

## ***Interactive comment on “Lightning-produced NO<sub>2</sub> observed by two ground-based UV-visible spectrometers at Vanscoy, Saskatchewan in August 2004” by A. Fraser et al.***

### **Anonymous Referee #1**

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#### General Comments

This paper presents NO<sub>2</sub> spectrometer measurements in Canada during the passage of a thunderstorm. The paper addresses scientific questions relevant to the scope of ACP. The measurements presented seem to be of good quality, however the determination of lightning-produced NO<sub>2</sub> and the following discussion need some improvements. Novel ideas or concepts of the paper are not pointed out clearly by the authors. The paper gives a brief but precise description of the performed work and used methods. The paper is well organized, well-written and fluent to read. The paper is suitable for publication in ACP after a major revision.

## Specific Comments

- It should be mentioned that descriptions of these kind of measurements are rather rare up to now in the literature. Why are they rare (difficult to perform)?
- It should be added that these measurements were performed in a remote region (?) with no local pollution (?), if this is true. This is an important statement otherwise the enhancement in calculated lightning-produced NO<sub>2</sub> could also be due to upward transport of anthropogenic NO<sub>2</sub>.
- It is assumed that the observed O<sub>3</sub> SCD increase is only due to multiple scattering in the thick cloud. However, it is mentioned that O<sub>3</sub> transported upward from the boundary layer could also enhance or decrease O<sub>3</sub> SCD measured in the thunderstorm (also known from airborne in situ measurements and cloud model simulations). Is there any possibility to quantify the amount of O<sub>3</sub> transported? Is it correct to assume that this contribution is negligible in comparison to the contribution from multiple scattering? You argue that the assumption is justified by the behaviour of the ozone to O<sub>4</sub> ratio.
- A big challenge is to determine which flashes, of all flashes in the monitored area (Fig. 4), contributed to the measured lightning-NO<sub>2</sub>. This should be pointed out in the paper with some discussion. Some uncertainty for the average flash rate of 2.87 flashes/min should be given (page 10068).
- For the estimate of the amount of excess NO<sub>2</sub> you integrate between 60° and 85° (3 hours?). You multiply this excess NO<sub>2</sub> with the area of the heavy-precipitation cell (30+3 km<sup>2</sup>). Is not the air mass area (with elevated NO<sub>2</sub>) that passed over the instrument during 3 hours much larger than this area (radius ~3 km)? The wind speed (multiplied with 3 hours) could perhaps give some better information on the size of the enhanced NO<sub>2</sub> area. I think that the area you use is not representative for the integrated NO<sub>2</sub> you estimated for 3 hours. A thunderstorm also has a complicated vertical structure as illustrated schematically in the Langford et al. (2004) paper.

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- In the introduction some information on O<sub>3</sub> and NO<sub>x</sub> is given, however it would also be useful to add something about O<sub>4</sub> (first mentioned on page 10068). Explain more in detail the reasons for using O<sub>4</sub> in the paper and add some references that introduced the O<sub>4</sub>-method to verify the influence of multiple scattering.

- Page 10070, Line 21-24: “The difference between the observed NO<sub>2</sub> slant column and the slant column calculated from the interpolated NO<sub>2</sub>/O<sub>4</sub> ratio is the amount of NO<sub>2</sub> attributed to production by lightning.” In this case you assume that the same kind of thick cloud is present before and after the large increase in NO<sub>2</sub> SCD (thunderstorm passage). Is this a correct assumption (I would expect to have less thick clouds before and after)?

- For the AMF calculation with the radiative transfer model you assume “a thick cumulus cloud near the surface, of optical depth 70, extending between 1 and 5 km”. The cloud opacity in Fig. 2 indicates a cloud top of 9-10 km that I also would assume from the strong radar reflectivity in the radar image (Fig. 3) and for the presence of a thunderstorm with lightning (Fig. 4). Include some sensitivity tests where you change the cloud depth and optical depth in your model to see how the AMF changes. Give uncertainties.

## Technical Corrections

### Abstract:

Page 10064: Line 11: Change to “The enhanced NO<sub>2</sub> columns are partly attributed to”

Line 18: Change 6.58 to 6.18 (compare to values in the conclusions).

Line 19-21: It is not common to give references in the abstract. Shorten last sentence to “These results are within the range of previous estimates.”

### 1. Introduction:

Since the paper has the focus on tropospheric and not stratospheric measurements,

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these stratospheric parts can be cut in the introduction (page 10064: line 25-26, page 10065: line 1-2, line 4-5, and line 19-22).

Page 10065, Line 2-4: High NO<sub>x</sub> causes ozone production, low NO<sub>x</sub> causes O<sub>3</sub> destruction. This statement is not completely correct, since if little sunlight is present in regions of high pollution (in winter or inside thick clouds), O<sub>3</sub> is destroyed (titrated by NO).

Line 13-15: Add some more recent estimates (year >2000), e.g. some of [Huntrieser et al., 2002; Martin et al., 2002; Tie et al., 2002; Ridley et al., 2004; Boersma et al., 2005; Beirle et al., 2006] Check order of the listed references (chronological).

## 2. Instruments:

Page 10067, Line 4: Write out DOAS.

## 3. Thunderstorm observations:

Page 10068, Line 8: Change “Figure 3 shows the precipitation occurring” to “Figure 3 shows the radar reflectivity”. The precipitation (rain rate) is calculated from the measured reflectivity and not measured.

Line 9: Change “A cell of heavy rain” to “A cell of heavy rain and probably also hail”. (The elevated values of the radar reflectivity indicate that hail is probably also prominent.)

## 4. Slant column measurements:

Page 10069, Line 12-14: Change to “The observed enhancements in ozone and partly in NO<sub>2</sub> are caused by increased path length through the atmosphere. In the case of NO<sub>2</sub>, the increase is partly also due to lightning-produced NO<sub>x</sub>.”

## 5. Derivation of lightning-produced NO<sub>2</sub>:

Page 10070, Line 3-5: Add some references for these two used methods.

Page 10071, Line 18-21: Why is the average of all ozonesondes during the campaign used and not a single ozone profile for the selected day used (would be more representative)?

#### 6. NO<sub>2</sub> flash production rate:

Page 10074, Line 11: In the original paper NO<sub>x</sub>/flash is stated and not NO<sub>2</sub>/flash.

Line 16-18: A more recent reference [Ridley et al., 2005] indicates that cloud-to-ground and intra-cloud flashes produce a similar amount of NO.

#### 7. Conclusions:

Page 10074, Line 22-24: Add “ground-based UV-visible spectrometers”.

Page 10075, Line 4: “the range of (6.18-7.45)E”, this is not the entire range of the estimates, just the range of the best estimates.

#### Acknowledgments:

Page 10075, Line 13: Change “Institute” to Institute.

#### References:

Page 10075, Line 27: Change to “Boccippio”.

#### Figures:

Fig. 1-8: Check that for all figures “day, time and location” are included in the text.

Fig. 2: What does first/second/third cloud height mean? Perhaps add that it is the cloud base height of different cloud layers.

Fig. 3 and 4: Add some information on longitude and latitude.

Fig. 3: Precipitation rate is not observed (only calculated from observed radar reflectivity).

Fig. 4: Add that the lightning flash data is superimposed on a GOES image (time?). Add “cloud-to-ground” lightning flash.

Fig. 5: Add a), b) and c).

## References

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Ridley, B. A., K. E. Pickering, and J. E. Dye (2005), Comments on the parameterization of lightning-produced NO in global chemistry-transport models, *Atmos. Environ.*, 39, 6184-6187.

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