



1 **The concentration, source apportionment and deposition**
2 **flux of atmospheric particulate inorganic nitrogen**
3 **during dust events**

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11 **Abstract.** To understand the impacts of long-range transport on particulate inorganic nitrogen
12 associated with dust in downwind areas, aerosol samples were collected in the Qingdao coastal region
13 on dust and non-dust (ND) days in spring from 2008 to 2011. The concentrations of water-soluble ions
14 were measured by ion chromatography, with metal elements detected using inductively coupled
15 plasma-atomic emission spectroscopy (ICP-AES) and inductively coupled plasma-mass spectrometry
16 (ICP-MS). Compared to atmospheric aerosols collected on ND days, samples from dust days exhibited
17 higher concentrations of particles and crustal elements. Total aerosol particle concentrations increased
18 by a factor of 5.9 on average dust days. On dust days, the average concentrations of crustal elements
19 (Sc, Al, Fe, Ca and Mg) increased by over a factor of four relative to those on ND days. The inorganic
20 nitrogen content increased 1.2 to 9.2-fold during some dust events in which storms were weak or slow
21 moving and reactions occurred during transport. By contrast, nitrate and ammonium exhibited very low
22 concentrations (<20% of ND samples) or decreased concentrations in some cases as a result of the
23 strong dilution effect of low-nutrient dust particles arising from their rapid transport in a strong dust
24 storm. If air masses traveled faster than 40.5 ± 9.9 km/h, the inorganic nitrogen content of most aerosol
25 samples decreased compared to that of ND samples because of the strong dilution effect. The
26 concentration of atmospheric particulate inorganic nitrogen was related to not only the transport path
27 and speed but also the local emissions and reaction conditions during transport. The positive matrix
28 factorization (PMF) receptor model results showed that the contribution of soil dust dramatically
29 increased from 23% to 36% (90% of the residuals < 3.0 and $r^2=0.97$) on dust days, while the
30 contributions of local anthropogenic sources decreased, especially that of secondary aerosols. The dry



31 deposition flux of atmospheric particulates increased from 2800 ± 700 mg/m²/month on ND days to
32 $16,800 \pm 15,900$ on dust days. The dry deposition flux of particulate inorganic nitrogen increased 1.1 to
33 5.8-fold under the weak dilution effects of dust events. The dry deposition flux of nitrate decreased by
34 46%-63%, while that of ammonium decreased by 14% or to ND levels when strong dilution occurred
35 during dust events. The atmospheric input of nitrogen to the ocean was not enhanced by dust events,
36 and dust deposition was an uncertain source of nitrogen to the ocean.

37 Keywords: aerosols, nitrogen, dust, source apportionment, dry deposition flux

38 **1 Introduction**

39 Nitrogen is a key element for marine phytoplankton growth. Nitrogen carried in dust particles can be
40 transported and deposited across vast distances at high wind speeds (Duce et al., 2008; Zhang et al.,
41 2010). Additionally, bioavailable nutrients, via dust particle deposition, may enhance phytoplankton
42 growth in some ocean areas (Shi et al., 2012; Guo et al., 2012; Liu et al., 2013). Tan and Wang (2014)
43 found that chlorophyll concentrations increased four-fold, and a phytoplankton bloom occurred 10-13
44 days after dust deposition. Banerjee and Kumar (2014) hypothesized that dust-induced episodic
45 phytoplankton blooms are important to the interannual variability of chlorophyll in the Arabian Sea
46 away from active winter convection. Therefore, the atmospheric particulate nitrogen delivered by dust
47 is important for obtaining a better understanding of the effects of dust inputs on marine primary
48 production.

49 Asian dust, one of the main components of dust worldwide, affects northern China and the eastern
50 area of the China Sea. Asian dust can also affect the northern Pacific Ocean via long-range transport by
51 west winds. During transport, the dust storm particles continually mix with anthropogenic gas and
52 particles from local sources along the pathway to carry desert dust and anthropogenic aerosols to
53 downwind areas. The physical and chemical characteristics of atmospheric particulates downwind of
54 dust weather are remarkably different from those normally present (Yang et al., 2002; Li et al., 2014;
55 Ma et al., 2012). Additionally, the concentrations and characteristics of atmospheric particulate
56 inorganic nitrogen species in downwind areas are greatly affected by dust weather.

57 Some researchers have found that inorganic nitrogen species in aerosols have high concentrations
58 during Asian dust events. The concentrations of atmospheric particulate NO₃⁻ and NH₄⁺ on dust storm



59 days were 2-5 times higher than those on non-dust storm (ND) days in Beijing (Liu et al., 2014; Liu
60 and Bei, 2016). Xu et al. (2014) found that particulate SO_4^{2-} and NO_3^- simultaneously increased in dust
61 storms on the northern boundary of the Tibetan Plateau because of the enriched dust including more
62 acidic species or anthropogenic aerosols. Compared to those on ND days, higher concentrations of
63 NO_3^- and NH_4^+ in aerosol particles were observed on dust storm days in northern China, and NO_3^- and
64 NH_4^+ showed lower concentrations during strong dust storm events than during weak dust events
65 (Zhang et al., 2010). Fitzgerald et al. (2015) found that nearly all Asian dust in Korea contains
66 considerable amounts of nitrate because pollution plumes mix with dust from the Gobi and Taklamakan
67 Deserts and are transported over the Asian continent.

68 However, some studies observed that the concentrations of inorganic nitrogen in aerosols decreased
69 on dust storm days. At Yulin, a rural site near the Asian dust source region, the concentration of NO_3^-
70 significantly decreased in aerosols on dust days as a result of the dilution effect (Wang et al., 2016).
71 The absolute abundances of NO_3^- and NH_4^+ were notably lower in dust plumes than in a polluted air
72 parcel because dust plumes are often separated by the arrival of a cold front in Shanghai, China (Wang
73 et al., 2013). Li et al. (2014) found that the concentrations of nitrate and ammonium in downwind
74 aerosol particles decreased on dust storm days, with a decreasing ratio of soluble inorganic ions to
75 $\text{PM}_{2.5}$ in the Yellow River Delta, China. When dust is rapidly transported from desert regions without
76 passing through a major urban area and lingers over the Yellow Sea, the concentrations and size
77 distributions of nitrate and ammonium have no significant variation in heavy Asian dust (AD) plumes
78 (Kang et al., 2013). Nitrate and ammonium also exhibited different concentration variations in other
79 desert regions. Jaafar et al. (2014) found that nitrate was more abundant than ammonium, which
80 showed no concentration variation in non-dust aerosol particles during dust episodes originating from
81 both the African and Arabian deserts.

82 The effect of dust events on the inorganic nitrogen concentration in downwind aerosols is
83 complicated because it involves many factors, such as the mixing state and transport pathway. The
84 effect of AD as a nitrogen source on biogeochemical cycles and marine ecology is not adequately
85 understood. Understanding the variations in the concentration and deposition flux of atmospheric
86 particulate nitrogen on dust days is essential to quantifying the impacts of nutrients on the marine
87 environment and primary production. To understand the influence of dust on atmospheric nitrogen, we
88 collected aerosol samples from the coastal region of the Yellow Sea in spring, when there is a high



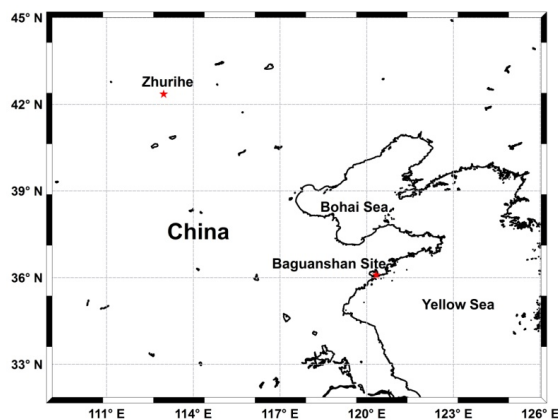
89 frequency of dust storms, from 2008 to 2011. Then, we analyzed the inorganic nitrogen concentrations
90 in the aerosol samples. The concentration, source apportionment and deposition flux of atmospheric
91 particulate inorganic nitrogen in dust events were examined.

92 2 Materials and methods

93 2.1 Sampling

94 As shown in Fig. 1, total suspended particles (TSP) were collected at the Baguanshan site in the
95 coastal region of the Yellow Sea. The samples were collected with quartz microfiber filters (Whatman
96 QM-A) using a high-volume air sampler (Model KC-1000, Qingdao Laoshan Electronic Instrument
97 Complex Co., Ltd.) with a flow rate of $1 \text{ m}^3/\text{min}$ on the roof of an office building ($36^\circ 6' \text{ N}$, $120^\circ 19' \text{ E}$,
98 77 m above sea level) approximately 1.0 km from the shore. The filters were heated at 450°C for 4.5 h
99 to remove organic compounds. Samples were collected on dust days and selected ND days in spring
100 from March 2008 to May 2011, with a sampling duration of 4 h for each sample. We refer to the ND
101 days as sunny and cloudy days before or after dust events in the following discussion.

102 The sand samples were collected at the Zhurihe site ($42^\circ 22' \text{ N}$, $112^\circ 58' \text{ E}$) in the Hunshandake Desert,
103 one of main Chinese sand deserts, in April 2012. After sand samples were packed in clean plastic
104 sample bags, the samples were stored below -20°C .



105
106 **Figure 1.** Location of the aerosol and dust sampling site

107 2.2 Analysis

108 The aerosol samples were balanced in a relative humidity- and temperature-controlled chamber until



109 the particle weights remained constant. The mass concentrations of TSP were calculated according to
 110 the particle masses and the sampling volume. The sample membranes were then cut into several
 111 portions for analysis.

112 One portion of each aerosol sample and 0.1 g of parallel sand sample were ultrasonically extracted
 113 with ultra-pure water in an ice water bath, and the concentration of inorganic water-soluble ions was
 114 determined via ICS-3000 ion chromatography (Qi et al., 2011). The parallel sand samples collected
 115 from the Hunshandake Desert were analyzed using the same procedure.

116 One portion of each aerosol filter was cut into 60 cm² pieces and digested with HNO₃+HClO₄+HF
 117 (5:2:2 in volume) at 160°C using an electric heating plate. A blank membrane was also analyzed using
 118 the same procedure to ensure analytical precision. Cu, Zn, Cr, Sc and Pb were measured by inductively
 119 coupled plasma-mass spectrometry (Thermo X Series 2), while Al, Ca, Fe, Na and Mg were detected
 120 by inductively coupled plasma-atomic emission spectroscopy (IRIS Intrepid II XSP). The metal
 121 concentrations were determined by calibrating the measured concentrations of samples using
 122 membrane blanks.

123 **Table 1.** Detection limits, precisions and recoveries of water-soluble ions and metal elements

Component	Measurement method	Detection limit (µg·L ⁻¹)	Precision (RSD%)	Recovery (%)
NO ₃ ⁻	IC	2.72	1.54	97
SO ₄ ²⁻		1.62	1.55	98
NH ₄ ⁺		0.4	1.10	97
Ca ²⁺		0.44	0.79	94
Cu	ICP-MS (Xin et al., 2012)	0.006	4.0	106
Zn		0.009	2.5	102
Cr		0.004	3.0	95
Sc		0.002	2.4	97
Pb		0.008	3.9	104
Al	ICP-AES (Lin et al., 1998)	7.9	0.6	103
Ca		5.0	1.2	99
Fe		2.6	0.7	104
Na		3.0	0.6	99
Mg		0.6	0.6	105
Hg	CVAFS	0.0001	6.6	105
As	CVAFS	0.1	5.0	98

124 One portion of each aerosol sample was digested with HNO₃ solution (10% HNO₃, 1.6 M) at 160°C
 125 for 20 min in a microwave digestion system (CEM, U.S.). Hg and As in sample extracts were analyzed



126 following the U.S. Environmental Protection Agency method 1631E (U.S. EPA, 2002) using cold vapor
127 atomic fluorescence spectrometry (CVAFS).

128 The detection limits, precisions and recoveries of water-soluble ions and metal elements are listed in
129 Table 1.

130 **2.3 Computational modeling**

131 To determine the origins of sampled air masses, the 72 h air mass back trajectories were calculated
132 for each TSP sample using TrajStat software (Wang et al., 2009) and the NOAA GDAS archive data
133 (<http://www.arl.noaa.gov/ready/hysplit4.html>). The air mass back trajectories were calculated at an
134 altitude of 1000 m.

135 The positive matrix factorization (PMF) receptor modeling method (Paatero and Tapper, 1993;
136 Paatero, 1997) was used to obtain the source apportionment of atmospheric particulates on dust
137 and ND days. The correlation coefficient between the predicted and observed concentrations was
138 0.97.

139 Dry deposition velocities were obtained using Williams' model and accounting for particle
140 growth (Qi et al., 2005). Relative humidity, air temperature and U10 from the National Centers for
141 Environmental Prediction (NCEP) were used in the model as the meteorological inputs. Surface
142 seawater temperature was collected from the European Centre for Medium-Range Weather
143 Forecasts (ECMWF). The climatic and seawater temperature data had a six-hour resolution.
144 According to a previously reported method (Qi et al., 2013), the dry deposition fluxes of the
145 particles and the nitrogen species were calculated for dust and ND days.

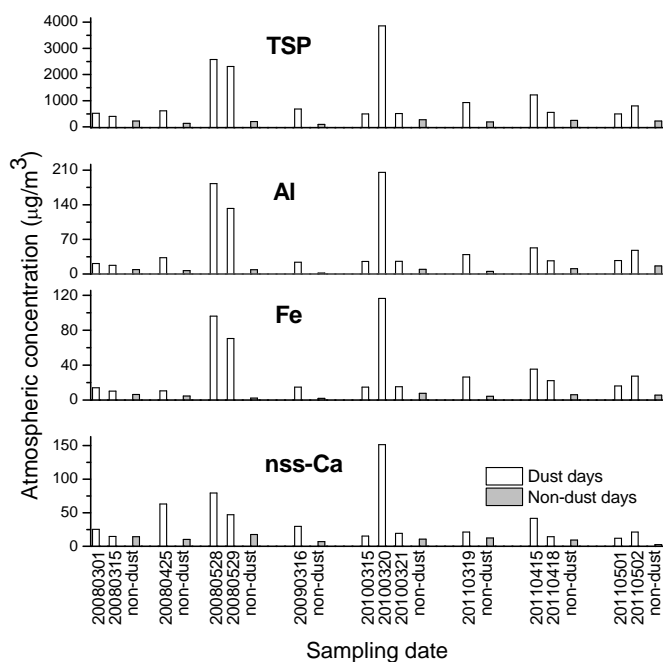
146 **2.4 Statistical analysis**

147 Meteorological data were obtained from the Qingdao Meteorological Administration
148 (<http://qdqx.qingdao.gov.cn/zdz/ystj.aspx>) and the Meteorological Information Comprehensive
149 Analysis and Process System (MICAPS) of the Meteorological Administration of China. NO₂ and air
150 quality index (AQI) data were downloaded from the Qingdao Environmental Protection Bureau
151 (<http://www.qepb.gov.cn/m2/>). Different weather characteristics, such as sunny days, cloudy days and
152 dust days, were defined according to MICAPS information.



153 **3 Results and discussion**

154 **3.1 Characterization of aerosol samples collected during dust events**



155

156 **Figure 2.** Concentrations of TSP, Al, Fe and nss-Ca in aerosol samples collected in the coastal region of the Yellow
157 Sea on non-dust and dust days from 2008 to 2011

158 To understand the impact of dust events on atmospheric particulate nitrogen in downwind areas, we
159 collected aerosol samples in the coastal region of the Yellow Sea in the spring from 2008 to 2011. We
160 examined the concentrations of particles and crustal and anthropogenic metals in aerosol samples.
161 Although metal concentrations in aerosols collected on different dust days varied, some
162 characterizations are highlighted below. The concentration of atmospheric particulate increased on dust
163 days. On non-dust (ND) days, aerosol particles varied in the range of 94-275 $\mu\text{g}\cdot\text{m}^{-3}$, with an average of
164 201 $\mu\text{g}\cdot\text{m}^{-3}$ (Fig. 2). Particle concentrations increased to 501-3857 $\mu\text{g}\cdot\text{m}^{-3}$ on dust days. The TSP
165 concentration on dust days was 1.8-14.0 times (mean: 5.9) higher than that on ND days. The crustal
166 elements increased considerably with the increasing particle concentrations when dust events occurred.
167 As shown in Table 2, the enrichment factors (EF) of Al, Fe, Ca, and Mg were lower than ten on ND
168 days and decreased to less than three on dust days. These data are indicative of the primarily crustal



169 origins of these elements. We found that the mean concentrations of Sc, Al, Fe, Ca and Mg increased
170 by over a factor of four as compared to those on ND days. Al concentrations in dust weather increased
171 1.7 to 21.9-fold (mean: 6.9) on ND days. The Al concentration of the “geometric mean \times 2GSD” (where
172 GSD is the geometric standard deviation) was used as a criterion to define major AD events in areas of
173 East Asia (Hsu et al., 2008). Al concentrations were higher than the criterion level in all dust samples,
174 which indicated that the samples we collected on dust days were truly affected by dust events. Fe was
175 10.3 times higher on dust days than on ND days. Additionally, nss-Ca, a typical dust index, increased
176 3.6-fold on dust days (Fig. 2). The EF of the anthropogenic metal elements decreased on dust days. Cu,
177 Pb, Zn, Cr, Hg and As had high EFs, much greater than 10, on ND days, which indicated that these
178 metal elements were mainly from anthropogenic sources. The concentrations of these anthropogenic
179 elements on dust days increased 1.1 to 7.2-fold on average compared to those on ND days. Additionally,
180 the EFs of these anthropogenic elements decreased on dust days. These data are consistent with the
181 very low EFs of these elements in dust particles. Thus, the influence of anthropogenic sources on
182 atmospheric particulates decreased on dust days.

183 **Table 2.** The average concentrations and EFs of metal elements on dust and non-dust days

Element	Concentration (ng/m ³)		EF*	
	Non-dust days	Dust days	Non-dust days	Dust days
Sc	1.11	13.90	-	-
Al	8.53 \times 10 ³	6.86 \times 10 ⁴	3.8	1.4
Fe	4.91 \times 10 ³	3.88 \times 10 ⁴	3.	1.2
Ca	1.05 \times 10 ⁴	4.29 \times 10 ⁴	14.0	2.1
Mg	1.62 \times 10 ³	1.58 \times 10 ⁴	3.5	1.1
Cu	50.2	124.5	36.3	6.1
Pb	127.9	221.0	389.4	56.1
Zn	340.0	457.7	248.9	20.6
Cr	33.8	244.0	44.0	11.1
Hg	0.26	0.36	176.0	13.8
As	25.5	27.4	707.2	43.9

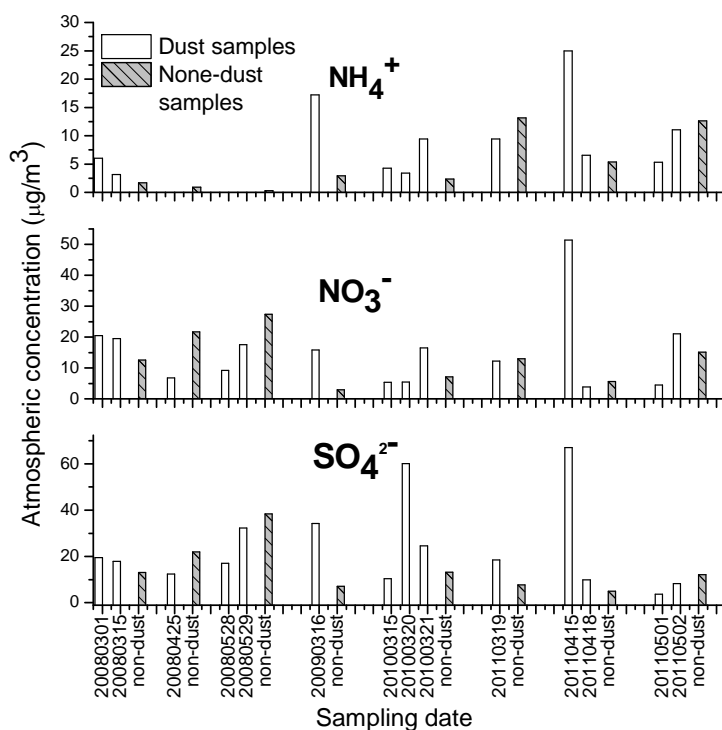
184 *The EF factor was calculated using Scandium as the reference element (Han et al., 2010).

185 3.2 Concentration distribution of inorganic nitrogen in dust events

186 As discussed above, the concentrations of TSP and metal elements increased on dust days compared
187 to those on ND days. However, the concentrations of inorganic nitrogen species NH₄⁺ and NO₃⁻
188 exhibited different variations in particles and crustal metal elements. The concentration of ammonium



189 increased by a factor of 1.2-5.7 on some dust days and decreased or had a very low concentration (less
190 than 20% of that on ND days) on other dust days (Fig. 3). Similar to ammonium, nitrate displayed two
191 different concentration variations on dust days. Nitrate concentrations increased by a factor of 1.4-9.2
192 on some dust days and decreased on other dust days. The secondary inorganic ion SO_4^{2-} exhibited
193 concentration variations that were similar to those of nitrate. Therefore, the influence of dust on these
194 secondary ions was different from that on crustal metal elements, and the effect of dust on inorganic
195 nitrogen differed during different types of dust events.



196

197 **Figure 3.** NH_4^+ , NO_3^- and SO_4^{2-} in aerosol samples collected in the coastal region of the Yellow Sea on non-dust
198 and dust days from 2008 to 2011

199 When we focused on inorganic nitrogen (IN), we found that IN concentrations could be grouped into
200 three cases (Table 3). IN concentrations were higher on dust days than on ND days for Case 1, while IN
201 was lower on dust days for Case 2. For Case 3, nitrate concentrations on dust days were less than on
202 ND days, while ammonium concentrations on dust days were slightly higher than those on ND days. To
203 understand the influence of dust on the nitrogen concentration, we compared the IN content in aerosols



204 from the coastal region of the Yellow Sea with sand particles and atmospheric aerosols from Duolun, a
 205 site very close to the Zhurihe Sand Desert. The Yellow Sea is mainly affected by dust storms from this
 206 sand source (Zhang and Gao, 2007). From Table 4, we found that nitrate and ammonium concentrations
 207 in the source sand particles were very low (less than 50 $\mu\text{g/g}$). Therefore, the dust particles in this
 208 source area that affect the Yellow Sea are nutrient poor. Although the IN content in aerosols at Duolun
 209 was higher than that in sand particles, the nitrate and ammonium concentrations were much lower than
 210 in the coastal region of the Yellow Sea. Therefore, we believe that the dust particles from the source
 211 have a dilution effect on atmospheric particulate nitrogen because of the low IN concentration in sand
 212 particles. When dust events occurred, the content of nitrogen per particle mass decreased because of the
 213 dilution of particulate nitrogen resulting from the increased number of nutrient-poor dust particles
 214 rapidly leaving the source area. The dilution effect depends on the intensity of dust events and the
 215 distance from the dust source. The stronger a dust storm is and the closer to the source, the stronger the
 216 dilution effect is.

217 **Table 3.** Average concentrations of inorganic nitrogen, TSP, NO_x, Relative Humidity (RH) and T for each case in
 218 aerosol samples in the coastal region of the Yellow Sea

	Sample numbers	TSP $\mu\text{g}\cdot\text{m}^{-3}$	NO ₃ ⁻ $\mu\text{g}\cdot\text{m}^{-3}$	NH ₄ ⁺ $\mu\text{g}\cdot\text{m}^{-3}$	RH %	T °C	NO _x $\mu\text{g}\cdot\text{m}^{-3}$	Summary
Case 1	080301, 080315, 090316, 100321, 110415, 110502	696	24.1	14.9	46.6	13.8	62.7	IN>ND
Case 2	080425, 080528, 080529, 110319, 110501	1199	3.1	2.6	29.2	19.8	52.3	IN<ND
Case 3	100315, 100320, 110418	1639	4.9	4.7	10.1	10.1	70.7	NO ₃ ⁻ <ND NH ₄ ⁺ ≅ND
Non-dust		212	5.5	4.6	42.2	13.7	59.7	

219 During transport, the concentration of IN will increase by reacting with gas emitted into the air under



220 appropriate reaction conditions or by mixing with polluted aerosol particles from a local source.
221 Therefore, IN concentrations will increase in aerosols in downwind areas because of reactions on the
222 dust surfaces or mixing with anthropogenic particles along the transport path. If no effective emission
223 or absorption and reaction occur during transport, the IN content per particle mass ($\mu\text{g/g}$) will decrease
224 in atmospheric aerosols in the downwind area. In Case 1, the particle concentration in the dust events
225 was less than $700 \mu\text{g/m}^3$, and the IN concentration in atmospheric aerosols increased by a factor of
226 more than three in dust events in the coastal region of the Yellow Sea, which might be a result of slow
227 transport or a weak dilution effect. High relative humidity (RH), low temperature and high NO_x
228 transport over a long distance and at a low speed are beneficial to the formation of nitrate and
229 ammonium. In Cases 2 and 3, the particle concentrations were very high, with an average higher than
230 $1100 \mu\text{g-m}^{-3}$, which indicated that the samples were affected by a strong dust storm or that the dust
231 might be transported quickly. Concentrations of IN in aerosols in dust events decreased at the
232 downwind site in Case 2 because the low RH, high temperature and low NO_x during rapid transport
233 were not advantageous to the formation of IN. In Case 3, the low IN content was a result of a strong
234 dilution effect and low RH. In addition, the transport path affected the IN content of aerosol particles in
235 the downwind area, and this influence will be discussed in Section 3.3.

236 **Table 4.** Comparison of the IN content in dust particles according to the dust source region (unit: $\mu\text{g/g}$)

Sands sampled in Zhurihe		Aerosols in Duolun*		Aerosols in the coastal region of the Yellow Sea	
NO_3^-	NH_4^+	NO_3^-	NH_4^+	NO_3^-	NH_4^+
25.46±22.87	4.21±1.03	1200	900	Non-dust: 28,200±24,819	Non-dust: 24,063±21,515
				Case 1: 34,892±9570	Case 1: 22,571±7016
				Case 2: 5542±5117	Case 2: 4758±5698
				Case 3: 6359±4697	Case 3: 7059±5591

237 * Adapted from Cui (2009)

238 3.3 Influence of transport on particulate inorganic nitrogen



239 To understand the influence of transport on atmospheric particulate IN, we analyzed the air mass
240 trajectory of each sample (Fig. 4). The results showed that all dust samples were collected from north
241 or northwest air masses. The reported threshold of wind speed for dust mobilization in the Gobi Desert
242 ranges from 10-12 m/s (Choi and Zhang, 2008). We estimated 40.5 ± 9.9 km/h as the average wind
243 speed during a dust storm according to Asia dust observations (He et al., 2008; Li et al., 2006;
244 Natsagdorja et al., 2003; Zhan et al., 2009). If air masses were transported faster than 40.5 km/h, we
245 found that the IN content in most atmospheric aerosol samples was lower on dust days than on ND
246 days because of the strong dilution effect. This effect was observed in samples 080528, 080529,
247 110319 and 100315 (Table 5). If an air mass was transported over the ocean for some distance (ratio of
248 oversea to total distance of at least 10%), no matter how fast the transport velocity, the IN content
249 decreased because of the input of clean marine air, such as in samples 080425, 100320, 110418 and
250 110501. If the air mass was transported slowly (less than 42.4 km/h) or transported only a short
251 distance over the sea, with an oversea to total distance ratio of less than 10%, the IN content increased
252 in samples collected in the downwind area, such as in samples 080301, 090316, 100321 and 110502.
253 However, there were exceptions, such as samples 080315 and 110415, which had high transport speeds
254 and passed over the sea. Therefore, the IN content is related to not only the transport path and speed but
255 also local emissions and reaction conditions during transport.

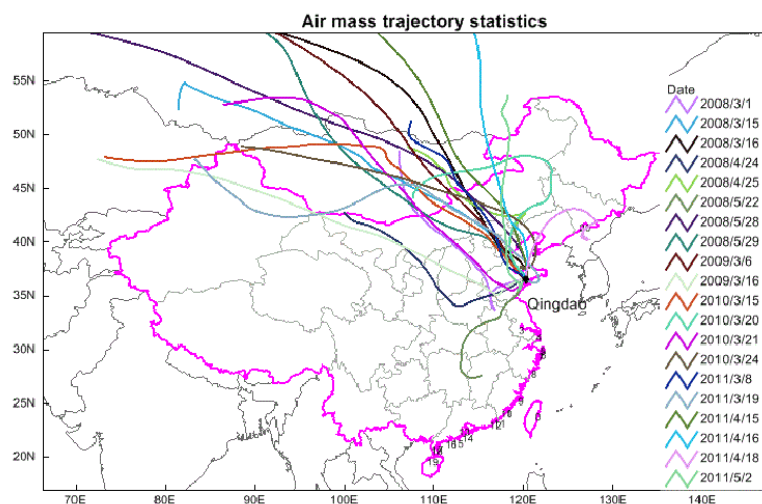
256 **Table 5.** IN content, RH, NO_x, transport speed and transport distance over the sea for atmospheric aerosol samples
257 on dust days

Group	Sample number	TSP ($\mu\text{g}/\text{m}^3$)	NO ₃ ⁻ ($\mu\text{g}/\text{g}$)	NH ₄ ⁺ ($\mu\text{g}/\text{g}$)	Speed (km/h)	Distance over the sea (km)	Ratio of the distance over the sea to the total distance (%)
Case 1 IN>ND	080301	527	38984	24107	35.1	0	0
	080315	410	47611	34130	53.4	262	6.8
	090316	688	23050	25012	59.5	0	0
	100321	519	31741	18155	49.2	0	0
	110415	1225	41970	20390	57.0	258	6.3
	110502	810	25995	13632	31.2	121	5.4
Case 2 IN<ND	080425	256	4089	372	29.3	253	12.0
	080528	2579	232	72	79.8	259	4.5



	080529	2314	26	166	78.0	229	4.1
	110319	939	13088	10067	55.4	404	10.1
	110501	502	8924	10631	30.6	298	13.5
Case 3	100315	501	10767	8515	58.7	183	4.3
NO ₃ ⁻ <ND	100320	3857	1418	884	38.0	254	9.3
NH ₄ ⁺ ≅ND	110418	558	6891	11778	23.0	380	22.9

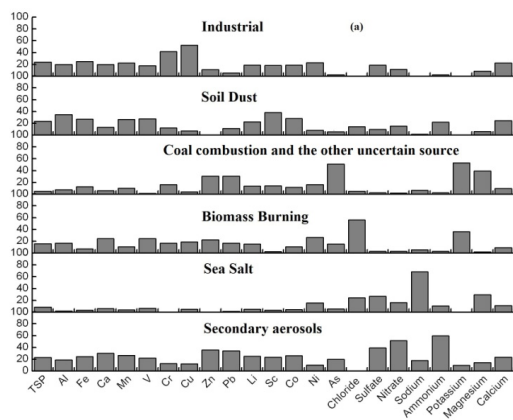
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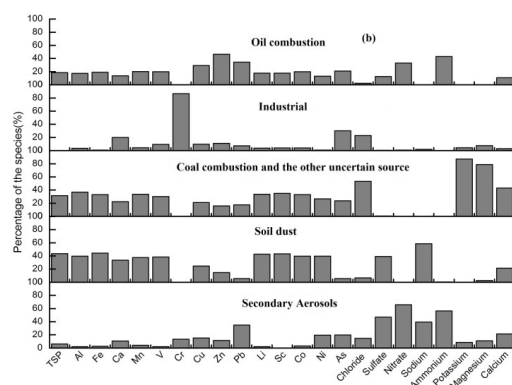
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260 **Figure 4.** The 72-h backward trajectories for non-dust and dust samples from 2008 to 2011

261 **3.4 Source apportionment of aerosols from dust and non-dust events**



262



263

264 **Figure 5.** Source profiles of atmospheric aerosol samples collected on non-dust (a) and dust (b) days using the
265 PMF model

266 The sources of atmospheric aerosols on dust and ND days were determined by running the PMF
267 model (Paatero and Tapper, 1993; Paatero, 1997). As shown in Fig. 5, atmospheric aerosols on ND days
268 were mainly from six sources: industry, soil dust, secondary aerosols, sea salt, biomass burning, coal
269 combustion and the other uncertain sources (90% of the scaled residuals between -3 and +3; $r^2=0.97$).
270 On dust days, the sources of aerosols differed from those on ND days, mainly including oil combustion,
271 industry, soil dust, secondary aerosols, coal combustion and other uncertain sources. We compared the
272 contributions of aerosol sources in dust and ND periods. As shown in Table 6, the contribution of soil
273 dust increased from 23% to 36% on dust days, which is consistent with the high concentrations of TSP
274 and crustal metals observed on dust days. Liu et al. (2014) also found that the contributions of dust
275 aerosols to PM_{10} increased by 31%-40% on dust days, which is greater than the 10%-20% contribution
276 of local soil dust on ND days. The contributions of local anthropogenic sources decreased, especially
277 those of secondary aerosols, which verified that the influence of anthropogenic sources on atmospheric
278 particulates decreased on dust days. Coal combustion emissions were mainly a mixture of coal
279 combustion and other pollutants emitted along the transmission path on dust days. Therefore, the
280 sources of aerosol particles changed on dust days. Dust events had a great impact on aerosol sources in
281 the downwind area. The influence of soil dust on aerosols and IN-loaded particles was greater than that
282 on local sources on dust days. In fact, the contribution of soil dust to aerosols was related to the
283 intensity of the dust storm and the transport path. However, we could not determine the contributions of
284 dust to aerosols for the different dust cases because of the limited number of samples.

285



286 **Table 6.** Sources and source contributions (expressed as %) calculated for aerosol samples collected during dust
287 and non-dust events

Dust event		Non-dust event	
Source	% of TSP	Source	% of TSP
Soil dust	36	Soil dust	23
Industrial	21	Industrial	24
Secondary aerosol	6	Secondary aerosol	23
Oil combustion	6	Biomass burning	16
Coal combustion and other uncertain sources	31	Coal combustion	5
		Sea salt	9

288 **3.5 Dry deposition fluxes of aerosol particles, particulate inorganic nitrogen and metals**

289 Dust events increased the concentration and deposition of aerosol particles during long-range
290 transport along the transport path. Fu et al. (2014) found that the long-range transport of dust particles
291 increased the dry deposition of PM₁₀ in the Yangtze River Delta region by 2398%. Some studies
292 observed enhancements in chlorophyll α following a dust storm event (Tan and Wang, 2014; Banerjee
293 and Kumar, 2014). The deposition magnitude of dust varied greatly among dust storms, and only some
294 dust episodes were followed by increases in chlorophyll (Banerjee and Kumar, 2014). The role of dust
295 deposition as a nutrient source leading to an increase in algal blooms has not been adequately
296 addressed. To understand the influence of dust weather on the nitrogen deposition flux, we calculated
297 the dry deposition fluxes of aerosols particles, IN and metal elements during dust and ND periods using
298 the measured component concentrations and modeled dry deposition velocities obtained from Williams'
299 model (Qi et al., 2005)(Table 7).

300 Compared to that on ND days, the dry deposition flux of atmospheric particulate increased on dust
301 days. On ND days, the dry deposition flux of particles was 2800 ± 700 mg/m²/month in the coastal
302 region of the Yellow Sea, and the particle flux varied over a wide range from 5,200-65,000
303 mg/m²/month under different dust conditions, with an average of 16,800 mg/m²/month. The results
304 verified that dust events enhanced the dry deposition flux of atmospheric particulates. However, the dry
305 deposition flux of IN showed variations with particles. In Case 1, the dry deposition flux of IN
306 increased by a factor of 1.1-5.8, and the flux of atmospheric particles increased by a factor of 1.8-6.3.
307 In Cases 2 and 3, the dry deposition flux increased 2.3 to 23.2-fold compared to that on ND days.
308 Except for ammonium in Case 3, the dry deposition flux of particulate IN decreased by an average of



309 41% in the case of high particle concentrations. The concentration of nitrate decreased 63% and 46% in
310 Cases 2 and 3, respectively. Additionally, the ammonium flux decreased by 14% in Case 2, while in
311 Case 3, ammonium was higher than that on ND days. We found that dust events sometimes led to an
312 increase in the nitrogen input to the ocean relative to that during ND events, but it did not always occur
313 depending on the chemical composition of the dust particles. As discussed, dust particles may carry
314 abundant reactive nitrogen when they travel through polluted continental atmosphere. However, the
315 relatively pure dust particles may be transported when no air pollution occurs along the dust transport
316 route to oceans.

317 The dry deposition flux of Fe in atmospheric particulates increased by a factor of 2-25 on dust days
318 compared to that on ND days. Atmospheric inputs of iron to the ocean can enhance primary production
319 in high-nutrient, low-chlorophyll regions (HNLC) (Jickells et al., 2005). However, except for Pb and
320 Zn in Case 2, the dry deposition fluxes of Cu, Pb and Zn increased with nitrogen and iron on dust days.
321 These trace metals were found to have a toxic effect on marine phytoplankton and inhibit their growth
322 (Bielmyer et al., 2006; Echeveste et al., 2012). In Case 3, dust was deposited in the ocean, the
323 atmospheric supply of nitrogen decreased, and the atmospheric inputs of Fe and some toxic metals
324 increased. Moreover, phytoplankton growth was affected by the addition of nutrient elements and toxic
325 elements. The overall effect of dust deposition on primary productivity was a combination of these two
326 effects. This is likely the reason why inhibition coexisted with the promotion of some phytoplankton
327 species in incubation experiments using additions of AD in the southern Yellow Sea in the spring of
328 2011 (Liu et al., 2013).

329 The contribution of dust events to marine nitrogen input and primary production will be
330 overestimated if the nutrient flux simply considers dust concentrations and a constant ratio of nutrients
331 to particles. The atmospheric input of nitrogen to the ocean on dust days depends on the 'dilution effect'
332 of a dust event. Dust subjected to long-range transport does not always increase the atmospheric input
333 of nitrogen. Long-term observations of dust events must be performed to evaluate the contributions of
334 dust to the biogeochemistry of nitrogen and primary production in oceans.

335
336
337
338



339 **Table 7.** Dry deposition of aerosol particles ($\text{mg}/\text{m}^2/\text{month}$), particulate inorganic nitrogen ($\text{mg N}/\text{m}^2/\text{month}$) and
340 some toxic trace metals ($\text{mg}/\text{m}^2/\text{month}$) on dust and non-dust days

	Dry deposition flux						
	Particles	NO_3^- -N	NH_4^+ -N	Fe	Cu	Pb	Zn
Case 1	9600±	87±53	25±13	650±340	2±1	0.3±0.2	6±3
IN>ND	4300						
Case 2	18000±	13±18	5±7	1300±1000	3±2	0.08±0.04	4±1
IN<ND	11,000						
Case 3	29,000±	26±6	17±8	2100±2200	6±1	0.20±0.02	5±3
NO_3^- <ND	31,000						
NH_4^+ ≅ND							
Non-dust	2800±	48±33	8±8	190±110	1±1	0.09±0.1	5±4
	700						

341 **4 Conclusion**

342 The concentration of particulate IN exhibited a large variation from event to event on dust days, and
343 a dust event did not simply increase nutrient concentrations. The effect of dust events on particulate
344 nitrogen in the downwind region was determined by the dilution effect of a dust event, which depends
345 on many factors, such as the dust storm intensity, transport speed and path, local source emissions
346 during transport, meteorological state and atmospheric reactions. Dust events affect the source
347 apportionment of aerosols. The contribution of soil dust to aerosols increased, while local
348 anthropogenic sources decreased during a dust event. The contribution of dust to aerosols must be
349 studied further under different IN conditions. Dust events enhance the input of atmospheric particulates
350 via dry deposition. However, the influence of dust events on the input of nitrogen to the ocean is still
351 uncertain. The dry deposition flux of IN on dust days decreased when a strong dilution effect was
352 present. The contribution of dust events to marine nitrogen inputs and primary production could be
353 overestimated if the dry deposition flux of nutrients is estimated using only particulate concentrations
354 on dust days.

355

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