



High surface ozone  
and strong influence  
of biomass burning

V. Sinha et al.

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# Chemical composition of pre-monsoon air in the Indo–Gangetic Plain measured using a new PTR-MS and air quality facility: high surface ozone and strong influence of biomass burning

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Abstract

One seventh of the world population lives in the Indo–Gangetic Plain (IGP) and the fertile region sustains agricultural food crop production for much of South Asia. Yet it remains one of the most under-studied regions of the world in terms of atmospheric composition and chemistry. In particular, the emissions and chemistry of volatile organic compounds (VOCs) that form surface ozone and secondary organic aerosol through photochemical reactions involving nitrogen oxides is not well understood. In this study, ambient levels of VOCs such as methanol, acetone, acetaldehyde, acetonitrile and isoprene were measured for the first time in the IGP. A new atmospheric chemistry facility that combines India's first high sensitivity proton transfer reaction mass spectrometer, an ambient air quality station and meteorological station, was used to quantify in-situ levels of several VOCs and air pollutants in May 2012 at a suburban site in Mohali (N.W. IGP). Westerly winds arriving at high wind speeds (5–20 ms<sup>-1</sup>) in the pre-monsoon season at the site, were conducive for chemical characterization of regional emission signatures. Average levels of VOCs and air pollutants in May 2012 ranged from 1.2–1.7 nmol mol<sup>-1</sup> for aromatic VOCs, 5.9–37.4 nmol mol<sup>-1</sup> for the oxygenated VOCs, 1.4 nmol mol<sup>-1</sup> for acetonitrile, 1.9 nmol mol<sup>-1</sup> for isoprene, 567 nmol mol<sup>-1</sup> for carbon monoxide, 57.8 nmol mol<sup>-1</sup> for ozone, 11.5 nmol mol<sup>-1</sup> for nitrogen oxides, 7.3 nmol mol<sup>-1</sup> for sulphur dioxide, 104 µg m<sup>-3</sup> for PM<sub>2.5</sub> and 276 µg m<sup>-3</sup> for PM<sub>10</sub>. By analyzing the one minute in-situ data with meteorological parameters and applying chemical tracers (e.g. acetonitrile for biomass burning) and inter-VOC correlations, we were able to constrain major emission source activities on both temporal and diel scales. Wheat residue burning activity caused massive increases (> 3 times of baseline values) for all the measured VOCs and primary pollutants. Other forms of biomass burning at night were also a significant source for oxygenated VOCs and isoprene ( $r^2$  with acetonitrile  $\geq 0.5$  for night-time data), which is remarkable in terms of atmospheric chemistry implications. Surface ozone exceeded the 8 h national ambient air quality limit of 100 µg O<sub>3</sub> m<sup>-3</sup> on a daily basis, except for 17 May 2012, when a se-

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

dous importance of in-situ data from the IGP, the region's atmospheric composition is still poorly constrained, with emission sources and the spatial and temporal variability of pollutant levels still not well understood. In particular, a considerable knowledge gap continues to exist with respect to levels of volatile organic compounds and primary gaseous pollutants due to lack of in-situ measurements within the region. Since INDOEX, much of the atmospheric research efforts in India and the IGP region have focused on constraining the optical and radiative properties of aerosols from a climate perspective (e.g., Satheesh and Moorthy, 2005; Moorthy et al., 2008), but in contrast very little data on hydrocarbons has been acquired. Trace gas research in India and the IGP has focused mainly on green house gases such as carbon dioxide, methane and in-situ ozone measurements (e.g., Lal et al., 2000, 2008a, b; Ghude et al., 2008). With regard to volatile organic compounds, only light C<sub>2</sub>–C<sub>5</sub> non methane hydrocarbons (e.g., Sahu and Lal, 2006) and BTEX aromatic compounds (e.g., Srivastava et al., 2005) have been measured, and no previous in-situ data on the ambient air levels of compounds such as acetonitrile, isoprene, methanol and acetaldehyde exists except for a recent case study on paddy residue burning emissions (Sarkar et al., 2013). Volatile organic compounds can be used as emission tracers (e.g. acetonitrile for biomass burning, (Holzinger et al., 1999), but more importantly due to their high reactivity with hydroxyl radicals, fuel the formation of secondary pollutants such as ozone and secondary organic aerosol through photochemical reactions involving nitrogen oxides (Ryerson et al., 2001). VOCs and NO<sub>x</sub> measurements are essential for understanding the photochemistry of ozone formation, and yet have seldom been measured simultaneously with ozone in studies conducted over the IGP and Indian region. Thus, mechanistic level understanding of the in-situ chemistry responsible for ozone production at sites over the IGP and Indian region is yet to be achieved, which is a pre-requisite for devising robust mitigation strategies. High surface ozone reduces crop yields, impairs regional air quality and causes respiratory ailments among the exposed population causing grave economic losses (Giles, 2005; Jerret et al., 2009).

**High surface ozone  
and strong influence  
of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Comprehensively instrumented measurement sites in other regions of the world such as the “Station for Measuring Forest Ecosystem–Atmosphere Relations” (SMEAR II) at Hyytiälä in Finland (e.g., Hari and Kulmala, 2005), and during field experiments such as CALNEX over Los Angeles (Holzinger et al., 2013) in California, MEGAPOLI (Dolgorouky et al., 2010) in Paris, and MCMA (Molina et al., 2007) and MILAGRO (Molina et al., 2010) in Mexico city have yielded important data on the chemistry of ozone and aerosol formation at several urban and suburban sites elsewhere in the world.

Here, we present results from a new ground based atmospheric chemistry facility at a regional suburban site (30.6794° N, 76.7289° E, 310 m a.s.l.) in the north west IGP in Mohali, Punjab. A comprehensive description of the site and experimental details of the facility which consists of a high sensitivity proton transfer reaction mass spectrometer (PTR-MS) (the first to be deployed within India and the IGP), an online ambient air quality station and a meteorological station are provided. Using high temporal resolution (1 measurement per minute) in-situ data acquired in May 2012 of a suite of volatile organic compounds (methanol, acetone, acetaldehyde, acetonitrile, isoprene, benzene, toluene, sum of C8 aromatics, sum of C9 aromatics), ozone, nitrogen oxides, carbon monoxide, sulphur dioxide, particulate matter and meteorological parameters, the impact on the ambient chemical composition due to regional emission activities typical of the pre-monsoon season in the IGP, such as open wheat residue burning, other types of biomass burning and a severe dust event were studied. The measured levels were compared with summertime measurements reported from other suburban/urban sites. Bivariate polar plots and inter VOC correlations, for example with acetonitrile (a chemical tracer for biomass burning), were employed to constrain the major emission sources of the compounds. The diel variability of different compounds was used to assess the role of emissions, boundary layer dynamics and photochemistry on the ambient levels of the measured species. Finally, a discussion of the new findings for future atmospheric chemistry and air quality studies is presented.

## 2 Experimental

### 2.1 Site description and prevalent meteorology during the pre-monsoon season

Figure 1 shows the location of the city of Mohali in the north west Indo–Gangetic plain in the Indian state of Punjab where greater than 80% land use is agricultural. The measurement facility is housed inside the campus of the Indian Institute of Science Education and Research (IISER), a suburban site (30.667° N–76.729° E, 310 m.a.s.l.) several kilometres away from the city centres of Mohali, Chandigarh and Panchkula which are the nearest neighbouring cities (see Fig. 2). The campus is built in an enclosed area of 1.25 km<sup>2</sup> and had less than 600 residents in May 2012. Local influences on the measurements from residences and hostel emissions are minimal and only expected to be significant at low wind speeds ( $< 1 \text{ m s}^{-1}$ ) as the buildings in the immediate vicinity of the site are downwind of the inlets and air masses typically arrived at the site at high wind speeds (3–27  $\text{m s}^{-1}$ ). Figure 2 gives the major land use surrounding the site in a 100 km × 200 km area. The image was derived by labeling the urban areas of the European Space Agency (ESA) 2010 land use map and fitting the map with a scale bar. As can be seen in Fig. 2, the largest urban area near the site is the planned city of Chandigarh (population of circa 1 million in 2012) which lies in the wind sector spanning north to east (0–90°). The land use in the wind sector east to south (90–180°) of the site is of mixed type as rural/agricultural land is interspersed by various small industries in areas such as Jagadhri (~ 110 km from site) and Ambala (~ 50 km from site). Glass manufacturing units, pharmaceutical industries, solvent industries and paint industries are the major small scale industries present in this sector. The capital of India, New Delhi lies about 300 km south of the site and the national highways (NH-1 and NH-2) which connect Delhi to north west India also lie in this wind sector. The land use in the wind sector spanning south to north-north-west (180–315°) is mainly rural and agricultural land for several hundred kilometres except for two wind directions: Patiala (~ 60 km from site at wind direction of ~ 215°) and Ludhiana (~ 100 km from site at wind direction of ~ 285°). During the pre-monsoon season, the north west Indo–Gangetic

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



plain is characterized by strong westerly wind flow. This can be clearly seen in Fig. 1 which shows the ensemble 72 h back trajectories of air masses that arrived at the site everyday from 1 May 2012–31 May 2012 at 14:30 LT (UTC + 05:30). Back trajectory ensemble calculations were performed using the NOAA HYSPLIT model and GDAS (Global Data Acquisition System) meteorology (Draxler and Rolph, 2013; Draxler and Hess, 1998) for air masses arriving at the site (30.667° N, 76.729° E) at 20 m above the ground (the approximate sampling height for all instruments). The trajectories were overlaid on the ESA land use map. Due to the fact that the model does not always resolve the terrain gradient accurately for simulations near the Himalayan foothills, we applied an additional filter and only trajectories consistent with the measurement site, Mohali, being located in the plain (< 400 m a.m.s.l.) and Shimla, a mountain site at 31.103° N, 77.172° E, being located in the mountains (> 400 m a.m.s.l.) were considered and are shown in Fig. 1. Evidence for rapid long range transport from the desert regions of Afghanistan and sometimes even from Turkey and Iran in just 3 days can clearly be noted from the back trajectories. Comparison with the back trajectories calculated at 04:30 LT (not shown in Fig. 1) revealed that the westerly transport indicated by back trajectory for 14:30 LT was representative of the the study period and persisted even at night.

Figure 3 shows the wind rose plot for May 2012 as measured at the site. Both the local wind direction (Fig. 3) and the back trajectories (Fig. 1) are consistent and point out that the dominant direction from which air masses arrived at the site in May 2012 was from the wind sector spanning west to north-north-west. The land use is predominantly rural and agricultural land in this sector (Fig. 2) and the most frequent wind direction was 300–330°. Moreover the wind speeds of air masses arriving from this sector typically ranged from 5–20 ms<sup>-1</sup>. The wind rose in Fig. 3 also shows that air masses from the urban-industrial wind sector (0–90°) was also appreciable at wind speeds typically lower than 10 ms<sup>-1</sup>. Due to the occurrence of mountain ranges about 30 km north to north east of the site, some westerly air masses were deflected off the mountain slopes, and arrive at the site from the north to east wind sector (0–90°), advecting emissions



range from below 0 °C in winter and up to 45 °C in summer), all the inlet lines are maintained at a temperature of 40 °C. The PTR-MS sampling inlet is co-located with the trace gas analyzer inlets but the instrument itself is housed in a separate room on the first floor of the building and has a longer Teflon inlet line (12 m). A membrane pump was used to sample air at a large flow rate from the main inlet and the PTR-MS then drew circa 300 mL min<sup>-1</sup> flow through a Teflon T-connector upstream of the membrane pump. Tests conducted by spiking the main inlet with acetone established that the total inlet residence and sampling time in the PTR-MS was less than 16 s.

### 2.2.2 Volatile organic compound measurements using Proton Transfer Reaction Mass Spectrometry (PTR-MS)

A high sensitivity proton transfer reaction quadrupole mass spectrometer (HS Model 11-07HS-088; Ionicon Analytik Gesellschaft, Austria) was used in the selected ion monitoring mode for online measurements of selected volatile organic compounds. Within the instrument, organic species with a proton affinity greater than water are chemically ionised by proton transfer with the reagent hydronium (H<sub>3</sub>O<sup>+</sup>) ions and the products are detected using a quadrupole mass spectrometer (Lindinger et al., 1998). The instrument has four main parts:

1. an ion source which produces a pure stream of reagent H<sub>3</sub>O<sup>+</sup> ions (> 95%) by plasma discharge of water vapour,
2. a reaction chamber in which soft chemical ionization of analyte VOCs takes place,
3. a quadrupole mass analyzer that separates the product ions based on their *m/z* ratio and
4. a secondary electron multiplier for detection and amplification of the ion signal.

The technique of proton transfer reaction mass spectrometry and its application for detection of VOCs in air samples at ppt level has been extensively validated in urban,

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



forested and rural regions of the world and has been described in two comprehensive reviews by de Gouw et al. (2007) and Blake et al. (2009) and for more details pertaining to the technique, the reader is referred to these works.

Here we describe the instrumental settings, calibration procedure and operational details of the instrument during its deployment in the present study. Throughout the measurement period, the primary ion count (signal due to the main reagent  $\text{H}_3\text{O}^+$  ions) in the instrument was always above 14 million ions enabling excellent detection sensitivity. Further, the impurity ion signals due to  $\text{NO}^+$ ,  $\text{O}_2^+$  were always less than 0.6% and 3.3% of the primary ion signal respectively. The pressure and temperature in the reaction chamber of the mass spectrometer were maintained at 2.2 mbar and 333 K, respectively during the measurements with an E/N ratio in the drift tube of  $\sim 135$  Townsend. The volatile organic compounds reported in this work namely methanol, acetonitrile, acetone, acetaldehyde, isoprene, benzene, toluene, sum of C8-aromatics (xylenes and ethyl benzene) and sum of C9-aromatics (trimethyl benzenes) were measured in the selected ion monitoring mode with a dwell time of 1 s each at the relevant  $m/z$  channel. The instrumental background was determined at all the relevant  $m/z$  channels at frequent intervals by sampling cylinder zero air (5.0 grade Synthetic air; Sigmas Gases, New Delhi) that had been passed through an activated charcoal scrubber (Supelpure HC, Supelco, Bellefonte) and a heated VOC scrubber catalyst (Ionimed Analytik) maintained at 350 °C in series to ensure good background measurements. Table 1 lists the compound specific  $m/z$ , sensitivity, detection limits and total uncertainty. These VOC- $m/z$  identifications are in keeping with extensive validation studies from diverse ecosystems around the world often in combination with slower but more specific gas chromatography techniques (de Gouw et al., 2007). Previous PTR-MS studies (de Gouw et al. 2007 and references therein) have shown that the contribution to  $m/z$  59 and  $m/z$  69 are predominantly from acetone and isoprene, respectively but contribution from isobaric species such as propanal to  $m/z$  59 and furan to  $m/z$  69 cannot be ruled out from combustion sources. Hence the measurements

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



reported here for acetone and isoprene should be regarded as upper limits for these compounds.

The instrument was calibrated thrice during the period of study on 6 May, 14 May and 23 May 2012 at different humidities ( $\sim 0\%$  RH, 25% RH, 50% RH, 75% RH and 90% RH) by dynamic dilution with zero air at four different mixing ratios (in the range of 3–20 nmol mol<sup>-1</sup>) for each of the VOCs shown in Table 1 using a standard gas calibration unit (GCU-s v2.1, Ionimed Analytik, Innsbruck, Austria). The flows of both the standard gas and zero air mass flow controllers were also measured independently before and after the calibration experiments using a NIST calibrated flow meter (BIOS Drycal definer 220). Figure 1 of the supplementary material shows example plots from the calibration experiments which showed excellent linearity over the entire dynamic range ( $r^2 = 0.99$ ). Table 1 of the supplementary material lists the details of the VOC gas standards. The first VOC gas standard (Apel-Riemer Environmental, Inc., Colorado, USA) contained VOCs at a mixing ratio of circa 500 nmol mol<sup>-1</sup> (stated accuracy better than 5%) while the second gas standard canister (Ionimed Analytik) contained VOCs at circa 1  $\mu$ mol mol<sup>-1</sup> (stated accuracy better than 6%) and reasonable agreement was found for the sensitivity factors derived from the two different gas standards for common compounds such as isoprene (8.2 and 9 ncpsppbV<sup>-1</sup>) benzene (12.3 and 13.5 ncpsppbV<sup>-1</sup>) and toluene (13.4 and 14.3 ncpsppbV<sup>-1</sup>) considering the overall uncertainty of each calibration experiment ( $\sim 10\%$ ). Furthermore, no perceptible drift was observed in the sensitivity factors obtained during the three different calibration experiments. It has been reported previously (e.g., de Gouw et al., 2007; Sinha et al., 2009) that the sensitivity of certain compounds such as methanol, benzene and toluene can be influenced strongly by the humidity of the sample. The humidity of sampled air affects the partitioning of the primary ion signal within the PTR-MS between the hydronium ion (H<sub>3</sub>O<sup>+</sup>  $m/z = 19$ ) and the hydrated hydronium ion water clusters (e.g. H<sub>3</sub>O<sup>+</sup> (H<sub>2</sub>O);  $m/z = 37$ ). In our calibration experiments, we observed that normalizing the measured ion signals to both  $m/z = 19$  and  $m/z = 37$  for our instrumental settings of  $\sim 135$  Td yielded a rather constant sensitivity beyond 25% RH for all the compounds

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



reported in Table 1. This is consistent with previous studies (de Gouw et al., 2007) where at such E/N ratios in the drift tube, the influence of humidity changes became less significant. The sensitivity in ncpsppbV<sup>-1</sup> determined at 75 % RH was therefore used for all the VOCs reported in this work. For the *m/z* 33 channel at which methanol is detected, a non zero background can arise from the contribution of O<sub>2</sub><sup>+</sup> ions made up of <sup>17</sup>O and <sup>16</sup>O isotopes, and we took this into account for both the calibrations and ambient measurements by subtracting the *m/z* 32 signal contributing to the *m/z* 33 signal (i.e. *m/z* 32 × 0.00078). The measured *m/z* ion signals in counts per second (cps) were then converted to normalized counts per second (ncps) using the following normalization equation (previously discussed in Sinha et al., 2009; Tani et al., 2004),

$$\text{ncps} = \frac{I(\text{RH}^+) \times 10^6}{I(\text{H}_3\text{O}^+) + I(\text{H}_3\text{O}^+(\text{H}_2\text{O}))} \times \frac{2}{\rho_{\text{drift}}} \times \frac{T_{\text{drift}}}{298.15} \quad (1)$$

The normalized signals were then converted to ppbV using the sensitivity factors derived from the calibration experiments. The total uncertainty in each case was less than 13 % (see Table 1). This was calculated using the root mean square propagation of individual uncertainties like the 6 % accuracy error inherent in the VOC gas standard concentration, the 2σ instrumental precision error while sampling 6 ppbV of the VOC and error in the flow reproducibility of the two mass flow controllers.

### 2.2.3 Measurements of O<sub>3</sub>, NO<sub>x</sub>, CO, SO<sub>2</sub>, PM<sub>2.5</sub> and PM<sub>10</sub> mass concentrations

Table 2 provides an overview of the technique, temporal resolution, detection limit, uncertainty and instrument model for the trace gas and aerosol mass concentrations performed at the IISER Mohali air quality station. Each instrument was comprehensively characterized for its stated linearity, accuracy, precision and the total uncertainty of the measurements over the relevant ambient range. Multi point calibrations in the dynamic range of 100–1500 nmol mol<sup>-1</sup> for CO, 25–125 nmol mol<sup>-1</sup> for O<sub>3</sub>, 10–50 nmol mol<sup>-1</sup> for both NO<sub>x</sub> and SO<sub>2</sub> were performed using custom ordered standards (Chemtron Science Laboratories Pvt. Ltd., Mumbai). The mixing ratios of NO, CO and SO<sub>2</sub> in the gas





**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



inlet line passes through a solenoid valve which divides and alternates the flow between two streams, one of which passes through MnO<sub>2</sub> based ozone scrubber and serves as reference. Both the cells containing sample and reference are illuminated by single lamp and the intensity of the UV light is measured by photo diode detectors at the end of the cells. With ageing the efficiency of both the UV light source and the detector can decrease and to account for this and characterize the instrument's accuracy, a U.V. Photometric Ozone (O<sub>3</sub>) Primary Source Calibrator (Model 49iPS; Thermo Fischer Scientific, USA, NIST traceable) was used for zero drift checks (on 8, 19 and 24 May with a drift range of 0.4 to 0.5 nmol mol<sup>-1</sup>) and multipoint span calibrations (on 8 and 19 May).

NO<sub>x</sub> (NO and NO<sub>2</sub>) measurements were performed by Chemiluminescence technique using a Thermo Fischer Scientific 42i trace level analyzer. The detection principle is based on the reaction of NO with O<sub>3</sub> and subsequent production of luminescence in the wavelength range of 600–3000 nm. Ozone is generated inside the instrument using an ozonator to provide excess ozone for the chemiluminescence reaction. For measurement of NO<sub>2</sub>, all the NO<sub>2</sub> in the sample stream is converted to NO prior to reaction with O<sub>3</sub>. This is accomplished by a NO<sub>2</sub> to NO molybdenum converter heated at 325 °C (Nunnermacker, 1990). This converter is known to convert other reactive nitrogen species (e.g. nitric acid (HNO<sub>3</sub>), nitrous acid (HONO), nitrate radical (NO<sub>3</sub>), dinitrogen pentoxide (N<sub>2</sub>O<sub>5</sub>), peroxyxynitric acid (HNO<sub>4</sub>), alkyl nitrates (RONO<sub>2</sub>), peroxyalkyl nitrates (ROONO<sub>2</sub>) and peroxyacyl nitrates (PAN) (R(O)O<sub>2</sub>NO<sub>2</sub>)) into NO as well (Wang et al., 2002). Hence the NO<sub>2</sub> measurements should be treated as upper limits for the value of NO<sub>2</sub>. Similar to other analyzers, zero drift checks (on 8, 17, 21, 25, 28 and 31 May with a zero drift range of -0.05 to 0.05 nmol mol<sup>-1</sup>) and multipoint span calibrations were carried out for NO<sub>x</sub> analyzer.

Particulate matter PM<sub>10</sub> and PM<sub>2.5</sub> mass concentrations were measured using separate Thermo Fischer Scientific 5014i beta continuous ambient particulate monitors working on the principle of beta attenuation. Ambient air was drawn into the analyzers at a constant flow rate of 16.67 L min<sup>-1</sup> and the particulates deposited onto a filter

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



tape sample spot. The mass loading on the filter tape is continuously measured using attenuation of  $\beta$  rays from a  $^{14}\text{C}$  source. The interference caused by  $\beta$  rays emitted by daughter nuclides of natural Radon ( $^{222}\text{Rn}$ ) is corrected by measuring their airborne natural activity and then subtracting from gross particulate count rate. Dynamic change in ambient relative humidity (RH) and temperature ( $T$ ) affect PM mass measurements, as the liquid water content of the aerosol and the mass of semivolatile species is a function of temperature and relative humidity. While the EPA off-line reference method prescribes that the filters be conditioned for a minimum of 24 h (at 30–40 % RH and 20–23 °C) prior to determining the PM mass, on-line methods determining the PM mass have to strike a balance between determining the dry mass and excessive loss of semivolatile species due to inlet heating. At IISER Mohali Air Quality station we followed the protocol described in Takahashi et al. (2008) of continuously measuring the ambient RH at inlet sampling system and then heating the sample inlet to maintain a constant RH of 40 %. However, the inlet is not fitted for cooling and whenever ambient RH < 40 % the aerosol mass will be determined under ambient conditions (> 90 % of the time in pre-monsoon season). Span calibration was carried out once for both PM analyzers during the measurement period (on 21 May).

### 3 Results and discussion

#### 3.1 General trends in time series of gas phase species and aerosol mass concentrations: influence of varied biomass burning and an intense dust event

Figure 5 shows the measured one minute temporal resolution data for the mixing ratios of methanol, acetaldehyde, sum of acetone and propanal (top panel), isoprene and acetonitrile (middle panel) and benzene, toluene, sum of C8-aromatics (xylene isomers and ethyl benzene), sum of C9 aromatics (isomers of trimethyl benzenes and propyl benzenes) (bottom panel). Similarly, Fig. 6 shows the measured one minute temporal

resolution data for the mixing ratios of ozone (top panel), NO<sub>2</sub> and NO (second panel), SO<sub>2</sub> (third panel), CO (fourth panel) and PM<sub>2.5</sub> and PM<sub>10</sub> (bottom panel).

It can be seen that extremely high levels were observed in the evening and early morning hours for both VOCs and the primary air pollutants during certain pollution events in May 2012 (e.g. 3 May 19:23–23:33 LT; 7 May 08:10–09:37 LT; 10 May 07:15–08:00 LT; 11 May 19:39–20:39 LT; 16 May 05:38–06:34 LT; 31 May 21:00–23:59 LT in both Figs. 5 and 6). Elevated levels of acetonitrile ( $> 6 \text{ nmol mol}^{-1}$ ), a specific chemical tracer for biomass burning (Holzinger et al., 1999), suggest that the air plumes (some of which arrived at wind speeds of  $\sim 4 \text{ ms}^{-1}$ ) had passed over regions where intense biomass burning was active. Using in-situ PTR-MS data at the same site in 2012, (Sarkar et al., 2013) showed that paddy residue burning in October–November released large amounts of carcinogenic benzenoids. Previous in-situ studies within India, have demonstrated that the rice wheat (RW) cropping system and associated crop residue burning activity releases vast amounts of CO, NO<sub>x</sub>, CH<sub>4</sub> and CO<sub>2</sub> (Gupta et al., 2004; Sahai et al., 2007). Analysis of the regional MODIS fire count data (see Fig. 3 of the Supplement) and air mass back trajectories for the plume events, revealed that wheat residue burning was quite widespread in the upwind regions and was the likely cause for the extremely high plume values of circa  $100 \text{ nmol mol}^{-1}$  for methanol,  $50 \text{ nmol mol}^{-1}$  for sum of acetone and propanal and  $30 \text{ nmol mol}^{-1}$  for acetaldehyde,  $7 \text{ nmol mol}^{-1}$  for acetonitrile,  $5 \text{ nmol mol}^{-1}$  for isoprene and  $20 \text{ nmol mol}^{-1}$  for each of the aromatics,  $80 \text{ nmol mol}^{-1}$  NO<sub>x</sub>,  $50 \text{ nmol mol}^{-1}$  SO<sub>2</sub> and  $2000 \text{ nmol mol}^{-1}$  CO observed in Figs. 5 and 6). Although no in-situ data on the release of volatile organic compounds from crop residue burning in India has been reported till date, VOC emissions due to crop residue burning in the months of May–April and October–November in the agricultural regions of India is expected to be significant (Venkataraman et al., 2006). Peaks were observed in the mixing ratios of methanol, acetontirile and CO (Fig. 5) even in periods that were not impacted strongly by wheat residue burning plumes with high values of circa  $40 \text{ nmol mol}^{-1}$  for methanol,  $4 \text{ nmol mol}^{-1}$  for acetonitrile and  $1000 \text{ nmol mol}^{-1}$  for CO. This indicates the possible presence of an industrial

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



source for methanol (which is examined further in Sect. 3.4) and occurrence of other types of biomass combustion activity. During these other “non-wheat residue burning” biomass burning peaks, aromatics and isoprene also appear to have been co-emitted. These peaks occurred primarily at night and in the morning hours and it is likely that activities involving biofuel (wood) combustion for domestic cooking (Venkataraman et al., 2005), periodic open burning of leaf litter and plastic on streets and burning for disposal of waste and assorted garbage by small shops after the day’s business were the main causes. Burning frequency especially for leaf litter is higher on days with low wind speeds so that the fire does not get out of control; however such stagnant conditions also create very favourable conditions for accumulation of the emissions under a stable nocturnal boundary layer resulting in high mixing ratios for the VOCs and trace gases. Overlaying such emission activity in the time series are the bimodal traffic related emission peaks, which contributed to the morning and evening rush hour peaks in the time series profile and diel profiles (discussed in Sect. 3.5) of the benzenoid compounds, NO<sub>x</sub> and CO.

Tropical vegetation dominates the global source of isoprene which is normally the major contributor to  $m/z$  69 in the PTR-MS (de Gouw et al., 2007). It is also known that isoprene is emitted from biomass burning (Christian et al., 2003; Warneke et al., 2011; Andreae and Merlet, 2001) and to a minor extent from traffic (Borbon et al., 2001). On the other hand, compounds such as furan which are emitted from combustion can also contribute to the detected  $m/z$  69 signal causing an overestimate of the true value of isoprene if the signal at  $m/z$  69 is attributed only to isoprene. Our group performed measurements with a proton transfer reaction time of flight mass spectrometer (PTR-TOF-MS) in Kathmandu Valley in December–January of 2012–13, a period when biomass burning is quite intense. In contrast to a proton transfer reaction mass spectrometer equipped with a quadrupole mass analyzer (the instrument used in this study), a PTR-TOF-MS is able to distinguish between furan and isoprene due to its higher mass resolution. In the Kathmandu study (Sarkar et al. personal communication) we found that the signal at  $m/z$  69 was mainly from isoprene and furan

**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



contributed only a maximum of 25 % to  $m/z$  69 even during peak evening combustion activity. Therefore, we are confident that our measurements of  $m/z$  69 are mainly due to protonated isoprene at all times with only minor contribution from isobaric ion species such as protonated furan. There are several eucalyptus forests in the N. W. IGP region so in the daytime, biogenic emission of isoprene could be a significant contributor to isoprene levels as well and we examine the strength of the biogenic and biomass burning sources of isoprene in Sect. 3.3 using inter VOC correlations with emission tracers such as acetonitrile. Irrespective of the source, the presence of high levels of isoprene has significant implications for ozone formation potential and hydroxyl radical mediated atmospheric oxidation (Lelieveld et al., 2008; Fuchs et al., 2013). In widely used state of the art global emission models for isoprene such as MEGAN (Guenther et al., 2006), due to unavailability of in-situ data from the IGP region thus far, the emission strength of isoprene from the IGP region is not well not constrained and these first measurements of isoprene in ambient surface air of the IGP suggest that there could be high emissions of isoprene in the region.

In Fig. 6 it can be seen that daytime ozone levels frequently reached values of  $80 \text{ nmol mol}^{-1}$  on a daily basis in May 2012 except for 17 May, a day affected by a massive dust storm. Being a downwind suburban site several kilometers away from the city centres, air masses arriving at the site especially from the predominant north west wind sector (see Fig. 3), would have had considerable time for photochemical production of ozone within the air mass from precursors such as  $\text{NO}_x$  and reactive VOCs contributing to the high ozone downwind. This was reflected in the fact that the 8 h national ambient air quality limit of  $100 \mu\text{g O}_3 \text{ m}^{-3}$  stipulated by the Ministry of Environment and Forests, India was exceeded on 29 out of 31 days in May 2012 at the site. The dry air (daytime relative humidity was generally below 20 % RH), high temperatures (daytime temperature ranged from  $32\text{--}44^\circ\text{C}$ ), strong solar radiation ( $600\text{--}800 \text{ W m}^{-2}$ ), availability of at least  $4\text{--}5 \text{ nmol mol}^{-1} \text{ NO}_x$  and significant levels of reactive VOCs such as acetaldehyde and isoprene during the daytime were obviously conducive for active in-situ photochemical production of ozone as well. Kirchner et al. (2001) and Sinha et al. (2012)

**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



have previously shown that an indicator  $\Theta$  defined as the ratio of  $\text{NO}_x$  OH Reactivity/VOC OH Reactivity, can be used to assess whether the peak ozone production falls within a  $\text{NO}_x$  or VOC limited regime. For a ratio of  $\Theta > 0.2 (\pm 0.1)$ , the ozone production regime is VOC limited, whereas for  $\Theta$  below 0.01 the ozone production regime would be primarily  $\text{NO}_x$  limited. The intermediate range,  $0.01 < \Theta < 0.2$ , indicates that the peak ozone production depends strongly on both  $\text{NO}_x$  and VOC levels. During May 2012, the 1 min daytime values of  $\Theta$  were generally in the intermediate range, suggesting that peak ozone production was limited by availability of both  $\text{NO}_x$  and VOCs. However since the OH reactivity was not measured directly in May 2012, the VOC OH reactivity value employed in the calculation could be an underestimate of the actual VOC OH reactivity value, in turn implying that  $\Theta$  could in fact be lower. Hence, based on the present data, the peak ozone production regime tendency is either VOC and  $\text{NO}_x$  limited ( $0.01 < \Theta < 0.2$ ) or  $\text{NO}_x$  limited ( $\Theta < 0.01$ ). Direct OH reactivity measurements at the site in future for example using the comparative reactivity method (Sinha et al., 2008) would help constrain the ozone production regime more accurately. At night during certain periods (e.g. 3, 4, 10 and 30 May), the ozone mixing ratios decreased to almost zero due to titration with high NO plumes. Generally however ozone levels far exceeded NO levels at night with average values of  $40 \text{ nmol mol}^{-1} \text{O}_3$  even at night. At  $40 \text{ nmol mol}^{-1} \text{O}_3$ , the lifetime of NO against oxidation by  $\text{O}_3$  at night to  $\text{NO}_2$  is of the order of few minutes. Hence it is not surprising that NO levels were typically very low at the site ( $< 0.3 \text{ nmol mol}^{-1}$ ), as transport times for air masses to the suburban site exceeded the chemical titration time. Only wheat residue burning plumes containing high NO that were advected to the site or very fresh traffic exhaust emissions reaching the site at night caused the occurrence of high nighttime NO observed in Fig. 6. Figures 5 and 6 also show that an intense dust event occurred from 03:00 17 May–09:00 18 May (shown as shaded rectangles in Figs. 5 and 6; times are local time). The 72 h back trajectories of the air masses indicated long range transport of dust from the west towards the site (Fig. 1; yellow trajectories). MODIS Terra imagery from the NASA Earth Observatory reveals that the dust originated in the Arabian Peninsula and





## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



acts equally for all primary pollutants and so nighttime data are optimal for constraining emission sources. In this analysis we used acetonitrile as a proxy for biomass burning sources such as residential biofuel cooking, garbage burning and leaf litter. This is reasonable since nighttime correlation of acetonitrile with  $\text{NO}_x$  which was mainly emitted by traffic exhaust was insignificant ( $r^2 < 0.2$ ). Aromatics such as toluene, and the C8 and C9-aromatics are known to be co-emitted in traffic exhaust and indeed were found to correlate with  $\text{NO}_x$  ( $r^2 \geq 0.5$  in each case) but again not significantly with acetonitrile ( $r^2 = 0.19$ ). Thus, strong correlation with acetonitrile and lack of correlation with toluene and the C8 and C9-aromatics for any VOC would imply that the VOC has a significant biomass burning source at night. Conversely, strong correlation with toluene and the C8 and C9-aromatics and lack of correlation with acetonitrile for any VOC implies that the VOC has a significant traffic/industrial emission source at night. It can be seen in Table 5 that oxygenated VOCs seemed to have a strong biomass burning source at night as all of the OVOCs correlated significantly with acetonitrile (correlation coefficients: acetaldehyde, acetonitrile ( $r^2 = 0.8$ ); methanol, acetonitrile ( $r^2 = 0.6$ ); acetone, acetonitrile ( $r^2 = 0.6$ )). Note also that the oxygenated VOCs had poor correlation with toluene and the C8 and C9-aromatics ( $r^2 < 0.3$  in each case) suggesting that traffic was at best a minor source for the oxygenated VOCs at night. It is known that both isoprene and benzene are emitted from traffic exhaust (Borbon et al., 2001) and biomass burning (Christian et al., 2003; Holzinger et al., 1999). Based on the observed correlations it appears that isoprene was emitted from biomass burning significantly at night ( $r^2 = 0.5$ ) with some contribution also from traffic emissions ( $r^2 = 0.4$ ) whereas benzene was emitted significantly from traffic emission sources at night ( $r^2 \geq 0.6$  with other aromatics) with some contribution from biomass burning sources as well ( $r^2 = 0.4$  with acetonitrile). Generally isoprene is considered only to be emitted biogenically during the day. Based on the results in this study however, significant levels of nighttime isoprene can also result from nighttime biomass burning and traffic emissions into a shallow nocturnal boundary layer. The influence on daytime photochemistry due to high isoprene has already been mentioned in Sect. 3.1. The presence of nighttime isoprene

could affect oxidation chemistry at night too and influence  $\text{NO}_x$  transport through  $\text{NO}_x$  sequestering by isoprene nitrates (Browne and Cohen, 2012).

### 3.4 Identification of point sources contributing to high methanol and isoprene using bivariate polar graphs

In Sects. 3.1 and 3.2, we saw that the average levels of methanol and isoprene observed at our site are generally higher compared to what has been reported in summertime surface air at most suburban and urban sites elsewhere in the world. This suggests that there could be strong point sources for these compounds upwind of our site that affect the levels of these compounds regionally. In order to examine this possibility we plotted bivariate polar graphs (Carslaw et al., 2009; Carslaw and Ropkins, 2012) wherein the measured concentrations were averaged in separate wind speed and wind direction bins corresponding to sizes of  $1 \text{ ms}^{-1}$  and  $10^\circ$ , respectively. The aforementioned binning process was then used for conversion to polar coordinates which were further interpolated using Kriging technique to make the bivariate polar plot (Carslaw et al., 2009). Figure 7 shows the bivariate polar plots for daytime methanol and isoprene. Daytime data was chosen as it had the largest fetch region upwind of the site due to high wind speeds. In Fig. 7a, it can be seen that the highest values of circa  $38 \text{ nmol mol}^{-1}$  methanol were observed from the urban and industrial land use areas located north to south east of the measurement site, at wind speeds of  $1\text{--}5 \text{ ms}^{-1}$ . As described in Sect. 3.1, several industries were present in this fetch region including pharmaceutical industries, solvent industries and paint industries as well as glass manufacturing units. In addition there are at least 4–5 large sewage and wastewater treatment plants in Chandigarh and Mohali, which could be a source of leaked methanol to the atmosphere. Methanol is used in water and sewage treatment plants as a biodegradable source of carbon for microbes which are employed in de-nitrification of the treated waste water and sewage (Tchobanoglous et al., 2003). In Fig. 7b, we note that in contrast to methanol for which the dominant sources lay in the urban and industrial sector (north to south east), for isoprene, the major sources were west of the site

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



in the mainly rural/agricultural sector. Significant contribution also appeared to come from the urban sector at wind speeds of 10–15 ms<sup>-1</sup> likely from the mosaic cropland and vegetation cover beyond Chandigarh (Fig. 2). The rural/agricultural land west of the site has large swathes of poplar and eucalyptus trees at the periphery of agricultural fields (due to mixed agro-forestry practices) and there are also a few plantations of poplar and eucalyptus trees. Eucalyptus and poplar trees have very high isoprene emission potential (Kesselmeier and Staudt, 1999; Sharkey and Singaas, 1995). It is also worth mentioning that the city of Chandigarh has significant eucalyptus tree cover lining roads (e.g. Sector 17 in Chandigarh city) and several patches of tree stands in the many parks that dot the city landscape. Thus the high daytime isoprene is likely due to biogenic emissions from mixed agro forestry and urban tree cover. For both methanol and isoprene in Fig. 7a and b, extremely high values were observed for the 280–290° wind direction at wind speeds of 20–25 ms<sup>-1</sup>. One of the limitations of the bivariate plot is that it can be biased strongly by statistical effects (a unique event with peculiar meteorology) for bivariate bins located near the border of the graph. Figure 3 shows that during May 2012 very few data periods (< 1%) were associated with wind speed bin of 20–25 ms<sup>-1</sup>. On examining the time series closely we found that these data points corresponded to 2 h of data between 07:25 LT and 09:25 h LT on 21 May 2012 (see Fig. 5), in which acetonitrile was also quite high (8 nmol mol<sup>-1</sup>) and hence are likely due to a biomass burning event. While there could certainly be strong regional sources of isoprene further upwind of the site in the north west direction, due to the short atmospheric lifetime of isoprene in the daytime (< 1.5 h), the impact of any isoprene source more than 80 km away from the measurement site would be marginal at best.

### 3.5 Diel box and whisker plots of volatile organic compounds, ozone, carbon monoxide, and nitrogen oxides

The diel variability of chemical species in the atmosphere is driven by atmospheric dynamics, emissions, oxidation and deposition processes. Thus diel box and whisker

plots of chemical species can be used to yield insights on emission activity, dynamics and the chemistry of the atmosphere.

Figure 8 shows the diel box and whisker plots for the ambient temperature, radiation, ozone, carbon monoxide, nitrogen monoxide and nitrogen dioxide derived from all measurements ( $n > 14\,300$  for each species) made in May 2012 at the site. In order to minimize the influence of severe pollution episodes and very local events, the time periods affected by wheat residue burning and periods when the wind speed was less than  $3\text{ ms}^{-1}$  were filtered out ( $\sim 25\%$  of data; see also Fig. 3).

The width of the box represents the interquartile range (25th and 75th percentile) while the whiskers represent the 10th and 90th percentile of the data. In each plot, the median is shown as a cross while the average is shown as a solid black circle. The time on the horizontal axis represents the start time of the corresponding hourly bin. It can be seen from the solar radiation profile, that sunrise typically occurred between 05:00 and 06:00 while sunset typically occurred between 18:00 and 19:00 LT. The average and median values for ozone, temperature and solar radiation were similar for all hours of the day and the range in the median hourly values for the ambient temperature and ozone were  $26\text{--}36\text{ }^{\circ}\text{C}$  and  $35\text{--}80\text{ nmol mol}^{-1}$ , respectively. Carbon monoxide and nitrogen dioxide showed a bimodal profile characteristic of surface sites with urban influence as early morning (05:00–07:00 LT) and evening emissions (18:00–21:00 LT), released into a shallow boundary layer result in peak mixing ratios for both species at these times of the day. Being a sub-urban site, rapid conversion of nitrogen monoxide to nitrogen dioxide due to the generally high ozone levels present regionally, resulted in median hourly nitrogen monoxide mixing ratios of only few hundred  $\text{pmol mol}^{-1}$ , and only when very fresh plumes arrived at the site sporadically during the night, the average hourly value of nitrogen monoxide sometimes exceeded  $1\text{ nmol mol}^{-1}$ . After 07:00 LT till 11:00 LT, decreases were observed in median profiles of all primary pollutants due to the dilution effect exerted by the growing day time boundary layer (the average ambient temperature increased from  $28\text{ }^{\circ}\text{C}$  to  $32\text{ }^{\circ}\text{C}$  in the same period). For ozone which is a secondary pollutant, photochemical production and entrainment of ozone rich air

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

from above resulted in a net increase of  $10 \text{ nmol mol}^{-1} \text{ O}_3 \text{ h}^{-1}$  from 06:00–11:00 LT. From 11:00 LT until sunset, both average and median ozone levels remained constant at a rather high value of  $80 \text{ nmol mol}^{-1}$ . NOAA HYSPLIT back trajectory model simulations for the evolution of the daytime boundary layer predicted an average mixed-boudary layer height of 850 m at 07:30 LT and 2450 m at 14:00 LT, indicating that despite an increase by a factor of 3 in the mixed layer height, chemical production of ozone, entrainment and transport of ozone rich air, sustained the high average ozone levels of  $80 \text{ nmol mol}^{-1}$  reached between 11:00–12:00 LT, until sunset.

Figure 9 shows the box and whisker plots for methanol, acetonitrile, acetaldehyde, isoprene, benzene and toluene. Bimodal profiles similar to those observed for primary pollutants such as carbon monoxide and nitrogen dioxide were observed for methanol, acetonitrile, acetaldehyde, isoprene and benzene. For toluene, the morning peak was less pronounced. Toluene appeared to be emitted mainly in traffic emissions with minor contribution from daily biomass burning (see Sect. 3.3 on inter VOC correlations). The diel profile of toluene therefore suggests that the strength of the morning biomass burning emissions was stronger relative to that of the morning traffic emissions, while in the evening time, emissions due to both activities were comparable. It is also interesting to note that the VOC mixing ratios decreased by a factor of two between 07:00–10:00 LT due to the dilution effect of the day time boundary layer growth for both isoprene ( $2 \text{ nmol mol}^{-1}$  to  $1.3 \text{ nmol mol}^{-1}$ ) and benzene ( $2 \text{ nmol mol}^{-1}$  to  $1 \text{ nmol mol}^{-1}$ ). If there was no day time emission of isoprene between 07:00–10:00 LT, considering that isoprene is oxidized 100 times faster than benzene by ambient hydroxyl radicals ( $k_{\text{isoprene}+\text{OH}} = 1.1 \times 10^{-10}$  at 298 K;  $k_{\text{benzene}+\text{OH}} = 1.2 \times 10^{-12}$  at 298 K; Atkinson et al., 2003), a sharper decrease would have been expected for isoprene relative to benzene between 07:00–10:00 LT. Since this is clearly not the case, the observed diel profiles provide further evidence (see Sect. 3.4) for day time emission of isoprene from biogenic sources even in the dry scorching summer heat. Differences in the diel profiles of benzene and toluene indicate that though they may have a strong and common traffic emission source (e.g. between 18:00–21:00 LT), some benzene is released from biomass

burning as well ( $r^2$  for benzene and acetonitrile in night time data = 0.4). Gustafsson et al. (2009), reported that the contributions of biomass and fossil fuel combustion are equally important for the formation of atmospheric brown clouds over South Asia in winter. Our in-situ measurements of a suite of VOCs and primary pollutants including acetonitrile, a biomass burning tracer suggest that even in summer, varied forms of regional biomass burning strongly impact the regional chemical composition of the atmosphere.

## 4 Conclusions

This study is the first ever to deploy a proton transfer reaction mass spectrometer (PTR-MS) for measurements of a suite of volatile organic compounds in India. A comprehensive description of the new measurement facility, regional nature of the measurement site and the pre-monsoon meteorology have been presented in this work. The suburban site experiences regional emission influences, rendering it extremely useful for chemical characterization of regional inflow and outflow of air masses depending on the prevalent meteorology and season.

The deployment of the PTR-MS in parallel with an air quality and meteorological station has resulted in a novel chemical dataset that highlights the key emission influences occurring in pre-monsoon air of the north west Indo-Gangetic Plain, one of the most densely populated but understudied regions of the world in terms of ambient chemical composition. Information regarding the in-situ levels of volatile organic compounds such as acetonitrile, methanol, acetone, acetaldehyde and isoprene in Indian air is available for the first time ever at high temporal resolution and was used to characterize the impact of emission activities on regional atmospheric composition in the N. W. Indo-Gangetic Plain. The occurrence of elevated levels of acetonitrile ( $> 6 \text{ nmol mol}^{-1}$ ) during May 2012, highlighted the strong influence of biomass burning in this region of the world even in summertime, when temperatures at night are amiable and no heating is necessary to counter cold nights. The emissions of regional wheat residue burning in

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





biogenic emissions from poplar and eucalyptus trees as well as biomass burning at night.

Studies on the chemistry of isoprene and recycling of hydroxyl radicals, the main atmospheric oxidant, in ambient field studies would be extremely interesting in such an environment. The results obtained in this study clearly show that the emissions and regime (high temperature, low RH) for photochemical reactions in the North West Indo–Gangetic Plain in the pre-monsoon summer season make it a unique “natural” setting to study high surface temperature and high mineral dust regime gas phase atmospheric chemistry, which could help understand global warming perturbed atmospheric chemistry better. The results of the present work also suggest that the North West Indo–Gangetic plain due to its diverse emission sources, in particular the varied biomass burning may be a very significant but hitherto poorly constrained source of volatile organic compounds. Future studies should focus on the spatial and seasonal variability of emissions and chemistry occurring in the IGP surface air and also address knowledge gaps in basic atmospheric chemistry processes pertaining to atmospheric oxidation by radicals and heterogeneous dust uptake and chemistry. Such scientific know how would be immensely useful for devising the best mitigation strategies from an air quality, health and climate perspective for the region.

**Supplementary material related to this article is available online at**  
**[http://www.atmos-chem-phys-discuss.net/13/31761/2013/  
acpd-13-31761-2013-supplement.pdf](http://www.atmos-chem-phys-discuss.net/13/31761/2013/acpd-13-31761-2013-supplement.pdf)**

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## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**High surface ozone  
and strong influence  
of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**High surface ozone  
and strong influence  
of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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**High surface ozone  
and strong influence  
of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

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**High surface ozone and strong influence of biomass burning**

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.

**Table 1.** Compound specific VOC  $m/z$  assignments, sensitivity, detection limit and uncertainty. The limit of detection is defined as  $2\sigma$  of the measured normalized signal while measuring zero air divided by the sensitivity.

VOC	Nominal protonated $m/z$ (Th)	Sensitivity (ncpsppbv <sup>-1</sup> )	Limit of detection (nmolmol <sup>-1</sup> )	Uncertainty
Methanol	33	10.1	0.46	12.3 %
Acetonitrile	42	20.7	0.05	9.1 %
Acetaldehyde	45	20.2	0.09	8.4 %
Acetone	59	22.8	0.06	8.8 %
Isoprene	69	9.0	0.12	10 %
Benzene	79	13.5	0.07	9.5 %
Toluene	93	14.3	0.08	8.6 %
<i>p</i> -Xylene	107	13.8	0.10	10.8 %
1,2,4-Trimethylbenzene	121	11.2	0.13	11.1 %

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


## High surface ozone and strong influence of biomass burning

V. Sinha et al.

**Table 2.** Species measured at the IISER Mohali air quality station with respective technique, temporal resolution, detection limit, uncertainty and instrument model.

Measured species	Technique	Time resolution	Detection limit	Uncertainty	Instrument model
CO	Non-Dispersive Infra Red (NDIR)	1 min	30 nmol mol <sup>-1</sup>	< 6 %	Thermo Fisher Model 48i
NO <sub>x</sub>	Chemiluminescence	1 min	150 pmol mol <sup>-1</sup>	< 6 %	Thermo Fisher Model 42i
O <sub>3</sub>	UV-Photometry	1 min	1 nmol mol <sup>-1</sup>	< 6 %	Thermo Fisher Model 49i
SO <sub>2</sub>	Pulsed UV-Fluorescence	1 min	1 nmol mol <sup>-1</sup>	< 6 %	Thermo Fisher Model 43i
PM <sub>10</sub>	β attenuation	1 min	< 4 μg m <sup>-3</sup>	< 8 %	Thermo Fisher Model 5014i Beta
PM <sub>2.5</sub>	β attenuation	1 min	< 4 μg m <sup>-3</sup>	< 8 %	Thermo Fisher Model 5014i Beta

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table 3.** Average, median and standard deviation of VOCs measured at IISER Mohali in May 2012 and comparison with average levels reported in summertime from other cities of the world. All values are reported in  $\text{nmol mol}^{-1}$ .

Measured VOC	Mohali average <sup>a</sup>	Mohali median	Paris <sup>b</sup> 2007	Beijing <sup>c</sup> 2005	Mexico City <sup>d</sup> 2003	Tokyo <sup>f</sup> 2002	Texas <sup>h</sup> 2000	London <sup>i</sup> 2006
Methanol	37.5 (17.9)	34.1	7.8	11.7	23.2	36.5 <sup>g</sup>	10.8	
Acetonitrile	1.4 (0.9)	1.1	0.7		0.7	0.5–5.5 <sup>j</sup>	0.5	0.3
Acetaldehyde	6.7 (3.7)	5.8	2.5	3.6	6.5	4.5–12 <sup>j</sup>	3.4	3.6
Acetone	5.9 (3.7)	5	7.1	4.4	8	10–50	3	1.6
Isoprene	1.9 (0.9)	1.6	0.6	0.2–1.1 <sup>j</sup>	0.4–2 <sup>e,j</sup>	0.1–0.7 <sup>j</sup>	0.3	0.2
Benzene	1.7 (1.5)	1.3	1.2		1.7	1–4 <sup>j</sup>	0.6	0.2
Toluene	2.7 (2.9)	1.8	5.2	1–4 <sup>j</sup>	7.2	3–9 <sup>j</sup>	0.8	1.9
C8-aromatics	2 (2.2)	1.3	5.5		3.3	1.3 <sup>g</sup>	0.6	0.2
C9-aromatics	1.2 (1.3)	0.9	4			0.3 <sup>g</sup>		

<sup>a</sup> Number in parenthesis represents the variability as the standard deviation.

<sup>b</sup> Gros et al. (2011).

<sup>c</sup> Shao et al. (2009).

<sup>d</sup> Lamb et al. (2004).

<sup>e</sup> Range was taken from Fortner et al. (2009).

<sup>f</sup> Kato et al. (2004).

<sup>g</sup> Daytime average values were taken from Yoshino et al. (2012).

<sup>h</sup> Karl et al. (2003).

<sup>i</sup> Langford et al. (2010).

<sup>j</sup> Only ranges were reported.

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table 4.** Average, median and standard deviation of measured parameters at air quality station at IISER Mohali in May 2012 and comparison with some previous works.

Measured species	Mohali monthly average <sup>a</sup>	Mohali monthly median	Ahmedabad May 1991–95 <sup>b,g</sup>	Pune summer 2004 <sup>c</sup>	Kolkata May 2004 <sup>d</sup>	Delhi May 1997–98 <sup>e,g</sup>
CO <sup>h</sup>	566.7 (293.2)	500	50–1000	372.1	600	2600–4400
SO <sub>2</sub> <sup>h</sup>	7.3 (7.1)	5.1				6–10
NO <sup>h</sup>	0.8 (3)	0.3				
NO <sub>2</sub> <sup>h</sup>	10.7 (7.7)	8.7				
NO <sub>x</sub> <sup>h</sup>	11.5 (9)	9.1	3–20	10.4	40–70 <sup>g</sup>	25–40
O <sub>3</sub> <sup>h</sup>	57.8 (25.4)	53.5	5–40	43	35–75 <sup>g</sup>	45–55 <sup>f</sup>
PM <sub>10</sub> μg m <sup>-3</sup>	276 (292.2)	220.2				
PM <sub>2.5</sub> μg m <sup>-3</sup>	104 (80.3)	89.1				
Ambient temperature (°C)	32.2 (5)	32.1				
Relative Humidity (%)	24.7 (10.2)	23.8				

<sup>a</sup> Number in parenthesis represents the variability as the standard deviation.

<sup>b</sup> Lal et al. (2000).

<sup>c</sup> Beig et al. (2007).

<sup>d</sup> Purkait et al. (2009).

<sup>e</sup> Aneja et al. (2001).

<sup>f</sup> Value was taken from Kulkarni et al. (2010).

<sup>g</sup> Only monthly average ranges were reported.

<sup>h</sup> Values are reported in nmol mol<sup>-1</sup>.

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

**Table 5.** Inter VOC correlation analysis of nighttime data.  $r^2$  values  $\geq 0.5$  are shown in bold.

	Methanol	Acetonitrile	Acetaldehyde	Acetone	Isoprene	Benzene	Toluene	C8-aromatics	C9-aromatics
Methanol	1	<b>0.6</b>	<b>0.5</b>	<b>0.6</b>	0.4	<b>0.5</b>	0.3	0.3	0.2
Acetonitrile	–	1	<b>0.8</b>	<b>0.5</b>	<b>0.5</b>	0.4	0.2	0.2	0.2
Acetaldehyde	–	–	1	0.4	<b>0.6</b>	0.4	0.2	0.2	0.2
Acetone	–	–	–	1	0.3	0.4	0.2	0.2	0.2
Isoprene	–	–	–	–	1	<b>0.5</b>	0.4	0.4	0.3
Benzene	–	–	–	–	–	1	<b>0.6</b>	<b>0.8</b>	<b>0.6</b>
Toluene	–	–	–	–	–	–	1	<b>0.7</b>	0.4
C8-aromatics	–	–	–	–	–	–	–	1	<b>0.7</b>
C9-aromatics	–	–	–	–	–	–	–	–	1

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

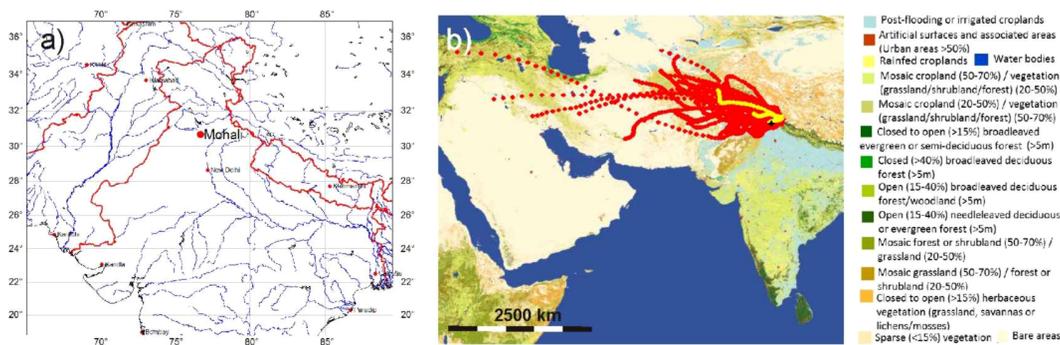
Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.



**Fig. 1.** (a) Location of the city of Mohali in the north west Indo–Gangetic plain (30.667° N, 76.729° E, 310 m a.s.l.) (b) 72 h back trajectories of air masses that arrived at the site (site shown as black marker) everyday from 1 May 2012–31 May 2012 at 14:30 LT (UTC + 05:30) superimposed on a land classification map (courtesy ESA GlobCover 2009 Project).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

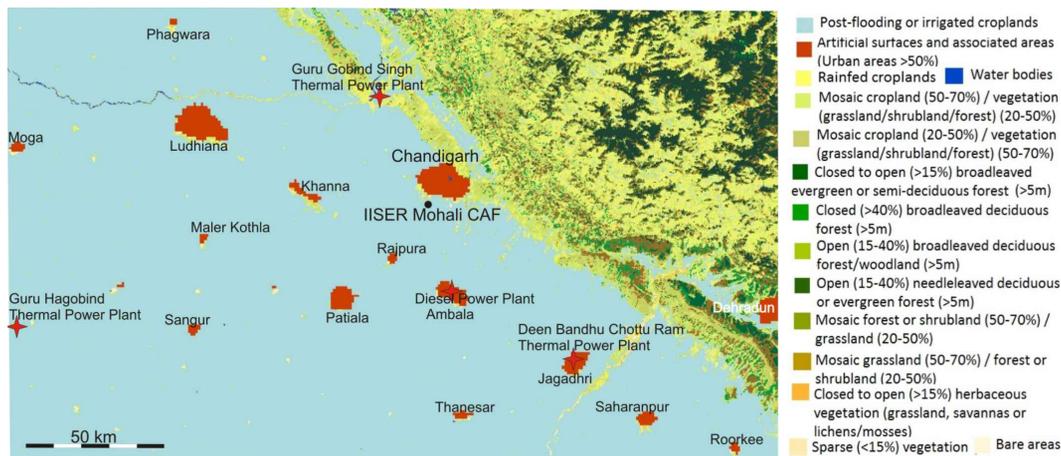
Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.



**Fig. 2.** Map of the land classification (courtesy ESA GlobCover 2009 Project) in a 100 km × 200 km area surrounding the measurement site (black dot).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



High surface ozone  
and strong influence  
of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

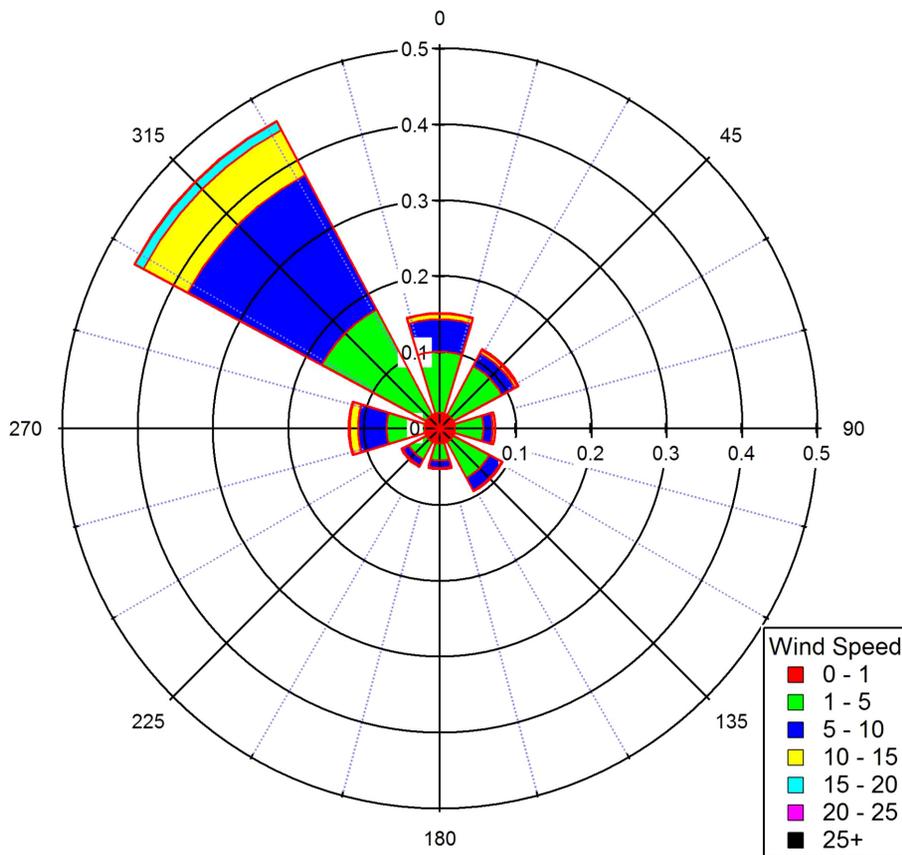
Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Fig. 3.** Wind rose plot derived from in-situ one minute wind direction and wind speed data at the measurement site ( $30.667^{\circ}$  N,  $76.729^{\circ}$  E, 310 ma.s.l.) in May 2012.



**Fig. 4.** Picture of the co-located instrument inlets and meteorological sensors ( $\sim 20$  m.a.g.l.) at the IISER Mohali atmospheric chemistry facility.

High surface ozone and strong influence of biomass burning

V. Sinha et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

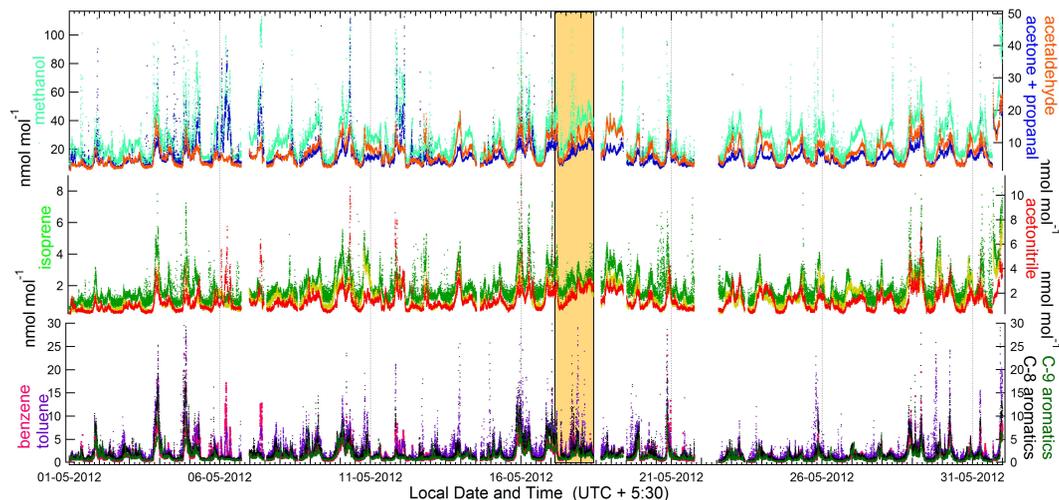
Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.



**Fig. 5.** Time series of the one minute data in May 2012 for the mixing ratios of methanol, acetaldehyde, sum of acetone and propanal (top panel), isoprene and acetonitrile (middle panel) and benzene, toluene, sum of C8-aromatics (xylene isomers and ethyl benzene), sum of C9-aromatics (isomers of trimethyl benzenes and propyl benzenes) (bottom panel).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

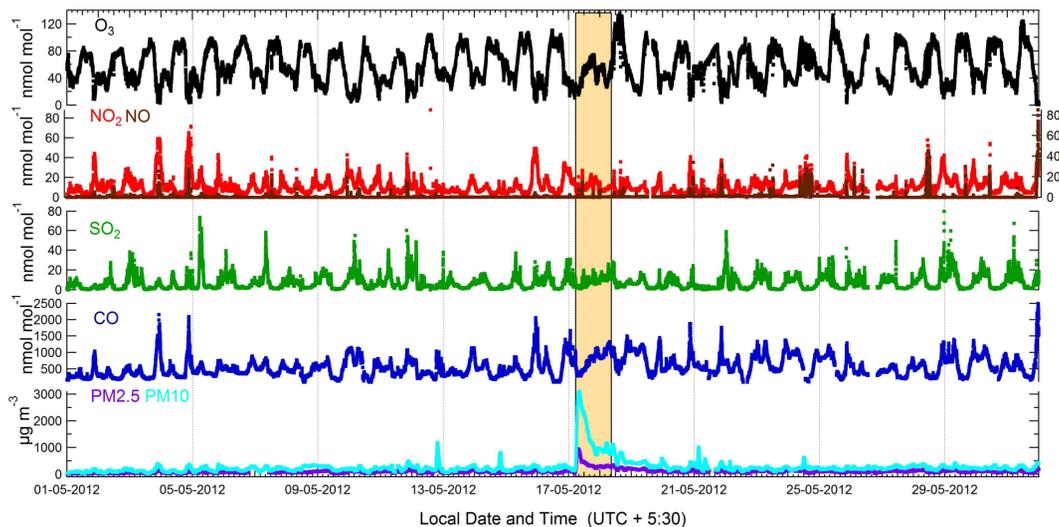
Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.

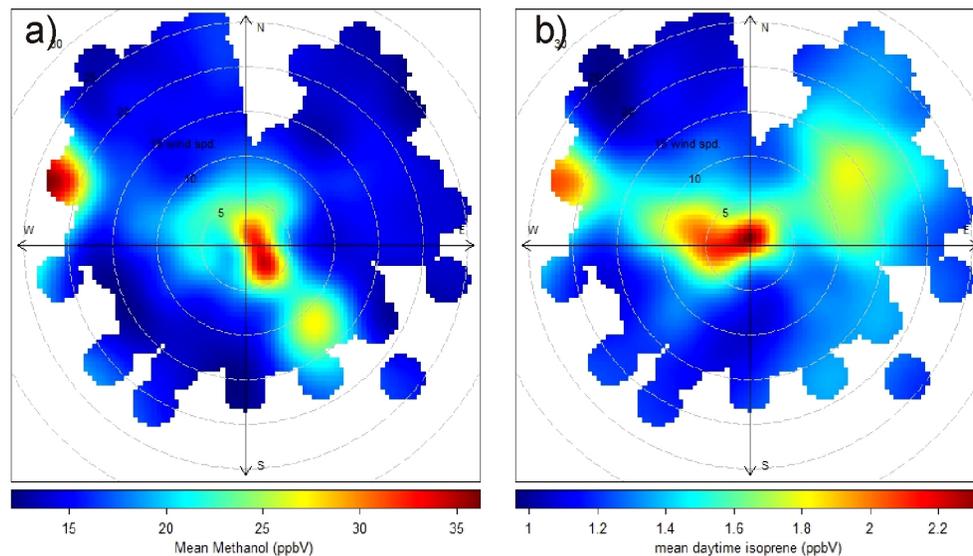


**Fig. 6.** Time series of the one minute data in May 2012 for the mixing ratios of ozone (top panel),  $\text{NO}_2$  and  $\text{NO}$  (second panel),  $\text{SO}_2$  (third panel),  $\text{CO}$  (fourth panel) and mass concentrations of  $\text{PM}_{2.5}$  and  $\text{PM}_{10}$  (bottom panel).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## High surface ozone and strong influence of biomass burning

V. Sinha et al.

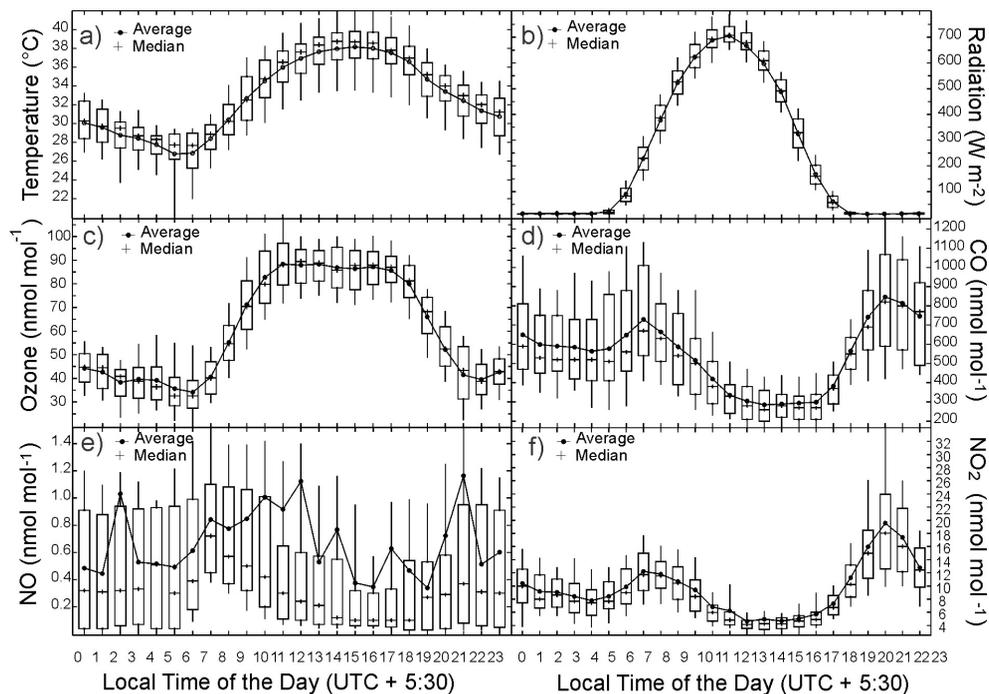


**Fig. 7.** Bivariate polar plots of the (a) daytime methanol data and (b) daytime isoprene data used for constraining location of major point emission sources.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## High surface ozone and strong influence of biomass burning

V. Sinha et al.



**Fig. 8.** Diel box and whisker plot of (a) the ambient temperature, (b) radiation, (c) ozone, (d) carbon monoxide, (e) nitrogen monoxide and (f) nitrogen dioxide derived from all measurements ( $n > 14300$  for each species) in May 2012 at the measurement site.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

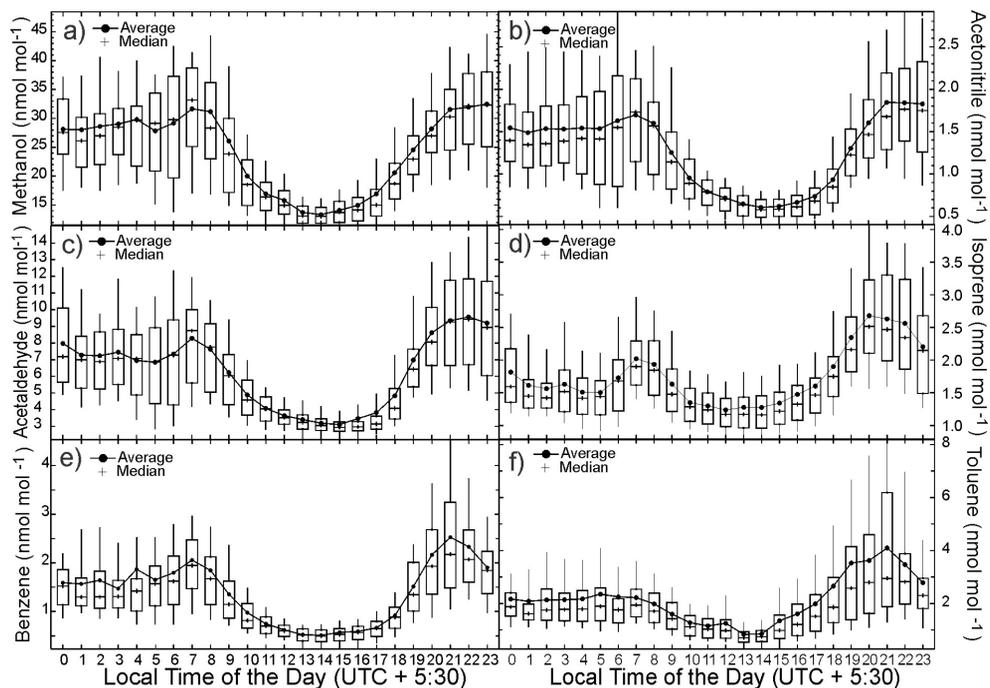
Printer-friendly Version

Interactive Discussion



## High surface ozone and strong influence of biomass burning

V. Sinha et al.



**Fig. 9.** Diel box and whisker plot of **(a)** methanol, **(b)** acetonitrile, **(c)** acetaldehyde, **(d)** isoprene, **(e)** benzene and **(f)** toluene derived from all measurements ( $n > 14\,300$  for each species) in May 2012 at the measurement site.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion