High spatial resolution aerosol retrievals used for daily particulate matter monitoring over Po valley, northern Italy

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Abstract

The Moderate Resolution Imaging Spectroradiometer (MODIS) Collection 5.1 Aerosol Optical Depth (AOD) data retrieved at 0.55 µm with spatial resolution of 10 km (MYD04) and the new 1 km Multi-Angle Implementation of Atmospheric Correction (MAIAC) algorithm from MODIS is investigated in this work. We focus on evaluating the ability of these two products to characterize the spatial distribution of aerosols within urban areas. This is done through the comparison with PM\(_{10}\) measurements from 126 of the Italian Regional Agency for Environmental Protection (ARPA) ground monitoring stations during 2012. The Po Valley area (northern Italy) was chosen as the study domain since urban air pollution is one of the most important concerns in this region. Population and industrial activities are located within a large number of urban areas within the valley. We find that the annual correlations between PM\(_{10}\) and AOD are \(R^2 = 0.90\) and \(R^2 = 0.62\) for MYD04 and for MAIAC respectively. When the depth of the planetary boundary layer (PBL) is used to normalize the AOD, we find a significant improvement in the PM–AOD correlation. The introduction of the PBL information is needed for AOD to capture the seasonal cycle of the observed PM\(_{10}\) over the Po valley and significantly improves the PM vs. AOD relationship, leading to a correlation of \(R^2 = 0.98\) for both retrievals when they are normalized by the PBL depth. The results show that the normalized MAIAC retrieval provides a higher resolution depiction of the AOD within the Po Valley and performs as well in a statistical sense as the normalized standard MODIS retrieval for the same days and locations.

1 Introduction

Particulate matter (PM), also defined as atmospheric aerosol, is one of the major pollutants studied and monitored sense it affects air quality in urban and rural areas worldwide. PM is the general term used to define a complex mixture of solid and liquid particles. These particles vary in size and composition, and remain suspended in the
air for different periods of time. The sources of the atmospheric aerosols include both natural activity, such as fire, sea salt, volcanic eruptions and windblown dust, and anthropogenic activity, such as combustion, traffic and industrial emissions. PM with aerodynamic diameter of 10 µm or less (PM$_{10}$) leads to serious human health effects. They can be inhaled into the respiratory system and so cause respiratory lung diseases and even premature death (Forastiere et al., 2005; Brunekreef and Forsberg, 2005; Pope et al., 2004). At local scale, urban pollution plays a significant role on issues related to health due to high urban population densities. Prior to the twentieth century, most urban air pollution problems arose from the burning of wood, coal and other raw materials without any emission controls. Such burning resulted in significant increases in health issues related to urban pollution (Jacobson, 2012).

The Po Valley, in the northern part of Italy, is the area with the most severe air pollution problems in the country as it is the largest industrial, trading and agricultural area with a high population density (Bigi et al., 2012). As shown by the MODIS true color image on 16 March 2012 (Fig. 1), a region of visible haze due to strong aerosol pollution is evident throughout the Po Valley. The pollution problems that affect the Po Valley are not only related to the presence of highly urbanized and industrial centers. In fact, the presence of the Alpine mountain chain at the North and West sides of the valley, and the Apennines to the South, act as a barrier to winds blowing from Northern Europe and the Mediterranean, favoring stagnation conditions and accumulation of pollutants (Mazzola et al., 2010; Putaud et al., 2004, 2010). Because of this, monitoring in this area requires data with high spatial resolution to better characterize the spatial variability of pollution within the Po Valley.

Due to health problems associated with urban air pollution, many environmental protection agencies have been developing capabilities for continuous monitoring and assessment of air pollution from ground-based stations and for improving sampling techniques. These ground-based measurements are necessary to guide studies of possible ways to reduce the air pollution problems. Yet, ground-based observations represent point measurements and do not have the necessary coverage to characterize the regional distribution of aerosols in the atmosphere. Moreover, the PM ground-based stations only provide information at the surface. The development of satellite remote sensing aerosol products since the launch of the Moderate resolution Imaging Spectroradiometer (MODIS) onboard the NASA Terra and Aqua satellites has permitted the exploration of new research techniques for monitoring global air quality (Fishman et al., 2008; Gupta et al., 2006). This alternative approach for air quality monitoring provides air quality data where the ground-based measurements are not available. The potential for using space-based sensors for the air quality monitoring was demonstrated using Aerosol Optical Depth (AOD) data in combination with the PM ground-based stations (Chu et al., 2003; Wang and Christopher, 2003). The use of MODIS aerosol products to investigate air pollution was demonstrated both in research fields (van Donkelaar et al., 2006; Tian and Chen, 2010; Gupta and Christopher, 2008) and for operational applications (Al-Saadi et al., 2005). But, the satellite AOD quantifies the presence of aerosols in an atmospheric column, while the surface PM mass concentration is needed for the assessment of air quality health impacts: it is not obvious what the relationship between these two quantities is for a particular region. Hoff and Christopher (2009) provide a detailed review of the literature on how satellite remote sensing provides an alternative way to monitor the surface PM mass concentration. As they point out, correlations between ground measurements and optical thickness are actively used and investigated. However, their correlation is often not high enough for AOD retrievals to be operationally incorporated in air quality monitoring procedures. The MODIS standard aerosol product spatial resolution is appropriate for application on regional to global scale. However, its nominal resolution at nadir (10 km) may be too coarse to resolve urban scale processes. Therefore, alternative aerosol retrieval algorithms have been developed using the MODIS data in order to produce a finer high-spatial resolution product. The MODIS NASA research team recently released a new MODIS product: MODIS Collection 6, which includes a global aerosol product at nominal 3 km, in addition to the standard MYD04 at 10 km (Remer et al., 2013; Munchak et al., 2013; Livingston et al., 2014). Recently, the Multi-Angle Implementation of Atmospheric Correction (MAIAC) algorithm
was developed for MODIS (Lyapustin et al., 2011a–c, 2012). The MAIAC algorithm performs a simultaneous retrieval of surface Bidirectional Reflection Distribution Function (BRDF) and aerosol properties at a resolution of 1 km, and represents an interesting alternative for characterizing spatial variability of aerosol within polluted and industrial urban areas (Chudnovsky et al., 2013a, b; Emili et al., 2011; Hu et al., 2014).

In the current study we extend the preliminary analysis of Arvani et al. (2013) by using the MODIS Collection 5.1 10 km and the MAIAC 1 km resolution retrievals to analyze the relationship between PM$_{10}$ and AOD within the Po Valley for an entire year (2012). We start with a direct comparison between MYD04 and MAIAC AOD retrievals and the surface PM$_{10}$ measurements. Then, we introduce an additional factor, the depth of the planetary boundary layer, to better characterize the seasonal variability of the particular matter over the Po Valley and improve the PM–AOD relationship. Finally, we investigate the impact of neighboring pixels and co-location radius on the correlation.

2 Data and methods

The region of interest, Po Valley in northern Italy, shown in Fig. 1, covers an area of approximately 40–50°N and 5–15°E. The Po Valley is divided into four administrative divisions (regions) from west to east, Piemonte, Lombardia, Emilia Romagna and Veneto. Although the study area is small, it is possible to identify specific features for each of the four administrative regions. Piemonte is characterized by major urban centers and is heavily industrialized, including the capital city of Turin. Turin has one of the main industrial centers in Italy. Most of the Italian Regional Agency for Environmental Protection (ARPA) network stations within the Piemonte region are located near the Alps. Lombardia is the most populated region. It also has major industrial sites, including the capital city of Milan. The Emilia Romagna region, similar to Piemonte, is characterized by major urban centers. Some of the ARPA sites within this region are located close to the Apennine mountain chain to the south, and near the coast of the Adriatic Sea to the east. There are only a few ARPA stations within the Veneto region.

2.1 Ground-level concentration of PM$_{10}$

Twenty-four hour average PM$_{10}$, mass concentration in µgm$^{-3}$, was considered for 126 air quality monitoring stations within the ARPA network during the year 2012: Piemonte (27 stations), Lombardia (59 stations), Emilia Romagna (37 stations) and Veneto (3 stations). The distribution of ground stations, and location within the Po Valley is highlighted in Fig. 2. Each regional ARPA network has a unique set of measurements with different uncertainties. This non-uniformity may affect the regional correlations. In particular, in ARPA Piemonte and Veneto PM$_{10}$ is measured using a Beta Attenuation Monitor (BAM) with an accuracy of 2 %. In ARPA Lombardia TEOM, TEOM-FDMS (Tapered Element Oscillating Micro-balance – Filter Dynamics Measurement System), or BAM are used, with an accuracy of ±2.5 µgm$^{-3}$. In ARPA Emilia Romagna, PM$_{10}$ data have been collected by the beta attenuator SWAM 5A RL by FAI Instruments with an uncertainty lower than ±10 %. All information related to the ARPA’s stations and instrumentation are available at the ARPA web sites of their respective regions, mentioned above.

2.2 Satellite data

In the present work, we use the standard MODIS Aqua Collection 5.1 (MYD04) aerosol product. This product has a nominal spatial resolution of 10 km at nadir, and increases by roughly four-fold at the edges of the swath. The MODIS AOD algorithm uses multispectral observed radiance and pre-computed look-up tables to retrieve AOD over ocean and land (Remer et al., 2005, 2006). The recently developed MAIAC aerosol retrieval algorithm (Lyapustin et al., 2011a, b) is also considered. This algorithm is part of MODIS Aqua Land retrieval and is provided over the land only. It is retrieved at a finer spatial resolution (1 km), making simultaneous use of BRDF parameters. This is accomplished by using the time series of MODIS measurements and simultaneous processing of groups of pixels. The MAIAC algorithm guarantees that the number of measurements exceeds the number of unknowns, a necessary condition for solving an
inverse problem without empirical assumptions, which are commonly used in current operational algorithms. The MAIAC time series approach also provides coverage at multiple (15) view angles for every surface grid cell, which is required for the BRDF retrievals from MODIS data. Moreover, MAIAC incorporates a Cloud Mask (CM) algorithm based on spatio-temporal analysis which augments traditional pixel-level cloud detection techniques (Lyapustin et al., 2008). In Fig. 3, a comparison between MODIS standard and MAIAC aerosol retrieval results between is shown. As immediately evident, the higher resolution MAIAC retrieval shows substantial spatial variability which is not captured by the standard 10 km retrieval.

The MODIS Collection 5.1 retrieval has already been cloud filtered and is used without additional quality control. On the contrary, the new aerosol retrieval algorithm MAIAC includes cloud and terrain in masks, incorporated into the AOD Quality Assurance (QA) parameter definition. The MAIAC Cloud Mask (CM) and the Land-Water-Snow (LWS) mask fields have been considered during the MAIAC run in order to avoid pixels where clouds, water or snow are detected. As example, 16 March 2012 is shown in Fig. 4. Comparison with the MODIS (Aqua) true color image (Fig. 1), shows an extensive area of clouds to the south-west, near the coast, that are well characterized in the MAIAC CM. The snow cover over the Alpine chain in north is also well characterized by the MAIAC LWS mask.

2.3 Satellite data and ground measurements co-location

The MODIS Collection 5.1 aerosol product, MAIAC retrievals and surface PM$_{10}$ were collocated in space and time for 2012 for each ARPA station considered. The spatial co-location of MODIS and MAIAC pixels with PM$_{10}$ ground-based stations was accomplished using two different methods. The nearest neighbor method calculates the distance between the pixel and PM$_{10}$ station and determines the minimum radius within a tolerance radius of 0.20” (about 20–25 km at the latitude of the Po Valley) for MODIS Aqua Collection 5.1, and a tolerance radius of 0.02” (about 2–2.5 km) for MAIAC AOD retrieval. The second approach, labeled average as in Gupta et al. (2006) paper, uses the average of all pixels within a tolerance radius of 0.20” and 0.02” for MODIS and MAIAC respectively. For both approaches, we accounted for more than one MAIAC coincidence per PM$_{10}$ station. This is due to the overlap of different Aqua orbits within the MAIAC local tiles.

2.4 AOD normalization

PM$_{10}$ and AOD represent two different measurements of the atmospheric loading of pollutants. The PM$_{10}$ is the dry mass, measured at ground level, at a specific geographic location. On the other hand, the satellite AOD represents total column aerosol loading averaged over a specific spatial area. As suggested by the literature (Gupta et al., 2006; Tsai et al., 2011), the PM–AOD correlation may be improved by considering meteorological data information or vertical distribution of aerosols. In this work, variations in the vertical distribution of aerosols are considered by introducing information on the Planetary Boundary Layer (PBL) depth. The PBL height (ZPBL) values derive from 6 hourly 0.5° × 0.5° analysis files from the NOAA National Center for Environmental Prediction (NCEP) Global Data Assimilation System (GDAS), downloaded from nomads.ncdc.noaa.gov. As mentioned previously, the Aerosol Optical Depth is an integration of the aerosol extinction, from the surface to the top of the atmosphere:

$$\text{AOD} = \int_{0}^{\text{TOA}} \sigma_{0.55\mu m}(z) dz$$

Under the hypothesis that most of the aerosols are confined and mixed homogeneously within the planetary boundary layer, the values of AOD normalized by PBL depth may be regarded as mean PBL extinction in km$^{-1}$ and may be more representative of the surface PM$_{10}$ concentration since the depth of the PBL is accounted for. The normalization was applied both for MYD04 and MAIAC retrievals.

The seasonal trend of PBL heights over Po Valley shows low values during the winter and high values during the summer (Fig. 5). This seasonal trend is reflected in the
AOD monthly mean values normalized by PBL depth and results in higher values in the winter and fall period (Figs. 7 and 8, panels c and d). The normalized MYD04 and MAIAC mean AOD values follow the PM trend throughout the whole year and shows the strong seasonal correlation between PM and PBL depths within the Po valley.

To verify the GDAS PBL depths we compare them to PBL depths obtained from CALIPSO measurements (Winker et al., 2003, 2009). The CALIPSO PBL depths are derived using a Haar wavelet analysis to detect boundaries in scattering ratio (i.e. a normalized backscatter) in Lidar observations that include the atmospheric boundary layer. The CALIPSO PBL heights are taken as altitude where the maximum amplitude average wavelet occurs computed over a range of Haar filter widths ranging from 0.9 to 1.65 km (Kuehn, R. E., 2013, personal communication). Figure 5 shows the monthly mean of the 6 hourly gridded GDAS PBL depths (green), as well as the mean of the GDAS PBL interpolated to the CALIPSO track (blue), and the CALIPSO PBL depths (red) over the Po Valley. The blue trend and the red one follow almost the same seasonal trend. The CALIPSO sampled GDAS PBL heights are comparable to the seasonal trend of the gridded GDAS PBL heights over Po Valley, except in January, where the CALIPSO sampling introduces a high bias. Comparison between the coincident CALIPSO and GDAS PBL depths shows very similar seasonal trends but CALIPSO PBL depths are systematically higher than the GDAS analysis.

3 Results and discussion

3.1 AERONET validation

Following Chu et al. (2002, 2003), the validation of satellite aerosol retrievals is done by comparing the satellite AOD to ground based AOD retrievals by AERONET Sun photometers (Holben et al., 1998). During 2012, three AERONET measurement sites located in Ispra, Modena and Venice (marked in black dots in Fig. 6) were available. Even though limited in number, the available AERONET stations are representative of the Po Valley conditions: the Ispra site is located in a pre-Alpine rural region dominated by forest, Modena (the middle site) is located in an urban environment in the middle of Po Valley, and the Venise site is at the Acqua Alta Oceanographic Tower (AAOT) in the northern Adriatic Sea, 8 nautical miles off the Venice Lagoon. Due to the limited amount of AERONET measurements at the Modena and Venise sites during 2012, only Ispra AERONET measurements are used for validation. Co-location of MODIS retrievals and AERONET direct Sun measurements requires at least two AERONET measurements within ±30 min of the MODIS overpass. The mean values of the temporal ensemble are then used in linear regression analysis. Moreover, the AERONET data are interpolated in logarithm of wavelength 0.55 and 0.47 µm for MYD04 and MAIAC respectively. Scatter plots for the collocated MAIAC, AERONET Level 2.0 and MYD04 AOD are shown in Fig. 6, with a determination coefficient of $R^2 = 0.84$ and $R^2 = 0.69$ for MODIS and MAIAC AOD respectively. MAIAC ($N = 32$) provides more data points than Collection 5.1 ($N = 25$), reflecting the higher spatial resolution of the MAIAC aerosol retrieval algorithm.

3.2 Time series analysis

The AOD–PM$_{10}$ analysis begins with the study of the 2012 monthly mean trend of PM$_{10}$ vs. AOD for both the spatial co-location approaches presented in Sect. 2.3. The results are reported in Fig. 7. The top panels show the mean daily value of PM$_{10}$ 24 h mass concentration (black line), for all 126 ARPA stations, while the red dots show the monthly mean PM$_{10}$ mass concentration. The AOD monthly mean values are represented on the graph by the blue and the green dots, for MYD04 and MAIAC, respectively. As immediately evident, the trends in PM and AOD are different during the winter and fall periods for both nearest-neighbor (Fig. 7a) and average (Fig. 7b) coincidence approaches. For both methods, a radius of coincidence equal to 0.02’ was used for MYD04 and MAIAC to allow for a more direct comparison. The disagreement is particular notable for the two last months of the year, where the PM monthly mean values increase, while AOD values decrease. The highest values of PM are recorded in this
period of the year due to the meteorological conditions that favor the buildup of near-surface pollutants, and regional environmental protection agencies are actively trying to limit air pollution problems (Mazzola et al., 2010). The same analysis was conducted considering a larger radius of coincidence and did not have significant impact on the final time series result (Fig. 8a and b).

### 3.3 Scatter plot and bin analysis: PM$_{10}$ vs. AOD relationship. Comparison of different method of coincidence

As second step of analysis, we proceed with the calculation of the determination coefficient between daily PM$_{10}$ mass concentration and AOD from both MYD04 and MAIAC to estimate quantitatively the linear relationship between the two data sets. In Sect. 2.3, two different approaches to find the spatial co-location of satellite pixel with the ground-based station have been presented, the nearest neighbor and average method.

Because of the large spread of PM and AOD values, it is more useful to derive a simple linear regression equation between PM$_{10}$ mass concentration and AOD by dividing the PM$_{10}$ (24 hrs) into 10 bins of 5 µg m$^{-3}$ intervals and comparing them with the mean AOD within each bin. So, the final value of AOD per each PM$_{10}$ bin is determined by calculating the average value of AOD between the values of AOD retrieved in each class of PM. This set of ten points is reported on the scatter plot as black dots in Fig. 9, using the nearest neighbor approach. The solid red line shows the linear regression line for these two data sets. White dots refer to median values of AOD at fixed value of PM$_{10}$. Yellow symbols represent the 25th and 75th percentile (first and third quartiles) respectively in AOD for a particular PM$_{10}$ bin. This simple statistical approach gives a robust estimate of the linear regression between the PM$_{10}$ and satellite data (Gupta et al., 2006). The correlation between bin-averaged AOD and PM$_{10}$ concentration is highest for MYD04, with $R^2 = 0.90$. The higher resolution MAIAC retrieval algorithm has a significantly lower $R^2$ of 0.62. As mentioned in the previous subsection, the PM$_{10}$-AOD correlation may be improved with the introduction of aerosol planetary boundary layer information. Therefore, the same analysis was applied, considering the PM vs. AOD/ZNBL data sets. The results are summarized in Fig. 9 in the right panels b and d. They show an $R^2$ of 0.98 for both MYD04 and MAIAC when AOD is normalized by the PBL depth. Looking closely at the results shown in Fig. 9 shows that while the PM–AOD/ZNBL correlation has a strong linear behavior with a slope close to 1, this is not true for the PM–AOD correlations. The median AOD is roughly linear with respect to PM for the first five bins. After the fifth bin, the median AOD shows less variation with PM, especially from MAIAC data set. This points to a more complex relationship between PM$_{10}$ and AOD than represented with a linear regression for the PM$_{10}$ vs. AOD relationship.

Regarding to the use of the average approach, if there is at least one missing pixel within the tolerance area we set the mean AOD value equal to a missing data. This is a conservative approach since it automatically excludes all the retrievals in the neighborhood of pixels identified by the terrain and cloud masks within a fixed tolerance area. Following this approach to determine the coincidences, the binned scatter plot results in $R^2 = 0.92$ and $R^2 = 0.76$ for MYD04 and MAIAC, respectively. If this conservative approach is not taken we find that the $R^2$ for the MAIAC AOD normalized by PBL depth has a much lower correlation with PM$_{10}$. On the other hand, we find that the impact of neighboring missing values is less of an issue for the standard MODIS retrieval, possibly due to a more conservative cloud mask. The difference between the two approaches is shown in Fig. 10. Looking at Fig. 10a and b, or Fig. 10d and e, the PM$_{10}$ ground-based sites located at the edge of the AOD retrieval map are detected as “coincidence” site in the nearest neighbor approach, while “no coincidence” in the average one. From the average coincidence results shown in Fig. 11, it is remarkable there are a different number of coincidences when the MYD04 or MAIAC AOD retrievals are considered. This is not surprising because of the different spatial resolution of the AOD retrieval, as better explain in Sect. 2.2.
3.4 Study of the impact of the radius of coincidence on $R^2$

Until now, the statistical analysis has been conducted using a fixed radius of coincidence of $0.20\degree$ for MYD04 and $0.02\degree$ for MAIAC based on the nominal spatial resolution of the two satellite aerosol retrievals used in this work. Changing the value of the radius of coincidence will change the final number of coincidences detected and likely the final $R^2$ value in this statistical analysis. Therefore, this section studies the impact of changing the value of the tolerance radius on $R^2$ value. Figure 12 shows a comparison between $R^2$ and radius of coincidence in km for MYD04 and MAIAC for each method of coincidence used in this study. For the sake of clarity, the radius of coincidence was converted from degrees to kilometers assuming $1\degree = 111.121$ km valid at the mean latitude of the study domain. For the standard MODIS retrieval, changing of the radius of coincidence value does not have a significant impact on $R^2$ value except for changing from about 2 to 5 km where $R^2$ increases from 0.78 to 0.90. Also, the standard MODIS PM–AOD/ZPBL relationship is less sensitive to the radius of coincidence than the PM–AOD correlation. The MAIAC results, on the other hand, have completely different behavior. The nearest neighbor approach, as expected, does not depend on the radius of coincidence, and maintains a constant offset between the PM–AOD and PM–AOD/ZPBL correlations. The average approach shows a significant dependence on the radius of coincidence. The $R^2$ value in the average method for the PM–AOD correlation is maximum (0.87) for radius equal to 12 km, and minimum (0.76) for radius equal to 2 km. On the other hand, in the PM–AOD/ZPBL relationship, the $R^2$ value is roughly constant. Figure 13 shows the direct comparison between the MYD04 and MAIAC number of coincidences (N) and radius for both the methods of coincidence analyzed in this work. For MYD04, the number of coincidences shows a significant dependence on the radius, especially for the first three values ($0.02, 0.05$ and $0.08\degree$) where N increases rapidly when the radius threshold is increased. For larger coincidence radius, the N trend is roughly constant. For MAIAC, there is a much more rapid decrease in the number of coincidences for the average approach than with MYD04 due to the high spatial resolution of the retrieval. MAIAC shows a nearly constant N for the nearest neighbor approach.

4 Conclusions

Until recently, the MODIS Collection 5.1 satellite AOD data product, with 10 km resolution, was the main source of global satellite aerosol data used by the air quality community. The new MAIAC AOD product, at 1 km resolution, provides a significantly higher resolution product that may be more appropriate for urban air quality studies. This paper analyzed the effect of spatial resolution on the correlation between remotely sensed AOD and ground-based PM$_{10}$ concentration measurements. One year (2012) of MODIS Aqua retrievals over the Po valley domain were used. The major findings of this study are summarized below.

1. Time series analysis of the monthly mean AOD values shows a notable disagreement with surface PM$_{10}$ in the fall and winter period, days characterized by meteorological conditions that favor the buildup of near-surface pollution, for both retrievals. This disagreement is reduced by normalization of AOD by the PBL height.

2. A direct comparison between coarse MYD04 10 km AOD and high-resolution MAIAC 1 km AOD for all collocated PM vs. AOD pairs for the same period of analysis and sites show that both retrievals have low correlations.

3. As often noted by the literature, the link between surface measurements and AOD data alone is limited and often not suitable for quantitative analysis. However, the normalization by the PBL depth significantly improves the $R^2$ due to seasonal changes in the PBL depth over the Po Valley. $R^2 = 0.98$ for both standard MODIS and MAIAC retrievals was obtained during 2012.
4. Analysis of changing the tolerance radius on $R^2$ indicates a significant impact if the average method is applied on the PM–AOD correlation for the MAIAC retrieval. The standard MODIS retrieval was found to be largely independent of the range of radii considered in this study. This may be due to a more conservative cloud mask in the standard MODIS retrieval.

The results presented in this paper show that although the MAIAC 1 km retrieval provides higher resolution information on aerosol optical depth within the highly industrialized Po Valley of northern Italy although the correlation between PM and AOD is worse than standard MODIS retrieval. However, if both retrievals are normalized by the PBL depth, the MAIAC results are nearly identical to the coarser resolution standard MODIS retrieval. The results reported in this work were obtained but considering just one factor that may affect the relationship between the ground-based and the satellite remote-sensing measurements. Future work will focus on improving the understanding of the satellite-retrieved AOD and PM$_{10}$. Specifically, we intend to investigate the use of higher resolution PBL estimates obtained from regional NWP over Italy (Kukkonen et al., 2012; Baldauf et al., 2011; Barthlott et al., 2010) and explore the relationship for each administrative district over Po valley separately. The aim would be to investigate if the use of finer PBL depth and satellite-retrieved AOD (MAIAC) helps to characterize the spatial variability of aerosol pollution within the Po Valley and study the impact of industrialized regions on PM vs. AOD relationships.

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References


Figure 1. Geographic study domain. MODIS (Aqua) satellite True Color RGB image –1 km, 16 March 2012 (http://lance-modis.eosdis.nasa.gov/imagery/subsets/?project=aeronet&subset=Ispra). The blue circle represents the geographic location of the Ispra AERONET station in the Po Valley.

Figure 2. Geographic study domain. The dots mark locations of the ARPA PM$_{10}$ ground station sites. The dots are grouped into different color, one per each administrative division considered. The black dots are located position of AERONET stations, mentioned in Sect. 3.1.
As it is immediately evident, the higher resolution data reveal a substantial spatial variability of AOD which cannot be captured using a coarse 10 km scale.

Figure 3. MYD04 10 km (on the left) and MAIAC 1 km (on the right) AOD for 16 March 2012.

Figure 4. Understanding of the Cloud mask (CM) field and Land-Water-Snow (LWS) mask field. CM on the left; LWS on the right.
Figure 5. Seasonal (GDAS) PBL height trend, calculated over Po Valley domain, compared to both CALIPSO and Interpolated GDAS seasonal one.

Figure 6. Validation scatter plot of (a) MODIS (Aqua) Collection 5.1 ($\lambda = 0.55\, \mu m$) and (b) MAIAC AOD ($\lambda = 0.47\, \mu m$) vs. AERONET Level 2.0 AOD for Ispra station. The statistic for the validation is given above the plot.
Figure 7. Trend of PM$_{10}$ ($\mu$g m$^{-3}$) compared to MODIS Aqua and MAIAC respectively, over the Po valley domain. In the panels (a), (b) the relationship PM-AOD is considered, while in the panels (c), (d) are reported the results for PM-AOD/ZPBL relationship. Radius of coincidence equal to 0.02°.

Figure 8. Trend of PM$_{10}$ ($\mu$g m$^{-3}$) compared to MODIS Aqua and MAIAC respectively, over the Po valley domain. In the panels (a), (b) the relationship PM-AOD is considered, while in the panels (c), (d) are reported the results for PM-AOD/ZPBL relationship. Radius of coincidence equal to 0.20°.
Figure 9. Bin scatter plots results for the Po valley domain, for the nearest neighbor approach. In the upper side of the panel, (a) and (b), the MYD04 results are shown. In the lower side, (c) and (d), the comparison with the MAIAC results are presented.

Figure 10. MYD04 10 km (on the left) and MAIAC 1 km (on the right) AOD for 16 March 2012, for the two methods to find coincidences analyzed in this work. The dots mark locations of the ARPA PM10 ground station sites, where the red ones identify a coincidence site, while the black one a non-coincidence site.
Figure 11. In scatter plots results for the Po valley domain, for the average approach. In the upper side of the panel, (a) and (b), the MYD04 results are shown. In the lower side, (c) and (d), the comparison with the MAIAC results are presented.

Figure 12. Coefficient of determination $R^2$ dependence on different values of radius of coincidence, where on the left are reported the results obtained from MYD04 data sets and on the right the results obtained from MAIAC.
Figure 13. Number of coincidence dependence on different values of radius of coincidence, where on the left are reported the results obtained from MYD04 data sets and on the right the results obtained from MAIAC.