Interactive comment on “Modeling chemistry in and above snow at Summit, Greenland – Part 2: Impact of snowpack chemistry on the oxidation capacity of the boundary layer” by J. L. Thomas et al.

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COMMENT: The coupled 1-D snow chemistry and atmospheric boundary layer model MISTRA-SNOW is used to study the impact of snow chemistry on the oxidising capacity of the boundary layer at Summit/Greenland. Model output is compared with observations in ambient and partially also in firn air over a three day period in summer 2008 (GSHOX campaign). Within the measurement uncertainties the model agrees well with observed oxidant levels, i.e. OH and the sum of HO2+RO2 radicals and most of the time with O3. In order to investigate impact of snow chemistry on oxidant concentrations, runs without snow bromide and without snow nitrate are compared to the base case. The main result is that OH levels can only be matched by including snow emissions of NOx and halogens. The model indicates that NOx emissions contribute to ozone production, which is reduced but not offset by catalytic destruction via snow sourced bromine chemistry. The mismatch with HONO is interpreted as a potential missing HONO source.

This is a nice model study investigating the coupled snow-atmosphere system in order to understand better the chemical composition of air above snow. The finding that snow photochemistry influences the HOx radical budget at Summit is not a novelty as such, as this has been shown previously by Chen et al. (2007). However, difficulties in that study remained in reproducing radical levels. The present work is clearly a step forward in that it employs a truly coupled approach and matches observations of hydroxyl radicals with only temperature as a constraint. In addition, links and feedbacks between NOx and halogen chemistry and how they relate to the oxidant budget above snow are illustrated. The MISTRA-SNOW appears to be a valuable tool for the interpretation of atmospheric chemistry studies above snow.

The main concern, which needs to be addressed/discussed further before going ahead in ACP, is that the model probably underestimates the impact of snow emissions of HOx precursors on the respective atmospheric budget and on total oxidation capacity.

RESPONSE: The authors thank the reviewer for the helpful comments, which have improved the scientific content of the manuscript. We address the specific comments below.

COMMENT: The authors state that model predicted HOx precursors H2O2 and CH2O compare well with observations in summer 2000. They find that H2O2 snow emissions do not contribute much to a) the atmospheric budget of H2O2 and b) to primary production of OH radicals. Both statements are quite different to what has been found previously at the same site or other polar sites. For example, regarding a) Chen et al.
(2007) showed that in summer 2003 the net snow contribution of the gross chemical production of H2O2 is about 25%. Re:b) the same study showed that primary HOx sources were O(1D)+H2O (41%) and the photolysis of snow emitted precursors H2O2 and CH2O (40%); for the latter H2O2 dominated with 37% (3% from CH2O).

The conclusions regarding model performance & oxidant budget might be different had the authors not chosen to use H2O2 and CH2O from summer 2000 (due to lack of measurement during GSHOX). This was a summer with atypically low concentrations of both species at Summit, e.g. mean H2O2 in the 16-Jun to 7-July 2000 period was 0.65 ppbv (Jacobi 2002) whereas average levels measured in 4-20 June 1996 (Hutterli, 2001) and 27-June to 12-July 2003 (Frey, 2009) were with 1.4 ppbv about twice as high. The same applies to CH2O with 0.12 ppbv in summer 2000 (Jacobi et al., 2002) compared to 0.23 ppbv in June 1996 (Hutterli et al., 1999) or >0.15 ppbv in July 2003 (Chen et al., 2007). Thus, uncertainties in modelled oxidant levels & production/loss pathways from less accurate model description of H2O2/CH2O snow-air exchange need further discussion & revision of corresponding statements.

RESPONSE: A model run with additional H2O2 and HCHO in the gas and condensed phases has been completed to investigate how this changes the conclusions in the paper. For the run with higher H2O2 and HCHO concentrations, we have initialized the surface HCHO equal to 230 pptv and H2O2 equal to 1.4 ppbv. We also initialized the LLL concentration with the appropriate Henry's law equilibrium concentration for these higher concentrations. The OH concentrations change very little with these increased primary radical sources (shown in green in Fig. 1).

Even though the levels themselves don’t change, the relative contribution of H2O2 to the primary OH production does change. In the updated run, which does not sustain these high levels of HCHO or H2O2 for the entire model run, we only use the first day for the purposes of constructing an OH budget. The relative importance of H2O2 compared to O1D+H2O changes significantly as pointed out by the reviewer, and doubles the contribution of H2O2 to the primary OH sources (22% in the base run compared to 41% in the run with more H2O2 and HCHO). This does not result in significant additional OH in the boundary layer because the budget is largely dominated by recycling reactions (e.g. NO + HO2).

We have added the results from this sensitivity run into a supplementary information document that is available online and have added a discussion of these results to the manuscript.

COMMENT: As acknowledged by the authors, the model has some shortcomings such as lack of organic snow chemistry or still rather simplistic snow physics, e.g. the representation of the liquid-like layer (LLL). The main tuning parameter is the bulk-LLL ratio of nitrate and bromide and is adjusted to match NO and BrO in ambient air. How sensitive are modelled NOx and BrO mixing ratios to this parameter?

RESPONSE: This is an important point because the LLL initialization and partitioning are clearly important. This is shown in the sensitivity runs included in the manuscript with varying snow nitrate and bromide concentrations. A systematic sensitivity study of different LLL initializations is out of the scope of the current paper.

COMMENT: And finally, a more detailed comparison between modelled and observed firn air concentrations would be beneficial for model validation, but certainly another study in its own right. However, discussing some more of the published firn air data at Summit would yield insight into where the model needs development, e.g. firn air concentrations of ozone (Peterson et al., 2001), OH radicals (Beyersdorf et al., 2007) or possibly organic peroxides (Frey et al., 2009).

RESPONSE: Peterson et al. 2001 shows there is significantly more ozone destruction in interstitial air than predicted in the model (Thomas et al., 2011). It is currently not known what causes rapid ozone destruction in firn air, but is likely due to catalytic destruction of ozone by halogen cycling in the snowpack. If that is the case, then the model under-predicts halogen concentrations and the corresponding ozone destruction in interstitial air. However, it is also possible there is another (currently unknown) mech-
anism causing ozone destruction in interstitial air. To address this, the following text has been added to the manuscript: "Significant ozone destruction in firn air measured by Peterson et al. (2001) is not reproduced in the model, suggesting that the ozone destruction may not have occurred during the campaign. Alternatively, the model could under-predict halogen concentrations in interstitial air and the corresponding ozone destruction. It is also possible there is another (currently unknown) mechanism causing ozone destruction in interstitial air."

Beyersdorf et al. (2007) showed using hydrocarbon gas ratios in a transparent flow chamber with undisturbed surface snow that the lower limit to peak OH radical concentration on 10-12 July 2003 was \( \sim 3.0 \times 10^6 \) (lower than ambient concentrations reported in Sjostedt et al. (2007)). Beyersdorf et al. calculated OH radical concentrations using an inlet placed below the snow surface at a depth of 5 cm with a sample flow rate of 0.5 L min\(^{-1}\). The values shown from the model in the paper at a depth of 10 cm are higher than the values suggested by Beyersdorf et al. (2007). There are two factors that determine interstitial air OH concentrations. First, there is decreasing sunlight with depth, which should decrease the OH concentration. Second, the additional NOx and bromine in interstitial air should increase the concentrations of OH significantly. We have shown earlier in the paper that even a small amount of NO (\( \sim 10 \) pptv) increases the OH in the boundary layer by a factor of 2. So, it is not surprising that in the interstitial air, the large NOx concentrations overwhelm the decrease in sunlight and that the model calculated OH concentrations are higher than in the boundary layer. There are several possibilities to explain the discrepancy, the first is that the model NOx concentrations are not representative interstitial air NOx concentrations during the campaign. The second is that in sampling NMHC Beyersdorf et al. (2007) sampled concentrations at multiple depths, therefore there may be a contribution from interstitial air with very low OH concentrations deeper in the snowpack. Beyersdorf et al. (2007) also state that their measurement is a lower limit, so the model may not be inconsistent with the measurements.

To address this we have added the following text to the manuscript: "In an initial study, Beyersdorf et al. (2007) showed using hydrocarbon gas ratios in a transparent flow chamber with undisturbed surface snow that the lower limit to peak OH radical concentration on 10-12 July 2003 was \( 3.0 \times 10^6 \) molec cm\(^{-3}\). Given that their measurement was a lower limit, this may not be inconsistent with the model predicted OH concentrations. There are several potential explanations for the discrepancy between Beyersdorf et al. (2007) and the present study. The model may over-predicting OH, due to over prediction of the NOx concentrations in the interstitial air. But, without measurements to compare with it is difficult to determine if this is the cause of the discrepancy. It is also possible that during sampling Beyersdorf et al. (2007) also sampled hydrocarbons from deeper in firn air, which contained lower OH concentrations."

Frey et al. (2009) compared measurements of organic peroxides from different years at Summit. We have added a short discussion of the variability of observations by adding the following to the paper: "Frey et al. (2009) presented H\(_2\)O\(_2\) measurements at Summit from 2003 and 2004 which ranged from 1.4 ppbv in June/July to 100-300 pptv in March-early May. Summer 2000 H\(_2\)O\(_2\) levels were lower, as shown by Jacobi et al. (2004). Suggesting both seasonal and inter-annual variability in radical sources may be important. But, given that H\(_2\)O\(_2\) and H\(_2\)CHO were not measured during GSHOX the comparison with prior years can only be considered qualitative."

COMMENT: Abstract: l16-19 The impact of BrO is a lot smaller than that of NOx. It would be instructive to include the fractional contributions of NOx and BrO to the OH increase at noon. Credit should be given (here or in the summary of previous model work) that Chen et al. (2007) demonstrated previously the impact of snow chemistry on the HOx budget above snow at Summit, although not with a coupled snow-air model and with different conclusions (see above).

RESPONSE: We have added to the abstract: "Bromine chemistry increases the OH concentration by between 10-18% (10% at noon), while snow sourced NOx increases
We have added to the discussion of prior work: "Prior modeling work using a constrained box model showed the importance of these species for OH budget at Summit (Chen et al., 2007)."

RESPONSE: To clarify this point we modified the abstract to read: "Other model predicted HOx precursors, H2O2 and HCHO, compare well with measurements taken in summer 2000, which had lower levels than other years."

COMMENT: p5560, I16-23: A brief note would be helpful regarding H2O2 & CH2O snow grain distribution and initial snow profiles.

RESPONSE: The text has been updated to address this: "The base case model run is initialized with gas and aqueous phase concentrations as described in Thomas et al. (2011), Table 2. Briefly, the initial gas phase mixing ratios in the interstitial air are vertically uniform for species with long atmospheric lifetimes (for example CO) and decay with snow depth for reactive species (for example O3, H2O2, and HCHO)."

COMMENT: p5564-5566 (section 3): A summary table would be helpful, e.g. parameter (mean, range), uncertainty & reference and also including a quantitative measure of model performance such as the rms error for each parameter & run (base, noNO, noBr).

RESPONSE: A summary table would duplicate Table 2 in Thomas et al. (2011). A full error analysis of how the input parameters impact the output of the model is outside the scope of the present manuscript.

COMMENT: p5570, I14-21 & Fig.6: The differences in O3 mixing ratios between base case and noBr are surprisingly small and would not explain a difference of >15 ppbv between firn and ambient air previously observed by Peterson et al., (2001). Is this due to model inconsistencies or are perhaps other halogens at play?

RESPONSE: The answer to this question is not known. Either the model underpredicts halogen concentrations in the interstitial air, or there is another process involved in ozone destruction (as discussed in an earlier comment). A measurement of BrO concentrations in firn air is needed to evaluate if the model is reasonable, but is currently unavailable. The following has been added to address this: "The model does not predict strong ozone depletion in the snowpack, which has been observed at Summit (Peterson and Honrath, 2001). This is because of the relatively small halogen concentrations in interstitial and ozone production from NOx in the interstitial air offsets most ozone destruction that does occur. Measurements of halogens and NOx in the interstitial air are needed to compare with the model. However, measurements of undisturbed interstitial chemical concentrations are difficult and will likely require developments in measurement techniques."

COMMENT: p5571 I15-16: See above. In addition, the model underestimate of CH2O in firn air might be even more significant due to a low bias in measured CH2O from...
ventilation effects during sampling. RESPONSE: We have updated the text to: "In addition, the measurements are likely a lower limit of interstitial air concentrations due to dilution while sampling. In the future, model improvements will be needed to fully represent this chemistry."

COMMENT: p5572,l8: see above, choice of H2O2 data set for comparison p5572,l15: not Jacobi 2002,2004, the correct citation would be Hutterli et al., 2001, where June 1996 H2O2 data were first described and interpreted

RESPONSE: We have changed the text to read: "Snow sourced hydrogen peroxide, H2O2, can contribute to the OH budget and the boundary layer oxidizing capacity over snow upon photolysis. Multiple years of measurements of hydrogen peroxide have been made at Summit, Greenland (e.g. Hutterli et al., 2001). As no H2O2 measurements were made during GSHOX we also compare H2O2 modeled in the base case with prior measurements during summer 2000 (Jacobi et al., 2004)."

"Measured H2O2 fluxes in June 1996 suggested a daytime H2O2 release from the surface snow and redeposition at night (Hutterli et al., 2001)"

COMMENT: p5572, l18-19: here most of ambient H2O2 production is from HO2 self reaction; however, the model seems to underestimate the H2O2 snow source if typical Summit values are taken into account and the budget would change accordingly.

RESPONSE: This has been addressed via the sensitivity runs now included in the supplementary information provided with the manuscript, which contains an OH/HOx budget for a run with higher H2O2 concentrations.

COMMENT: p5572, l20-22: needs revision with above comments

RESPONSE: We have updated the text to: "gas phase chemistry in the boundary layer can largely explain H2O2 levels measured in summer 2000."

COMMENT: p5572, l27: to avoid confusion define oxidation capacity, e.g. as the sum of HOx, H2O2 and O3 or similar.

RESPONSE: We have updated the text to: "boundary layer oxidation capacity, which for the purposes of this study we define as the sum of OH, HO2, and RO2."

COMMENT: p5574, l21-24: in contrast to the SUM03 model study (Chen et al., 2007), see previous comments.

RESPONSE: The text has been updated to: "only a small portion of this is a primary OH source (from snow emissions of H2O2). In the sensitivity run presented in the supplementary information shows that for a case with higher hydrogen peroxide concentrations the contribution of H2O2 to primary OH is consistent with the study of Chen et al. (2007). These higher concentrations likely result from snowpack emissions of hydrogen peroxide."

COMMENT: p5575,l24-27: An important terminal loss of HO2 radicals identified by the model is the formation of organic peroxides (see also Fig.13). This could be further supported/evaluated by comparison with existing observations of ROOH at Summit (Frey et al., 2009).

RESPONSE: We believe this comparison is best left for a separate model study.

COMMENT: p5579,l11-22: revision according to comments above.

RESPONSE: We have updated the text to: "Model predicted HCHO and H2O2 are in agreement with prior measurements taken in 2000 at Summit. However, measurements during this year were lower than other campaign measurements (Hutterli et al., 2001, Frey et al., 2009)."

COMMENT: TECHNICAL CORRECTIONS Figure 12d. title should be 'HO2 destruction'

RESPONSE: This figure has been updated to show only the first 24 hours of the model run so that it's clear what dominates HOx cycling. The label has also been corrected.
Fig. 1. Model predicted OH for the base case and for the sensitivity run with higher HCHO and H2O2. Model results show that the additional primary radical sources have a small impact on OH concentrations.