Interactive comment on “Where do winds come from? A new theory on how water vapor condensation influences atmospheric pressure and dynamics” by A. M. Makarieva et al.

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Reply to the eight questions of Dr. Nick Stokes: What independent physics is contained in Eq. 34

We thank Dr. Nick Stokes for his continued interest in our work. Here we answer his eight questions concerning Eq. 34 and its derivation in our previous comment (hereafter M10-C1). By doing so we also attend to the first out of the four issues outlined by Dr. Judith Curry in her review: the need to clarify the meaning and derivation of Eq. 34. Here we abbreviate the questions and re-arrange their order and encourage readers to refer directly to Dr. Stokes comment regarding his precise wordings.

1. What is \( N_v \)?

Eq. (34), see also Eq. (8) in M10-C1 and below, is based on the premise that the deviation \( k_v \) of the vertical distribution of water vapor (or indeed any other condensable component) from equilibrium is due to condensation alone. (This premise is empirically testable: where condensation is absent, the vertical water vapor distribution should have the same scale height as the non-condensable gases and moist air as a whole.) This relaxes the need to consider \( N_v \) as the saturated vapor concentration and ensures that Eq. (34) is generally applicable. When \( k_v = 0 \), the condensation rate is zero independent of whether water vapor is saturated or not. When \( k_v \neq 0 \), \( N_v \) is saturated water vapor by formulation.

5, 8. Linearity of \( S \) over \( N_v \)

The assumption of linearity is based on consideration of condensation as a first-order reaction over saturated molar density \( N_v \) of the condensing gas. The rate of first-order reactions is directly proportional to the molar density of the reagent, with the proportionality constant having the dimension of inverse time: \( S = C N_v \), where \( C \) (dimension \( s^{-1} \)) is in the general case independent of \( N_v \).

This linearity is a natural theoretical assumption justified by the particular physical nature and stoichiometry of condensation, with gas turning to liquid. Note, for example, that the reverse process (evaporation) should be a zero-order reaction over \( N_v \). Experimental validation of the linearity of condensation rate over \( N_v \) can be made by

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\(^1\) Condensation rate and hydrostatic equilibrium of moist air, http://www.atmos-chem-phys-discuss.net/10/C10922/2010/

\(^2\) http://www.atmos-chem-phys-discuss.net/10/C11235/2010/

\(^3\) http://www.atmos-chem-phys-discuss.net/10/C11046/2010/ p. C11049.
considering condensation of water with different isotopic composition.

1. The limit \( N_v \to 0 \)

Coefficient \( C \) in chemical kinetics depends on temperature and the molecular properties of the reagent as follows from the law of mass action. Since the saturated concentration \( N_v \) of condensable gas depends on temperature as dictated by the Clausius-Clapeyron law, we can ask what the proportionality coefficient \( C \) physically means in this case. Different substances have different partial pressures of saturated vapor at any given temperature – this is controlled by the vaporization constant \( L \) and the molecular properties of the substance. Note too that for any given substance (like water) the saturated concentration depends on various additional parameters including the curvature of the the liquid surface and availability of condensation nuclei. Therefore, a range of saturated concentrations is possible at any given temperature. This allows one to consider \( C \) and \( N_v \) as independent variables in the space of all possible combinations of \( C \) and \( N_v \), see also the discussion below.

2. What is equilibrium?

The notion of equilibrium and deviation from it is key to determining the rate of any reaction. For example, in the case of evaporation the deviation from equilibrium is measured by the water vapor deficit: the deviation of relative humidity from unity. Atmospheric condensation is peculiar in being physically associated with air movement in a particular direction – water vapor condenses as the air moves vertically towards the lower temperature.

Here by equilibrium we understand the vertical distribution that the water vapor would locally take in the absence of condensation, all other conditions being equal. Let us denote the inverse scale height of such an equilibrium distribution for \( k_E \). Since in equilibrium there is no condensation, we now assume that the condensation rate \( S \) is proportional to the first order deviation \( k_v \) of the observed vertical distribution of water.
vapor from the equilibrium:

\[ k_v = -\frac{1}{N_v} \frac{\partial N_v}{\partial z} - k_E. \]  

(2)

The physics of Eq. (2) consists in the fact that the character of equilibrium distribution is not affected by condensation. For example, for the case of hydrostatic equilibrium any gas having molar mass \( M \), temperature \( T \) and finding itself on a planet with acceleration of gravity \( g \) in the presence of a temperature gradient \( \partial T/\partial z \) will have a distribution of its molar density following \( -\partial N/\partial z = k_E N \), where \( k_E = Mg/RT + (1/RT)\partial T/\partial z \).

3.4. What are \( k, k_v, \) and \( k_d \)?

As follows from Eq. (2), the expression for condensation rate depends on the nature of the equilibrium under consideration. The case considered in M10 corresponds to the case when it is moist air as a whole that is in equilibrium. This means that \( k_E = -(1/N)\partial N/\partial z \equiv k \). Then \( k_v \) is obtained from Eq. (2) with \( k_E = k \) and \( k_d \equiv -(1/N_d)\partial N_d/\partial z - k \) is obtained from the relationship \( N = N_v + N_d \).

From consideration of the Clausius-Clapeyron law and hydrostatic equilibrium (see also Eqs. (20), (24) and (25) in our paper, pp. 24024, 24026) one can see that

\[ k_v = \frac{L \Gamma}{RT^2} - \frac{Mg}{RT}, \]  

(3)

where \( L \) is molar vaporization constant, \( \Gamma \equiv -\partial T/\partial z \) is temperature lapse rate, and \( M \) is molar mass of air.

The value of \( k_v \) is controlled by temperature lapse rate \( \Gamma \) – keeping all other variables constant, changing \( \Gamma \) it is possible for \( k_v \) to take any value, \(-\infty < k_v < \infty \). This validates our assumption that \( k_v \) can be kept independent of \( N_v \) when investigating the limit behavior \( N_v \to 0 \) in Eq. (1): for any \( N_v \) (e.g., set by ambient temperature) any value of \( k_v \) can be prescribed by changing \( \Gamma \).

The independent physics of Eq. 34

The answer to the last (unnumbered) question of Dr. Stokes, what extra physics makes Eq. 34 independent of Eqs. 32 and 33, is as follows (note that it is also given on p. 24034 in M10).

Eqs. 32 and 33, taken together, contain the information that it is water vapor and not dry air that undergoes condensation. Eq. 34 contains information about the magnitude of deviation from equilibrium that causes condensation. Jointly considered, these facts are sufficient to determine the horizontal pressure gradient produced by condensation.

We emphasize the point we made in our paper, p. 24033, and M10-C1: if it is dry air that is in equilibrium, i.e. \( k_E = -(1/N_d)\partial N_d/\partial z \), the same consideration of the same equations gives \( \partial N_d/\partial x = 0 \) instead of \( \partial N_d/\partial x = S/u \) as in the case when it is moist air that is in equilibrium. This profound impact of the nature of equilibrium on the dynamics associated with condensation has never been expored in meteorological theory.