINN; Wiedinmyer et al., 2011) estimated that 48 Tg of CO emissions occurred by August 2012 which was already more than double the biomass burning emission for the entire 2010 season and second only to the 72 Tg from the 2003 season (NASA, 2012).

Cottle et al. (2014) examined the impact of this event in the Lower Fraser Valley (LFV) in British Columbia, Canada using lidar measurements taken at the University of British Columbia. HYSPLIT trajectory analysis confirmed that aerosol layers subsiding over the LFV had largely originated from over wildfire source areas in Siberia at least 6 days earlier. Aerosol backscatter measurement and low depolarization volume ratios during the event showed the progressive entrainment of smoke into the LFV through July 10th, 2012 which coincided with the high PM_{2.5} observed by the region's fixed air quality monitoring network. This study expands on the Cottle et al. (2014) work in a number of significant ways: it analyzes potential air quality impacts over a much greater geographical area encompassing large parts of British Columbia and Washington State; it uses detailed air quality measurements at a high elevation background site to provide insight into plume chemistry; and it makes use of photochemical modelling to establish baseline air quality conditions in the absence of any wildfire emissions to determine the smoke plume's contribution to degraded air quality and exceedances of regional air quality objectives and national standards of O₃ and PM_{2.5}.

2. Methods

2.1 Monitoring Data

2.1.1 Ambient Air Quality Monitoring Data

Ambient O₃ and PM_{2.5} data were obtained from several air quality monitoring networks (Fig. 1 c and d), including 35 stations from the British Columbia Ministry of Environment (BCMoE) network (BC MoE, 2014a), 22 air quality monitors within the Lower Fraser Valley (LFV) Air Quality Monitoring Network (Metro Vancouver, 2014), and 15 stations from the Washington State Monitoring Network (WSMN) (WA ECY, 2012). The air quality monitoring network across British Columbia is, in part, co-managed by the National Air Pollution Surveillance Network (NAPS) (Dabek-Zlotorzynska et al., 2011). Study sites also included two marine background sites (Amphitrite Point, Ucluelet, site C, 18 m above sea level (ASL)) on Vancouver Island (McKendry et al., 2014); and Cheeka Peak, site D, 480 m ASL in Washington State (Weiss-Penzias et al., 2004) and two elevated background sites (Whistler Peak, site 1, 2182 m ASL, located 120 km north of Vancouver (Takahama et al., 2011; Macdonald et al., 2011); and Mt. Rainier Jackson Visitor Centre, site F, 1782 m ASL, located in Pierce County, Washington State). A description of the data analyzed and sampling methodologies for the various networks and sampling sites is summarized in Table 1.

2.1.2 Lidar Sites

Two lidars were used in this study: one located on the University of British Columbia Point Grey campus in Vancouver (site B, 80 m ASL) and a second located at the north end of the Whistler village (site A, 650 m ASL), approximately 8 km north of the Whistler Peak High Elevation (WHI) (Gallagher et al., 2012; Strawbridge, 2013). The lidars use a Continuum Inlite III

laser with an approximate output power of 1.5W at 10 Hz and operates simultaneously at wavelengths of 532 and 1064 nm. The system measures the return signals through 3 detection channels: one at 1064 nm, and two at 532 nm (one for each polarization). Backscatter information provided by the upward pointing setup has a vertical resolution of 3 m from near the ground up to 15 km and uses a 10s averaging period. Additional details on the lidar systems can be found in Strawbridge
5 (2013).

2.1.3 Whistler Peak High Elevation site

Air quality measurements were collected at the Whistler Peak High Elevation site (see Table 1 for details). CO was measured with Thermo Environmental Instruments Inc., Model 48C-Trace Level analyzer and O₃ was measured with a Thermo Environmental Instruments Inc., UV absorption monitor (TECO 49C). Particle chemistry was provided both
10 through integrated filter samples and continuously with an Aerosol Chemical Speciation Monitor (ACSM). Inorganic ions (Cl⁻, NO₃⁻, Na⁺, NH₄⁺, K⁺, Mg²⁺, and Ca²⁺) were analyzed on 48-hour filter samples by ion chromatography (IC) and ICP-MS. OC, POC, and EC measurements were also obtain on a Sunset lab OC-EC aerosol analyzer (Sunset Laboratory Inc.,
2016) using the EnCan Total-900 thermal method (Huang et al., 2006) with an 8-day sampling period (July 3-11) split between day and night. The ACSM provided particle chemical composition for SO₄²⁻, NO₃⁻, Cl⁻, NH₄⁺, organics on a 30-min
15 time resolution. Optical measurements of black carbon (rBC) were acquired using the Single Particle Soot Photometer (SP2). During the study, the SP2 instrument had a steady 6% loss in its laser power; however internal diagnostics suggest that the power loss is unlikely to significantly affect the readings because heating aerosol during intake was unaffected. SP2 readings were also affected by a leak in the inlet plumbing noticed after the event. Particle size distributions were measured
20 by an optical particle counter (OPC) over the (0.25 to 32 um) for the entire period and also from a Scanning Mobility Particle Sizer (SMPS) (14-572 nm diameter) from July 9 onwards. Details of the trace gas and particle measurements at Whistler Peak are discussed by Macdonald et al. (2011) and Takahama et al. (2011).

2.2 Trajectory and Dispersion Modelling

A detailed analysis by Cottle et al. (2014) using HYSPLIT (Hybrid Single-Particle Lagrangian Integrated Trajectory) trajectories and high resolution smoke dispersion models generated with the Bluesky framework (Sullivan et al., 2008)
25 concluded that wildfire sources within North America did not significantly influence air quality within the LFV during the July 6-10, 2012 period.

A similar analysis was done using the CMC Trajectory Model (D'Amours et al., 2015) with meteorological fields from the GEM-LAM 15 km model at 15 minute intervals for nearby North America wildfires (Long Draw, Oregon and Waldo
30 Canyon, Colorado). Forward trajectories were integrated for a 72 hour period at release heights of 10 m, 100 m, and 1000 m above ground level (AGL). Fig. S1 shows the forward trajectories for the Long Draw and Waldo Canyon wildfires released at 00 UTC on July 9th, 2012 and confirmed that smoke from these sources were unlikely to have contributed to the degraded

ambient air quality observed over the larger geographical extent in this study. Trajectory modelling was not done for the few small wildfires within the Interior of British Columbia. Overall, the effects of these sources are expected to be localized and with only minor air quality enhancements until after July 9th, 2012.

2.3 Analysis of O₃ and PM_{2.5} impacts

5 Based on measurements collected at over 70 air quality monitoring stations in the Pacific Northwest, the air quality impacts of the Siberian wildfire plume were assessed by examining the O₃ and PM_{2.5} concentrations at various averaging periods of 1-hr, 8-hr and 24-hr, depending on the pollutant. For each monitoring station, severe air quality episodes were identified using the regional objective and national standards, applicable at the time of the event (Table 2), as thresholds, and were compared to the average historical July daily maxima values from 2000 to 2010. It should be noted that although some of
10 these standards, for example the Canada Wide Standard (CWS) for O₃ (CCME, 2014), are based on multi-year statistics, all standards were compared to hourly observations of their respective averaging periods.

To further characterize the timing and spatial extent of these impacts, baseline concentrations of O₃ and PM_{2.5} over the Pacific Northwest, in the absence of any wildfire emissions from either Siberia or from within North America, were simulated using the AURAMS (A Unified Regional Air-quality Modelling System) air quality transport model (Gong et al.,
15 2006 with updates from Kelly et al., 2012) and were compared with observations.

2.3.1 AURAMS baseline simulation

AURAMS was run over a 12 day period (July 5-15, 2012) using a nested configuration of 12 and 4 km grid spacing with the inner (4 km) domain covering southern British Columbia and northern Washington State (Fig. 1 a and c). Meteorology for the simulations was provided by Environment and Climate Change Canada's (ECCC) Global Environmental Multiscale
20 (GEM) Limited-Area Model (LAM) weather forecast model (Côté et al., 1998) run at a 2.5 km resolution and then interpolated to the AURAMS domains. GEM-LAM was run in a series of 30-hour simulations starting from 00Z with the first six hours of each simulation discarded as "spin-up" in order to allow model fields to reach steady-state. The model was run with anthropogenic emissions based on the 2010 Canadian and 2008 U.S. emission databases. Emission totals for both databases were adjusted to 2012 levels using Metro Vancouver forecasted and backcasted LFV emission estimates (GVRD,
25 2007). The model used biogenic emissions that calculated using the Biogenic Emissions Inventory System (BEIS) version 3.0.9 emissions algorithms and Biogenic Emissions Landuse Database 3 (BELD3) data. AURAMS employed the ADOM-II gas-phase chemical mechanism (Stockwell et al., 1989) and used a chemically-speciated 12-bin sectional distribution to characterize particulate matter from 0.01 to 40.96 µm in diameter.

Lateral boundary conditions to the AURAMS model along its outer (12 km) domain were supplied by seasonally
30 representative climatological values for some species and for O₃ an additional adjustment was made to the climatological boundary conditions in response to the local tropopause height (Makar et al. 2010). Lateral boundary conditions to the inner

4 km domain were supplied from the 12 km model output. AURAMS runs did not include any wildfire emissions from either Siberia or from within North America.

The reliability of the AURAMS baseline simulation (Table S2) was examined by comparing the 1-hr observed and modelled O_3 and $PM_{2.5}$ during non-event days (July 5 and from July 12-15). Within the 4 km domain, the model bias in 1-hr O_3 and $PM_{2.5}$ is relatively low (2.35 ppbv and 1.14 $\mu\text{g}/\text{m}^3$, respectively). Over the 12 km domain, the O_3 and $PM_{2.5}$ performance degraded, with O_3 bias and RMSE of 14.3 ppbv and 20.86 ppbv, respectively, and $PM_{2.5}$ bias and RMSE of 5.84 $\mu\text{g}/\text{m}^3$ and 10.21 $\mu\text{g}/\text{m}^3$, respectively. However, it should be noted that the majority of the study sites fall within the 4 km domain. Overall, the 1-hr model performance of the baseline simulation is comparable to the assessed reliability of the AURAMS model from other studies (Makar et al., 2014).

10 2.3.2 LRT enhancements analysis

LRT enhancements were identified when ambient air quality conditions in the Pacific Northwest departed from the baseline conditions as modelled by AURAMS, beyond a reasonable margin of error. As the model bias can affect the magnitude of the LRT enhancement estimates, the uncertainty in these enhancements was quantified in a conservative manner in this study. This was done by examining the range of differences between the 8-hr and 24-hr observed and modelled values for O_3 and $PM_{2.5}$, respectively, during non-event days (July 5 and from July 12-15) at each monitoring location. Table S1 provides uncertainty estimates in 8-hr O_3 and 24-hr $PM_{2.5}$ for selected sites. For reporting purposes, enhancements for multiple monitoring locations were aggregated, when appropriate, to reflect the average contribution of the Siberian plume across the different regions in the Pacific Northwest.

3. Results and Discussion

20 3.1 July 2012 smoke event overview

On June 29th, 2012, smoke originating from wildfires over Eastern Russia, as noted by the NASA Earth Observatory group, drifted eastward across the Sea of Japan, the Sea of Okhotsk, and Kamchatka (Fig. 1 b). Moderate Resolution Imaging Spectroradiometer (MODIS) Aerosol Optical Depth (AOD) imagery between July 1st and July 6th, shown in Fig. S2 (a to f), illustrates the progression of the aerosol plume ($AOD > 1$) across the Pacific to western North America. Long-range transport of aerosols was aided by strong zonal flow established as a developing storm tracked from the Bearing Sea to the Gulf of Alaska. Subsidence ahead of the subtropical Pacific high is thought to have contributed to the descent of the plume over the study area. In western North America, subsidence and mountain wave activity are found to be important factors in bringing mid-tropospheric dust layers in range of boundary layer entrainment process (McKendry et al., 2001; Hacker et al., 2001).

On July 7th, the MODIS True Colour imagery (Fig. 2) shows the smoke plume oriented northeast to southwest across the southern British Columbia and Washington State. At this time, degraded air quality conditions were observed inland over the Interior of British Columbia, where both the 8-hr O₃ and 24-hr PM_{2.5} exceedances occurred at Kamloops (site 8). From July 8-9, poor air quality conditions continued to spread across the Interior, coinciding with the northward and inland progression of the plume as collaborated by AOD imagery (Fig. S2 g), and caused air quality exceedances at several communities (see Table 3). Over in the coastal region, elevated O₃ and PM_{2.5} also occurred on July 8th, with isolated 8-hr O₃ exceedances at Chilliwack (site 3) and Enumclaw (site 5). MODIS AOD on July 10th (Fig. S2 h) show the plume remnant had shifted southeastward and out of the Pacific Northwest domain.

Figure 3 provides an overview of the maximal enhancements to 8-hr O₃ and 24-hr PM_{2.5} estimated based on the differences between the ambient air quality observation and the baseline AURAMS air quality modelling for the event. Overall, the baseline modelling suggests that the greatest air quality impacts from the Siberian wildfire smoke were in the Interior of British Columbia and at the Whistler High Elevation monitoring site with peak increases of 34-44 ppbv for 8-hr O₃ and 10-32 µg/m³ for 24-hr PM_{2.5}. The coastal portions of British Columbia and Washington State saw lesser enhancements of 10-14 ppbv and 4-9 ug/m³, respectively. The regional disparity in the enhancement may be attributable to difference in subsidence across the region. Coastal regions had stable atmospheric conditions, leading to stagnant conditions under a strong thermal inversion between 500 and 1000 m ASL (Fig. 4 a). Stable atmospheric conditions likely limited the entrainment of the plume over the coastal area and caused a buildup of local anthropogenic emissions. Over the Interior of British Columbia, dry and less stable conditions existed with daytime mixing heights over 3 km ASL (Fig. 4 b). The variations in 8-hr O₃ and 24-hr PM_{2.5} enhancements (with uncertainty estimates) across the Pacific Northwest are listed in Table 4 and are discussed in more detail in the following sections.

3.2 Plume chemistry at Whistler High Elevation Monitoring site

The Whistler Peak High Elevation air chemistry monitoring site is situated to monitor free tropospheric background air pollutants and to study long-range transport events across the Pacific (Leitch et al., 2009; Macdonald et al, 2011). On July 6th, significant increases in 1-hr O₃, CO, and PM_{2.5} mark the onset of the smoke event at Whistler Peak (Fig. 5) and coincide with the timing of elevated aerosol layers detected by the Whistler lidar. Aerosol backscatter ratios show that the plume persisted over Whistler from July 6 14:00 PST (WHI1) to July 8 06:00 PST (WHI2) at an elevation of 2-3.6 km ASL. Backscatter ratios also indicate the entrainment of the aerosols deeper into the Whistler Valley as hourly observed O₃, PM_{2.5}, and CO reached peaks of 86 ppbv, 24 µg/m³, 276 ppbv, respectively. The maximum 1-hr O₃ and CO values were approximately 50 and 167 ppbv above the average July background levels as reported in Macdonald et al. (2011) for 2002-2006. O₃ conditions prior to the event were consistent with July background levels.

The ratio of O₃ to CO enhancements ($\Delta O_3/\Delta CO$) provides some insight into the relative photochemical production of O₃ in wildfire plumes (Jaffe and Wigder, 2012) and is strongly influenced by the travel time between the fire and measurement location (Mauzerall et al., 1998). For this study, the plume age is estimated between 7-10 days old based on a Siberian

origin of June 29th, 2012. The $\Delta O_3/\Delta CO$ ratio was estimated using observations from WHI1 to WHI2 and excluding cases where CO was lower than the July background (110 ppbv). The regression slope of O_3 and CO for the event ranged from 0.15 to 0.24 with a mean value of 0.26 ($n=20$, $r^2=0.80$) and is comparable to other similarly aged Siberian plumes with excess ratios of 0.22-0.36 (6-10 days; Bertschi et al., 2004) and 0.15-0.84 (7-10 days, $n=5$; Bertschi and Jaffe, 2005).

5 As shown in Figure 5d, during the smoke event (WHI1 to WHI2), the $PM_{2.5}$ mass was dominated by submicron aerosol with an average $PM_1/PM_{2.5}$ ratio of 0.96. The particle composition as determined from the ACSM (Fig. 6) during the event was on average approximately 87.6% organic mass, 3.7% NO_3^- , 5.9% SO_4^{2-} , 2.5% NH_4^+ , 0.3% Cl^- . PM_1 mass is estimated from the ACSM, since the size distribution data indicate that most of the mass is below 0.7 μm ; in Figure S3, an example is shown for July 9th, 2012 when the SMPS became operational again. Further, a comparison of the mass concentrations estimated from the SMPS, based on the assumptions of spherical particles and a particle mass density of 1.2 g/cm^3 as a result of the dominant organic composition during that period (Fig. 6), with the ACSM for July 9-31 is illustrated in Figure S4. The comparison is strong over the 518 1-hr averaged data: $r^2=0.95$; slope = 1.10. As in the example in Figure S3, the $PM_{2.5}$ mass is estimated with the addition of the OPC data.

The evolution of PM concentrations in wildfire plumes is largely affected by the rate of secondary organic aerosol (SOA) formation (Akagi et al., 2012). Using the regression methods described in Wigder et al. (2013), $\Delta PM_1/\Delta CO$ is estimated at 0.08 ($n=20$, $r^2=0.94$) during the WHI1 to WHI2 period. This ratio is comparable to those of long-range transported plume (>740 km) of 0.03 and 0.06 found in Weiss-Penzias et al. (2007) and Wigder et al. (2013), respectively. Compared to the median ratios of 0.17, 0.29, 0.29, and 0.19, corresponding to distances of <140, 140-340, 340-540, 540-740 km, respectively, as reported by Wigder et al. (2013), the ratio found in this study is relatively low, suggesting that the net PM_1 loss exceeds SOA formation during transport as the plume ages. A recent wildfire study (Jolleys et al., 2012) found little SOA production in aged biomass burning plumes; however, it was noted that other aircraft campaigns had found a large variability in OA enhancements.

The OC/EC ratio is also considered as its increase may indicate a contribution from SOA formation to the OC as a smoke plume ages. From the 8-day integrated PM_1 quartz filter samples, separated by day and night, the value of OC/EC was 3.8 (day) and 3.4 (night) at Whistler Peak for the period July 3-11. During the WHI1-WHI2 event, an OC/EC ratio of 3.7 was also estimated by scaling the average ACSM OM and rBC to the 8-day integrated OC and EC sample, respectively. In the week prior to July 1st, 2012, a large variation in OC/EC (approximately 1.2 to 7) was observed in the suburb of Novosibirsk, approximately 400-800km south to southwest of the main sources of the Siberian wildfire (Smolyakov et al., 2014). Due to this large variation in OC/EC close to the source as attributed to variations in burning phase and fuel types, no further inferences of SOA formation in the plume were made. The ratios of m/z44 and m/z43 to OM derived from the ACSM, indicate that the organic aerosol during the height of the BB plume (WHI1 to WHI2) was more oxygenated than at any other time during July 5-11 (Fig. S5). If SOA production during the plume transport was insignificant, then this observation implies either a relatively high level of oxygenation near the source, heterogeneous oxidation of the OM during plume transport, or the loss of less oxygenated components of the OM during transport.

The overall contribution of the July 6-10 Siberian Plume event to the air quality standards for O₃ and PM_{2.5} at Whistler Peak (site 1) were estimated using AURAMS baseline modelling (also shown in Fig. 5). Generally, the model depicts relatively clean air quality conditions (approximately 30 ppbv and 1.0 µg/m³ for O₃ and PM_{2.5}, respectively) on July 5th with an increase trend in baseline O₃ (between 5 to 10 ppbv from July 5-8). It is estimated that the Siberian wildfires contributed an additional 44 ppbv to the daily maximum 8-hr O₃ and 10 µg/m³ to the 24-hr PM_{2.5} concentration (Table 4).

3.3 Lower Fraser Valley (LFV)

The UBC lidar (site B) data indicated a plume arrival time of 12:00 PST on July 6th (LFV1, Fig. 7a) and a plume elevation of between 2 – 3.6 km ASL over Metro Vancouver which is similar to that found by the Whistler lidar (site A) but with generally weaker aerosol backscatter ratio readings (Fig. 7a). The plume had a shorter overall persistence over Metro Vancouver with the plume's trailing edge shifting northward by 03:00 PST on July 8th (LFV2). It appears that the conditions favourable for the transport of Siberian smoke to LFV were also conducive to O₃ production (e.g. light winds, clear skies, and high temperatures). This makes the attribution of Siberian smoke impacts on local air quality slightly more challenging in the LFV than at Whistler. Figure 7 (b,c) depicts the increasing O₃ and PM_{2.5} across the LFV during July 6-10 with the network average 1-hr O₃ and 1-hr PM_{2.5} rises to 54 ppbv and 18 µg/m³, respectively, on July 8th. Peak O₃ condition of 79 ppbv observed at Hope Airport (site 4) and PM_{2.5} of 26 µg/m³ at Port Moody Rocky Point (site 2) arose as inland temperature reached a maximum of 31°C at Chilliwack (site 3) on July 8th. Inland temperatures and O₃ conditions moderated subsequently on July 9th following a brief period of elevated convection in the early morning hours that affected parts of Washington State and LFV. High O₃ concentrations were reported over the high elevation (>1000m ASL) site in North Vancouver, located close to the coast over the norther portion of the LFV, by a separate field campaign (Bart et al., 2014).

The AURAMS model was used to assess the underlying baseline air quality (Fig. 7 b and c) and to assess the local impacts through a site-by-site analysis of differences with LFV monitoring network data (Fig. 8 a and b). For July 6-8, the AURAMS baseline simulation without wildfire emissions captured the rise in daytime O₃ on average across the LFV network (Fig. 7 b) and on a site-by-site basis (Fig. 8 a). In contrast, the baseline simulation was consistently lacking in nighttime O₃ titration. Previous air quality studies have shown that these discrepancies do not significantly affect the rest of the baseline AQ simulation (Makar et al, 2014). As such, the estimates of O₃ enhancement within the LFV are based on daytime hours only. The 1-hr O₃ enhancements occurred on July 9-10 appears, based on the lidar data, to be about 18 - 27 hours after the plume advected north of the LFV (Fig. 8 a). The average enhancement of peak 8-hr maximum O₃ across the LFV was estimated to be 10 ppbv (Table 4; hourly variability shown as shaded regions in Fig. 8a). Spatial maps of 8-hr O₃ (Fig. S6) show widespread enhancements across the LFV on July 9th lingered in the western portion on July 10th. The lack of O₃ enhancement estimated throughout the LFV on July 8th (Fig. 8a and Fig. S6) suggests the exceedance-level O₃ episode in Chilliwack (Fig. 8c) may have been due to local O₃ production and not the Siberian wildfire.

The PM_{2.5} enhancements associated with the LRT event developed in the morning of July 8th, much earlier than with any O₃ enhancements, and then persisted through July 10th. The Siberian plume is estimated to have enhanced the 24-hr PM_{2.5} by a

maximum of $9 \mu\text{g}/\text{m}^3$ for sites across the LFV. Spatially, the largest enhancements to $\text{PM}_{2.5}$ appear to be concentrated over the northern portion of the LFV (Fig. 3 and by day in Fig. S6). This suggests that entrainment of the plume may have been facilitated by slope flows along the northern mountainous terrain, yet appears not to have reached the floor of the LFV, as evidenced by O_3 data from the fixed monitoring network.

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3.4 Washington State

In general, the air quality concentrations in Washington State (WA) during the episode were similar to those of the LFV. The average hourly $\text{PM}_{2.5}$ across WA rose to a maximum of $\sim 15 \mu\text{g}/\text{m}^3$ on July 9th and the peak 8-hr O_3 of 72 ppbv (Fig. S7) occurred at Enumclaw on July 8th. The LRT smoke enhancements in WA (Table 4) were similar to those in the LFV for O_3 with a 10 ppbv increase to the 8-hr O_3 and a weaker contribution to the 24-hr $\text{PM}_{2.5}$ of $4 \mu\text{g}/\text{m}^3$ (Fig. 3 and by day in Fig. S6). However, the timing and geographical extent of the impacts of $\text{PM}_{2.5}$ and O_3 differed considerably. In particular, the impacts on $\text{PM}_{2.5}$ were minimal over WA until July 10th, while O_3 underwent some modest enhancement on July 9th at a limited number of monitoring locations (Fig. S7). It is likely that some of the discrepancies (between the Canadian and US portions of the LFV in the estimated plume impacts) arise from model errors stemming from different methodologies used to create the 2010 Canadian and 2008 US emissions inventories.

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3.5 The Interior of British Columbia

The character of the Siberian plume and its impacts over the Interior of British Columbia differed considerably from the enhancements seen over the LFV and WA. The air quality for two communities (Fig. 9), Kamloops (site 8) and Williams Lake (site 9), reflect the dramatic response seen across the Interior. The arrival of the plume at these communities, 19:00 PST on July 6 for Kamloops (site 8) and 12:00 PST on July 8th for Williams Lake (site 9), resulted in a sudden shift in hourly $\text{PM}_{2.5}$ and elevated O_3 on the subsequent day. Based on the AURAMS modelling, it is estimated that the plume contributed an additional 18 and 38 ppbv at Kamloops (site 8) and Williams Lake (site 9), respectively, to the daily maximum 8-hr O_3 , and an additional $32 \mu\text{g}/\text{m}^3$ and $30 \mu\text{g}/\text{m}^3$, respectively, to the 24-hr $\text{PM}_{2.5}$ (Table 4). The suddenness and magnitude of these impacts suggests that the plume became well mixed into the boundary layer after its transition across the Coast Mountains, which also corroborates with the entrainment signatures noted on the Whistler lidar that shows the plume mixed down to the base of the Whistler Valley (Fig. 5).

Unfortunately, the 12 km model study domain did not extend north sufficiently to provide baseline estimates of the air quality conditions to Quesnel (site 10), Prince George (site 11), and the Northern Interior (site 12 to 14), which also experienced conditions that exceeded national and provincial air quality standards. Instead, a baseline air quality level was inferred from the observations on July 6th and the enhancements at these sites (listed in Table 4) were found to be consistent with those experienced at Williams Lake (site 9; based on a July 6th baseline, as well as the 12km domain) except for slightly weaker 24-hr $\text{PM}_{2.5}$ enhancements of 22-24 $\mu\text{g}/\text{m}^3$.

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On July 10th, as the bulk of the Siberian plume shifted northeastward out of British Columbia, remnants of the plume shifted over the Southern Interior and Whistler; however, attribution of the plume's impacts was impeded due to local wildfire activity (Matthew Creek; approx. 100 km west of Kelowna; 155 acres). Overall, the maximal air quality enhancements due to smoke from the Siberian wildfires in the Southern Interior was approximately 15 ppbv and 15 $\mu\text{g}/\text{m}^3$ for 8-hr O_3 and 24-hr $\text{PM}_{2.5}$, respectively, based on the impact analysis conducted at Vernon (site 7) for the July 6-9 period (Fig. S10).

4. Conclusions

The extreme 2012 Russian boreal fire season resulted in the trans-Pacific transport of a large smoke plume across the Pacific in early July which had substantial, yet varied, impacts across large part of British Columbia and Washington State. Air quality concentrations in excess of the Canada Wide Standard (65 ppbv for 8-hr O_3) and the British Columbia Air Quality Objective (25 $\mu\text{g}/\text{m}^3$) of 24-hr $\text{PM}_{2.5}$ were experienced on July 6-10 at several municipalities across British Columbia.

The aged smoke plume (6-10 days old), as supported from data from Whistler Peak, had a mean normalized enhancement ratio ($\Delta\text{O}_3/\Delta\text{CO}$) of 0.26 and a PM_1/CO ratio of 0.08, consistent with other long-range transport smoke studies of similar origin. Fine organic aerosol mass (based on the ACSM measurements) accounted for the majority (~88% by mass fraction) of the 1-hr and 24-hr $\text{PM}_{2.5}$ enhancements of 23 $\mu\text{g}/\text{m}^3$ and 10 $\mu\text{g}/\text{m}^3$, respectively. Analysis of OC/EC did not provide evidence for significant SOA production during transport as the plume ages. However, analysis of the ratios of m/z44 and m/z43 to OM derived from the ACSM, the level of oxygenation of organic material in this plume was much higher than during flanking times. If SOA formation during transport was minimal, then the relatively high level of oxygenation may have resulted from three factors: the OM near the source was relatively oxygenated; heterogeneous oxidation of the OM during plume transport; or removal of less oxygenated components of the OM during transport.

Baseline air quality modelling provided by the AURAMS model was used to delineate the scope of the LRT enhancements to air quality under varying atmospheric conditions. Generally, the largest enhancements to 8-hr O_3 of 34-44 ppbv and 24-hr $\text{PM}_{2.5}$ of 10-32 $\mu\text{g}/\text{m}^3$ occurred at Whistler Peak (site 1), Kamloops (site 8) and across the Central and Northern Interior of British Columbia (site 9 to 14). In contrast, lesser LRT enhancements of approximately 10 ppbv to 8-hr O_3 and 4-9 $\mu\text{g}/\text{m}^3$ to 24-hr $\text{PM}_{2.5}$ were estimated for coastal British Columbia and Washington State. Stable atmospheric conditions along coastal area of British Columbia and Washington State likely limited the initial entrainment of the plume. The AURAMS modelling approach accounted for enhancements in anthropogenic air quality due to the smoke plume in the majority of cases with the exception of July 8th at Chilliwack (site 3) and Enumclaw (site 5), where elevated levels may have been due to other factors.

Long range transport smoke events, such as the 2012 Siberian forest fire plume, highlight the far reaching impacts of wildfires and how air quality impacts can vary with downwind meteorology, plume age and entrainment pathways. This demonstrates the need to include wildfire emissions within chemical transport model, and to derive effective parameterizations for the lofting and subsequent reactions in modelling ambient air quality conditions. Lateral boundary for chemistry used in regional air quality modeling application should be responsive to the long-range transport of pollutants.

Wildfires sources pose an increasing concern as global climate change is expected to increase wildfire frequency and severity in many regions (IPCC, 2007; Gillett, 2004; Liu et al., 2013). As such, wildfire smoke is expected to contribute more to the ambient levels for O₃, PM_{2.5}, CO, and other pollutants and can potentially increase the frequency of extended periods of degraded air quality.

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Data availability

Data sources used in this paper are listed in the supplemental material (Table S3). For ECCC data not readily accessible, please contact the corresponding author (Andrew.Teakles@canada.ca) for assistance.

Supplementary material related to this article is available online at:

10 **Author contribution**

A. Teakles designed the study, developed the methodology and prepared the manuscript with contributions from co-authors.

R. So consolidated and analyzed the data and provided contributions to the methodology.

J. Baik provided assistance in developing R code to generate figures.

B. Ainslie generated the AURAMS model run.

15 S. Hanna, A. Bertram and L. Huang provided data and their expertise on data and instrumentation comparability.

K. Strawbridge provided the lidar data and his expertise in interpreting this data.

AM. Macdonald and R. Leaitch provided and analyzed data collected at Whistler Peak High Elevation site.

D. Toom assisted with the ACSM data and supplementation analysis on OA oxygenation.

I. McKendry provided his expertise on the LFV analysis.

20 R. Vingarzan reviewed and provided contributions to the methodology and manuscript composition.

C. Schiller reviewed the manuscript and provided her chemistry expertise to the data interpretation.

R. Nissen performed the synoptic analysis.

D. Jaffe contributed to the enhancement ratios analysis and the U.S. analysis.

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Table 1. Instrument list for various monitoring networks used in the study.

Network/ Site	Measured Parameters	Instrument(s)	Reference
<i>British Columbia</i>			
British Columbia Ministry of Environment (BCMoe) network & Lower Fraser Valley Air Quality Monitoring Network	O ₃	Thermo Environmental Instruments Inc., UV absorption monitor (TECO 49C)	Environment Canada (2013); Metro Vancouver (2014)
	PM _{2.5}	Thermo Scientific TEOM 1400ab with sample equilibration system Met One BAM-1020 Thermo Scientific Model 5030 SHARP	
Amphitrite Point, Ucluelet (site C)	O ₃	Thermo 49i O ₃ analyzer	McKendry et al. (2014)
	PM _{2.5}	Thermo Sharp 5030 particulate monitor	
UBC Lidar (site B) Whistler Village Lidar (site A)	aerosol backscatter	Lidar	Strawbridge (2013)
Whistler Peak High Elevation site (site 1)	CO	Thermo Environmental Instruments Inc., Model 48C-Trace Level analyzer	Macdonald et al. (2011)
	O ₃	Thermo Environmental Instruments Inc., UV absorption monitor (TECO 49C)	Macdonald et al. (2011)
	Refractory Black Carbon	Droplet Measurement Technologies Inc., Single Particle Soot Photometer (SP2)	Takahama et al. (2011)
	Particle Chemistry < 1 micron (SO ₄ ²⁻ , NO ₃ ⁻ , Cl ⁻ , NH ₄ ⁺ , Organics)	Aerodyne Aerosol Chemical Speciation Monitor	Takahama et al. (2011)
	Particle size distributions from 0.25 to 32 µm diameter	GRIMM portable aerosol spectrometer 1.109	Grimm Technologies Inc. (2010)
	Particle size distributions in range ~ 14 nm to 572 nm diameter	TSI 3081 Scanning Mobility Particle Sizer (SMPS) with a TSI 3775 Condensation Particle Counter (CPC) operated at a low flow setting	McKendry et al. (2008)
	Integrated Inorganic Particle Chemistry (in this study: SO ₄ ²⁻)	Teflon filters; 48-hr samples	McKendry et al. (2008)
	Integrated Organic Chemistry (OC and EC)	Sunset lab OC-EC aerosol analyzer (EnCan Total -900 Thermal method); 8-days sample	Sunset Laboratory Inc. (2016); Huang et al. (2006)
<i>Washington State</i>			
Washington State Monitoring Network (WSMN)	O ₃	Teledyne-API T400	WA ECY (2008 & 2012)
	PM _{2.5}	Nephelometer	

Cheeka Peak (site D)	O ₃	Teledyne-API T400	WA ECY (2008 & 2012)
	PM _{2.5}	Radiance Research M903	

Mt. Rainier Jackson Visitor Center (site F)	O ₃	Teledyne-API T400	WA ECY (2008 & 2012)
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Table 2: 2012 Ambient air quality standards and objective for O₃ and PM_{2.5}.

Standard/ Objective	Pollutant	Averaging period	Value	Reference
Canada NAAQS	O ₃	1-hr	82 ppbv	Health Canada (1999)
Canada CWS			65 ppbv	CCME (2014)
BC provincial and Metro Vancouver	O ₃	8-hr	65 ppbv	BC MOE (2014b) Metro Vancouver (2014)
U.S. EPA NAAQS			75 ppbv	U.S. EPA (2014)
Canada CWS			30 µg/m ³	CCME (2014)
BC provincial and Metro Vancouver	PM _{2.5}	24-hr	25 µg/m ³	BC MOE (2014b) Metro Vancouver (2014)
U.S. EPA NAAQS			35 µg/m ³	U.S. EPA (2014)

Table 3. Summary of ambient O₃ and PM_{2.5} conditions for stations that had severe (in bold) air quality degradation (exceeding one of the values listed in Table 2) between July 6th and July 11th. Values indicated reflect the maxima of 1, 8, and 24-hr averaging periods. Values in bracket denote historical July average daily maxima of 1 and 8-hr average periods and the historical July average of 24-hr running mean.

Station(s)	Station(s) ID	O ₃ (ppbv)		PM _{2.5} (µg/m ³)	
		8-hr	1-hr	24-hr	1-hr
<i>Whistler</i>					
Whistler Peak High Elevation Site	1	83	86	12	24
<i>Lower Fraser Valley</i>					
Port Moody Rocky Point Park	2	45 (28)	55 (34)	19 (7)	26 (13)
Chilliwack Airport	3	67 (35)	74 (41)	14 (6)	21 (12)
Hope Airport ^a	4	64.8 (39)	79 (43)	18 (6)	21 (11)
<i>Washington State^b</i>					
Enumclaw	5	72	80	NA	NA
<i>Southern Interior</i>					
Kelowna College	6	74 (40)	78 (44)	15 (5)	30 (13)
Vernon Science Centre	7	67 (32)	73 (36)	19 (5)	30 (12)
Kamloops Fire Station	8	65 (39)	72 (43)	36 (5)	44 (12)
<i>Central Interior</i>					
Williams Lake ^c	9	80 (34)	84 (37)	34 (5)	46 (16)
Quesnel ^c	10	76 (31)	92 (35)	31 (6)	57 (20)
Prince George ^c	11	NA	NA	31 (6)	49 (17)
<i>Northern Interior (N. INT)</i>					
Burns Lake Fire Centre	12	NA	NA	28 (3)	38 (9)
Houston Firehall	13	NA	NA	27 (3)	38 (9)
Telkwa	14	NA	NA	26 (3)	36 (11)

5 ^a air quality conditions at Hope Airport are noted here as it nearly exceeded the regional 8-hr ozone objective

^b encompasses coastal stations in northwestern Washington State as shown in Figure 1

^c multiple stations in the area of interest

Table 4. Maximum daily enhancement (observed – model) of daily maximum 8-hr rolling average O₃ and 24-hr rolling average PM_{2.5} based on various air quality baselines (4km and 12km AURAMS and historical values) from July 6th to July 11th, 2012.

Area of Interest	Baseline	No. of Sites	O ₃ (ppbv)		PM _{2.5} (µg/m ³)		
			enhancement	Uncertainty range	No. of Sites	enhancement	Uncertainty range
Whistler Peak		1	43.7	(15.8, 49.2)	1	10.3	(7.7, 10.6)
LFV ^{a,b}		20	9.9	(1.6, 22.0)	17	8.7	(6.0, 12.8)
Vancouver Island	AURAMS (4km)	6	13.7	(1.7, 27.0)	9	9.4	(7.1, 13.1)
Ucluelet		1	12.4	(0.1, 14.9)	1	6.6	(0.9, 5.8)
U.S. ^{a,b,c}		10	10.4	(0, 23.4)	10	3.8	(4.1, 12.4)
Kelowna ^a		1	5.6	(5.5, 19.6)	1	7.7	(9.6, 11.0)
Vernon ^a	AURAMS (12km)	1	14.6	(16.7, 32.6)	1	14.7	(9.8, 13.1)
Kamloops ^a		1	18.2	(22.0, 33.3)	1	32.1	(29.2, 33.1)
Williams Lake		1	37.5	(20.3, 57.0)	2	30.3	(23.9, 31.3)
Williams Lake		1	38		2	31	
Quesnel	Historical	1	43		3	24	
Prince George	(July 6th)	NA	NA	-	2	23	-
Northern Interior		1	34		4	22	

^a estimated enhancement excluding July 10th and July 11th

5 ^b O₃ enhancements were based on daytime values

^c encompass northwest coastal stations in Washington State as shown in Figure 1

Table S1. Station information and modelling uncertainty for select sites used in this study.

Station	Station #	Elevation (m)	Lat.	Long.	8-hr O ₃ uncertainty *	24-hr PM _{2.5} uncertainty*
<i>Whistler</i>						
Whistler Peak High Elevation Site	1	2182	50.06	-122.96	(-5.6, 27.9)	(-0.3, 2.5)
Whistler Lidar	A	660	50.13	-122.95		
<i>Lower Fraser Valley</i>						
Port Moody Rocky Point Park	2	17	49.28	-122.85	(-11.6, 7.9)	(-3.3, 5.3)
Chilliwack Airport	3	12	49.16	-121.94	(-22.6, 1.4)	(1.2, 2.4)
Hope Airport	4	39	49.37	-121.50	(-18.2, 8.8)	(-4.0, 10.0)
UBC Lidar	B	80	49.26	-123.25		
<i>Vancouver Island</i>						
Ucluelet Amphitrite Point	C	14	48.92	-125.54	(-2.6, 12.3)	(0.9, 5.7)
<i>Washington State</i>						
Enumclaw	5	402	47.14	-121.93	(-22.8, 7.3)	
Cheeka Peak	D	480	48.30	-124.62	(-3.1, 4.9)	(-11.5, 2.4)
Quillayute	E	59	47.94	-124.56		
Mt Rainier Jackson Ctr	F	1782	46.78	-121.74	(-21.5, 12.7)	
<i>Southern Interior</i>						
Kelowna College	6	347	49.86	-119.48	(-14.0, 0.1)	(-3.3, -1.9)
Vernon Science Centre	7	476	50.23	-119.28	(-18.1, -2.2)	(1.6, 4.9)
Kamloops Fire Station	8	381	50.70	-120.39	(-15.2, -3.8)	(-1.0, 2.9)
Kelowna Airport	G	344	49.96	-119.38		
<i>Central Interior</i>						
Williams Lake Columneetza School	9	631	52.14	-122.15	(-19.5, 17.2)	(-1.1, 6.4)
Williams Lake CRD Library		609	52.13	-122.14		(-0.9, 6.5)
Quesnel Maple Drive	10	614	52.96	-122.50		
Quesnel Senior Secondary		490	52.98	-122.49		
Quesnel West Correlieu School		478	52.97	-122.52		
Prince George Gladstone School	11	617	53.86	-122.76		
Prince George Plaza 400		588	53.91	-122.74		
<i>Norther Interior</i>						
Burns Lake Fire Centre	12	710	54.23	-125.76		
Houston Firehall	13	602	54.40	-126.65		
Telkwa	14	515	54.69	-127.05		

*AURAMS model uncertainty (observed-model) during non-event days

Table S2. 1-hr AURAMS model performance for O₃ and PM_{2.5} at 4 and 12km domain.

	y-intercept	Slope	r	Bias	MAE	RMSE
<i>Current Study</i>						
O ₃ (4km)	8.5	0.66	0.69	2.35	7.57	9.55
PM _{2.5} (4km)	5.68	0.35	0.35	1.14	3.71	4.97
O ₃ (12km)	16.46	0.89	0.6	14.3	16.32	20.86
PM _{2.5} (12km)	9.75	0.44	0.24	5.84	7.79	10.21
<i>Makar et al. (2014)</i>						
O ₃ (All; Table 4a)	15.32	0.7	0.64	8.48	12.53	16.17
PM _{2.5} (All; Table 4b)	3.36	0.45	0.28	1.2	5.28	7.72
PM _{2.5} (LFV sites; Table 5)	7.95	0.24	0.23	2.82	6.77	10.53

Table S3. Source information and data availability

Site	Data	Source
<i>British Columbia</i>		
BCMoE network	O ₃ , PM _{2.5}	BCMoE (http://envistaweb.env.gov.bc.ca/)
Lower Fraser Valley Air Quality Monitoring Network	O ₃ , PM _{2.5}	BCMoE (http://envistaweb.env.gov.bc.ca/)
Amphitrite Point, Ucluelet	O ₃ , PM _{2.5}	ECCC ,contact info: Andrew.Teakles@canada.ca
UBC & WHI Lidar	Lidar	ECCC, contact info: Andrew.Teakles@canada.ca
Whistler Peak High Elevation Site	Particle size distribution , O ₃ , CO, Filter packs IC (tSO ₄), Black Carbon, TOT EC & OC, ACSM, Lidar	ECCC, contact info: Andrew.Teakles@canada.ca
BC surface air quality stations	O ₃ , PM _{2.5} (climatology)	ECCC NAPS website (annual summaries: http://maps-cartes.ec.gc.ca/rnspa-naps/data.aspx)
<i>Washington State</i>		
WSMN	O ₃ , PM _{2.5} (neph)	Puget Sound Clean Air website (http://airgraphing.pscleanair.org/)
Cheeka Peak	O ₃ , PM _{2.5}	Department of Ecology, State of Washington, https://fortress.wa.gov/ecy/enwiwa/
Mt. Rainier Jackson Visitor Center	O ₃	Department of Ecology, State of Washington, https://fortress.wa.gov/ecy/enwiwa/
<i>Air Quality Modelling</i>		
All sites	AURAMS O ₃ , PM _{2.5}	ECCC, contact info: Andrew.Teakles@canada.ca
<i>Satellite</i>		
World	MODIS AOD	http://modis.gsfc.nasa.gov/data/
Pacific	MODIS true colour	http://visibleearth.nasa.gov/view.php?id=78406
<i>Others</i>		
Quillayute & Kelowna	Radiosondes	http://weather.uwyo.edu/upperair/sounding.html
Elevation data (for fig 1)	GLOBE Elevation	http://www.ngdc.noaa.gov/mgg/topo/globe.html

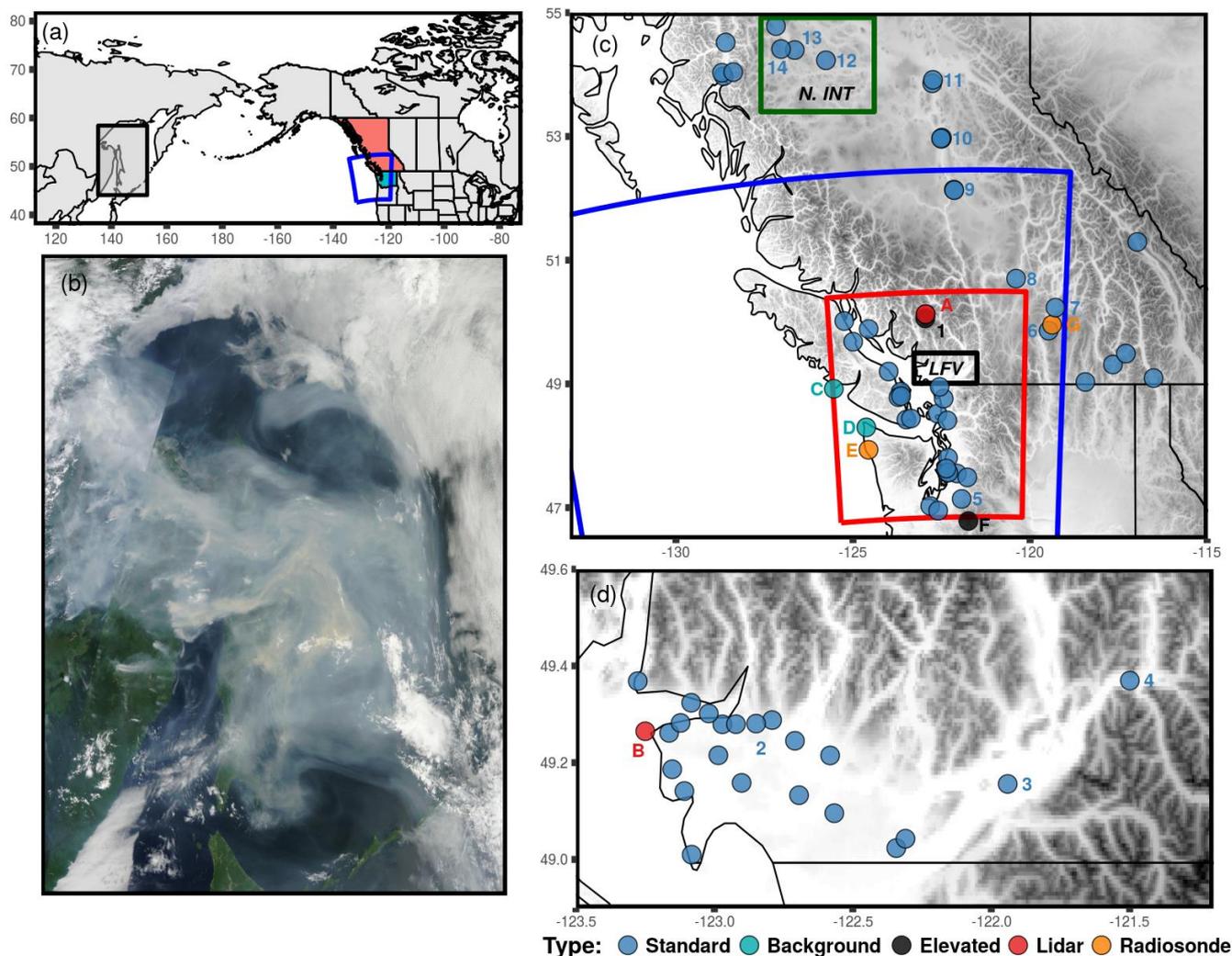


Fig. 1. Map of the monitoring network data and AURAMS modelling domain used in the study. The pink and blue regions in (a) highlights the two geographical regions of interest, British Columbia and Washington state, respectively. The black outlined area in (a) highlights the geographical location of the MODIS true colour Image on June 29th, 2012 from the NASA earth Observatory site (b). The blue and red outlined areas in either (a) and/or (c) highlight the 12 km and 4 km AURAMS grid domains, respectively. The green outline in (c) highlights the Northern Interior sites that fall outside the gridded domain. The Lower Fraser Valley (LFV; black outline in c) inset (d) displays the Metro Vancouver air quality monitoring network and the location of the UBC lidar. Abbreviations used in this figure are defined in Table S1 and throughout the text. Topography shaded in panels (c) and (d).

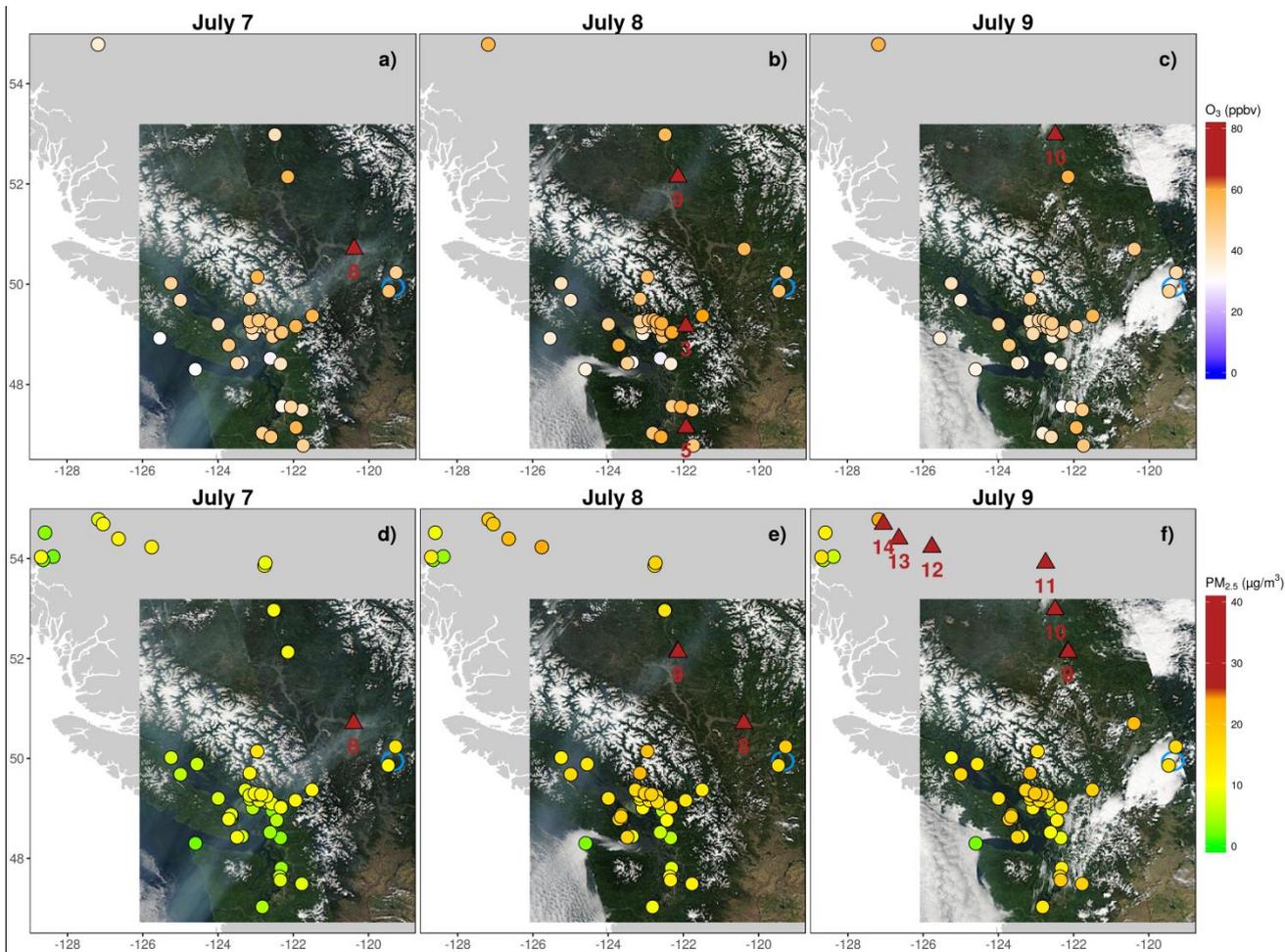
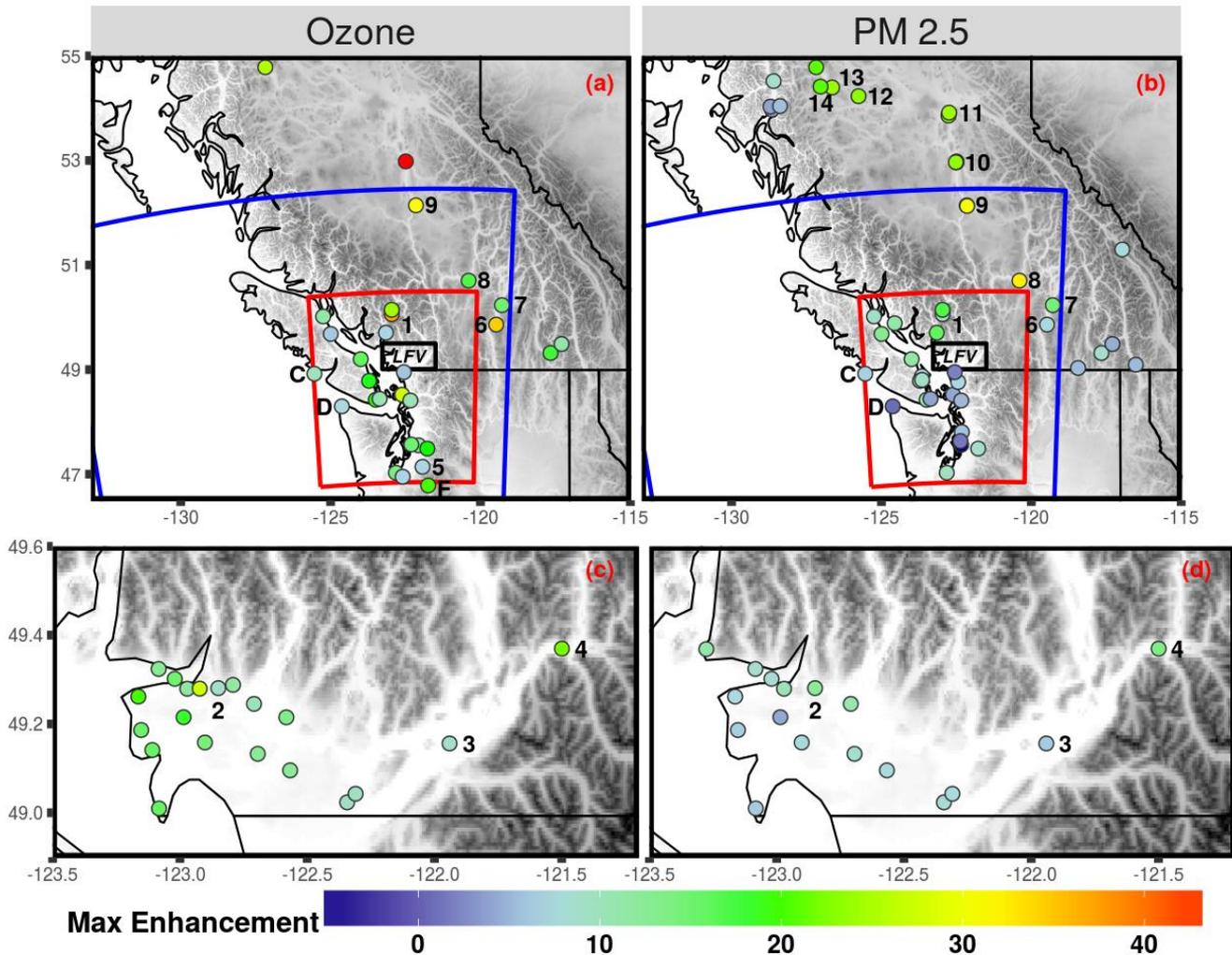


Fig. 2. MODIS true colour Image composite maps for July 7, 8, and 9, 2012 overlaid with daily maximum of 8-hr O₃ (a, b, and c, respectively) and of 24-hr PM_{2.5} (d, e, f, respectively) for the study sites. The locations where observed concentration exceeded the national or regional air quality standards for either O₃ or PM_{2.5} are indicated in red triangles in the appropriate panel and described in Table 3. The light blue circle on the MODIS source imagery indicates the location of the Kelowna Airport (site G) on all panels. Abbreviations used in this figure are defined in Table S1 and throughout the text.



5 **Fig. 3.** Overview of the maximal enhancements to 8-hr O₃ and 24-hr PM_{2.5} estimated for the study area (a and b, respectively) and the LfV (c and d, respectively) based on the differences between the ambient air quality observation and the AURAMS baseline modelling for the July 6-10, 2012 smoke event. Abbreviations used in this figure are defined in Table S1 and throughout the text.

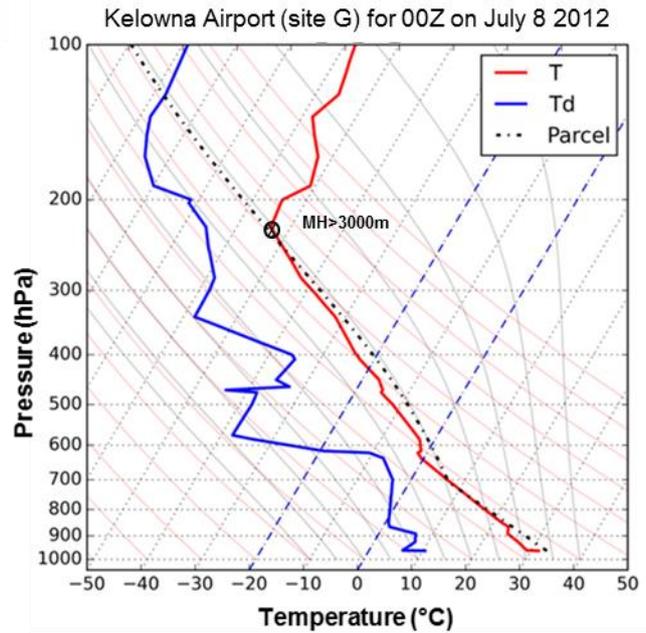
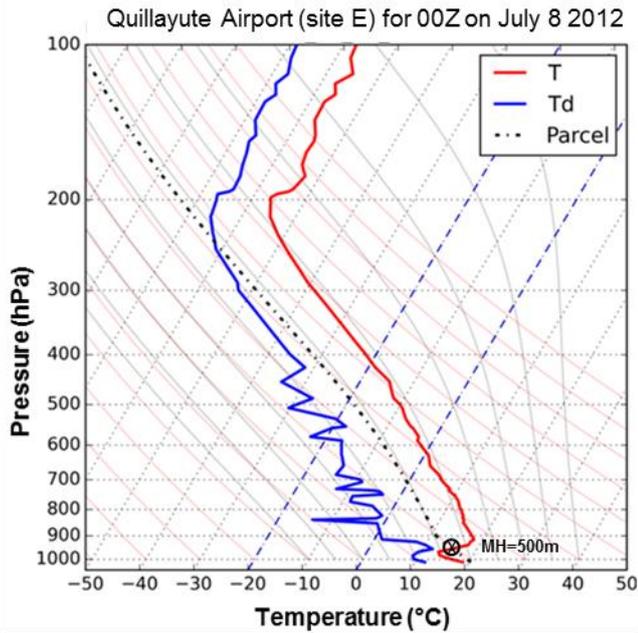
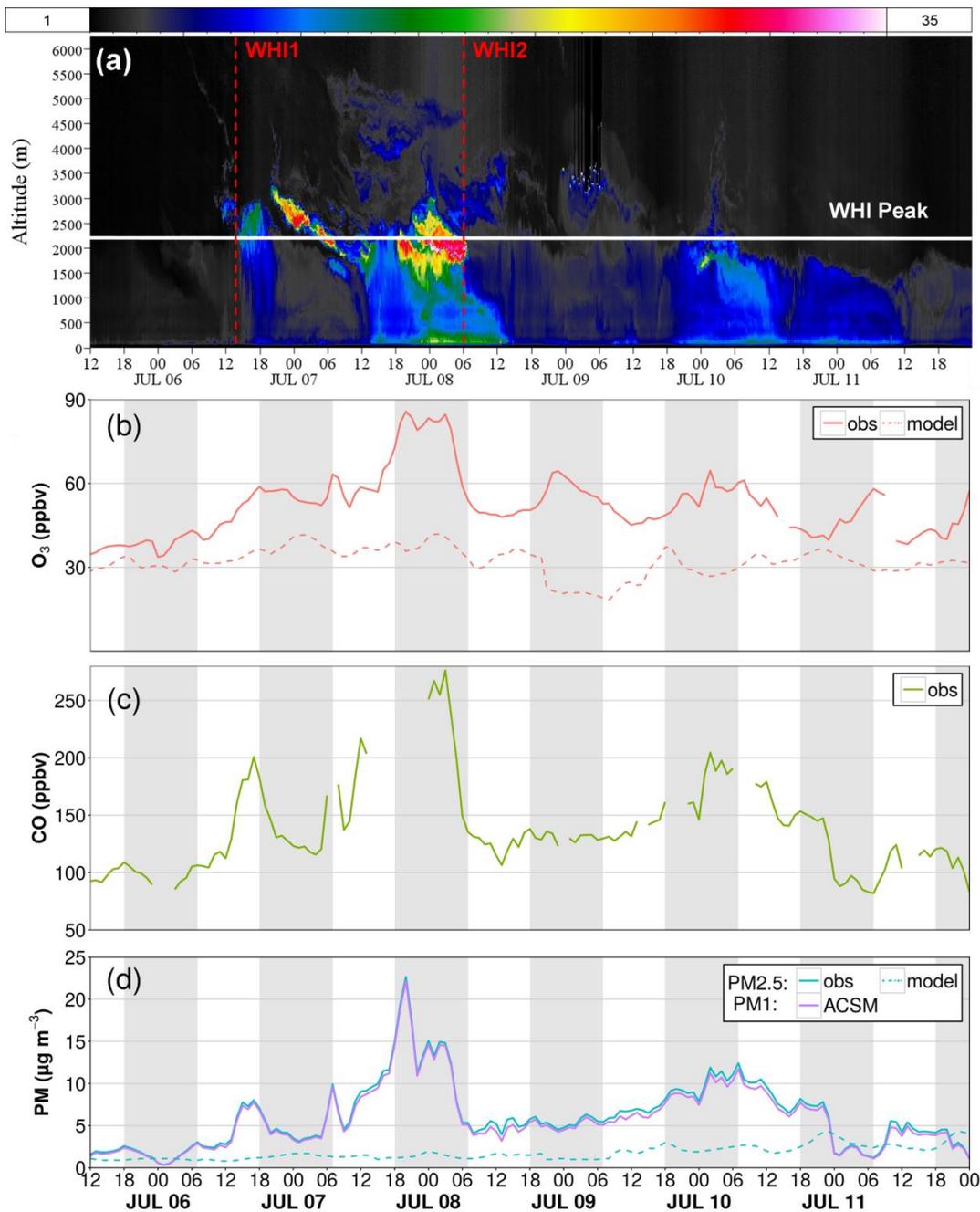


Fig. 4. Skew-T atmospheric radiosondes for Quillayute (site E) and Kelowna Airport (site G) at 00Z on July 8th, 2012. The atmospheric dry bulb temperature (T) and dewpoint temperature (Td) are shown from the surface up to 100 mb. Black dashed line is a trace of a surface parcel lift to estimate the mixed layer depth at the height where it intersects the environment dry bulb temperature.

5



5 **Fig. 5.** Whistler lidar (site A, 650 m ASL) measured backscatter ratio of the 1064 nm channel to that of clear air for the July 2012 Siberian smoke plume (a) and observed (solid) 1-hr O₃ (b), 1-hr CO (c), and 1-hr PM₁ (based on the ACSM) and PM_{2.5}. (d) at Whistler Peak High Elevation station (site 1, 2182 m ASL, indicated as a horizontal white line in panel a). AURAMS baseline (dashed) 1-hr O₃ and 1-hr PM_{2.5} are shown in panel b) and d), respectively. Elevated aerosol backscatter in panel (a) are noted from July 6 14:00 PST (WHI1) to July 8 06:00 PST (WHI2) by red dashed lines. The shaded regions indicate night time hours between 18:00 PST to 07:00 PST.

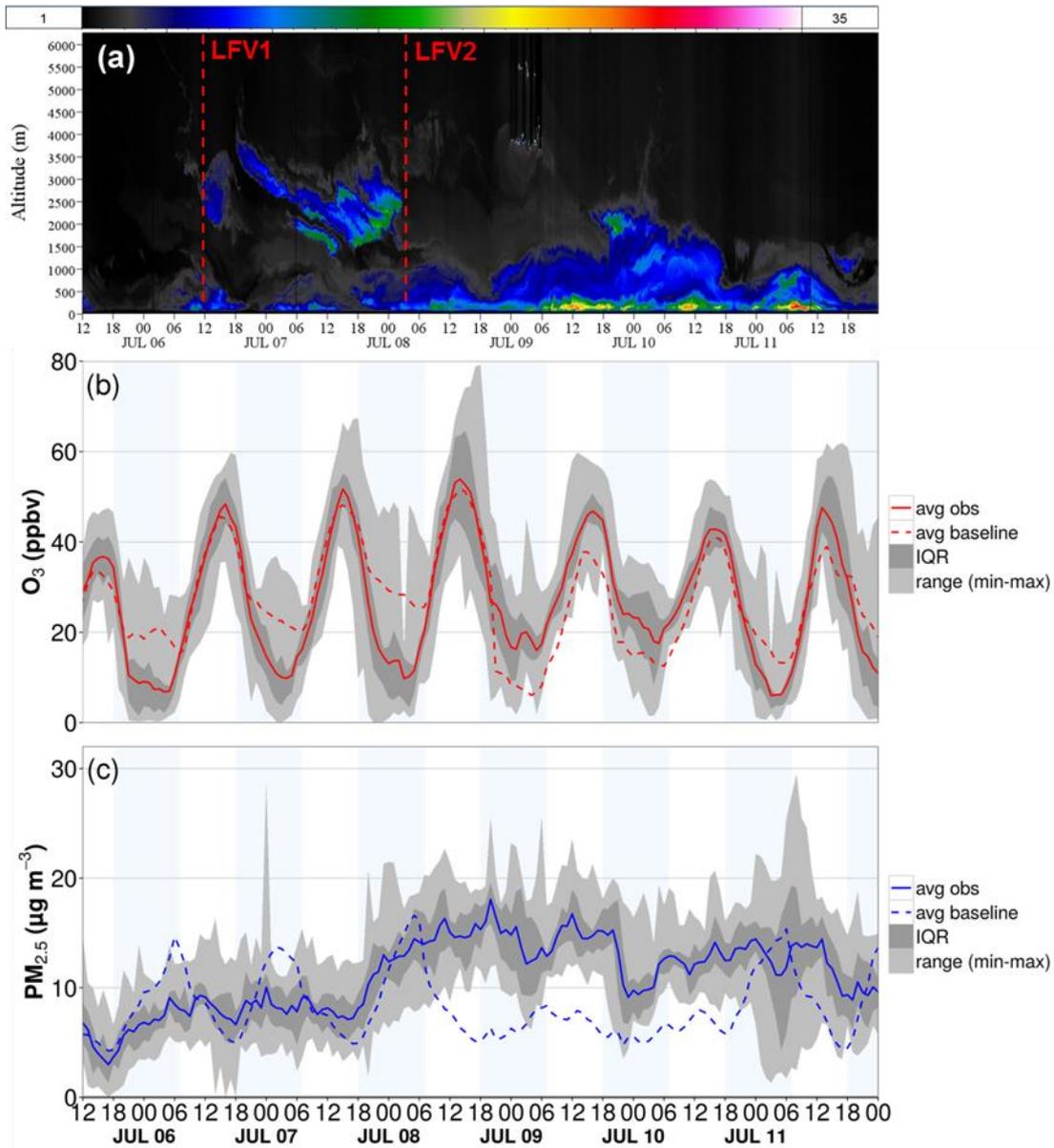


Fig. 7. UBC lidar (site 2) observations and air quality conditions across the Lower Fraser Valley network between July 6th and 12th, 2012. Panel (a) shows the ratio of measured backscatter of the 1064 nm channel to that of clear air. Panel (b) and (c) indicates the mean observed (solid) and AURAMS baseline without wildfire emissions (dashed) 1-hr O₃ (ppbv) and 1-hr PM_{2.5} (µg/m³) for 19 and 17 sites, respectively, across the LFV air quality monitoring network. The light and dark grey shading represents the range and inter-quartile range, respectively, across the network at a particular hour. The vertical shaded regions indicate night time hours between 18:00 PST to 07:00 PST.

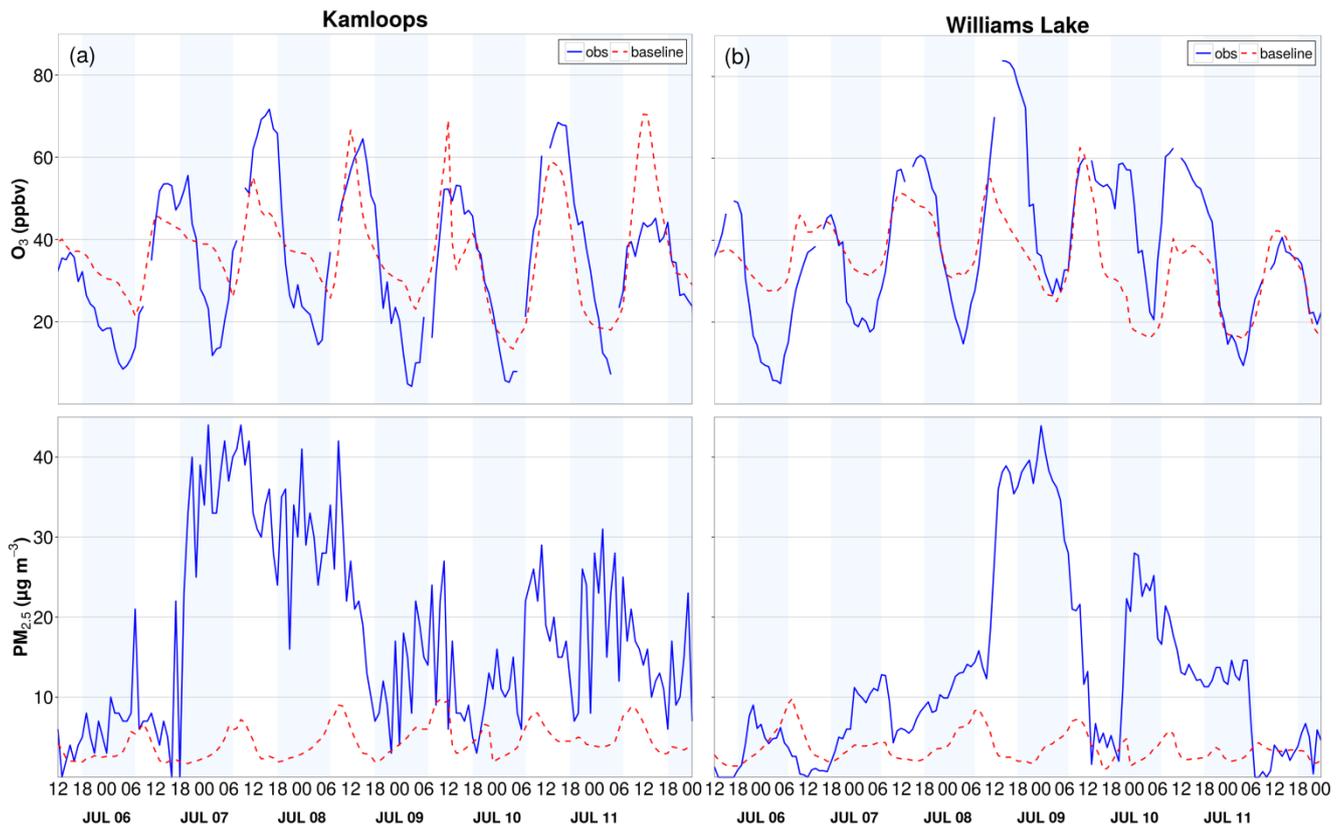


Fig. 9. The hourly observed (blue) and baseline AURAMS (red, dashed) without wildfire emissions O₃ (top) and PM_{2.5} (bottom) for Kamloops (site 8) (a) and Williams Lake (site 9) (b). The shaded regions indicate night time hours between 18:00 PST to 07:00 PST.

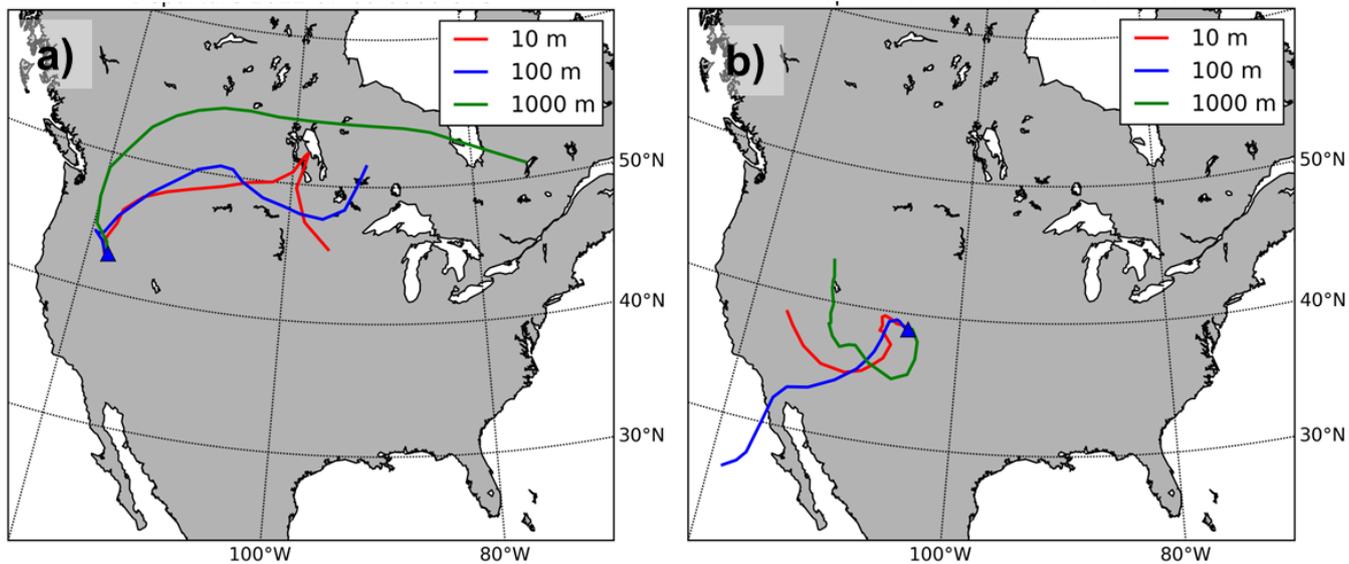


Fig. S1. Five day forward trajectories using the CMC Trajectory model from the Long Draw, Oregon (a) and Waldo Canyon, Colorado wildfires (b) released at 00UTC on July 9th, 2012 at heights of 10 m, 100 m, and 1 km AGL.

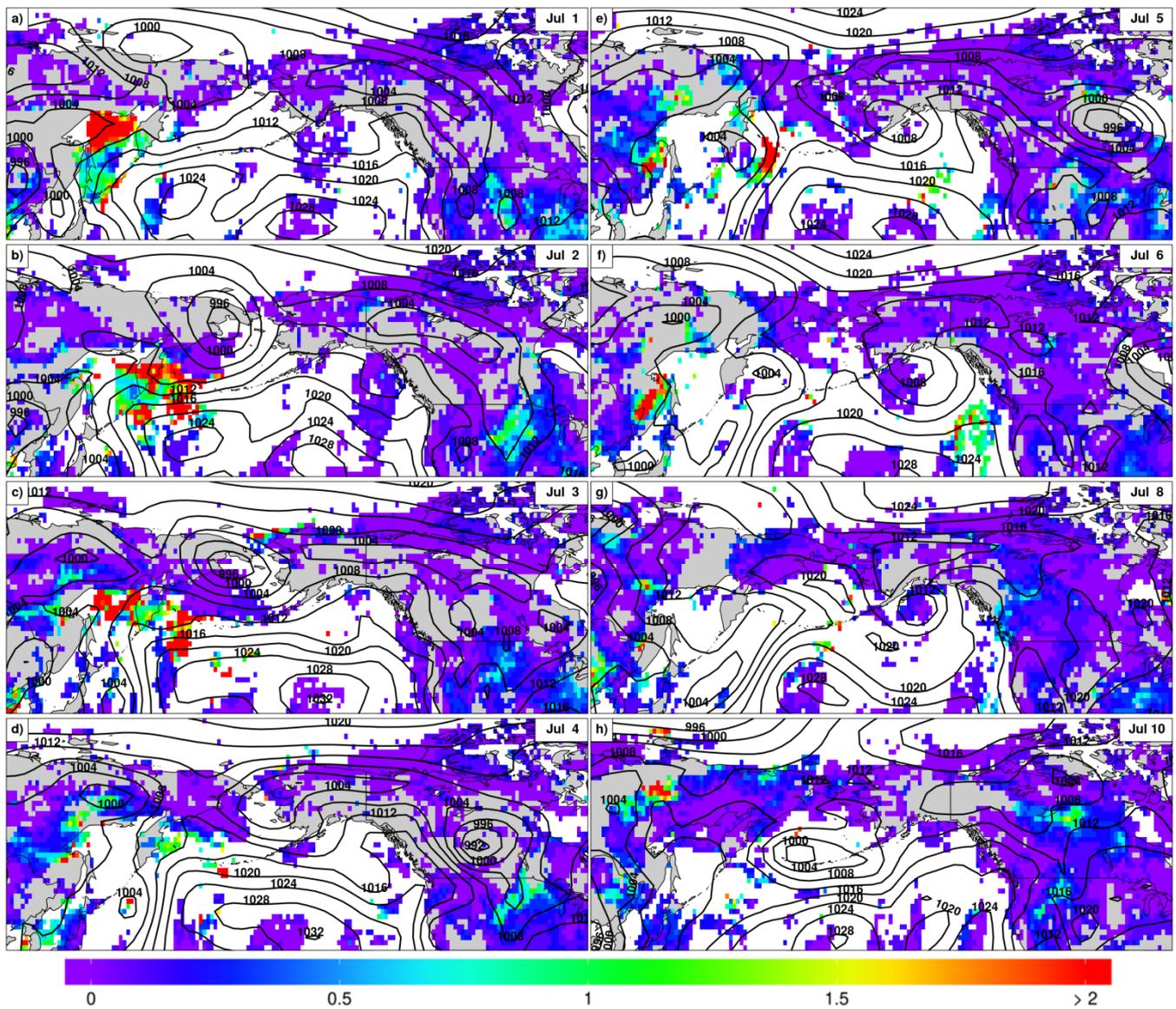


Fig. S2. The daily average MODIS Aerosol Optical Depth product for July 1-6 (a to f) illustrates the eastward progression of the Siberian Fire plume in early July onto the Western Pacific prior to its arrival off the coast of Washington and Oregon States on July 6th, 2012 (f). Enhanced AOD values spread across South and Central British Columbia by July 8th, 2012 (g) then shift southeastward out of the Pacific Northwest domain on July 10th, 2012 (h). Mean sea level pressure is contour (in solid black) on all panels.

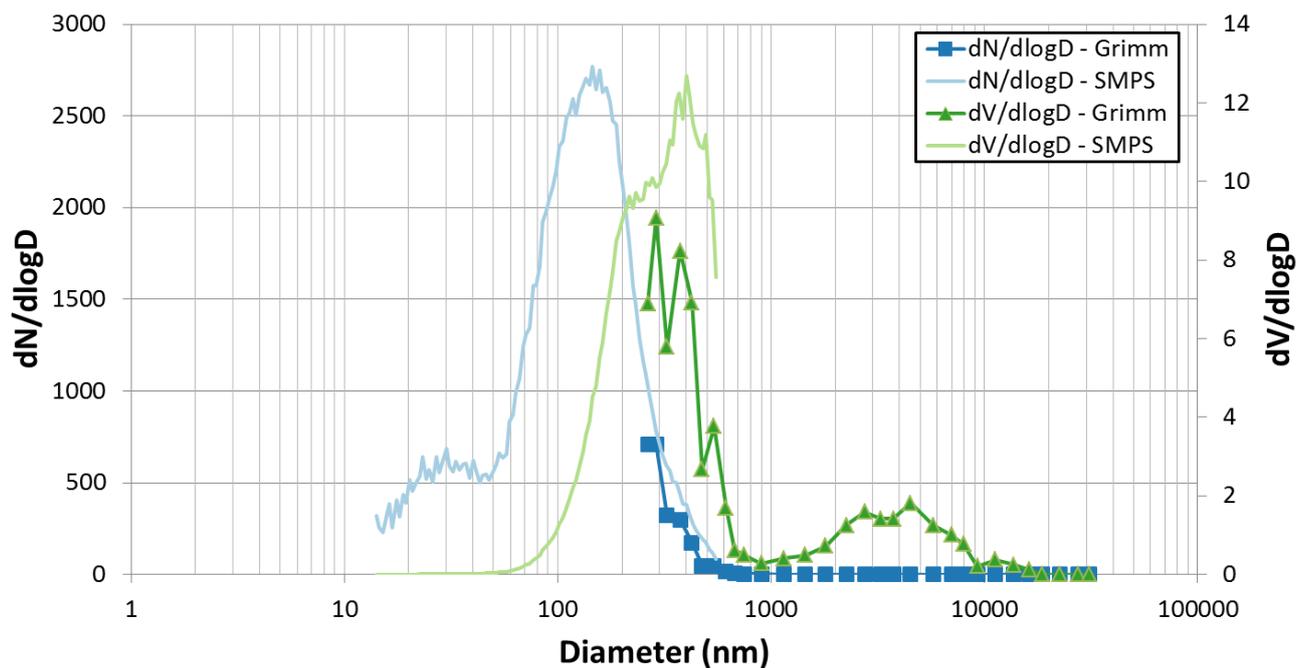


Fig. S3. Comparison of SMPS and GRIMM aerosol number and volume distributions for July 9th, 2012 16h30 -20h30 PST.

5

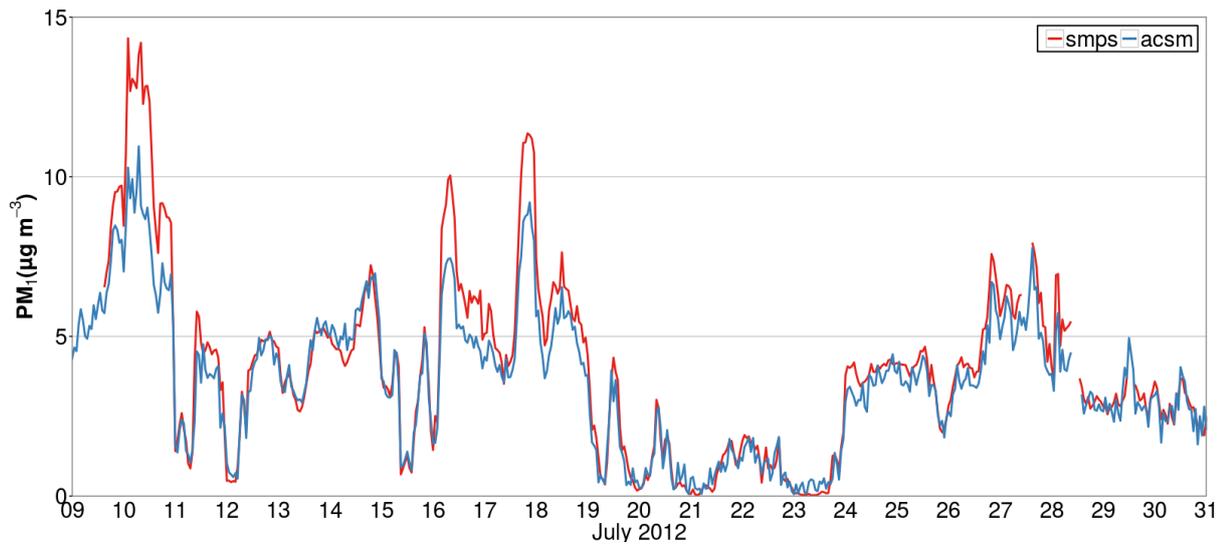
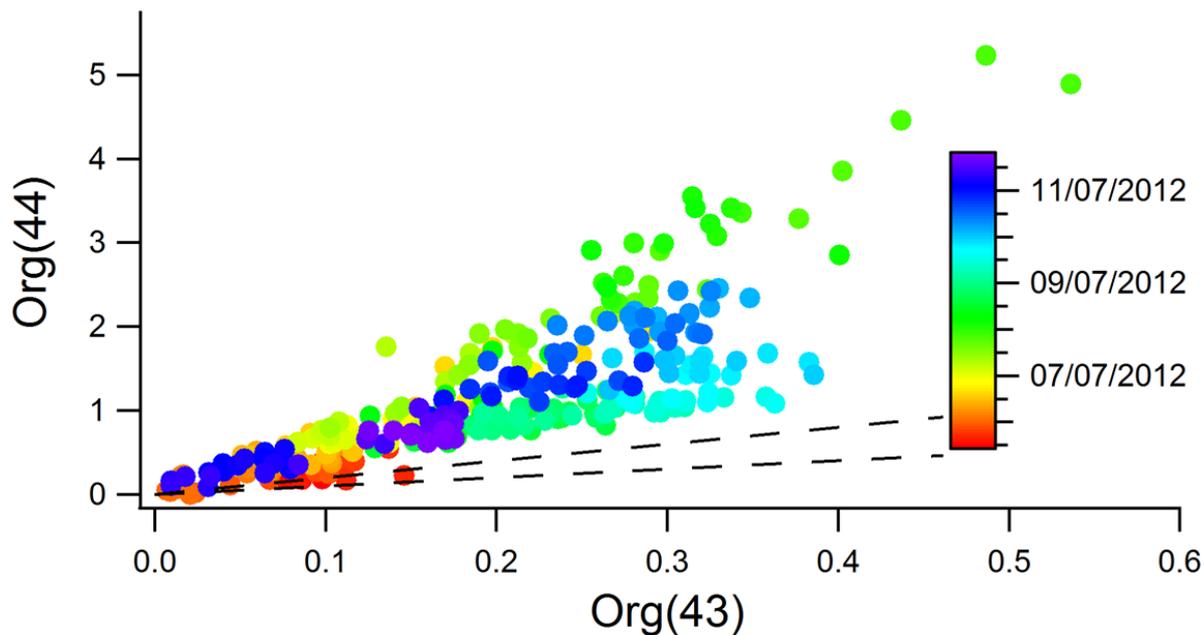


Fig. S4. Timeseries of SMPS PM₁ and ACSM readings between July 9th and July 31st, 2012. The ACSM data have
 5 been adjusted for a collection efficiency of 0.5.



5 Fig. S5. The $m/z44$ and $m/z43$ components of the organic aerosol spectrum derived from the ACSM spectra. The ratio of organic signal at $m/z44$ (Org43) vs. $m/z43$ (Org44) is compared to show that the organic aerosol was more oxygenated during the LRT smoke event (July 6th 14:00 PST to July 8th 06:00 PST) than at any other times during the 6 day period from July 5th 12:00 to 11th 18:00 PST. Dashed black lines mark the 1:1 and 2.5 ratio lines for reference.

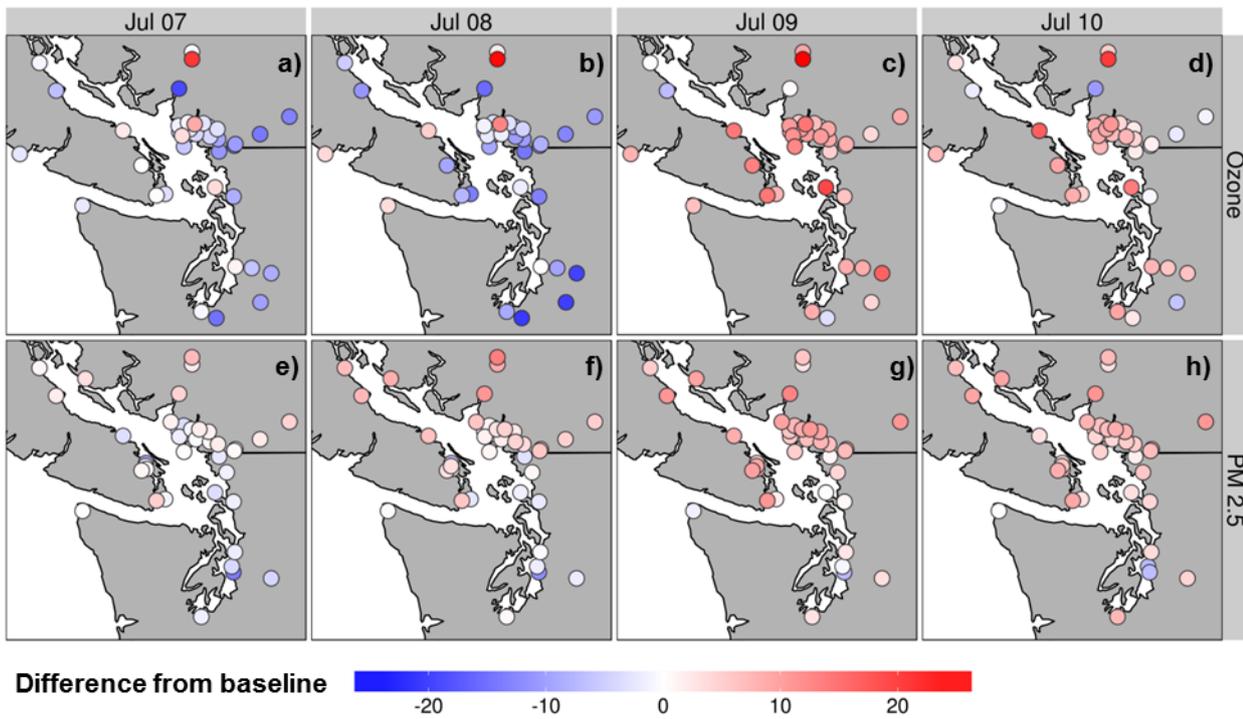
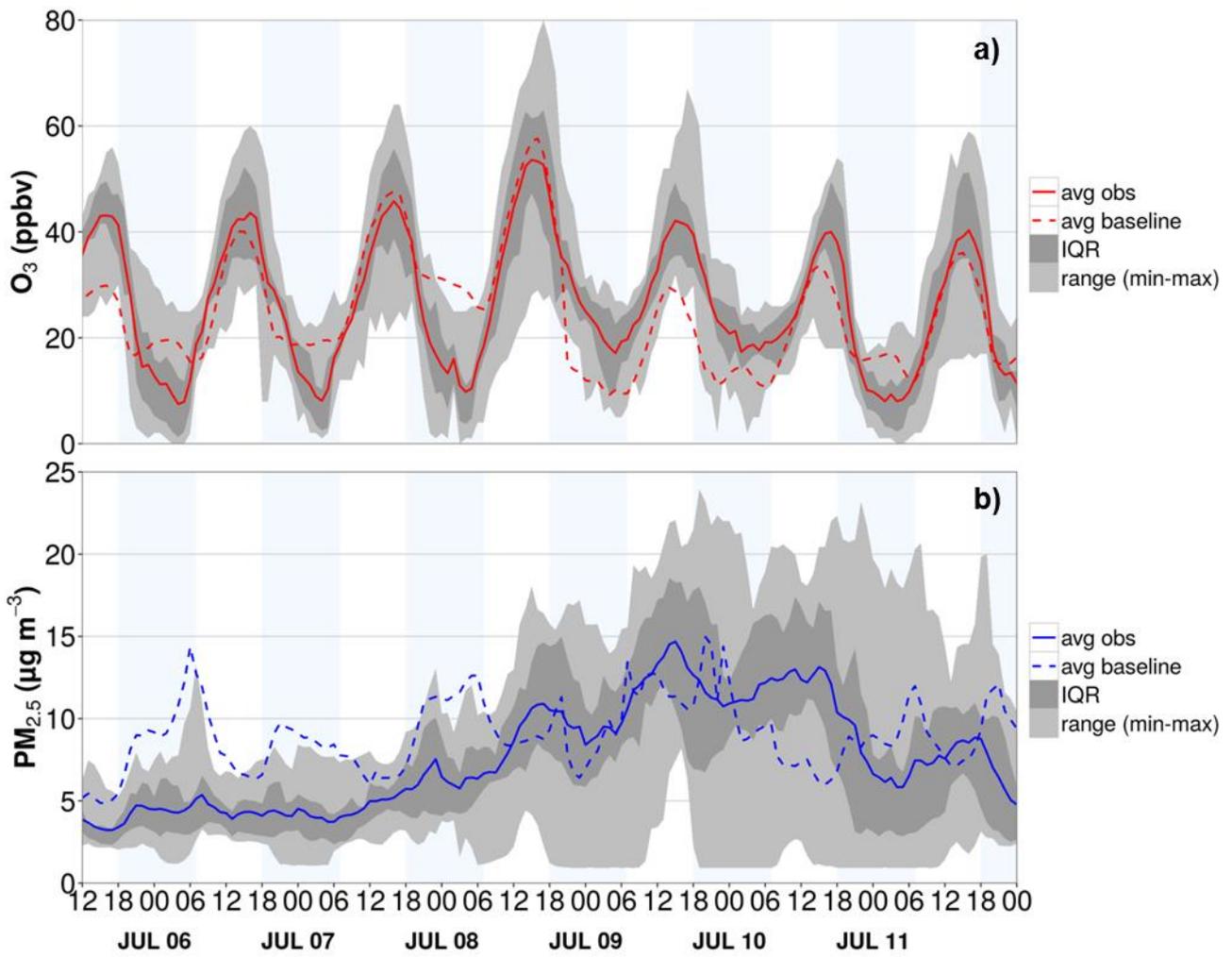


Fig. S6. Spatial map of the differences of the AURAMS baseline without wildfire emissions to daytime averaged 8-hr O₃ (ppbv, panels a, b, c, d) and of the daily averaged PM_{2.5} (µg/m³, e, f, g, h) from July 7th, 2012 to July 10th, 2012.



5 **Fig. S7. Hourly average observed (solid) and AURAMS baseline (dashed) without wildfire emissions: hourly O_3 (a) and $PM_{2.5}$ (b) for sites across the Washington State air quality monitoring network. The light grey and dark grey shading represents the range and inter-quartile range (IQR) across the network at a particular hour. The shaded regions indicate night time hours between 18:00 PST to 07:00 PST.**

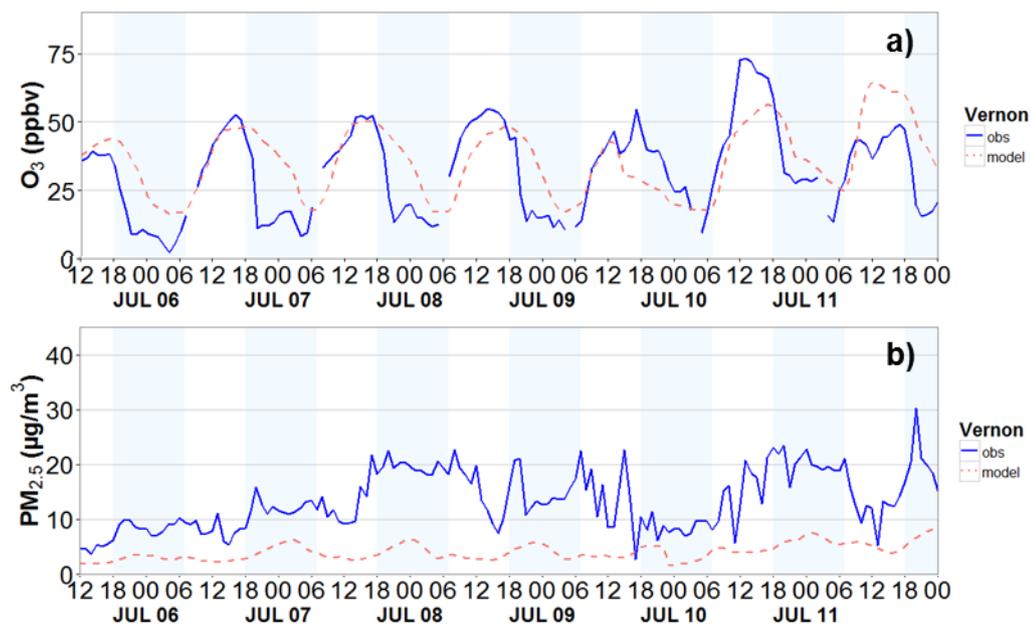


Fig. S8. The observed and modelled baseline hourly O_3 (a) and $PM_{2.5}$ (b) for Vernon (site 7). The shaded regions indicate night time hours between 18:00 PST to 07:00 PST.