Qualitative and quantitative determination of water in airborne particulate matter

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Abstract

This paper describes the optimization and validation of a new simple method for the quantitative determination of water in atmospheric particulate matter (PM). The analyses are performed by using a coulometric Karl-Fisher system equipped with a controlled heating device; different water contributions are separated by the application of an optimized thermal ramp (three heating steps: 50–120 °C, 120–180 °C, 180–250 °C).

The analytical performance of the method was verified by using standard materials containing 5.55 % and 1 % by weight of water. The recovery was greater than 95 %; the detection limit was about 20 µg. The method was then applied to NIST reference materials (NIST1649a, urban particulate matter) and to real PM$_{10}$ samples collected in different geographical areas. In all cases the repeatability was satisfactory (10–15 %).

When analyzing the reference material, the separation of four different types of water was obtained. In real PM$_{10}$ samples the amount of water and its thermal profile differed as a function of the chemical composition of the dust. Mass percentages of 3–4 % of water were obtained in most samples, but values up to about 15 % were reached in areas where the chemical composition of PM is dominated by secondary inorganic ions and organic matter. High percentages of water were also observed in areas where PM is characterized by the presence of desert dust.

A possible identification of the quality of water released from the samples was tried by applying the method to some hygroscopic compounds that are likely contained in PM (pure SiO$_2$, Al$_2$O$_3$, ammonium salts, carbohydrates and dicarboxylic acids) and by comparing the results with those obtained from field samples.

1 Introduction

The interaction of water with atmospheric particulate matter (PM) has been described in many research papers, mainly focused on the study of the relationship between relative humidity and water adsorbed on PM and on the effects the interaction with
water may exert on the physico-chemical behavior of PM. Because of this interaction, the original physical and chemical features of the atmospheric particles may undergo significant variations, mainly in their size distribution (Swietlicki et al., 1999), optical properties and residence time in the atmosphere (Sloane and Wolfe, 1985; Pilinis et al., 1995; Kreidenweis et al., 2001). Moreover, it is well known that high values of relative humidity (RH) favor the occurrence of acid-base reactions leading to the formation of secondary inorganic salts (Baek et al., 2004; Squizzato et al., 2012). Finally, from the point of view of PM sampling, the amount of adsorbed water affects the solid-vapour equilibria of secondary ammonium salts collected on the filters, a mechanism that is among the main responsible for sampling artifacts (Vecchi et al., 2009). Although in most cases the artifact is negative (loss of NH$_4$NO$_3$ due to release of NH$_3$ and HNO$_3$) (Chow et al., 2005; Yuc et al., 2006; Pathak et al., 2009), in the presence of very high RH values it becomes positive, as the formation of particulate NH$_4$NO$_3$ from gaseous NH$_3$ and HNO$_3$ is favored (Gysel et al., 2007; Khlystov et al., 2009; Hu et al., 2011).

Of interest is also the contribution of water to the mass concentration of PM. The presence of considerable amounts of water in PM samples, in fact, causes an increase of PM mass concentration that might be responsible for exceedances of the regulatory limit values. It is worth noting that the knowledge of this contribution may be of interest for a correct evaluation of the health effects of PM: water is harmless in itself, but it may cause a relevant variation of the aerodynamic diameter of the particles and thus of their ability to penetrate into the respiratory tree.

Moreover, in many papers attempting PM mass closure, the unidentified mass (difference between gravimetric determination and reconstruction from single chemical analyses) is generally attributed to the presence of water and/or to the difficulty in determining an adequate conversion factors to calculate organic matter (OM) from the measurement of organic carbon (OC) (Balasubramanian et al., 2003; Harrison et al., 2003; Tsyro, 2005; Almeida et al., 2006; Sasaki and Sakamoto, 2006; Perrino et al., 2007, 2009; Perrone et al., 2012). A quantitative determination of the amount of water
collected on the filter would make it possible to distinguish the weight of these two factors and to improve the definition of OC/OM conversion factors.

Finally, the interaction with water seems to play a role in the different performance exhibited by Teflon and quartz membrane filters in sampling PM (Perrino et al., 2012).

In spite of the relevant role played by water in the study of atmospheric PM, a quantitative determination of adsorbed water was attempted only in a few papers. Water content was generally determined by indirect methods, consisting of the differential determination of particle dimension (Dick et al., 2000; Hu et al., 2001; Rees et al. 2004; Stanier et al., 2004; Kitamori et al., 2009) or collected mass amount (Speer et al., 1997, 2003) before and after the exposure to controlled RH conditions. These studies showed that water may constitute up to 20–30% of the total PM mass and indicated a dependence of the water uptake on the dimensions and chemical composition of the particles. Water adsorption resulted to be relevant for fine particles, characterized by high surface area, and for particles containing water-soluble inorganic salts, mainly ammonium sulphate (Speer et al., 1997; Stanier et al., 2004; Kitamori et al., 2009), and hygroscopic organic species such as dicarboxylic acids (Ansari and Pandis, 2000; Dick et al., 2000; Decesari et al., 2001; Speer et al., 2003). In the paper of Rees et al. (2004), the addition of the water content, measured by Dry-Ambient Aerosol Size Spectrometer (DAASS), to the amount of PM macro-components allowed a significant improvement of the mass closure. It is worth noting, however, that these differential techniques evaluate only the amount of surface-adsorbed water and are not able to give information about further contributions due to structurally-bound water.

To our knowledge, only two papers reporting a direct analytical determination of water on PM filters by the Karl-Fisher method have been published up to now (Ohta et al., 1998; Tsai et al., 2005). The first one was based on the thermal desorption of water at 150°C from Teflon filters sampled for two weeks; the results showed that a percentage of the total PM$_{2.0}$ mass ranging from 0.4% to 3.2% could be attributed to the crystallization water of (NH$_4$)$_2$SO$_4$, NH$_4$NO$_3$ and NaCl. In the second paper, the water contained on quartz filters conditioned at 60% RH was extracted by using anhydrous
methyl alcohol; in these conditions water resulted to be the individual component at highest concentration (about 30% of the total PM mass). In spite of the different operative conditions employed in these two studies, these results demonstrate that the Karl-Fisher technique is suitable for determining water in PM samples; the described methods, however, are not appropriate for a routine use in intensive monitoring campaigns because of too high detection limits (Ohta et al., 1998) and/or method complexity (Tsai et al., 2005).

The present work reports the optimization and validation of a new method for the determination of water in atmospheric PM samples, easily applicable to routine field campaigns. The method is based on the use of a thermal ramp for the selective desorption of different water contributions, which are analyzed by coulometric Karl-Fisher (KF). This approach is based on the results of a previous work, where three different water contributions, released at different temperatures, had been identified by analyzing the Certified Material NIST 1649a by thermogravimetry (TG) (Perrino et al., 2012). In that study, however, it was not possible to carry out a quantitative determination of the three contributions because other volatile species were released in the same temperature ranges.

The proposed method has been applied to pure hygroscopic materials that are commonly contained in atmospheric PM and to real PM_{10} samples collected in various geographical areas, so as to highlight the relationship between the chemical composition of PM and the thermal behavior of water.

2 Experimental

2.1 Materials

For the validation of the method we used two standard materials: HYDRANAL Water Standard KF-Oven 220–230°C (HYD; Fluka Analytical) and Water Standard Oven
1 %, (WSO; ACS Merck KGaA), containing 5.55 ± 0.05 % and 1.0 ± 0.03 % water, respectively.

SiO$_2$, Al$_2$O$_3$, NH$_4$HSO$_4$, NH$_4$Cl, NH$_4$NO$_3$, (NH$_4$)$_2$SO$_4$, glucose, lactose, fructose, sucrose, levoglucosan, tartaric acid, citric acid, succinic acid, oxalic acid, phenylalanine were analytical grade (ACS Merck KGaA).

The other materials (Saharan dust, road dust, soot from vehicular exhaust, skin, bacteria, moulds, pollens) were taken directly from the environment and analyzed without any pre-treatment.

### 2.2 Karl-Fisher analysis

All determinations of the water content were carried out by using a 831 KF Coulometer (Metrohm AG, Herisau, CH). The instrument is equipped with a programmable temperature oven that allows the user to adjust the heating ramp (874 Oven Sample Processor; Metrohm AG, Herisau, CH). For transferring the sample from the oven to the measurement cell we used ambient air at the flow rate of 20 ml min$^{-1}$, previously filtered and dried by using molecular sieves (0.3 nm pore size, Metrohm AG, Herisau, CH). Karl-Fisher reagent was used in the titration cell (100 ml; Hydranal-Coulomat AG-Oven, Sigma-Aldrich); minimum value of the voltage between indicator platinum electrodes was set at 30 mV; the limit value of the drift was set at 1 µg min$^{-1}$; extraction time was 3000 s.

Before each measurement the sample vials (6 ml, Metrohm AG, Herisau, CH) were kept in oven at 250 °C for 12 h, then placed in a desiccator and transferred inside an AtmosBag (Sigma-Aldrich) filled with Argon dried by using molecular sieves. Samples were weighted by using an analytical balance Gibertini E505 (sensitivity 0.01 mg; Gibertini Elettronica, Novate Milanese, Milan, I) and put inside the AtmosBag, where they were inserted into the vials.

To avoid clogging of the gas lines due to the condensation of species desorbed from PM (e.g. ammonium salts) the transfer lines and the needle were accurately cleaned every day.
The instrument is equipped with an automatic system that identifies the end point of the titration according to the return of the baseline drift to its initial value. When using a thermal ramp, it is necessary to exclude the end point identification, as each return of the drift to the initial value indicates the separation of a different water contribution. The software calculates the amount of desorbed water by subtracting the contribution due to the baseline drift from the final water amount; the baseline drift contribution is calculated by extrapolating its initial value to the end of the analysis. In our case, the use of the thermal ramp caused the increase of the analytical time with a consequent increase of the drift contribution to the final water amount, which becomes comparable with the amount of water in the PM samples. Also, the drift value is variable, and the influence of this variability on the repeatability of the analytical results also increases with the duration of the analysis. In our conditions, the intra-day and inter-day variations were 0.05 µg min⁻¹ and 0.9 µg min⁻¹, respectively (10 replicate analyses of the operative blank). To reduce the effect of the drift variability as much as possible, the measurement of the operative blank was carried out every two samples, and the KF curve of each sample (water vs. time) was obtained by subtracting the blank value to each point of the graph (see Fig. 1). This process allows a direct reading of the water amount on the ordinates of the graph; also, this method allows a direct subtraction of the water contribution due to the humidity of air inside the vial, the adsorption on the vial walls and the content of the plastic vial cap.

The limit of detection (LOD) and limit of quantification (LOQ) were calculated as three times and ten times the standard deviation of the blank (10 replicates). For a typical intra-day variability we obtained the following results: LOD = 7 µg; LOQ = 25 µg.

2.3 Real PM₁₀ samples

Twenty pairs of equivalent real PM samples were collected during the period November–December 2011 in three different geographical areas. The location in Rome (Central Italy, 41°54’03.69” N, 12°30’44.93” E) was a traffic site, about 50 m from the nearest road (8 pairs of samples); the location in Ferrara (Po valley, Northern Italy, 27373
44°50′55.44″ N, 11°33′40.96″ E) was an industrial site, about 5 km from the city center and 1 km from the nearest industrial plant (10 pairs); the location in Tel Aviv (Israel, 32°06′54.16″ N, 34°48′16.74″ E) was the roof of the Tel Aviv University (2 pairs).

Daily PM$_{10}$ samples were collected on 47 mm diameter PTFE membranes, 1 µm pore size (PALL Corporation, USA) by means of dual-channel samplers (HYDRA Dual Sampler, FAI Instruments, Fontenuova, Rome, I) equipped with two independent PM$_{10}$ sampling heads compliant with EN 12341 (1998). In Rome and Tel Aviv the two channels were operated at the flow rate of 2.3 m$^3$ h$^{-1}$. In Ferrara, which is characterized by very high relative humidity, the sampling heads were modified in order to reduce the flow rate to 1.15 m$^3$ h$^{-1}$ and to avoid clogging of the sampling membranes.

An additional collection of six parallel samples was carried during the period 14–20 December 2011 at the Ferrara site, by using three HYDRA Dual Samplers placed side-by-side. These samples were used to evaluate the effect of sample conservation. Teflon filters were allowed to equilibrate for two days at 20°C and 50% RH before weighting. After sampling, filters were placed in Petri dishes and stored at 5°C. Different procedures were used to evaluate the effect of storage conditions. Before the analysis all samples were kept again at 20°C and 50% RH for two days; the plastic rings of the filters were cut off by using a steel scalpel.

3 Results and discussion

3.1 Method optimization

The proposed method was optimized by using the reference material NIST1649a (National Standard Institute of Technology – USA), consisting of urban atmospheric particulate matter. Although its water content is not certified and it show some differences with respect to real PM$_{10}$ samples (Canepari et al., 2006), this material makes it possible to perform the optimization phases on a rather representative homogeneous material.
The first optimization phase concerned the heating ramp and was aimed to obtain a satisfactory separation of the different water contributions to atmospheric PM. Figure 2 shows some examples of thermal ramp (left panels) and of KF curves obtained when applying the ramp to NIST 1649a (right panels). The total amount of water was 100 ± 6 g kg⁻¹, irrespective of the selected ramp. Instead, the curve profile strongly depends on the thermal ramp, as shown by the drift graph (grey lines in Fig. 2). When the heating temperature is constant, the operational mode generally used in this type of instruments, we obtain only one signal that is relative to the total water amount (Fig. 2, upper panels). The use of a linear thermal ramp allows the differentiation of several water contributions that are released from the sample at increasing temperatures (Fig. 2, middle panels). In these conditions, however, the contributions are insufficiently separated. It is worth noting that this linear ramp was used in previous scientific works regarding the TG analysis of NIST 1648 (Matuschek et al., 2004; Perrino et al., 2012). These papers reported two well-defined weight losses in the range 80–120 °C and 150–180 °C, which were attributed to particle-bound water, and one wider loss in the range 225–310 °C that was due to the release of several organic and inorganic species, with a possible contribution of crystallization water. Although the results obtained by KF and TG are qualitatively similar, the separation of the water contributions by the KF method is lower, probably because of a higher thermal inertia in the heat transfer to the sample.

The lower panels of Fig. 2 reports the optimized thermal ramp that was used to perform all subsequent analyses (5 min at 50 °C, from 50 °C to 120 °C at 14 °C min⁻¹, 5 min at 120 °C, from 120 °C to 180 °C at 12 °C min⁻¹, 2 min at 180 °C, from 180 °C to 250 °C at 14 °C min⁻¹, 20 min at 250 °C). In these conditions, we obtained a significant improvement in the analytical resolution, and a profile very similar to that obtained by TG. The first peak in the drift graph, which had not been detected by the TG analysis, identifies water that can be desorbed at low temperature (50 °C), probably due to moisture weakly adsorbed on the particles. The following three peaks (in the range 400–700 s, 700–1100 s, 1100–2300 s) are in very good agreement with the results of
the TG analysis and confirm the presence of different water contributions, bound to atmospheric particles with different strength.

### 3.2 Method validation

#### 3.2.1 Recovery

To calculate the exactness, we used two standard materials (HYD e WSO) containing different amounts and qualities of water. Figure 3 reports the KF curves and the drift graphs of the two certified materials. HYD (upper graph) shows only one water contribution, desorbing over 200 °C, that is due to crystalization water. WSO (lower graph) shows two different contributions: the first one, weakly retained, is released at the beginning of the analytical run at about 50 °C, the second one is released in the range 500–800 s, at the temperature of about 120 °C.

Table 1 shows the recovery of different amounts of the two standard materials. We did not consider amounts lower than 5 mg in order to avoid the introduction of a high uncertainty contribution due to the gravimetric determination and transfer of very small sample amounts. In all cases the recovery was close to 100 %, while the repeatability was better for HYD (high amounts: about 3 %; low amounts: about 6 %) than for WSO (high amounts: about 9 %; low amounts: about 11 %), probably because of the lower stability of the water contained in WSO.

A further validation phase concerned the recovery from silica and alumina, which are among the hydrophilic species contained in PM. These compounds are thermally stable and available on the market at high purity level; it is thus possible to carry out a gravimetric determination of the water that is desorbed when heating in oven (105 °C for 24 h). We also tested the molecular sieves used to dry the carrier gas of the KF (silicate with inorganic binder), in order to verify the influence of the material porosity on the profile of the KF curve. The KF curves obtained for these materials are shown in Fig. 4. All three materials show a contribution in the range 0–500 s, probably due to moisture weakly adsorbed on the particles; the following part of the curves, instead, are
quite different. The curve of silica (upper graph) shows only a further single contribution in the range 500–1000 s. Molecular sieves (middle graph) show a very similar behavior in the first part of the curve; with respect to silica, however, molecular sieves release a relevant amount of water also in the second part of the analysis (1000–1500 s), likely due to the release of water adsorbed inside pores. Alumina shows the same profile in the range 500–1000 s (lower graph), but also two other more retained contributions, not completely separated, between 1000 and 2300 s, probably due to the presence of active sites of different strength.

Table 2 reports the recovery values for these materials. Also in this case the method shows good performance, with recoveries higher than 95% and repeatability around 6%. It is worth noting that the results reported in Table 2 refer to very variable sample amounts (5–100 mg) and to measured water amounts in the range 0.3–7 mg. The shape of the curves and the recovery values do not depend on the sample weight, indicating robustness and wide linearity range of the proposed method.

### 3.2.2 Interferences

As well known, the KF method suffers from the interference of some classes of compounds, both organic and inorganic, some of which are likely found in PM samples (aldehydes, ketones, carbohydrates, Fe(III) and Cu(II) salts) (EPA Method 9000, 2007). Most of these species are not able to interfere in analytical techniques based on heating due to their very low vapor pressure (levoglucosan, glucose, fructose, sucrose) (Chow et al., 2007). As far as carbonyl compounds are concerned, the species that reach high atmospheric concentration are mostly in the vapor phase (Levart and Veber, 2001). Some Authors highlight the presence of some semi-volatile ketones (n-alcan-2-ones and o-Hydrocarbons Polycyclic Aromatics) (Schnelle-Kreis et al., 2005, 2007), but the atmospheric concentration of these species is very low (a few tenths of ng m$^{-3}$ for 6,10,14-trimethylpentadecan-2-one, the most abundant species) and their possible interference is well below the quantification limit of the method.
The possible positive interference of iodine reduction reaction due to ammonia is also highlighted in the EPA Method 9000 (2007). In the temperature range of the proposed KF method, ammonia is released from secondary ammonium salts, which constitute a relevant fraction of atmospheric PM (Perrino et al., 2012). To verify this hypothesis, the method was applied to pure NH$_4$NO$_3$, (NH$_4$)$_2$SO$_4$ and NH$_4$Cl salts. Measured amounts of water were 0.9±0.07 % for chloride, 1.6±0.1 % for nitrate and 8.5±0.4 % for sulphate, values that are consistent with the hygroscopicity of these salts. As reported in Perrino et al. (2012), at the maximum operative temperature of the oven (250 °C) the release of ammonia is almost quantitative from NH$_4$Cl and NH$_4$NO$_3$, while it is about 20 % from (NH$_4$)$_2$SO$_4$. To evaluate the interference of ammonia we considered NH$_4$Cl, that is the species that yields the lowest signal. Considering that all the ammonia content of the salt is transferred in the cell, we obtain a maximum signal equivalent to a water amount of 0.9 % of the salt weight.

Considering a total ammonium amount on the sampled filter as high as 500 µg (very high values, corresponding to an ammonia concentration of about 10 µg m$^{-3}$ for 24-h samplings at the flow rate of 2.3 m$^3$ h$^{-1}$) even if the measured signal were totally attributed to interfering ammonia (anhydrous salt) we would obtain a signal equivalent to only 14 µg of water. We can thus conclude that ammonia interference, if present, would in any case be negligible.

3.3 Real samples

Some preliminary tests addressed the analysis of the most common sampling media (Teflon and quartz) and the effect of the sampled filters conservation procedure on the water content.

The profile of KF curves from the analysis of quartz fiber filters resembles the profile of molecular sieves, in agreement with the common siliceous composition. Blank values, as expectable, were quite high (about 1400 µg per filter). The profile of Teflon membrane filters, instead, was indistinguishable from the blank, both before and after conditioning at 50 % RH, in agreement with their hydrophobic characteristics. Teflon...
filters were thus considered as suitable media for PM sampling aimed at water content determination.

The influence of the sample conservation procedure was studied by analyzing six groups of six equivalent PM$_{10}$ samples sampled side-by-side. One filter of each group was removed from the sampler immediately after the end of the sampling, conditioned at 20°C and 50 % RH for 48 h and then analyzed. The other five filters were left inside the unloader of the sampler until the end of the last sampling period. A second element of each group was directly conditioned and analyzed, while other three filters were placed inside Petri dishes, sealed with parafilm and kept for three months at ambient temperature, 5°C or −18°C, respectively. The last element of each group was kept for six months at 5°C, before being conditioned and analysed. No significant differences were observed among the six profiles of each group of samples; the analytical repeatability was about 10 %, indicating a good stability of the water retained by the particles. Storage temperature was then set at 5°C.

The validation of the method on real samples was completed by analyzing 20 pairs of PM$_{10}$ samples collected at very different locations: a traffic site in Rome, an industrial site in Ferrara (Po Valley, Northern Italy), an urban site in Tel Aviv (Israel). The collected amount of dust was variable (range: 0.5–2.9 mg); the wide differences in the emission sources and meteo-climatic conditions at these three sites should also assure a variety in the chemical composition of the collected dust.

According to EN 14902 (2005), the repeatability was calculated as follows:

\[
r_{rel} = \frac{r}{\bar{X}} \cdot 100; \quad r = \sqrt{\frac{\sum_{i=1}^{N} (m_{iA} - m_{iB})^2}{2N}}; \quad \bar{X} = \frac{\sum_{i=1}^{N} (m_{iA} + m_{iB})}{2N}
\]

where \(m_{iA}\) and \(m_{iB}\) are the water amount determined on each component of the filter pair, and \(N\) is the total number of PM$_{10}$ sample pairs (\(N = 20\)). The obtained value (\(r_{rel} = 10.2\%\)) was satisfactory and consistent with the repeatability values determined for NIST 1649a and for pure materials.
Figure 5 shows the scatter plot of the results obtained from each filter pair. The regression parameters (slope: 0.98; intercept: 0.15 µg m\(^{-3}\); Pearson’s coefficient: 0.96) confirm the good analytical performance of the method also when applied to real samples.

It is worth noting that the water content of our samples was very variable and reached remarkably high values, especially in Ferrara and Tel Aviv. In general, water constituted about 3–4 % of the total PM mass of the samples collected in Rome, while in Ferrara and in Tel Aviv we obtained percentages over 10 %, with a contribution of more than 20 µg m\(^{-3}\) to the environmental concentration of PM\(_{10}\).

In the left panel of Fig. 6 we report examples of the KF curves obtained from pairs of samples collected at the three sites. The KF profiles of the two elements of each pair are alike; the three examples, however, show very different profiles.

The first water contribution, in the range 0–500 s, is due to weakly adsorbed moisture and is contained in all samples; the amounts are in the range 50–150 µg. From 500 s on, the profiles become characteristics of the sampling site, with small inter-day variations. The release of water from the samples collected in Rome occurs in many steps, during the whole analytical run. The samples collected in Ferrara show a sharp signal at about 1500 s, superimposed to a broader signal in the interval 1000–2000 s. Tel Aviv samples show, instead, two well-defined contributions, in the range 700–1100 s and after 1500 s.

Since these differences are very probably due to a different chemical composition of PM at the three sites, we carried out some exploratory qualitative analyses of some hydrophilic components of PM. The most interesting results are reported in the right panels of Fig. 6. The profiles of soot (sampled inside the exhaust silencer of a diesel car) and of road dust (sampled at the kerbside of a high-density traffic road) were very similar to the profile of PM\(_{10}\) sampled in Rome, in agreement with the relevance of local traffic sources in the composition of PM in this area (Canepari et al., 2009; Perrino et al., 2009). Water contained in secondary salts and in some hydrophilic species (carbohydrates and dicarboxylic acids) is detected mainly in the range 1000–2000 s.
is likely that the curve obtained by analyzing the samples collected in Ferrara is due to water associated to these compounds. In the area of Ferrara, in fact, particularly during the winter, the formation of organic and inorganic secondary species is favored by the occurrence of frequent and long-lasting atmospheric stability conditions (Matta et al., 2003). Saharan dust (collected at Erfoud, Morocco) in agreement with its composition, shows a profile close to those of SiO$_2$ and Al$_2$O$_3$, with a sharp contribution between 700 and 1100 s. This contribution is very similar to one of the contributions detected in the filters collected in Tel Aviv, where the dust intrusions from the nearby desert are very frequent (Ganor, 1994). It is worth noting that the last contribution in these filters resembles the road dust and soot profile, as expected for an urban site.

4 Conclusions

The described method, employing a coulometric Karl-Fisher systems, allows a direct determination of water in atmospheric particulate matter. It shows a good analytical performance, does not require any filter pre-treatment and is applicable to routine field studies.

The method was applied to a series of PM$_{10}$ samples collected in different geographical areas. The results of these field studies show that the water content of PM samples is variable and can reach values over 20 µg m$^{-3}$. Such an amount of water constitute a relevant fraction of the total mass of PM and may seriously affect the exceedance of the regulatory concentration limit values.

The use of a thermal ramp for the desorption of water allows the separation of different water contributions that are bond to atmospheric particles with different strength. The profile of the curve resulted to be characteristic of the sampling site; the amount and variety of water is then probably linked to the chemical characteristics of the dust.

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Table 1. Recovery of water from standard materials HYDRANAL (HYD) and Water Standard Oven (WSO). $N = 10$.

<table>
<thead>
<tr>
<th></th>
<th>Certified</th>
<th>Measured</th>
<th>Recovery</th>
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<tr>
<td></td>
<td>µg</td>
<td>µg</td>
<td>%</td>
</tr>
<tr>
<td>HYD</td>
<td>Low (5 mg)</td>
<td>278 ± 3</td>
<td>272 ± 15</td>
</tr>
<tr>
<td></td>
<td>High (50 mg)</td>
<td>2775 ± 25</td>
<td>2817 ± 67</td>
</tr>
<tr>
<td>WSO</td>
<td>Low (10 mg)</td>
<td>100 ± 3</td>
<td>101 ± 8</td>
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<td></td>
<td>High (50 mg)</td>
<td>500 ± 15</td>
<td>497 ± 28</td>
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Table 2. Comparison of gravimetric and Karl-Fisher determination of the water content. $N = 10$.

<table>
<thead>
<tr>
<th></th>
<th>Water content g kg$^{-1}$</th>
<th>Recovery %</th>
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<tr>
<td></td>
<td>Gravimetry</td>
<td>Karl-Fisher</td>
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<tr>
<td>SiO$_2$</td>
<td>46 ± 1</td>
<td>45 ± 3</td>
</tr>
<tr>
<td>Molecular sieves</td>
<td>165 ± 4</td>
<td>167 ± 6</td>
</tr>
<tr>
<td>Al$_2$O$_3$</td>
<td>50 ± 2</td>
<td>48 ± 2</td>
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Fig. 1. Karl-Fisher analysis of the reference material NIST1649a. Left panel: original sample signal and blank signal. Right panel: sample signal after blank subtraction.
Fig. 2. Different thermal ramps (left panels) and resulting Karl-Fisher curves (right panels) for the analysis of reference material NIST1649a.
Fig. 3. Karl-Fisher curves of about 50 mg of standard materials HYD (upper panel) and WSO (lower panel).
Fig. 4. Karl-Fisher curves of SiO$_2$ (upper panel), molecular sieves (middle panel) and Al$_2$O$_3$ (lower panel).
Fig. 5. Scatter plot of the water content of PM$_{10}$ samples collected side-by-side (20 pairs).
Fig. 6. Examples of the Karl-Fisher curves of PM$_{10}$ paired samples collected side-by-side at the three sites (left panels) and drift profiles of some single hydrophilic components of PM (right panels).