



On biases in atmospheric CO inversions assimilating MOPITT satellite retrievals

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Abstract

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Carbon monoxide (CO) inverse modelling studies have so far reported significant discrepancies between model concentrations optimized with the Measurement of Pollution in the Troposphere (MOPITT) satellite retrievals and surface in-situ measurements. Here, we assess how well a global chemistry-transport model (CTM) fits a large variety of independent CO observations (surface and aircraft air sample measurements and ground-based column retrievals) before and after assimilating MOPITTv6 total column (X_{CO}) retrievals to optimize CO sources/sinks. Consistent negative prior biases to all types of observations in all sensitivity tests suggest an underestimation of current surface emissions in the Northern hemisphere. In contrast, prior simulations fit the surface air sample observations well in the Southern hemisphere but underestimate CO in the free troposphere and on average in the column. Positive biases in MOPITT retrievals are identified in the Northern mid- and high latitudes, highlighting the importance of proper bias-correction of those satellite retrievals. Biases in representing vertical CO profiles are found over the ocean and most significantly in the Southern hemisphere, suggesting errors in the vertical distribution of CO chemical sources/sinks or in the vertical mixing to be improved in future modelling studies. Varying model-data differences are found in the vertical between CTM and MOPITT retrieved vertical profiles after having assimilated MOPITT X_{CO} ; these bias structures indicate that the posterior model differences to in-situ observations would be even larger if the near surface retrievals were assimilated instead of X_{CO} . In addition, given the higher long-term stability of the MOPITT X_{CO} retrievals as opposed to divergent temporal bias drifts found in the vertical profile retrievals and a lower sensitivity to model errors in the total column quantity than at a certain altitude, we recommend assimilating the column rather than the profiles or the
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40 partial profiles at the current stage.



1 Introduction

Carbon monoxide (CO) is a pollutant, an ozone precursor, and a key driver of the oxidation capacity of the atmosphere. Its direct natural and anthropogenic sources at the Earth's surface are mainly from incomplete combustion processes. It is also produced in the atmosphere by the photochemical oxidation of methane (CH₄) and non-methane volatile organic compounds (NMVOCs) (Shindell *et al.*, 2006). Being the major sink of hydroxyl radical (OH) in the troposphere, CO influences the tropospheric oxidation capacity and thus the lifetime of CH₄, NMVOCs, and other trace gases whose major sink path is OH (Logan *et al.*, 1981). With an average lifetime of about 1-2 months depending on latitude, CO also serves as a useful proxy tracer for emission and transport of other species that are co-emitted with CO such as CO₂ from fossil fuel (Worden *et al.*, 2012) and biomass burning (Yin *et al.*, 2016).

Related to its importance for air quality and climate, CO is one of the best-observed trace gases with air sample measurements from surface stations (Larsen *et al.*, 1999; Novelli *et al.*, 2003), aircraft (Wofsy *et al.*, 2011; Zbinden *et al.*, 2013), and with ground-based total column retrievals from stations in the Total Carbon Column Observing Network (TCCON) (Wunch *et al.*, 2011). CO has also been observed globally by satellite from several space-borne instruments, like the Measurement of Pollution in the Troposphere (MOPITT) (Deeter *et al.*, 2014) or the Infrared Atmospheric Sounding Interferometer (IASI) (Clerbaux *et al.*, 2009), with some coarse vertical profiling capability for some of the products. The official MOPITT CO products, now reaching version 7, has played a pivotal role in the study of the spatial-temporal distribution and variability of CO sources and sinks at a global scale (Arellano *et al.*, 2004; Duncan *et al.*, 2007; Fortems-Cheiney *et al.*, 2011; Gaubert *et al.*, 2016; Hooghiemstra *et al.*, 2012; Kopacz *et al.*, 2010; Shindell *et al.*, 2006; Stein *et al.*, 2014; Worden *et al.*, 2013; Yin *et al.*, 2015).

The chemical transport models (CTMs) that link the source/sink processes to atmospheric mixing ratios are key to the interpretation of the CO measurements, particularly in the global atmospheric inversion approach that infers the surface emissions from observed atmospheric mole fraction gradients. Although CTM performances have improved significantly over recent decades, model results are still plagued by systematic uncertainties due to a complex interplay of source distributions, transport and chemistry, with most CTMs showing negative biases to surface and satellite observations in the Northern hemisphere when prescribed with current emission inventories (Naik *et al.*, 2013; Patra *et al.*, 2011; Shindell *et al.*, 2006).

Satellite retrievals are partial or total column-integrated concentrations, with different vertical sensitivities at different altitudes depending on the instruments. Assimilating surface in-situ measurements or MOPITT CO retrievals using the same CTM and inversion configuration could result in considerable differences in the posterior surface fluxes (Hooghiemstra *et al.*, 2012). MOPITT-based atmospheric inversions were also shown to be biased high when compared to independent in-situ surface observations in the boundary layer (Gaubert *et al.*, 2016; Yin *et al.*, 2015). Assimilating MOPITT CO profile retrievals for different pressure levels resulted in significant differences in the posterior CO budgets, suggesting some inconsistency in the information provided to the inversion systems on vertical CO distributions



90 between the model and the observation (Jiang *et al.*, 2015).

In order to diagnose these inconsistencies that could result from model errors or satellite retrieval biases or both, this study evaluates the results of a global MOPITT CO total column assimilation using LMDz-SACS (as described by Yin *et al.* 2015) against various independent measurements for the period 2009-2011. Three categories of CO measurements are used for the evaluation: i) independent ground-based TCCON total column retrievals, ii) independent in-situ observations from surface networks and aircraft, and iii) MOPITT vertical profile retrievals at ten pressure levels. The observations used here have diverse spatial footprints on surface CO sources and different sensitivities to horizontal and vertical transport and chemical processes. We also include a series of sensitivity tests (described in Section 2.2 and 2.3) to discuss model uncertainties.

This paper is organized as follows. Section 2 describes the datasets and methods used in this study, covering successively (i) the MOPITT satellite retrievals, (ii) our CTM, LMDz-SACS, (iii) the prior/posterior simulations and sensitivity tests, and (iv) observations used for model evaluation. In section 3, we first evaluate the model with MOPITT X_{CO} that is used for assimilation, then we present the results of evaluation against independent CO observations from TCCON total column retrievals, surface station measurements, and aircraft vertical profile measurements. Section 4 represents the assessment against MOPITT vertical profile retrievals before and after assimilating the MOPITT total column retrievals. In section 5, we summarize all evaluation results and discuss potential biases in the model and in the satellite retrieval, as well as their implications for CO inverse study.

2 Data and Methods

2.1 MOPITT satellite retrievals of CO total column and vertical profiles

The MOPITT instrument flies in a sun-synchronous polar orbit at an altitude of 705 km, onboard NASA's Terra spacecraft launched in December 1999 (Deeter, 2003). It has been providing the longest continuous temporal coverage of global CO measurements from space with a spatial horizontal resolution of $22 \times 22 \text{ km}^2$ since 2000 (with some interruptions in 2001 and 2009). "Multispectral" retrieval products since version 5 combined near infrared (NIR, around $2.3 \mu\text{m}$) and thermal infrared (TIR, around $4.7 \mu\text{m}$) radiances and thus offer higher vertical resolution and higher sensitivity to the lower troposphere that are closer to the surface sources (Deeter *et al.*, 2013). In this study, we refer to MOPITT TIR/NIR CO retrieval version 6. It benefits from various improvements given some issues identified in version 5, including a correction of geolocation bias due to instrument viewing angle, an updated *a priori* CO profile, a higher quality meteorological field for water vapour and temperature, and a radiance bias correction to reduce the latitude-dependent upper troposphere retrieval bias found in V5 (Deeter *et al.*, 2014).

The MOPITT-retrieved total column integrated dry air CO molecule number, noted as X_{CO} hereafter, are calculated in a Bayesian way as

$$C_{rtv} = C_{prior} + A'(X_{mod} - X_{prior}) \quad (1)$$



where $\mathbf{C}_{\text{prior}}$ and \mathbf{C}_{rtv} represent the prior and the satellite retrieved total column, expressed as molecule/cm²; $\mathbf{X}_{\text{prior}}$ and \mathbf{X}_{mod} respectively represent the prior and the modelled vertical CO profiles for the satellite retrieval (here, with \mathbf{X}_{mod} based on information gain from radiance transfer model). The vertical profiles are resolved at 10 pressure levels, including surface pressure and successively from 900 hPa to 100 hPa with 100 hPa intervals. The vector quantities \mathbf{X} are expressed as log₁₀ of the corresponding volume mixing ratio. \mathbf{A}' is the retrieval averaging kernel matrix, which quantifies the sensitivity of both the instrument and the retrieval to the abundance of CO at different pressure levels (Deeter *et al.*, 2013).

The MOPITT CO volume mixing ratios are also retrieved at the 10 pressure levels for the satellite products, expressed as

$$\mathbf{X}_{\text{rtv}} = \mathbf{X}_{\text{prior}} + \mathbf{A}(\mathbf{X}_{\text{mod}} - \mathbf{X}_{\text{prior}}) \quad (2)$$

where $\mathbf{X}_{\text{prior}}$ and \mathbf{X}_{mod} represent the same quantities as in Equation (1), and \mathbf{X}_{rtv} represent the retrieved CO profile at the ten pressure levels. The method for covering \mathbf{A} and \mathbf{A}' is described in the MOPITT Version 4 User Guide (Deeter, 2009).

We use MOPITT X_{CO} for the global inversion. To calculate the model equivalent X_{CO} to MOPITT, we use Equation (1) with \mathbf{X}_{mod} being CTM concentrations interpolated at the ten MOPITT pressure levels, together with associated prior CO profiles and averaging kernels. MOPITT retrievals with a solar zenith angle larger than 70°, with latitudes within 25° from the poles, or with surface pressures less than 900 hPa are excluded from the inversion as detailed in Yin *et al.*, (2015). The MOPITT vertical profile retrievals are used for model evaluation as described later in Section 2.4.4.

2.2 Atmospheric Chemical Transport Model

The transport model used in this study is the global general circulation model LMDz (Hourdin *et al.*, 2006), which is part of the Earth system model of the Institut Pierre-Simon Laplace (IPSL). It is coupled with a simplified chemistry module SACS (Pison *et al.*, 2009), derived from the full chemistry model – Interaction with Chemistry and Aerosols (INCA) (Hauglustaine, 2004). SACS represents the main species in the chemical oxidation chain of CH₄–CH₂O–CO; it also represents methyl chloroform (MCF), with OH being the common sink path to relate all the tracers (Yin *et al.*, 2015).

The reference version of LMDZ-SACS, used to perform the inversion (briefly described below in Section 2.3), has a vertical resolution of 19 eta-pressure levels and a horizontal resolution of 2.5° (latitude) × 3.75° (longitude), noted as medium resolution (**MR**) hereafter.

We also include models with higher resolution for sensitivity tests, noted as **HR** (High Resolution). They are associated with a higher vertical resolution of 39 eta-pressure levels and a finer horizontal resolution of 1.89° (latitude) × 3.75° (longitude). Version **HR1** has the same physics configuration as MR, using the deep convection scheme of (Tiedtke, 1989) and the boundary layer mixing scheme of (Louis, 1979). **HR2** is an updated version that uses the deep convection scheme of (Emanuel, 1991) and the boundary layer mixing parameterization from (Deardorff, 1966). HR2 corresponds to the version called standard physics (SP) in (Locatelli *et al.*, 2015),



190 where more detail regarding these configurations are described. HR2 has been shown to represent better the deep convection and inter-hemisphere exchange than HR1 as suggested by Radon (^{222}Rn) and sulfur hexafluoride (SF_6) observations (Locatelli *et al.*, 2015).

195 For all versions of LMDz-SACS here, sea surface temperature and sea ice coverage from ERA-Interim climate reanalysis produced by the European Centre for Medium-Range Weather Forecasts (ECMWF) are used as boundary conditions; horizontal winds are nudged towards the ERA-Interim wind fields with a relaxing time of 3 hours.

2.3 Simulation configurations

200 In this study, we analyze two sets of global CO simulations for the period from 2009 to 2011: forward simulations prescribed with prior CO sources/sinks and posterior simulations after assimilating MOPITT X_{CO} . All simulations are listed in **Table 1**.

2.3.1 Prior forward simulations

205 All the three LMDz-SACS versions described in Section 2.2 are used to perform the prior forward simulations, noted as **F**. The prior CO sources include: i) bottom-up emission inventories of CH_4 , CH_2O , and CO, and ii) 3D chemical production fields of CH_2O from the oxidation of NMVOCs emissions simulated by INCA. The surface emission maps are aggregated from their original inventories to corresponding model resolution (details available in Yin *et al.*, 2015). The 3D chemical CH_2O productions are modelled with MR and HR versions of the LMDZ-INCA full chemistry model respectively, given the same emissions and climate forcing (updates from Folberth *et al.*, 2006). The spatial distributions of CH_2O production between MR and HR results are similar. One scale factor is applied to the HR results over the globe to conserve the global mass budget to be consistent with the reference version MR.

215 For CO sink, two different OH fields are used: INCA-OH and TransCom-OH. INCA-OH is produced by the LMDZ-INCA full chemistry model in the same simulation as for chemical CH_2O production using MR and HR respectively (updates from Folberth *et al.*, 2006). TransCom-OH was originally prepared for the international TransCom-CH₄ experiment of Patra *et al.*, (2011). The TransCom-OH fields are associated with a horizontal resolution of 1 degree and a vertical resolution of 60 levels; they are linearly interpolated to the model resolution of MR and HR respectively. INCA-OH concentrations are generally higher over the Northern Hemisphere (NH) but lower in the high latitudes of Southern Hemisphere (SH) compared to TransCom-OH. Thus, the two OH have very different North-to-South inter-hemisphere ratios, with ~ 1.2 for INCA and ~ 1.0 for TransCom. The vertical structures between the two prior OH fields are also different: INCA-OH has higher OH concentrations in the lower troposphere below 700 hPa over the tropics and subtropics, but much lower concentrations above the tropopause compared to TransCom-OH (for more information see Figure 3 in Yin *et al.*, (2015)).

2.3.2 Posterior simulations

The posterior CO sources/sinks are optimized using the MR version assimilating



235 simultaneously the MOPITT X_{CO} and surface observations of CH_4 and MCF as
described in Yin *et al.*, (2015). The HR versions are used for evaluating the posterior
fluxes obtained with MR. The posterior simulations are noted as **P**. Two MR
inversions associated with INCA and TrasCom OH are performed respectively, as OH
fields were shown to induce the largest uncertainty in our prior forward simulations
and in previous modelling studies.

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For the MR inversions, for each weekly window, the surface emission maps of
 CH_4 , CH_2O , and CO are optimized for each model grid point, the 3D secondary CH_2O
sources are optimized by a scaling factor for the entire column at each model grid
point, and the 3D OH fields are adjusted by scaling factors over six large latitude
245 bands. That is to say, there are no direct modifications to the vertical distribution of
the chemical CO sources in the atmosphere or to the OH fields in the inversion, and
the horizontal OH gradients within a large latitude band are not adjusted. More detail
regarding the inversion configurations are given in Yin *et al.*, (2015). The MR-
optimized surface emissions and scaling factors for CH_2O productions and OH fields
250 are then spatially interpolated to the HR resolution, with global budgets being
conserved.

In summary, there are four simulations for each model version (MR/HR1/HR2):
prior and posterior simulations associated with INCA or TransCom OH respectively.
255 It is emphasized that posterior HR simulations are prescribed with optimized surface
emissions and scaling factors obtained by the MR model with the same OH type. The
vertical distribution of CH_2O and OH are not exactly the same due to changes in
model resolution. The posterior model performances of HR1 and HR2 are not a direct
evaluation of the model ability to fit independent observations, but rather, they
260 demonstrate the model sensitivity due to resolution/parameterization.

2.4 Observations for evaluation and statistical analysis

2.4.1 Ground-based CO total column measurements

265 We use ground-based X_{CO} retrievals from the TCCON
(<http://www.tcon.caltech.edu/>) (Wunch *et al.*, 2011), a network of Fourier transform
spectrometers (FTSs) from near-infrared (NIR) solar absorption spectra, designed to
retrieve precise total column abundances of CO_2 , CH_4 , N_2O and CO. It is noted that
the MOPITT NIR/TIR retrievals, combining information from both TIR and NIR,
270 have generally higher sensitivity to the lower troposphere compared to TCCON.
Besides, their footprints are also different. The TCCON stations are listed in **Table 2**,
with their location indicated in **Figure 1**. The TCCON prior profile and averaging
kernel associated with each observation are used to calculate corresponding model
retrievals, in a similar Bayesian way as the MOPITT X_{CO} .

2.4.2 Surface station measurements

275 In-situ CO measurements are accessed from the World Data Centre for
Greenhouse Gases (WDCGG) database (<http://ds.data.jma.go.jp/gmd/wdogg/>,
[accessed on Dec. 2015](#)). We include CO surface measurements from 88 stations to
evaluate modelling results within the boundary layer, calibrated to WMO scale. The
280 geographical distributions of the stations are shown in **Figure 1**, and a list of stations
is available in **Table S1**.



2.4.3 Aircraft measurements

We use aircraft in-situ CO measurements of vertical profiles from two projects:

285 1) MOZAIC/IAGOS (Measurements of Ozone and water vapour by in-service
Airbus aircraft / In-service Aircraft for a Global Observing System)
(<http://www.iagos.org>), which contain 20 years (1994 to 2015) of in-situ CO, O₃ and
NO_x measurements, sampled from multiple commercial flight during take-off
290 (ascending) and landing (descending) to achieve quasi-vertical profiles (Marengo *et al.*, 1998; Petetin *et al.*, 2015). The observations are thus close to international airports
and hence represent an environment of high anthropogenic emissions, in particular in
the layers close to the surface.

295 2) High-Performance Instrumented Airborne Platform for Environmental
Research (HIAPER) Pole-to-Pole Observations (HIPPO) large-scale aircraft
campaigns, which sampled the atmosphere from the North Pole to the coastal waters
of Antarctica over the Pacific Ocean (https://www.eol.ucar.edu/field_projects/hippo)
(Wofsy and HIPPO Science Team and Cooperating Modellers and Satellite Teams,
2011). The flights took air samples from the surface up to 14 km on a campaign basis,
300 spanning different seasons in different years (I: Jan. 2009, II: Nov. 2009, III: Mar.-
Apr. 2010, IV: Jun.-Jul. 2011, V: Aug. – Sep. 2011).

Individual aircraft profiles are assigned to certain model grid points given their
geographic location and pressure levels, and all measurements are then averaged per
305 model grid point at a 30-minute resolution to compare with corresponding model
value. Only observations over land are kept for the MOZAIC/IAGOS measurements
for further analysis; while only observations over the ocean are kept for the HIPPO
measurements. Their spatial coverages are shown at the model resolution in **Figure 1**.

310 2.4.4 MOPITT CO vertical profile retrievals

The MOPITT vertical profile retrievals are used for evaluation. As the MOPITT
total columns are assimilated, the profiles are not independent of the posterior
simulations. The analysis of posterior models would thus show how well the model
agrees with the MOPITT profile when fitting the X_{CO}. The MOPITT-equivalent
315 model CO vertical profiles are retrieved in a consistent manner as the satellite
retrievals using Equation (2) described in section 2.1, in which the X_{mod} become the
CTM concentrations interpolated at the then MOPITT pressure levels to be convolved
with corresponding MOPITT prior CO profiles and averaging kernels.

320 2.4.5 Model sampling and statistical analysis

Model values are sampled according to the time and location of the
measurements at a temporal resolution of 30 minutes and the model associated spatial
resolution. In the case of surface and aircraft measurements, no horizontal or vertical
interpolation among nearby model grid points is made to smooth the sampling. In the
325 case of interpolating the model vertical profile to other instrumental levels to obtain
X_{mod} at the instrument vertical resolution, e.g. from the model 39 vertical levels to the
10 MOPITT pressure levels, the total column of the model state integral is always
conserved to compensate uncertainty from vertical resolution change on the CTM side.



330 The same applies for TCCON X_{CO} retrievals as well. For each dataset, the observed and simulated values are averaged per month for further analysis in this study, e.g. the model-data biases are calculated using monthly averages of model-equivalences minus corresponding observations.

335 3 Results

3.1 Evaluation against the MOPITT total column (X_{CO})

340 The latitudinal mean bias of simulated CO compared to MOPITT X_{CO} is shown in **Figure 2a-b**. The prior forward model results show some significant underestimation of CO concentrations compared to the MOPITT X_{CO} , particularly in the northern hemisphere, irrespective of the model version or of the OH field (shown by dotted lines). Logically, the posterior model results agree with MOPITT X_{CO} much better than the forward run before optimization, with a mean bias of less than 0.5 ppb for both OH in the reference MR version (shown by solid red lines). Small positive biases using the HR models are noticed (shown by solid green/blue lines). The statistics of model biases for four latitudinal bands are summarized in **Table 3**.

350 For the reference transport model version MR, the mean X_{CO} bias of the prior forward modelling (shown by dotted-red lines in **Figure 2a-b**) is -12.1 ± 7.6 ppb and -9.4 ± 3.7 ppb for INCA-OH and TransCom-OH respectively (with the range showing 1-sigma standard deviation of the mean biases of all model grids, the same applies hereafter if not specified otherwise). INCA-OH produces a larger negative bias to MOPITT in the Northern high latitudes than TransCom-OH, because of its higher OH concentration in the NH. Accordingly, the inversion using INCA-OH gives $\sim 17.9\%$ higher NH emissions, but $\sim 18.4\%$ lower SH emissions than the inversion using TransCom-OH, resulting into $\sim 6.1\%$ higher global emissions in the INCA-OH based inversion. It is noted that although we optimize OH together with surface emissions, the system only scales slightly the big-region OH state vectors (in total six big regions over the globe) and thus the inverted surface emissions are sensitive to the prior OH fields.

365 The two HR models (shown in green and blue in **Figure 2a-b**) produce slightly larger X_{CO} values than the MR model when being forced by either prior or optimized CO fluxes. The CO surface sources are conserved in global mass between different resolutions when emitted to the atmosphere, but with the change in resolution, the CO sink via the reaction with OH (associated with different resolutions) may differ. For the global average, the HR models generate ~ 1.0 ppb ($\sim 1.5\%$) and ~ 2.0 ppb ($\sim 3\%$) higher X_{CO} compared to the MR, using INCA-OH and TransCom-OH respectively.

370 The difference between the HR and MR X_{CO} results (~ 1.5 ppb) is of a much smaller magnitude than the differences between the prior and the posterior MR simulations (~ 10.4 ppb); it is also of a smaller magnitude than that induced by the two OH fields in the prior forward simulations when the CO sources are identical (~ 2.8 ppb for global average with some cancelling effect between the NH and SH). The differences in modelled X_{CO} between HR1 and HR2 are not significant at a global scale.



3.2 Evaluation against TCCON column measurements

380 The average biases of modelled X_{CO} compared to TCCON measurements along
latitude are shown in **Figure 2c-d**. The prior forward simulations show general
underestimation of X_{CO} compared to TCCON measurements (shown in light colours
in **Figure 2c-d**), consistent with the negative bias against MOPITT X_{CO} (dotted lines
in **Figure 2a-b**). This negative bias is of -16.0 ± 6.0 ppb with INCA-OH and -9.3 ± 4.0
385 ppb with TransCom-OH on global average. The posterior MR results (shown in red
dots in **Figure 2c-d**), however, overestimate X_{CO} in the northern mid and high
latitudes compared to TCCON, resulting in a positive bias of 7.4 ± 6.4 ppb on average
in NH using INCA-OH (7.3 ± 6.0 ppb using TransCom-OH). They also slightly
overestimate X_{CO} in SH, resulting in a smaller positive bias 2.5 ± 3.1 ppb using INCA-
OH (3.0 ± 2.2 ppb using TransCom-OH).

390 The average prior forward model difference in TCCON equivalent X_{CO} between
the two OH is ~ 7 ppb, inducing the largest uncertainty among factors tested here
(difference between **Figure 2c** and **2d**). The HR models (shown in green and blue)
produce ~ 2.8 ppb ($\sim 3.6\%$) higher concentrations than the MR (shown in red), when
395 using the same type of CO sources and OH profiles (but different in resolution). The
relative differences of the column mixing ratios between HR1 and HR2 are not
significant, suggesting that modelling the column is not very sensitive to the choice of
deep convection and vertical diffusion schemes, however the model resolution does
have a visible impact (HR vs. MR). Statistical analysis of model biases and model-
400 data correlations for individual stations are shown in **Table S2**.

3.3 Evaluation against independent surface in-situ measurements

The average biases of model CO dry air mole fraction [CO] compared to surface
in-situ measurements along latitudinal gradient are shown in **Figure 2e-f**. In the prior
405 forward simulations (shown in light colours), the reference model (MR) generally
agrees well with the surface observations in the SH, with only a small negative bias in
the SH sub-tropics. The mean MR prior biases in SH are -2.5 ± 6.5 ppb and -4.6 ± 6.8
ppb with INCA and TransCom OH respectively, thus different from the significant
negative prior bias compared to MOPITT and TCCON X_{CO} . However, negative prior
410 biases increase with latitude in the NH, peak around 50°N , and remain relatively high
in the NH high latitudes. The mean MR prior biases in NH are -30.6 ± 12.6 ppb and
 -20.1 ± 12.7 ppb with INCA and TransCom OH respectively (**Figure 2e-f; Table 2**).

The posterior MR models slightly overestimate in-situ [CO] in the SH mid-to-
415 high latitudes, with a change of sign and a degradation of the bias to surface station
data compared to the prior simulations (mean posterior bias in SH: 7.5 ± 8.1 ppb with
INCA-OH and 8.2 ± 8.3 ppb with TransCom-OH). The posterior modelled [CO] agrees
quite well with ground observations in the NH sub-tropics, but show significant
overestimation North of 30°N , i.e. a change in the sign and a slightly smaller
420 amplitude of bias (mean bias in NH: 7.6 ± 39.2 ppb with INCA-OH and 18.2 ± 23.8 ppb
with TransCom-OH), as compared to the forward simulations (**Figure 2e-f**). The
average difference between the HR and MR model is 3.9 ± 4.1 ppb ($\sim 4\%$) when other
setups being the same; here, the 1σ standard deviation showing the spread across
prior/posterior simulations and between the two HR models. HR2 simulates slightly



425 higher surface CO concentrations than HR1 (1.4 ± 0.6 ppb, $\sim 1\%$). Larger differences
of model results among different model configurations are found in the surface [CO]
than in the quantity of the column integral, suggesting a higher sensitivity in
modelling boundary layer mixing ratios.

430 In the prior forward simulations, significant correlations between the model-data
difference in surface [CO] and in MOPITT X_{CO} at the same horizontal model grid are
found for all three models in the NH ($r=0.72$ for INCA-OH and $r=0.63$ for TransCom-
OH on average, **Figure 3a**). This result suggests that both surface and satellite
435 observations give consistent information on the errors of the prior CO sources/sinks in
the NH, regardless of the model version or OH type. No such significant correlations
are found in the posterior simulations over the NH (**Figure 3b**), suggesting no
consistent systematic errors of the MOPITT X_{CO} assimilation compared to surface
[CO] measurements. On the contrary, in the SH prior forward simulations, no
440 significant correlations between the model-data difference in surface [CO] and in
MOPITT X_{CO} are found (**Figure 3c**), suggesting different patterns of the prior model-
data deviations when compared to in-situ observations in the boundary layer and to
the satellite X_{CO} retrievals. However, a negative correlation ($r=-0.49$) is noticed for
the posterior simulations using HR version (not for the MR version) associated with
INCA-OH, suggesting contradicting model-data agreements at the surface and in the
445 column likely due to error in the vertical CO distribution (**Figure 3d**).

3.4 Evaluation against independent aircraft measurements

3.4.1 MOZAIC measurements over large airports

450 MOZAIC aircraft profiles and corresponding model values are shown for four
latitudinal bands on the left panel of **Figure 4** (results with INCA-OH are shown).
Globally averaged model biases to MOZAIC measurements and variations of model-
data correlations in the vertical profiles are shown on the left panel of **Figure 5**
(results with both OH fields). The forward model profiles show negative biases
455 are larger in the NH than in the SH, consistent with the different in prior model
column-integrated X_{CO} against MOPITT and TCCON.

The negative model-data biases in the prior model are well corrected in the
460 posterior simulations (presented by solid lines). The zonal vertical gradients are
generally well captured by the model (**Figure 4**). Compared to MOZAIC profiles, the
HR models represent the vertical gradients and boundary layer better than the MR
with a two-fold higher vertical resolution, i.e. the HR versions produce sharper
vertical gradients from the surface up to around 800 hPa, in better agreement with the
profile observation.

465 Apart from the general good representation of the MOZAIC profiles, modelled
CO concentrations are considerably lower in the near ground layers compared to the
MOZAIC measurements. This could be partly explained by non-modelled local
contamination of MOZAIC data in the boundary layer over large international airports,
470 given the limited resolution of our global model to resolve city scale emissions.
Another model deficiency compared to MOZAIC profiles is a positive bias in the
upper troposphere above 400 hPa (**Figure 5**), which becomes larger to the upper
levels, mostly contributed by the NH, in particular, the high latitudes (**Figure 4**).



475 Along with the increase of model biases in the vertical, model-data correlations also
degrade significantly above 500 hPa, with the lowest correlations at around 200 hPa.
HR versions show larger decreases in the model-data correlations in the upper
troposphere compared to MR. The vertical distributions of model-data correlations
between the two OH fields are similar, but INCA-OH simulates slightly smaller
positive biases in the upper troposphere than TransCom-OH (**Figure 5**).

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3.4.2 HIPPO measurements over the Pacific Ocean

The HIPPO aircraft profiles and corresponding model values (with INCA-OH)
are shown for four latitudinal bands on the right panel of **Figure 4**. Similar to that of
the MOZAIC measurements, the global average biases to HIPPO profiles and vertical
485 variations of model-data correlations are shown on the right panel of **Figure 5**. The
forward models (shown in dotted lines) significantly underestimate the CO
concentrations nearly across the entire vertical profile in the NH against HIPPO.
However, only small negative prior biases, limited to the upper troposphere, are found
in the SH (**Figure 4**). This pattern is consistent with the comparison to surface
490 network measurements and column integrated X_{CO} mixing ratios.

The posterior simulations show improvement of the overall agreement with
HIPPO profiles in the NH, but they overestimate the lower troposphere concentrations,
which is in agreement with the comparison to surface [CO] measurements. However,
495 over the SH, the forward models capture the lower tropospheric (from the surface to
800 hPa) CO concentrations better than the posterior models (**Figure 4**), consistent
with the small negative prior bias but significant positive posterior biases to surface
observations (**Figure 2e**). The posterior models also overestimate the top layers above
~300 hPa, similar to the feature found in comparison with MOZAIC. The model-data
500 correlations also drop significantly around 200 hPa, with a larger deterioration with
TransCom-OH than with INCA-OH (**Figure 5**). This pressure level is around the
tropopause depending on the latitude and season. The HR models simulate slightly
higher concentrations across the entire profile, consistent with a higher column
integrated mixing ratio equivalent to MOPITT or TCCON; they also represent the
505 boundary layer better.

Note that the HIPPO measurements were made over the Pacific Ocean, where
surface CO emissions are small and the CO sources are from secondary chemical
sources or transport. This is also partly the reason for the observed increasing
510 gradients from the surface toward the upper troposphere, up to ~500 hPa in the NH
and to ~300 hPa in the SH (**Figure 4**). Such bias in representing the oceanic vertical
profiles suggests error in the vertical distribution of CO source/sink over ocean or in
the vertical mixing.

515

4 Evaluation against MOPITT vertical profiles

After evaluating the model with independent observations, we now assess the
simulations against MOPITT CO vertical profile retrievals. The latitudinal average
520 MOPITT CO profiles and equivalent model retrievals (MR and HR1 using INCA-OH)
are shown in **Figure 6**. Compared to the MOPITT profiles (**Figure 6a**), the forward



simulations (**Figure 6b-c**) significantly underestimate the CO mixing ratios across the vertical profiles in the NH mid-high latitudes, consistent with the comparison to TCCON X_{CO} retrievals, and in-situ measurements from surface stations and MOZAIC aircrafts. The posterior vertical model profiles (**Figure 6e-f**) show better agreement with MOPITT profiles than the forward models, but there are still visible differences.

The average zonal profiles of the observation (solid lines), forward (dotted lines) and posterior (dash-dot lines) model results for the reference model (MR) are shown in **Figure 6d**. We focus here on the posterior model results, after assimilation of the MOPITT column X_{CO} thus not independent from the MOPITT profiles. The MOPITT vertical profiles are well reproduced by the model north of 30°N, but considerable deviations are found for the vertical structure over the tropics and subtropics. The vertical gradient of the CO decrease from the surface to the upper atmosphere is larger in the MOPITT profiles than in the MR model over the tropics and subtropics, in particular in the SH tropics. In the SH mid-high latitudes, the model produces lower concentrations at the surface but higher in the upper troposphere compared to MOPITT, contrary to the difference between the model states and HIPPO profiles. Significant CO enhancements around pressure level 200 hPa are noticed in the MOPITT profiles over the Tropics (**Figure 6a**), which persist throughout the year (**Fig. S1**). Such large CO enhancements in the upper troposphere are not represented by any version of the model; they are also not very visible in the HIPPO or MOZAIC profiles (**Figure 4**).

Comparing model results against MOPITT profile data at the surface level (higher than 900 hPa), the absolute biases in the prior simulation are significantly reduced after assimilating MOPITT X_{CO} . The prior surface layer zonal biases (MR) are -23.0 ppb, -4.4 ppb, -1.5 ppb, and -8.1 ppb for the four latitudinal bands from the north to the south, whereas the posterior biases (MR) are reduced to -6.5 ppb, +2.5 ppb, 1.6 ppb, and -5.4 ppb, with the signs changed over the tropics (**Figure 7**). The zonal average posterior model difference to MOPITT profiles remain negative for the mid-high latitudes in both NH and SH from the surface to ~700 hPa; On the contrary, the posterior model difference to MOPITT profiles become positive in the tropics (**Figure 7**).

For regional detail, the relative bias and the model-data correlation with MOPITT profiles for each pressure level are shown as pseudo Taylor diagrams in **Figure 8**. For all regions over land, the posterior simulations (shown by filled markers) agree better with MOPITT profiles than the prior (shown by empty markers). The posterior model-data correlation coefficients are higher than 0.9 for all regions for both X_{CO} and the near-surface level of the profiles. The model-data correlations decrease significantly with increasing altitude, suggesting that the model ability to represent the spatial distribution of MOPITT-based CO concentrations degrades towards the upper atmosphere. The poorest model-data correlations are found around 300 hPa, except for 200 hPa over the tropical ocean, in agreement with the comparison to aircraft measurements.

The posterior model CO profiles match the MOPITT profiles reasonably well over North America, Europe, Boreal Eurasia, Middle East, and Australia (**Figure 8**). Significant divergences of the posterior profiles vs. the MOPITT profiles are found over South America, Africa, Temperate Asia, and Indonesia, where the model



overestimates CO concentrations in the lower part of the atmosphere (900-600 hPa) but underestimates CO concentrations in the upper atmosphere. The posterior models also produce some positive biases over ocean, from the surface up to 400 hPa, although the column integrated X_{CO} are well fitted. Such misfit is most prominent over the tropical ocean, which is related to the MOPITT CO enhancement around 200 hPa as mentioned above.

The largest differences due to model configurations are induced by different OH fields, more than by model resolution or the choice of a convection scheme (**Figure 7**). The relative differences between the HR and MR retrieved MOPITT-equivalent CO profiles are not very significant at the regional scale (shown by different markers in **Figure 8**). The HR versions generally produce higher near-surface concentrations and stronger vertical gradients, thus being more in agreement with the structure of the MOPITT profile. However, these differences induced by model configuration are small; all model versions are associated with the vertically varying posterior model difference to the MOPITT profiles when MOPITT X_{CO} are reasonably fitted (**Figure 7**). The relative change of model differences to MOPITT vertical profiles between HR and MR or between the two HR versions are consistent no matter with OH field is used.

5 Discussion and conclusions

5.1 Sensitivity to model configurations

Using both our prior and MOPITT X_{CO} -assimilated CO sources/sinks, we evaluated the distribution of modelled CO mixing ratio in the atmosphere against independent observations from TCCON total column X_{CO} , surface measurements in the boundary layer, and aircraft profile measurements using six sets of model configurations. The choice of the prior OH seems to produce the largest differences in the simulated CO concentrations at a global scale. As the two prior OH field tested here have contrasted North-to-South distribution, the inversion using INCA-OH derives ~17.9% higher NH emissions, but ~18.4% lower SH emissions than the inversion using TransCom-OH, result in ~6.1% higher global emissions associated with INCA-OH. Although OH is optimized jointly with the surface emissions and atmospheric productions of CO, the system does not apply a strong correction to the big-region OH state vectors. The posterior OH fields stay very close to the prior ones, with scaling factors ranging from 0.97 to 1.01. Thus, the different N-S gradients between INCA and TransCom OH remain similar in the posterior.

TransCom-OH, with a North-to-South ratio of around 1, produce slightly smaller surface bias over the NH (**Figure 2e-f**) and also a smaller bias in the modelled seasonal cycle amplitude at the surface (**Figure S2**). The North-to-South gradient in TransCom-OH is also closer to a recent observation-derived near equal N/S OH distribution (Patra *et al.*, 2014). However, larger biases above 300 hPa are found with TransCom-OH, likely caused by inconsistencies between the imported climatological TransCom-OH (produced by another model) and the LMDz model physics, e.g. the difference in tropopause height between the model physics and that implied by the TransCom-OH profiles simulated by a different CTM (**Figure 7**). Our results,



620 therefore, highlight the importance of both horizontal and vertical OH structures in
modelling the CO vertical profiles, with upper atmospheric OH concentrations having
a larger impact on tracers with longer lifetimes.

625 Compared to the reference version (MR), the HR versions, with a nearly two-fold
increase in the vertical levels and a factor of one-third increase in latitudinal
resolution, simulate 2% higher MOPITT-equivalent- X_{CO} , 3.6% higher TCCON-
equivalent- X_{CO} , and 3.9% higher surface CO mixing ratio at sites from the surface
network when averaged globally. The differences between the two convection
630 schemes (HR1 vs. HR2) are not significant for the column integrated X_{CO} ; but for the
boundary layer, HR2 simulates slightly higher mixing ratios, in agreement with
previous studies showing that the HR2 scheme represents deep convection more
rigorously, agreeing better with ^{222}Rn observations (Locatelli *et al.*, 2015). The
differences between the two HR schemes associated with different physics and
convection schemes are much smaller than the differences between the HR and the
635 MR version, illustrating the importance of an appropriate vertical resolution to
represent vertical mixing and the boundary layer.

No significant correlations are found between the relative differences in surface
[CO] and column X_{CO} , for either HR vs. MR or HR2 vs. HR1 (**Figure S3**), suggesting
640 that the model does not systematically produce higher surface [CO] when it produces
higher X_{CO} . The small but consistent increase in the modelled X_{CO} from MR to HR
indicates that assimilating MOPITT X_{CO} with a higher resolution model is likely to
yield a slightly smaller surface emission compared to the MR results.

645 **5.2 Possible biases in the prior CO sources**

Despite the uncertainty due to model configurations, particularly OH, our
analysis shows that the prior simulations appear to underestimate the CO
concentrations persistently in the NH at the surface and at all altitudes, thus result in a
negative model bias for the total column (**Figure 2** and **Figure 4**). We also find
650 consistent model-data deviations in the surface and in the MOPITT X_{CO} vary in space
(**Figure 3a**), suggesting an underestimation of NH CO emissions in the current
inventories, in line with previous model studies. For instance, Stein *et al.*, (2014)
showed that a higher winter traffic emissions from North America and Europe and a
lower dry deposition rate could improve the agreement between simulated and
655 observed CO during winter and spring.

In the SH, however, prior simulations fit the surface in situ observations well, but
show negative biases to the aircraft profiles and X_{CO} (**Figure 2**, **Figure 3c**, and **Figure**
660 **4**), suggesting model errors in representing the vertical distribution of CO source/sink,
particularly over the ocean. This could be partly caused by errors in the vertical
distribution of tropospheric secondary chemical CO sources, as the surface emissions
are generally low but important sources are from chemical oxidation of hydrocarbons
and from long-distance transport (Zeng *et al.*, 2015). From this aspect, assimilating
 CH_2O would in principle improve the representation of chemical CO sources
665 (Fortems-Cheiney *et al.*, 2012), future studies to improve the vertical distribution of
secondary CO sources would improve the modelling of CO profiles.

Another possible source of errors in modelling the SH vertical CO profiles could



670 be a misrepresentation of the emission injection height of biomass burning by thermal
plumes. In our current model, pyrogenic emissions are released in the surface layer of
the model, in the same way as fossil fuel emissions, and the thermal impact of air
uplift is not particularly considered. Although most of the injection heights (>80%)
were observed to remain within the boundary layer, some occasional injections could
675 exceed 8 km and may contribute to some of the positive model bias at the surface
(Gonzi *et al.*, 2015; Labonne *et al.*, 2007), which may cause some errors in
representing the effects of fire emissions on the vertical structure of CO, as well as of
other CO precursors that are co-emitted from fires. For instance, the MOZAIC
measurements (mainly sampled over Africa as shown in **Figure 1**) show much higher
680 CO concentrations between 900 hPa and 800 hPa than the model in the sub-tropical
land in SH (0-30S) (**Figure 4**), where CO plumes are primarily from biomass burning.

5.3 Possible biases in the MOPITT retrievals

Although the prior model biases to the above-mentioned independent
observations are significantly reduced in most cases after assimilating the MOPITT
685 X_{CO} , there are noticeable remaining model biases in the posterior with a change in the
sign from negative in the prior modelling to positive in the posterior. The posterior
simulations show, on the one hand, consistent positive biases to independent
observations including TCCON X_{CO} , surface observations, and HIPPO vertical
profiles in the mid- and high latitudes (**Figure 2** and **Figure 4**); on the other hand,
690 negative biases to MOPITT vertical profile retrievals for the near surface levels (from
surface to 700 hPa) are seen over these regions (**Figure 7**). Such discrepancy indicates
positive biases in the MOPITT retrievals for the total column and at the near surface
levels in the mid- and high latitudes. As for the tropics and subtropics, the CO
enhancements around 200 hPa in the MOPITT vertical profile retrievals are not found
695 in the HIPPO or MOZAIC aircraft measurements.

A recent study directly compared the MOPITT retrievals against HIPPO aircraft
measurements and suggested positive X_{CO} biases in the mid-latitudes (40-70°N and
30-60 °S) and negative biases in the tropics (**Figure 1** in (Jiang *et al.*, 2016)). Their
700 results also showed a similar latitudinal bias pattern in the retrieved CO mixing ratios
in the lower troposphere as in the X_{CO} , associated with an opposite latitudinal bias
pattern in the upper troposphere retrievals. It is however noted that these evaluations,
in agreement with our model-data bias to HIPPO, were based primarily on
observations over the ocean made during reduced periods of time, while the
705 comparison of posterior models to MOZAIC profiles does not show such a significant
bias in our analysis (**Figure 4**). Therefore, more in-depth analysis is needed to bias-
correct the satellite data well.

Compared to the IASI data (v20100815), MOPITT total columns (V5T) are
710 generally higher over land (ranging from 0-13 %) for the overlapping period of the
two instruments from 2008 to 2013; Significant differences are due to a higher prior
CO in the MOPITT retrievals, but biases remain when reprocessing the MOPITT data
using the same a priori constraints as those used for IASI (George *et al.*, 2015). The
differences between these two instruments are significantly larger than that between
715 MOPITT v5 and v6 [Deeter *et al.*, 2014]. More study to reconcile recent satellite CO
retrievals that have improved quality and the combined use of multiple instruments as
explored in (Kopacz *et al.*, 2010) would benefit our understanding regarding CO



source/sink distributions.

720 **5.4 Assimilating satellite total column vs. vertical profile retrievals**

725 Significant divergences of the differences between the model and MOPITT
retrieved vertical profiles are found in the vertical among the ten MOPITT product
pressure levels (**Figure 7**, **Figure 8**, and **Figure S1**). Such a spread would certainly
result in significant differences in the inverted CO budgets when assimilating one
730 retrieval level at a time, i.e. as suggested in (Jiang *et al.*, 2015) , if there is no change
in the model configuration that could impact the vertical CO structures (e.g. CO
source/sink vertical distribution). With remaining negative posterior model biases to
MOPITT profiles in the lower troposphere over the NH and SH mid- and high
latitudes where significant positive model bias to surface and TCCON observations
735 are found, we anticipate that assimilating those profile retrievals of near surface levels
(from surface up to 700 hPa) would results in even larger positive biases (**Figure 7**).
Similar divergences were reported in previous studies using other CTMs to assimilate
MOPITT profiles or near-surface levels (Gaubert *et al.*, 2016; Hooghiemstra *et al.*,
2012; Jiang *et al.*, 2015). Jiang *et al.*, (2016) introduced a 4-order polynomial curve to
740 correct the latitudinal dependent bias for the MOPITT profiles at each retrieval
pressure level before the assimilation, comparing to HIPPO measurements as
described above in section 5.1. This approach could reduce the bias for a certain
pressure level, but it may degrade the overall column consistency and is not well
constrained.

745 In terms of long-term stability, the bias drift for the MOPITT V6 TIR/NIR
product varies from negative in the lower troposphere ($-1.3\% \text{yr}^{-1}$ at 800 hPa) to
positive in the upper troposphere ($1.6\% \text{yr}^{-1}$ at 200 hPa) as compared to aircraft
measurements; yet, the bias drift in the total column is negligible ($0.003\%/\text{year}$)
(Deeter *et al.*, 2014). Therefore, the column measurements are more suitable for
750 studies associated with long-term CO dynamics. In the meantime, the number of
retrieved layers exceeds the number of independent pieces of information available
vertically and hence the profiles are only weakly resolved, not representative of the
vertical resolution of the observation. The error covariance among different layers of
the profile retrievals is described in the full averaging kernel matrix, however we do
not have information about the profile error covariance of the CTM, *let alone* the fact
that the measurement error correlations are commonly ignored in inversion systems
for technical reasons.

755 From the CTM perspective, the evaluation against aircraft measurements reveals
significant model errors in representing the vertical CO gradient, in particular over the
ocean (**Figure 4**). Assimilating satellite retrievals at a particular pressure level would
be more prone to model biases in the vertical profiles, which are sensitive to the
surface pressure, the vertical distribution of chemistry-related source and sink terms,
760 the vertical mixing, and the lateral transport at a given height. Rayner and O'Brien,
(2001), Olsen and Randerson (2004), and Yang *et al.*, (2007) argued similarly for CO₂
that the assimilation of column-integrated data is less susceptible to uncertainty in
model transport than the surface network. The argument may be even stronger for
CO, because its relatively short lifetime makes the column integrated concentration
765 more informative about regional emissions.



On either the satellite retrieval side or the model side, representing the vertical profile correctly seems to be a grand challenge. Therefore, we argue that for the purpose of top-down estimates of CO emissions, in which the model cannot directly correct vertical model biases, it is more robust to assimilate the column than a particular pressure level retrieval, a partial profile, or the entire profile without the full AKs at the current stage. In terms of bias-correction, the use of a bias-correction scheme based on differences between satellite retrievals and a CO reanalysis assimilated with surface and aircraft measurements would be an alternative approach. More in-depth study with independent observations from various platforms, as well as from other satellite instruments, for instance IASI as explored in (George *et al.*, 2015), would be needed.

The model uncertainty that we discuss in this study is limited to the model representation of vertical profiles. Other aspects like lateral transport or NH-SH air exchange rate are not addressed. The horizontal and vertical distributions of OH induce the largest uncertainty in the modelling of global CO and thus its atmospheric inversion. The vertical distribution of chemical CO sources is the other key factor toward reconciling model agreement with X_{CO} and in-situ air samples. Future model studies would benefit from the use of multiple tracers, given that more than half of the CO sources are from secondary chemical oxidation of CH_4 and NMVOCs in the troposphere, and O_3 , NO_x are chemically related with CO through OH (Miyazaki *et al.*, 2015). The vertical distributions of all the tracers, measured from aircraft or satellite, would provide important information to improve our understating regarding the CO source/sinks, as well as the uncertainties of transport and chemistry in CTMs.



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810 We thank the entire HIPPO team for making these measurements available via <http://hippo.ucar.edu/>. The authors are very grateful to the many people involved in the surface and aircraft measurement and in the archiving of these data.



815 Tables and Figures

Table 1. List of simulations in this study. We include three model versions (MR/HR1/HR2), and for each version there are prior forward and posterior simulations (F/P) associated with two different OH fields INCA or TransCom (in/tc).

820

Model version	Physics	Simulation	CO sources	OH
MR* (2.5°x3.75°xL19)	Tiedtke (1989)	MR_F_in	Prior	INCA
		MR_F_tc		TransCom
		MR_P_in	Posterior_in	Posterior_in
		MR_P_tc	Posterior_tc	Posterior_tc
HR1 (1.895°x3.75°xL39)	Tiedtke (1989)	HR1_F_in	Prior	INCA
		HR1_F_tc		TransCom
		HR1_P_in	Posterior_in	Posterior_in
		HR1_P_tc	Posterior_tc	Posterior_tc
HR2 (1.895°x3.75°xL39)	Emanuel (1991)	HR2_F_in	Prior	INCA
		HR2_F_tc		TransCom
		HR2_P_in	Posterior_in	Posterior_in
		HR2_P_tc	Posterior_tc	Posterior_tc

* The only model version used for inversion to obtain optimized CO sources and sinks with INCA or TransCom OH respectively.

825

Table 2 TCCON measurement sites used for the evaluation.

Station	latitude	longitude	Altitude (m)	Reference
Bialystok, Poland	53.2	23.0	180	(Deutscher <i>et al.</i> , 2014)
Bremen, Germany	53.1	8.9	27	(Notholt <i>et al.</i> , 2014)
Darwin, Australia	-12.4	130.9	30	(Griffith <i>et al.</i> , 2014a)
Eureka, Canada	80.1	-86.4	610	(Strong <i>et al.</i> , 2014)
Garmisch, Germany	47.5	11.1	740	(Sussmann and Rettinger, 2014)
Izana, Tenerife	28.3	-16.5	2370	(Blumenstock <i>et al.</i> , 2014)
JPL, USA	34.1	-118.1	230	(Wennberg <i>et al.</i> , 2014a)
Kalsruhe, Germany	49.1	8.4	116	(Hase <i>et al.</i> , 2014)
Lamont, USA	36.6	-97.5	320	(Wennberg <i>et al.</i> , 2014b)
Lauder, New Zealand	-45.0	169.7	370	(Sherlock <i>et al.</i> , 2014a, 2014b)
Orleans, France	48.0	2.1	130	(Warneke <i>et al.</i> , 2014)
Park Falls, USA	48.5	2.4	60	(Wennberg <i>et al.</i> , 2014c)
Reunion Island	-20.9	55.5	87	(De Mazière <i>et al.</i> , 2014)
Saga, Japan	33.2	130.3	7	(Kawakami <i>et al.</i> , 2014)
Sodankyla, Finland	67.4	26.6	188	(Kivi <i>et al.</i> , 2014)
Tsukuba, Japan	36.1	140.1	30	(Morino <i>et al.</i> , 2014)
Wollongong, Australia	-34.4	150.9	30	(Griffith <i>et al.</i> , 2014b)

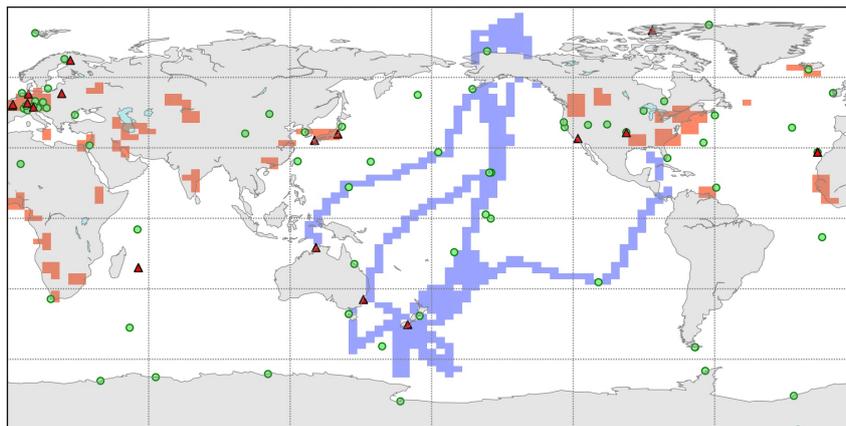


Table 3. Summary of zonal mean biases of the forward and posterior simulations compared to MOPITT X_{CO} and to ground measurements. NH represent 30-90°N, Ntr represent 0-30°N, Str represent 0-30°S, and SH represent 30-90°S. Sub-columns in and tr represent results with the two OH: INCA or TransCom. The MOPITT X_{CO} results here are sampled at corresponding ground stations instead of the entire latitudinal bands.

	MR						HR1						HR2						
	X _{CO} (ppb)			Surface (ppb)			X _{CO} (ppb)			Surface (ppb)			X _{CO} (ppb)			Surface (ppb)			
	F	P	P	F	P	P	F	P	P	F	P	P	F	P	P	F	P	P	
NH	in	-23.6±9.9	-0.7±4.4	-36.0±20.7	27.6±39.2		-20.4±9.0	2.0±4.7	-34.9±30.5	27.6±45.7		-20.9±9.2	1.4±5.1	-33.2±30.2	32.5±47.2				
	tr	-14.9±7.2	-0.8±4.1	-24.9±21.4	18.2±23.8		-11.8±6.9	1.9±4.3	-24.0±30.7	16.6±33.8		-12.1±7.1	1.9±4.5	-22.1±30.5	21.2±34.5				
Ntr	in	-11.7±10.4	-0.6±5.4	-21.8±13.4	-0.3±11.2		-9.3±10.2	1.2±6.3	-14.8±17.2	4.9±15.3		-9.6±10.4	0.6±6.3	-14.7±17.5	5.2±15.8				
	tr	-7.0±8.9	-0.2±5.2	-13.8±11.4	0.3±12.0		-4.2±9.1	2.6±6.0	-6.4±16.2	5.3±16.2		-4.5±9.4	2.4±5.9	-6.3±16.9	5.9±16.8				
Str	in	-7.1±7.7	-0.5±4.5	-7.6±10.1	3.3±15.5		-8.3±9.2	-1.3±6.0	-5.3±21.9	1.8±23.1		-7.5±8.8	-0.7±5.7	-4.8±22.8	2.0±23.1				
	tr	-8.3±9.1	-0.4±4.6	-8.9±9.7	5.5±17.8		-8.0±10.8	0.1±5.8	-4.8±21.9	4.0±24.1		-7.4±10.5	0.9±5.4	-4.5±23.0	4.3±24.1				
SH	in	-5.8±4.0	0.2±2.2	0.5±12.1	9.9±13.8		-5.2±3.9	0.7±2.5	2.5±6.8	11.8±8.4		-5.3±3.9	0.4±2.4	2.4±6.8	11.2±8.4				
	tr	-7.1±4.5	0.3±2.2	-2.1±12.8	9.4±15.4		-4.9±4.3	2.2±2.4	2.6±7.1	14.3±9.3		-5.2±4.4	2.0±2.4	2.1±7.1	13.8±9.4				



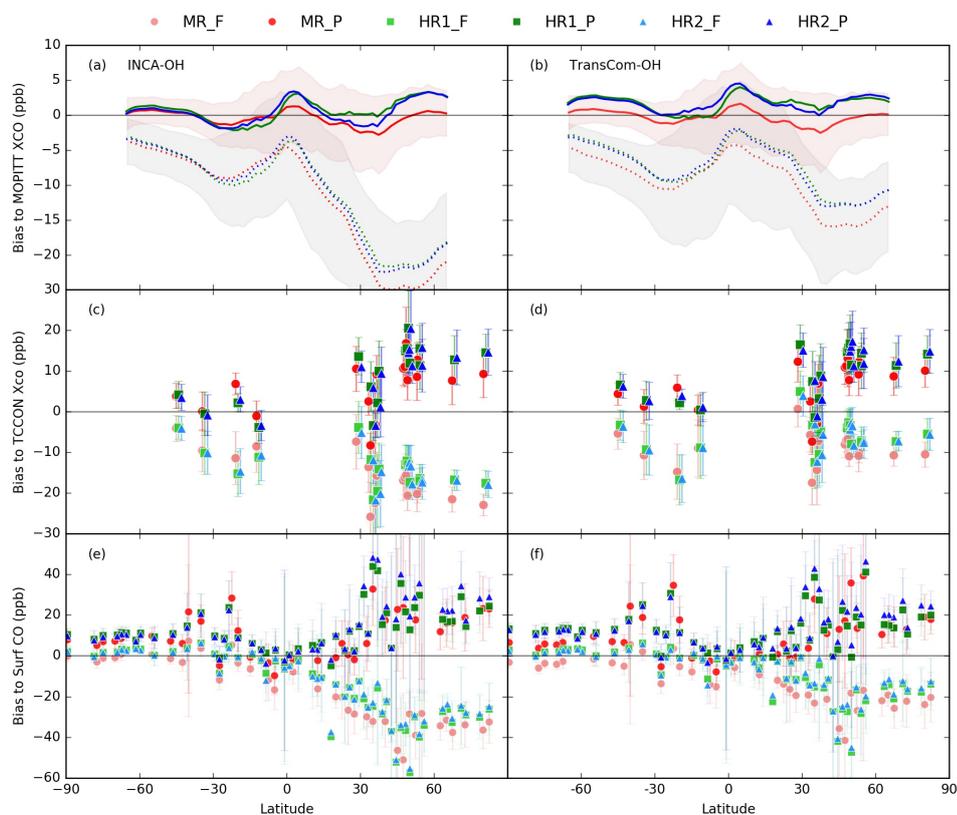
- 835 **Figure 1. Spatial distribution of the independent measurements.** The TCCON stations are shown by red triangles, the surface stations are shown by green circles, the MOZAIC/IAGOS and HIPPO profile projections on horizontal map are shown at the model resolution in orange and blue respectively.



840



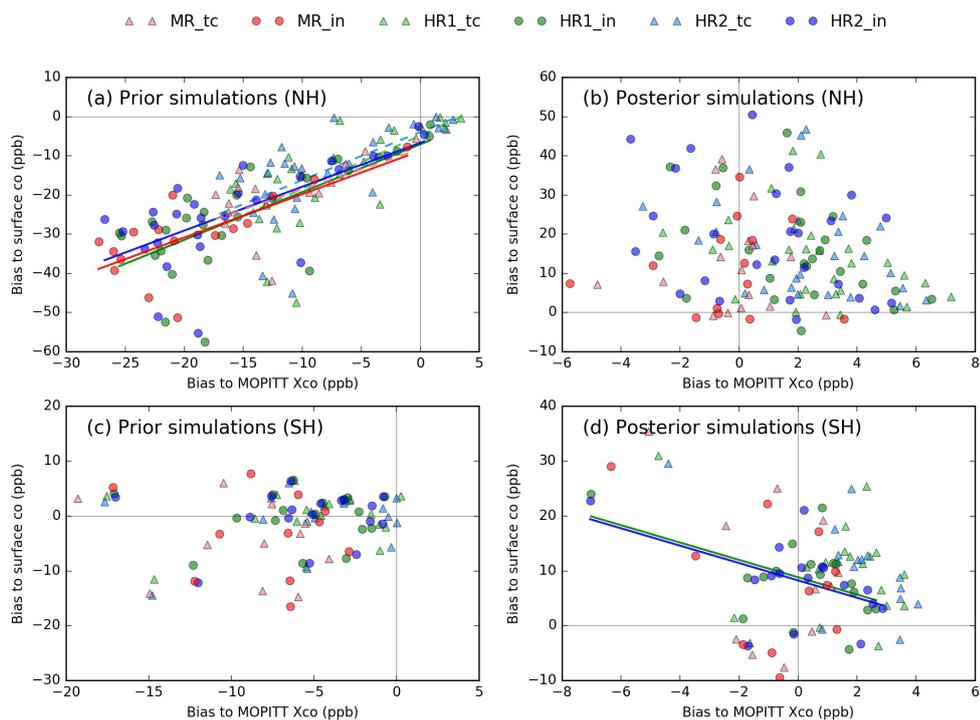
Figure 2. Model-data biases along the latitudes. (a) and (b) present model biases to MOPITT Xco using INCA-OH and TransCom-OH, respectively. Dashed and solid lines represent forward and posterior model results, respectively, while color codes correspond to model versions shown in the legend. (c) and (d) present model biases to TCCON Xco measurements using INCA-OH and TransCom-OH respectively. (e) and (f) show model biases compared to ground measurements from surface network using INCA-OH and TransCom-OH respectively. Note the ordinates are different for the three observation types.



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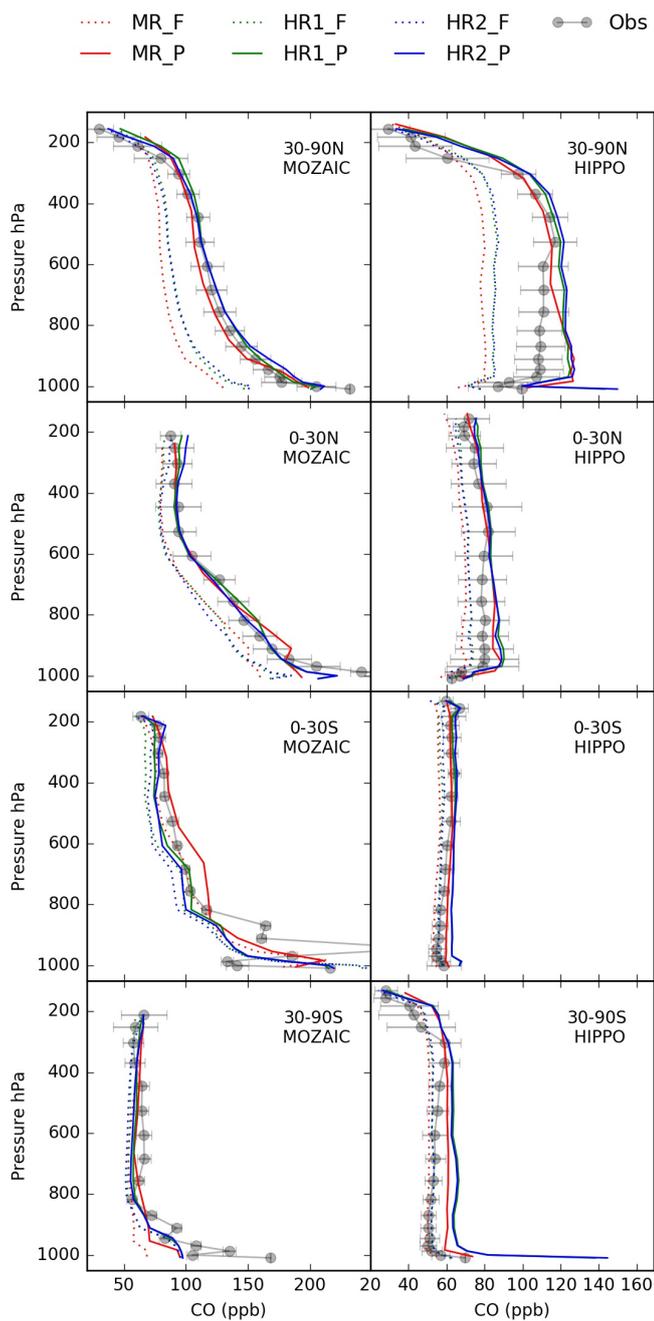
855 **Figure 3. Covariation of model-data differences in the surface [CO] and in the MOPITT X_{CO} .** (a) and (b) show prior and posterior model results in the NH, (c) and (d) show prior and posterior model results in the SH.



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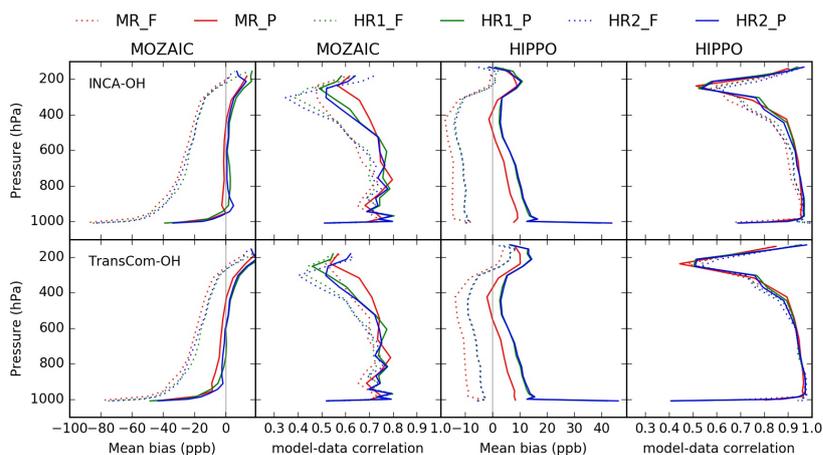


Figure 4. Zonal averages of the vertical CO profiles. The left column shows results against MOZAIC measurements, and the right column shows results against HIPPO measurements. The error bars indicate the standard variation of the measurements. Results associated with INCA-OH are presented.



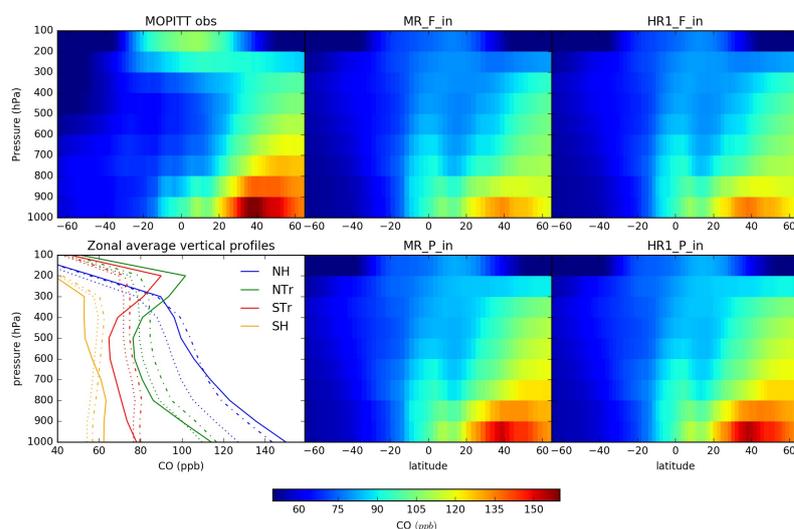


870 **Figure 5. Globally-average model-data biases of vertical profiles and vertical distribution of model-data correlation coefficients compared to MOZAIC and HIPPO measurements respectively.** The upper row shows results with INCA-OH, and the lower row shows results with TransCom-OH.





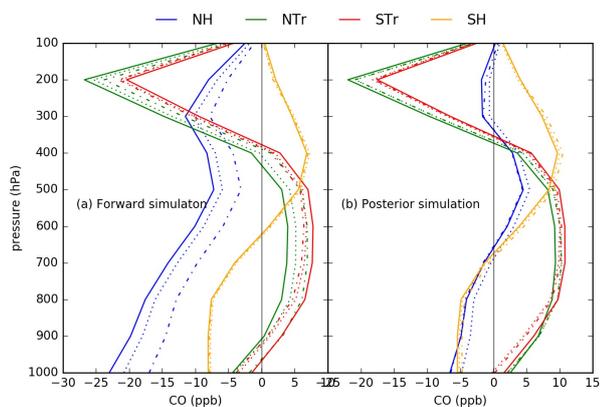
875 **Figure 6. Zonal averages of the vertical CO profiles at the MOPITT retrieval pressure**
levels. The upper panel successively shows the latitudinal average of the vertical profiles of (a)
the MOPITTv6 retrievals and prior forward model retrievals using (b) 19-level (MR_F_in) and (c)
39-level versions (HR1_F_in) with INCA-OH. The corresponding posterior model results are
shown below for visual comparison (e) MR_P_in and (f) HR1_P_in. The lower left subplot
880 shows (d) the zonal average of the vertical profiles of four bands, NH - North of 30°N, NTr - 0°-
30°N, STR - 0°-30°S, and SH - South of 30°S, where the solid lines represent the MOPITT profile
retrievals, the dotted-lines represent the forward simulation, and the dot-dash lines represent
the posterior model using INCA-OH. The results with TransCom-OH are of similar pattern, thus not
shown here.



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890 **Figure 7. Zonally-averaged model differences to MOPITT profile retrievals in (a) prior forward simulations and (b) posterior simulations.** Color codes for sub-regions are the same as Figure 6d. Solid lines represent the reference model with INCA-OH (MR_in), dotted lines represent the HR1 model with INCA-OH, dash-dotted lines show the HR1 model with TransCom-OH. The differences between HR1 and HR2 given the same OH are not significant, and thus not shown.

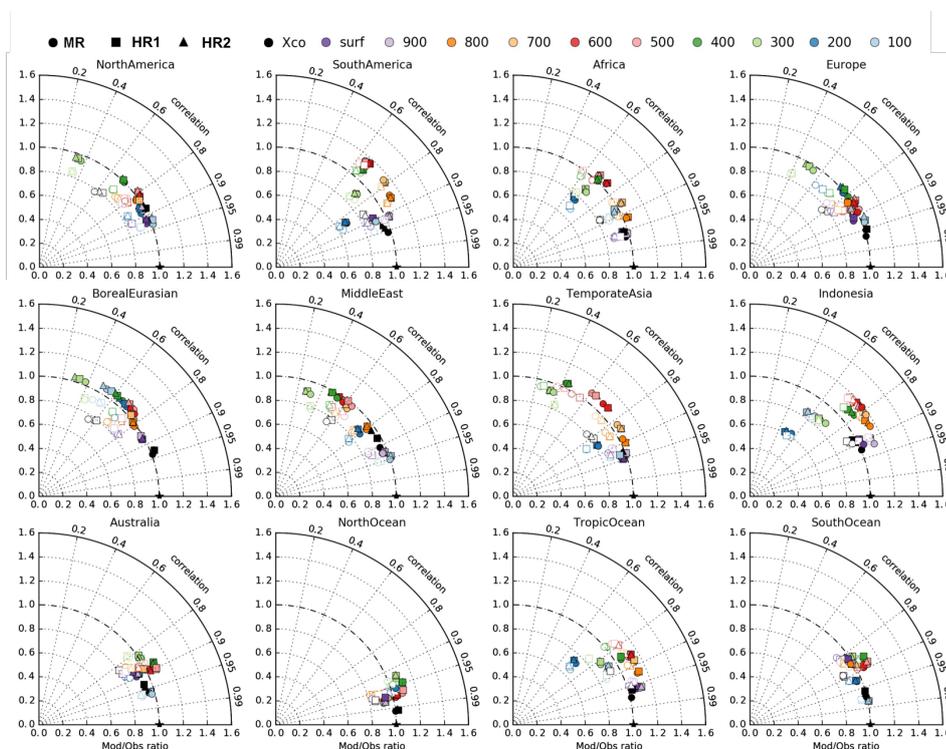


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900 **Figure 8. Pseudo Taylor diagrams of model agreement with MOPITT profiles for each region.** In each subplot, the radial distance indicates the relative bias (not standard deviation), with 1.0 being equal to the observation as shown by the dashed lines; the azimuthal angle is proportional to the model-data correlation coefficient, with higher correlations close to the horizon; the distance from to the reference data does not have an explicit statistical meaning, but qualitatively speaking the closer to the reference, the better it fits the observations. The hollow and filled markers represent forward and posterior simulations respectively, the color codes represent pressure level, the marker shapes represent model version.

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