

We would like to thank referee #1 for very helpful comments and insightful suggestions.

Reply to comments by Anonymous Referee #1

General Comments:

In this study, parameterizations for sea salt aerosol (SSA) emissions from blowing snow and frost flowers are implemented in the GEOS-Chem global chemical transport model. Model to measurement comparisons for SSA mass concentrations are presented for three Arctic and two Antarctic sites, as well as for an Arctic cruise. The authors conclude that blowing snow is a dominant SSA source during winter in the polar regions, with a smaller contribution from frost flowers. The paper is well written and addresses scientifically important questions regarding the sources for SSA in the polar regions. The related parameterizations are challenging to develop because there are several uncertainties involved. The impact of these uncertainties on the conclusions could be discussed and examined more explicitly as outlined in the following comments. The manuscript should be suitable for publication if the following concerns can be satisfactorily addressed.

Specific Comments

1) P1, L12: The model with open ocean emissions alone underestimates the SSA mass concentrations by factors of 2-10. What is the uncertainty in the measurement SSA mass concentrations? Is it possible that the model and measurements could agree within the measurement uncertainty?

- The reported uncertainties in the measurements of Na^+ mass concentrations are 11% (relative uncertainty, $0.01 \mu\text{g}/\text{m}^3$) at Barrow (Quinn et al., 2000), 5-10% ($0.01 \mu\text{g}/\text{m}^3$) at Alert and Zeppelin (WMO/GAW, 2003), 5-11% at Neumayer (Weller et al. 2008), and 5% at Dumont d'Urville (Legrand et al. 2012). This should have been included in the original manuscript, thank you for pointing out this omission. As observed wintertime Na^+ concentrations are $\sim 0.3\text{-}1 \mu\text{g}/\text{m}^3$, factors of 2-10 disagreement are well outside the observational uncertainties. We have added a discussion of measurement uncertainties in section 2.3:

“The Na^+ mass concentrations are determined by ion chromatography with uncertainties of 5%-11% ($0.01 \mu\text{g}/\text{m}^3$ in absolute uncertainty).”

2) P2, L17-19: Please consider adding a sentence here to describe why the first two mechanisms lead to depletion of sulfate relative to sodium through the precipitation of mirabilite.

-We have clarified this in the revised manuscript:

“The first two of these mechanisms lead to depletion of the sulfate to sodium ratio relative to bulk sea water as mirabilite ($\text{Na}_2\text{SO}_4 \cdot 10\text{H}_2\text{O}$) precipitates from brine at temperatures below -8°C during sea ice and frost flower formation (Alvarez-Aviles et al., 2008).”

3) P4, L9: These lines note that the simulation with open water emissions alone of SSA overestimates the summertime SSA. Is it possible that the model could have errors in the summertime removal of SSA in precipitation in the polar regions (particularly if the model neglects aerosol removal by summertime drizzle). If the summertime removal is too inefficient, are you able to justify this suppression of the emissions? Are there any previous studies that have examined emissions for temperatures below 5 C? Fig. S1 was instructive since this seems to indicate a low sensitivity to this assumption about the emissions for the polar winter. Is there is a concern that this assumption might introduce a low bias to the open ocean emissions in an effort to compensate for errors in the removal scheme?

- We agree with the referee that an alternative explanation for the model overestimate of summertime SSA could be related to an underestimate in the removal of SSA during summer. This point has been noted in the revised manuscript. There is support for our cold temperature suppression assumption from the laboratory study by Mårtensson et al. (2003). They examined marine aerosol production from seawater at -2°C and 5°C, finding a 50-60% decrease in the number of particles produced for $r > 0.1 \mu\text{m}$ when temperatures decrease from 5°C to -2°C. This is consistent with our assumption of suppression at very cold temperatures.

As noted by the referee, this modification to open ocean emissions does not affect our wintertime results, as the five polar sites are distant from open ocean, with very small amounts of open ocean SSA transported to those sites. For example, at Barrow and Alert wintertime SSA concentrations due to the open ocean are $< 0.1 \mu\text{g}/\text{m}^3$, more than an order of magnitude lower than observations (1-3 $\mu\text{g}/\text{m}^3$).

We have added the following:

“This is consistent with the laboratory study of Mårtensson et al. (2003), who report a 50-60% decrease in aerosol production (for $r > 0.1 \mu\text{m}$) when seawater temperature decreased from 5°C to -2°C. We note that another potential explanation for the summertime overestimate in SSA mass concentrations is inefficient wet removal from low-intensity summer precipitation in GEOS-Chem (Croft et al., 2016).”

4) P4, L14-15: Are there any uncertainties related to the dry deposition parameterization? How might this affect your analysis, particularly in winter? As well, are there any uncertainties related to the wintertime removal by precipitation from mixed-phase and ice clouds and how might this affect your analysis?

- The dry deposition velocity for aerosols over snow and ice is assumed to be 0.03 cm/s in the model, which is based on the estimates from measurements in the Arctic. As discussed in the manuscript (p10, L2), for cold clouds (mixed-phase and ice clouds with temperature $< 258 \text{ K}$) SSA is only removed by below-cloud precipitation and we assume no in-cloud scavenging in the model. If in-cloud scavenging were to be an efficient process for SSA, it would lead to decreases in our calculated SSA mass concentrations and require a larger source from sea ice to match observed concentrations. Previous studies using the GEOS-Chem model found reasonable agreement with ground-based and aircraft observations of sulfate, ammonium, black carbon, and organic aerosols

during winter and spring (Fisher et al., 2011; Wang et al., 2011). This suggests that the removal processes during winter and spring are reasonably represented in the model. This is now noted in the revised manuscript.

“Detailed comparisons of GEOS-Chem black carbon and organic aerosol (Wang et al., 2011) as well as sulfate and ammonium aerosol (Fisher et al., 2011) to ground-based and aircraft observations over the Arctic during winter and spring suggest that transport and removal processes are reasonable captured by the model. “

5) P4, L18-19: Please consider clarifying here how the sodium bins are related to the SSA bins in your parameterization. Are these sodium bins additional tracers in the model? As well, please specify where you mean by ‘see below’.

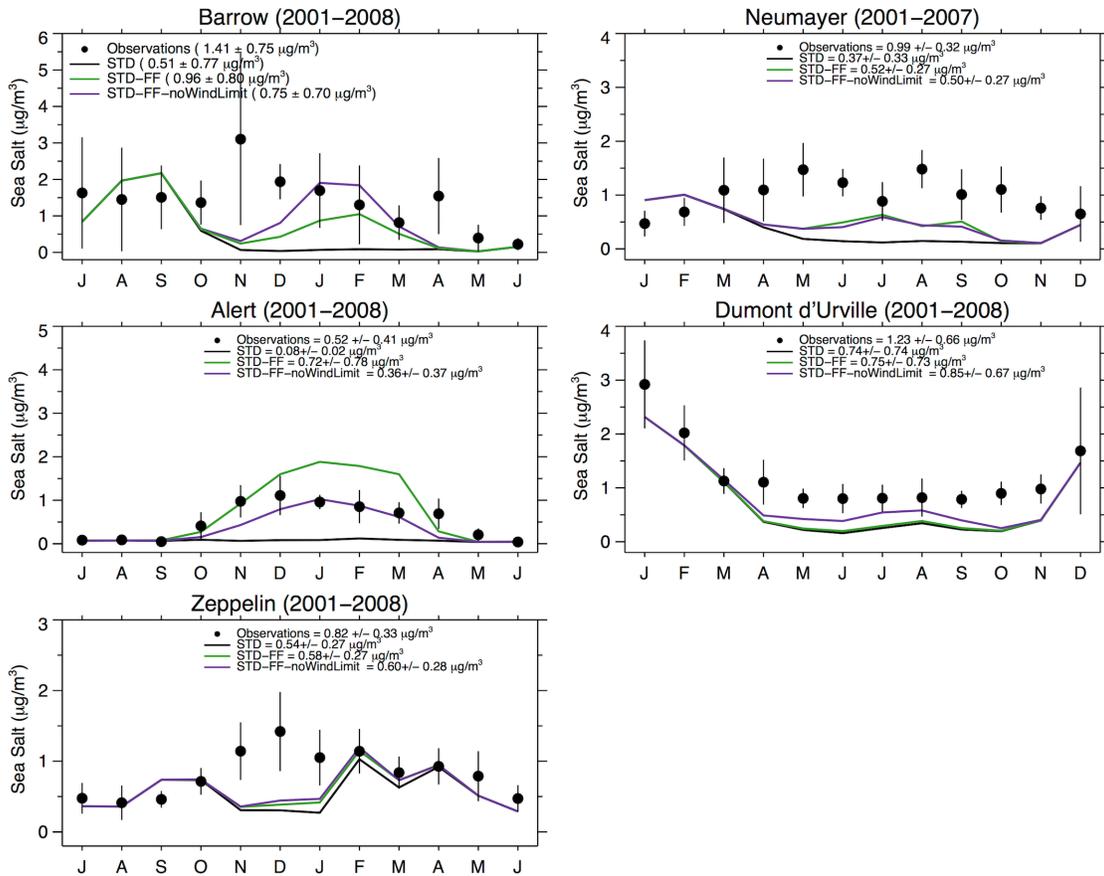
- We agree that this was unclear and have clarified this in the revised manuscript:

“For this work, we track SSA mass in two size bins: accumulation mode ($r_{\text{dry}} = 0.01\text{--}0.5 \mu\text{m}$) and coarse mode ($r_{\text{dry}} = 0.5\text{--}4 \mu\text{m}$), except in the comparison to in situ mass concentrations of SSA for which we use $r_{\text{dry}} = 0.01\text{--}0.3 \mu\text{m}$ and $r_{\text{dry}} = 0.3\text{--}3 \mu\text{m}$ (see section 2.3). “

6) P5, L26: In the frost flower parameterization, SSA emissions are only allowed if the wind speed is $< 5 \text{ m s}^{-1}$. However p. 2, lines 31-32 suggest that higher wind speeds are needed in order to break the frost flowers. There appears to be two processes here – the frost flowers form under low wind speeds, but do the emissions actually occur at greater wind speeds? Could this assumption that the frost flowers only emit SSA at low wind speed cause a low bias in the emissions from frost flowers? How do the emissions change if the frost flowers are allowed to emit SSA at larger wind speeds than 5 m s^{-1} , and how does this affect your conclusions?

- We recognize that there is likely to be a distinction between frost flower formation under low windspeed and lofting at high windspeed. Given the uncertainty and contradictory results in the literature regarding frost flower formation and lofting (see our section 1), we have conducted a frost flower simulation removing the wind speed limit and reverting to the formulation proposed by Xu et al. (2013). This also enables us to address this referee’s comment #15 about the more recent study of Xu et al. (2016). We compare in the figure below our original frost flower simulation (STD-FF green line) to the one with no wind speed limit (STD-FF-noWindLimit purple line). As we use the same scaling factor as Xu et al. (2013), the new emissions from frost flowers are slightly lower than in our STD-FF simulation. We find good agreement at Alert for Dec-Mar, and at Barrow in January-February, but for all the other stations our results remain unchanged, with very small influence of frost flowers at Zeppelin, Neumayer and Dumont d’Urville, and therefore we find that frost flower emissions are insufficient to explain the wintertime SSA enhancements in the observations at these sites. Thus overall our conclusions remain unchanged.

For simplicity, we replace our original FF simulation with this new updated one, without any windspeed limit. The figures and text have been changed to reflect this.



7) P6, L12: A factor of 3.256 is used to convert all observed Na^+ mass concentrations to SSA for comparison with the simulations. Is there any uncertainty in this factor that might make an apples-to-apples comparison between the measurements and model difficult?

- This mass ratio for Na^+ to SSA (3.256) is based on the standard seawater mass ratio at a salinity of 35 psu. There can be some small changes to the ionic composition of SSA from frost flowers and snow, but these are small (<0.5%). Based on the chemical composition of frost flowers reported in Rankin et al. (2000), this ratio is 3.237 for frost flowers. Krnavek et al. (2012) find that this mass ratio is 3.278 for snow on thick first year sea ice and 3.24 for those on thin ones. Therefore, we think it is reasonable to use a constant factor (≈ 3.256) to convert observed Na^+ to SSA for comparison with the model SSA.

We have added this to the revised manuscript:

“For comparison between the GEOS-Chem model and the observations, we convert observed Na^+ mass concentrations to SSA mass concentrations using a factor of 3.256

based on the mass ratio of Na^+ in seawater (Riley and Chester, 1971). For frost flowers, this ratio is 3.237 (Rankin et al., 2000). Krnavek et al. (2012) find a ratio of 3.24-3.278 for snow on first year sea ice. As this ratio varies by less than 0.5% for these different SSA sources, we use a constant factor of 3.256.”

8) P6, L12: ‘use two size bins in the model’ – are these bins for sea salt aerosol or sodium? Why are these bin limits different than in Table 1 and also different than P6, L21? As well, how are the observed Na^+ concentrations in the submicron and supermicron ranges apportioned between the two bins used by the model for the purpose of the model-measurement comparison? What are the size limits for the observed submicron and supermicron aerosol?

- Yes, this indeed requires some clarification. These bins are dry sea salt aerosol radii. For simplicity, we now use one single set of size bins ($r_{\text{dry}}=0.01-0.3 \mu\text{m}$ and $0.3-3 \mu\text{m}$) for comparison to the ground sites and ICEALOT cruises and justify this in the revised text. Most previously published studies use ($0.01-0.5 \mu\text{m}$; $0.5-4 \mu\text{m}$), so this is why we use these bins in Table 1.

“The reported aerodynamic cut-off diameters of the measurements are $1 \mu\text{m}$ and $10 \mu\text{m}$ at Barrow and during ICEALOT (Quinn et al., 2002), and $7-10 \mu\text{m}$ at the other sites (Wagenbach et al., 1998; WMO/GAW, 2003; Weller et al., 2008). In order to compare to model simulations, we need to convert these aerodynamic diameters to dry geometric radii. This conversion depends on aerosol density, relative humidity during sampling, and whether the particle is spherical (Seinfeld and Pandis, 2006). For example a $10 \mu\text{m}$ aerodynamic diameter could correspond to a dry geometric radius of $2.3 \mu\text{m}$ (80% RH, 1.2 g/cm^3 pure NaCl solution, factor of 2 growth between dry and 80% RH), $3 \mu\text{m}$ (dry cubical NaCl particle, 2.2 g/cm^3 , Lewis and Schwartz, 2004), $3.8 \mu\text{m}$ (30% RH, ammonium sulfate and sea salt aerosol, 1.7 g/cm^3 , Quinn et al., 1996). Thus for comparison to observations we conduct a simulation with two size bins: $r_{\text{dry}} = 0.01-0.3 \mu\text{m}$ and $r_{\text{dry}} = 0.3-3 \mu\text{m}$.”

9) P6, L23: Figure 1 shows a maximum in the measured SSA mass concentrations in the Arctic in November-December, whereas the simulations have a maximum Jan-Feb. What factors contribute to this model-measurement discrepancy?

- Given the simplified nature of the blowing snow simulation, a number of potential factors could explain this offset, such as seasonal and spatial variability in snow mobility and/or salinity. Exploring these factors is beyond the scope of our manuscript.

10) Figure 1: The blowing snow simulation alone does appear to improve agreement with observations. However, in reality, both blowing snow and frost flowers might be expected to contribute together to the SSA concentrations. Have you conducted simulations with both of these sources implemented at the same time? Figure 1 seems to suggest that if the model included both sources simultaneously, then the SSA would agree more closely with the observations in November/December in the Arctic and in April/May at Neumayer, but the model would strongly over predicts the observations in

subsequent winter months. Please consider adding some related discussion. Does this suggest that the simulation might over predict the blowing snow emissions? As well, Fig. 3 seems to indicate that the model over predicts the SSA during the blowing snow event near 17 April.

- It is indeed possible that both blowing snow and frost flowers contribute together to SSA concentrations. We have included a new figure in the supplemental material (Fig S5) where we compare a simulation with all 3 sources (open ocean, frost flowers and blowing snow) to observations. The resulting simulation overestimates observations, especially at Barrow and Alert, where modeled SSA is a factor of 2-3 too high in January-March. At the other 3 sites the influence of frost flowers is small, so the differences are not as large.

We have added this following discussion in P9, L1-3.

“It is possible that both blowing snow and frost flower emissions act together. However, when we add the contributions from both sources, we find that modelled SSA mass concentrations are a factor of 2-3 too high compared to observations at Barrow and Alert (Fig. S5). In particular, the frost flower simulation leads to a peak in SSA in February at Barrow, which is not observed.”

11) P6, L24: As noted here, the advance and retreat of sea ice affects SSA. How well does the model simulation of sea ice agree with observations?

- As described in Section 2.1, the sea ice in MERRA is set as a boundary condition from the NOAA Optimal Interpolation (OI) v2 gridded dataset (Reynolds et al., 2002). This dataset is based on SSM/I satellite observations, so should have a reasonable representation of sea ice cover. This has been clarified in section 2.1:

“The sea ice concentration boundary conditions in MERRA are derived from the weekly product of Reynolds et al. (2002), which is based on Special Sensor Microwave Imager (SSM/I) instruments on Defense Meteorological Satellite Program (DMSP) satellites.”

12) P7, L8-11: Why was the normalized mean bias chosen as the metric for the model evaluation as opposed to mean fractional bias? The latter metric has the advantages of not allowing a few data points to dominate the metric and allows for some error in measurements (Boylan and Russell 2006).

-We use the normalized mean bias ($NMB = (\overline{Model} / \overline{Obs} - 1) \times 100$) here as we assume that the observations are the absolute truth for evaluating the performance of different model simulations. We are using multi-year (7-8 years) observations, which helps decrease the influence of a few daily/weekly points.

13) P9, L1-2: Blowing snow and frost flowers are noted to have a larger impact on the 0.01-0.5 μm SSA mass concentrations than the open ocean source. What is the physical

mechanism for relatively greater emission into this size range? For the case of blowing snow emissions, are there any uncertainties related to how the emitted mass is distributed between the two SSA bins using the assumed size distribution? Has the model been used for sensitivity studies to examine the potential impact of related uncertainties?

- The physical mechanisms for SSA emissions from open ocean emissions (breaking waves and bubble bursting) are different than emissions from blowing snow and frost flowers (saltation of fallen snow and suspension) or frost flowers (saltation of broken frost flower crystals and suspension). The original crystalline form of snow particles/frost flower fragments are expected to be shattered by repeated impacts with the ground and other particles during saltation. For blowing snow, the size of the dry SSA particle depends on the original snow particle size, snow salinity, and number of SSA particles produced per snow particle (Yang et al., 2008). The blowing snow particle size distribution is constrained by observations of blowing snow events (Mann et al., 2000). As discussed in the text, there are few observations of salinity of snow on sea-ice so we used uniform values of salinity for the entire Arctic and Antarctic snow salinity. This is certainly something that can be improved as more measurements become available. Assumptions of the number SSA particles produced per snow particle are highly uncertain, and we include a sensitivity study in the supplement (Fig. S2).

We have added the following in the revised manuscript:

“This difference in size distributions is related to the different physical mechanisms for SSA emissions from open ocean emissions (breaking waves and bubble bursting) compared to blowing snow (saltation of fallen snow and suspension) or frost flowers (saltation of broken frost flower crystals and suspension). The original crystalline form of snow particles/frost flower fragments are expected to be shattered by repeated impacts with the ground and other particles during saltation. Sublimation of ice from these particles leads to relatively small SSA compared to bubble bursting in the open ocean.”

14) P9, L24: ‘they are not efficient ice nuclei’ – if possible, please add a reference to observations that support this statement. This seems in contradiction to some studies (e.g. DeMott et al. 2016). Or do you mean to indicate that SSA is not an efficient ice nucleus in the model?

- Excellent point. This statement is indeed based on what is assumed in the GEOS-Chem model for scavenging. We clarify this in the text and add references to DeMott et al. (2016) and Wise et al. (2012), and note that in-cloud scavenging is not currently included in GEOS-Chem.

“Recent laboratory studies have shown that SSA could act as ice nuclei by deposition freezing (Wise et al., 2012) and immersion freezing (DeMott et al., 2016), and might thus undergo in-cloud scavenging in mixed and ice clouds. This process is not currently included in GEOS-Chem.”

15) P10, L20: How do these frost flower simulations compare to the recent work of Xu et al. (2016)?

- Thank you for bring this work to our attention. Both the magnitude of emissions and spatial distribution of our frost flower simulation are similar to the results Xu et al. (2016) for the Arctic. This has been added in the revised manuscript:

“Our emissions from frost flowers over the Arctic (0.21 Tg/yr) are consistent the accumulation model emissions reported by Xu et al. (2016) (0.24 Tg for November-February, their Table 2). We also find a similar geographic distribution.”

16) P10, L28-30: As noted here, there are substantial uncertainties associated with these parameterizations. Since this is a model-based study, please consider whether the presentation of a few sensitivity study results related to the key uncertainties in the emissions parameterizations (salinity, size distribution for blowing snow and wind conditions for frost flower emissions) might be of help to the reader in interpreting the statement in the abstract that ‘blowing snow is likely to be the dominant SSA source during the winter’. As well, this could help in interpreting the presented values for the SSA emissions from blowing snow and frost flowers since there seems to be some evidence that the blowing snow parameterization over predicts the observations.

- Yes, we agree with the referee that there are substantial uncertainties in the parameterizations and have clarified the text, as well as added sensitivity simulations (wind threshold for frost flowers) in addition to the sensitivity simulation for the blowing snow particle number. We now note in the revised text:

“One key uncertainty in our simulations is snow salinity. Indeed, SSA emissions from blowing snow have a near-linear dependence on the salinity of snow. Thus a doubling of the assumed salinity would lead to a doubling in SSA emissions from blowing snow. Furthermore, we assume a uniform salinity of snow over both first-year and multi-year sea ice. This likely overestimates the contribution of blowing snow SSA over the western Arctic, which is dominated by multi-year sea ice. More extensive observations of surface snow salinity at multiple locations over both first-year and multi-year sea ice can help further refine these assumptions. Sampling of SSA size distributions during blowing snow events can help determine the number of particles per snowflake, which we determined empirically in this study. This number will not affect to total SSA emissions, but will change the relative importance of submicron and supermicron SSA emissions.”

We have added a sensitivity simulation with/without wind conditions of frost flower emission is detailed in the supplementary material.

The overestimate of SSA in spring by the blowing model may be due to the seasonal variability of the surface snow salinity which is not taken into account in this study. This should be further constrained with observations of surface snow salinity at different location and times which are limited at this moment.

17) Are you able to provide any recommendations to modelers about the implementation of blowing snow and frost flower parameterizations into global models?

- Our implementation of blowing snow follows Yang et al. (2008) with modification of salinity, age of snow and number of particles. And the implementation of frost flower emission follows the emission scheme in Xu et al. (2013). Details on implementation of these sources in the model can be found in section 2.2, Yang et al. (2008), Xu et al. (2013) and references therein.

18) P11, L9-10: For the potential impacts of wintertime SSA, would you expect any effect on mixed-phase and ice clouds?

-Yes, this is indeed a very good point. We have added the following statement:

“Improved process-based understanding of these emissions would also lead to better constraints on the potential climatic impact of wintertime SSA on clouds, in particular mixed-phase and ice clouds, which have a strong influence on downward longwave radiative forcing. Indeed, recent studies have shown the role of SSA as ice nuclei (Wise et al., 2012; DeMott et al., 2016). Thus over the Arctic and Antarctic regions, where the abundance of other ice nuclei such as dust or black carbon are low, SSA from sea ice sources could influence the formation, radiative forcing, and precipitation of mixed-phase and ice clouds.”

Reference

Alvarez-Aviles, L., Simpson, W. R., Douglas, T. A., Sturm, M., Perovich, D., and Domine, F.: Frost flower chemical composition during growth and its implications for aerosol production and bromine activation, *J. Geophys. Res.*, 113, D21304, doi:10.1029/2008JD010277, 2008.

Croft, B., Martin, R. V., Leitch, W. R., Tunved, P., Breider, T. J., D'Andrea, S. D., and Pierce, J. R.: Processes controlling the annual cycle of Arctic aerosol number and size distributions, *Atmos. Chem. Phys.*, 16, 3665-3682, doi:10.5194/acp-16-3665-2016, 2016.

DeMott, P. J., Hill, T. C. J., McCluskey, C. S., Prather, K. A., Collins, D. B., Sullivan, R. C., Ruppel, M. J., Mason, R. H., Irish, V. E., Lee, T., Hwang, C. Y., Rhee, T. S., Snider, J. R., McMeeking, G. R., Dhaniyala, S., Lewis, E. R., Wentzell, J. J. B., Abbatt, J., Lee, C., Sultana, C. M., Ault, A. P., Axson, J. L., Martinez, M. D., Venero, I., Santos-Figueroa, G., Stokes, M. D., Deane, G. B., Mayol-Bracero, O. L., Grassian, V. H., Bertram, T. H., Bertram, A. K., Moffett, B. F., and Franc, G. D.: Sea spray aerosol as a unique source of ice nucleating particles, *Proc. Natl. Acad. Sci.*, 113, 5797–5803, doi:10.1073/pnas.1514034112, 2016.

Fisher, J. A., et al.: Sources, distribution, and acidity of sulfate-ammonium aerosol in the Arctic in winter-spring, *Atmos. Environ.*, 45, 7301-7318, doi :10.1016/j.atmosenv.2011.08.030, 2011.

Krnavek, L., Simpson, W. R., Carlson, D., Domine, F., Douglas, T. A., and Sturm, M.: The chemical composition of surface snow in the Arctic: Examining marine, terrestrial, and atmospheric influences, *Atmos. Environ.*, 50, 349–359, 2012.

Legrand, M., Gros, V., Preunkert, S., Sarda-Estève, R., Thierry, A.-M., Pépy, G., and Jourdain, B.: A reassessment of the budget of formic and acetic acids in the boundary layer at Dumont d'Urville (coastal Antarctica): The role of penguin emissions on the budget of several oxygenated volatile organic compounds, *J. Geophys. Res.*, 117, D06308, doi:10.1029/2011JD017102, 2012.

Lewis, E. R. and Schwartz, S. E.: Sea Salt Aerosol Production: Mechanisms, Methods, Measurements, and Models: A Critical Review, American Geophysical Union, Washington, D.C., 2004.

- Mann, G. W., Anderson, P. S., and Mobbs, S. D.: Profile measurements of blowing snow at Halley, Antarctica, *J. Geophys. Res.*, 105(D19), 24491–24508, 2000.
- Mårtensson, E., Nilsson, E., Leeuw, G., Cohen, L., and Hansson, H.: Laboratory simulations and parameterization of the primary marine aerosol production, *J. Geophys. Res.* 108(D9), 4297, doi:10.1029/2002JD002263, 2003.
- Quinn, P. K., Kapustin, V. N., Bates, T. S., and Covert, D.S.: Chemical and optical properties of marine boundary layer aerosol particles of the mid-Pacific in relation to sources and meteorological transport, *J. Geophys. Res.*, 101, 6931-6952, doi: 10.1029/95JD03444, 1996.
- Quinn, P. K., Bates, T. S., Miller, T. L., Coffman, D. J., Johnson, J. E., Harris, J. M., Ogren, J. A., Forbes, G., Anderson, T. L., Covert, D. S., and Rood, M. J.: Surface submicron aerosol chemical composition: What fraction is not sulfate?, *J. Geophys. Res.*, 105, 6785–6806, 2000.
- Rankin, A. M., Auld, V., and Wolff, E. W.: Frost flowers as a source of fractionated sea salt aerosol in the polar regions, *Geophys. Res. Lett.*, 27, 3469–3472, doi:10.1029/2000GL011771, 2000.
- Reynolds, R. W., Rayner, N. A., Smith, T. M., Stokes, D. C., and Wang, W.: An improved in situ and satellite SST analysis for climate, *J. Clim.*, 15, 1609–1625, 2002.
- Seinfeld, J. H. and Pandis, S. N.: *Atmospheric Chemistry and Physics*, Second Edition ed., John Wiley & Sons, Hoboken, 2006.
- Wagenbach, D., Ducroz, F., Mulvaney, R., Keck, L., Minikin, A., Legrand, M., Hall, J. S., and Wolff, E. W.: Sea-salt aerosol in coastal Antarctic regions, *J. Geophys. Res.*, 103, 10 961–10 974, 1998.
- Wang, Q., Jacob, D. J., Fisher, J. A., Mao, J., Leibensperger, E. M., Carouge, C. C., Le Sager, P., Kondo, Y., Jimenez, J. L., Cubison, M. J., and Doherty, S. J.: Sources of carbonaceous aerosols and deposited black carbon in the Arctic in winter-spring: implications for radiative forcing, *Atmos. Chem. Phys.*, 11, 12453-12473, doi:10.5194/acp-11-12453-2011, 2011.
- Wise, M. E., Baustian, K. J., Koop, T., Freedman, M. A., Jensen, E. J., and Tolbert, M. A.: Depositional ice nucleation onto crystalline hydrated NaCl particles: a new mechanism for ice formation in the troposphere, *Atmos. Chem. Phys.*, 12, 1121-1134, doi:10.5194/acp-12-1121-2012, 2012.
- WMO/GAW: *WMO/GAW Aerosol Measurement Procedures: Guidelines and Recommendations*. World Meteorological Organization, 2003.
- Weller, R., Woltjen, J., Piel, C., Resenberg, R., Wagenbach, D., König-Langlo, G., and Kriews, M.: Seasonal variability of crustal and marine trace elements in the aerosol at Neumayer station, Antarctica, *Tellus*, 60, 742–752, doi:10.1111/j.1600-0889.2008.00372.x, 2008.
- Xu, L., Russell, L. M., Somerville, R. C. J., and Quinn, P. K.: Frost flower aerosol effects on Arctic wintertime longwave cloud radiative forcing, *J. Geophys. Res.-Atmos.*, 118, 13282–13291, doi:10.1002/2013JD020554, 2013.
- Xu, L., Russell, L. M., and Burrows, S. M.: Potential sea salt aerosol sources from frost flowers in the pan-Arctic region, *J. Geophys. Res. Atmos.*, 121, 10,840–10,856, doi:10.1002/2015JD024713, 2016.
- Yang, X., Pyle, J. A., and Cox, R. A.: Sea salt aerosol production and bromine release: Role of snow on sea ice, *Geophys. Res. Lett.*, 35, L16815, doi:10.1029/2008GL034536, 2008.