



# How much information do extinction and backscattering measurements contain about the chemical composition of atmospheric aerosol?

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## 1 Abstract.

2 We theoretically and numerically investigate the problem of assimilating lidar observations of extinction and backscattering  
3 coefficients of aerosols into a chemical transport model. More specifically, we consider the inverse problem of determining  
4 the chemical composition of aerosols from these observations. The main questions are how much information the observations  
5 contain to constrain the particles' chemical composition, and how one can optimise a chemical data assimilation system to make  
6 maximum use of the available information. We first quantify the information content of the measurements by computing the  
7 singular values of the observation operator. From the singular values we can compute the number of signal degrees of freedom  
8 and the reduction in Shannon entropy. For an observation standard deviation of 10 %, it is found that simultaneous measure-  
9 ments of extinction and backscattering allows us to constrain twice as many model variables as extinction measurements alone.  
10 The same holds for measurements at two wavelengths compared to measurements at a single wavelength. However, when we  
11 extend the set of measurements from two to three wavelengths then we observe only a small increase in the number of signal  
12 degrees of freedom, and a minor change in the Shannon entropy. The information content is strongly sensitive to the observation  
13 error; both the number of signal degrees of freedom and the reduction in Shannon entropy steeply decrease as the observation  
14 standard deviation increases in the range between 1 and 100 %. The right singular vectors of the observation operator can  
15 be employed to transform the model variables into a new basis in which the components of the state vector can be divided  
16 into signal-related and noise-related components. We incorporate these results in a chemical data assimilation algorithm by  
17 introducing weak constraints that restrict the assimilation algorithm to acting on the signal-related model variables only. This  
18 ensures that the information contained in the measurements is fully exploited, but not over-used. Numerical experiments con-  
19 firm that the constrained data assimilation algorithm solves the inverse problem in a way that automatise the choice of control  
20 variables, and that restricts the minimisation of the costfunction to the signal-related model variables.



## 21 1 Introduction

22 Atmospheric aerosols have a substantial, yet highly uncertain impact on climate, they can cause respiratory health problems,  
23 degrade visibility, and even compromise air-traffic safety. The physical and chemical properties of aerosols play a key role in  
24 understanding these effects. The aerosol properties are determined by a complex interplay of different chemical, microphysi-  
25 cal, and meteorological processes. These processes are investigated in environmental modelling by use of chemical transport  
26 models (CTMs). However, modelling aerosol processes is plagued by substantial biases and errors (McKeen et al., 2007). It is,  
27 therefore, fundamentally important to evaluate and constrain CTMs by use of measurements.

28 Measurements from satellite instruments provide consistent long-term data sets with global coverage. However, it is notori-  
29 ously difficult to compare measured radiances to modelled aerosol concentrations. An alternative to using radiances is to make  
30 use of satellite retrieval products. For instance, one of the products of the CALIPSO lidar instrument (Cloud-Aerosol Lidar  
31 and Infrared Pathfinder Satellite Observations) is a rough classification of the aerosol types (i.e. dust, smoke, clean/polluted  
32 continental, and clean/polluted marine). This retrieval product is based on lidar depolarisation measurements (Omar et al.,  
33 2009). For the evaluation of aerosol transport models this provides us with a qualitative check for the chemical composition of  
34 aerosols. However, this is of limited practical use, since what we really need is quantitative information on the particles' chem-  
35 ical composition (which can be size-dependent). The most popular approach in evaluating and constraining aerosol transport  
36 models is the use of retrieved optical properties, such as aerosol optical depth, or extinction and backscattering coefficients.  
37 Yet another idea is to provide the particles' refractive index as a retrieval product (e.g. Mishchenko et al. (2007)). However, the  
38 use of such retrieval products still leaves us with the challenge of solving an ill-posed inverse problem, namely, of determining  
39 the particles chemical composition from their retrieved optical or dielectric properties.

40 A systematic class of statistical methods for solving this inverse problem is known as data assimilation. Recent studies  
41 have applied data assimilation to aerosol models with varying degrees of sophistication, ranging from simple dust mod-  
42 els (Khade et al., 2013) and mass transport models (Zhang et al., 2014) to microphysical aerosol models based on modal  
43 (Rubin and Collins, 2014) or sectional descriptions (Sandu et al., 2005; Saide et al., 2013) of the aerosol size distribution. The  
44 assimilation techniques that have been used comprise variational methods, such as 2D (Zhang et al., 2014), 3D (Kahnert,  
45 2008; Liu et al., 2011), and 4D variational methods (Benedetti et al., 2009), as well as ensemble approaches (Sekiyama et al.,  
46 2010). Assimilation of satellite products for trace gases is relatively straightforward, since observed and modelled trace gas  
47 concentrations are almost directly comparable. However, aerosol optical properties observed from satellites are not directly  
48 comparable to the modelled size distribution and chemical composition of the aerosols. Solving this problem amounts to reg-  
49 ularising a severely under-constrained inverse problem. Previous aerosol assimilation attempts have been mainly based on  
50 educated guesses about the information content of the observations. For instance, there have been studies on the assimilation  
51 of aerosol optical depth (AOD) in which all chemical aerosol components in all size classes and at all model layers were used  
52 as independent control variables (Liu et al., 2011). This is a rather bold approach that largely disregards the problems involved  
53 in inverse modelling. By contrast, it has been proposed to only allow for the total aerosol mass concentration to be corrected  
54 by data assimilation of AOD (Benedetti et al., 2009). This is a more prudent approach based on the plausible assumption that a



55 single optical variable only contains enough information to control a single model variable. There have also been intermediate  
56 approaches in which the total aerosol mass per size bin have been used as control variables (Saide et al., 2013).

57 In all such approaches the choice of control variables is based on ad hoc assumptions. Numerical assimilation experiments by  
58 Kahnert (2009) suggest that observations of several aerosol optical properties at multiple wavelengths may allow us to constrain  
59 more than just the total mass concentration, but certainly not *all* aerosol parameters. However, it is still an unsolved mystery how  
60 much information a given set of observations actually contains about the size distribution and chemical composition of aerosols,  
61 and exactly which model variables are related to the observed signals, and which ones are related to noise. Thus a prerequisite  
62 for assimilating remote sensing observations into aerosol transport models is to thoroughly understand the information content  
63 of the observations as well as the relation between the model variables and the signal degrees of freedom.

64 In a recent study, Burton et al. (2016) have investigated the information content of “ $3\beta + 2\alpha$ ” lidar measurements, i.e., obser-  
65 vations of backscattering at three wavelengths and extinction at two wavelengths, where the information content was analysed  
66 with regard to the refractive index and number distribution of the aerosol particles. As mentioned earlier, the refractive index  
67 is a very useful retrieval product of remote sensing observations. However, from the point of view of chemical transport mod-  
68 elling, the main quantities of interest are the concentrations of the different chemical species of which the aerosol particles  
69 are composed. Although the chemical composition determines the refractive index, the inversion of this relationship is still  
70 under-determined, hence an ill-posed problem. In this paper, we want to investigate the inverse problem that goes all the way  
71 from optical properties to the chemical composition of particles.

72 Thus the two main goals of this paper are (i) to apply a systematic method for analysing the information content of aerosol  
73 optical properties with regard to the particles’ chemical composition, and (ii) to test an algorithm for making an automatic  
74 choice of control variables in chemical data assimilation that relate to the signal degrees of freedom, while all other model vari-  
75 ables remain unchanged by the assimilation procedure. The focus will be on spectral observations of extinction and backscat-  
76 tering coefficients, which can be retrieved from lidar observations.<sup>1</sup> We will not restrict this analysis to any fixed choice of  
77 wavelengths, such as  $3\beta + 2\alpha$ . Instead, we will investigate the information content for varying combinations of the three main  
78 Nd:YAG wavelengths. However, it should be mentioned that extinction measurements at the lowest harmonic of 1064 nm can  
79 be difficult and plagued by high errors; in practice, this will affect the observation error, resulting in a low information content  
80 of this particular measurement.

81 The paper is organised as follows. Section 2 gives a rather concise introduction of the modelling tools and of the numerical  
82 approach employed to studying the information content of extinction and backscattering observations. Section 3 presents the  
83 main results of this study, and Sect. 4 offers concluding remarks. To make this paper self-contained, we included an appendix  
84 that gives a brief introduction to some essential concepts of data assimilation, and a detailed explanation of the methods we  
85 used for quantifying the information content of aerosol optical observables. It is advisable to read the appendix first before

<sup>1</sup>In addition to lidar measurements from ground-based and aircraft-carried instruments, there are currently two space-borne lidar instruments in orbit. The CALIOP instrument on-board the CALIPSO satellite has been launched in April 2006; it has three receiver channels, one at 1064 nm, and two channels at 532 nm to measure orthogonally polarised components. The CATS instrument on-board the International Space Station has been operational since January 2015; It measures backscattering at 355 nm, 532 nm, and 1064 nm, were the latter two have two orthogonal polarisation channels. It is also capable of performing high spectral resolution measurements at 532 nm. A third instrument is planned to be launched in 2018 (ATLID on-board EarthCARE).



86 reading the body of the paper. Readers that are not interested in the theory may leave out the appendix at the risk of missing  
87 some of the discussions in the main body of the paper.

## 88 **2 Methods**

89 This study consists of two parts. In the first part we quantify the information content of extinction and backscattering co-  
90 efficients at multiple wavelengths. In the second part we perform a numerical experiment to investigate to what extent the  
91 concentrations of different chemical aerosol components can be constrained by observations of extinction and backscattering  
92 coefficients. The modelling tools required for this study are (i) a chemical transport model; (ii) an aerosol optics model; and  
93 (iii) a data assimilation system.

### 94 **2.1 Multiple scale Atmospheric Transport and CHemistry modelling system (MATCH)**

95 We employ the chemical transport model MATCH, which is an off-line Eulerian CTM with flexible model domain. It has been  
96 previously used from regional to hemispheric scales. Here we use a model version that contains a photochemistry module with  
97 64 chemical species, among them four secondary inorganic aerosol (SIA), namely, ammonium sulphate, ammonium nitrate,  
98 other sulphates, and other nitrates. It also contains a module with 16 primary aerosol variables, namely, seasalt, elemental  
99 carbon (EC), organic carbon (OC), and dust particles, each emitted in four different size bins. Thus, the total number of aerosol  
100 model variables is 20.

101 The model reads in emission data, meteorological data, and land use data and computes transport processes, chemical  
102 transformation, and dry and wet deposition of the various trace gases and aerosols. As output, it provides concentration fields  
103 of gases and aerosols, the deposition of these chemical species to land and water-covered areas, as well as the temporal  
104 evolution of these variables.

105 We mention that there exists another model version that includes aerosol microphysical processes, such as nucleation, con-  
106 densational growth, and coagulation. In that model version the aerosol size distribution evolves dynamically. The model has  
107 20 size bins and seven chemical species (EC, OC, dust, seasalt, particulate sulphate (PSOX), particulate nitrate (PNOX), and  
108 particulate ammonium (PNHX)), although not all species are encountered in all size bins. The total number of model variables  
109 in the present setup is 82.

110 More complete information about the mass transport model can be found in Andersson et al. (2007). The seasalt module is  
111 discussed in Foltescu et al. (2005). The aerosol microphysics module is described in Andersson et al. (2015).

112 For the sake of simplicity we here use the mass transport model without aerosol microphysical processes (see next section).  
113 The model is set up over Europe covering  $33^\circ$  in the longitudinal and  $42^\circ$  in the latitudinal direction in a rotated lat-long  
114 grid with  $0.4^\circ \times 0.4^\circ$  horizontal resolution. In the vertical direction the model domain extends up to 13 hPa, using 40  $\eta$ -layers  
115 with variable thickness depending on the underlying topography. The meteorological input data are taken from the numerical  
116 weather prediction model HIRLAM (Undén et al., 2002). For the emissions we used EMEP data for the year 2007, where EC



117 and OC emissions were computed from total primary particle emissions based on the data in Kupiainen and Klimont (2004,  
118 2007).

## 119 2.2 Aerosol optics model

120 We have two different optics models coupled to MATCH, one to the mass transport module, and another to the aerosol mi-  
121 crophysics module. The former assumes that all aerosol species are homogeneous spheres, and that each chemical species  
122 is contained in separate particles. Under these assumptions the optics model is linear, i.e., the optical properties are linear  
123 functions of the concentrations of the chemical aerosol species. The latter model accounts for the fact that in reality different  
124 chemical species can be internally mixed, i.e., they can be contained in one and the same particle. That model also accounts  
125 for the inhomogeneous internal structure of black carbon mixed with other aerosol components, and for the irregular fractal  
126 aggregate morphology of bare black carbon particles (Kahnert et al., 2012a, 2013). Under these assumptions the optics model  
127 becomes nonlinear, which introduces additional complications in the inverse-modelling problem. This is the main reason why  
128 we chose to use the simpler mass transport optics model in this study. Much of the theory explained in the appendix relies  
129 on the assumption that the optics model is either linear, or that it is only mildly nonlinear, so that it can be linearised — see  
130 Eq. (B6). More information about the aerosol optics models implemented in MATCH can be found in Andersson and Kahnert  
131 (2016).

## 132 2.3 3-dimensional variational data assimilation (3DVAR)

133 Data assimilation is a class of statistical methods for combining model results and observations. The algorithm weighs these  
134 two pieces of information according to their respective error variances and covariances. As output the assimilation returns a  
135 result in model space of which the error variances are smaller than those of the original model estimate. Data assimilation is  
136 commonly employed for constraining model results by use of observations. However, one can also employ data assimilation  
137 as an inverse-modelling tool, i.e. for retrieving a model state from measurements. A brief summary of the theoretical basis of  
138 variational data assimilation is given in the appendix.<sup>2</sup>

139 The MATCH model contains a 3DVAR data assimilation module. This model uses a spectral method, i.e., the model state  
140 vector is Fourier-transformed in the two horizontal coordinates. All error correlations in the horizontal direction are assumed  
141 to be homogeneous and isotropic. Error correlations in the vertical direction and among different chemical species are *not*  
142 assumed to be separable. The background error covariance matrix of the model a priori is modelled with the NMC method  
143 (Parrish and Derber, 1992). A more complete description of our 3DVAR program can be found in Kahnert (2008).

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<sup>2</sup>One should actually distinguish between data *assimilation* and data *analysis*. The latter refers to post-processing the model output by statistically weighing model results and observations. The former refers to a process in which the data analysis is incorporated into the time-integration of the CTM. Our 3DVAR code can be used in either analysis or assimilation mode. However, in this study we only perform numerical experiments at a fixed point in time. Thus we use the 3DVAR code as a data analysis tool.



## 144 2.4 Analysis of the information content of aerosol optical parameters

145 The questions we ask are these.

146 1. Given  $m$  observations of, e.g.,  $m_1$  different parameters at  $m_2$  different wavelengths, so that  $m_1 \cdot m_2 = m$ , how many  
147 independent model variables  $\ell$  can we constrain to better than observation error? Obviously, the best we can achieve  
148 would be  $\ell = m$ ; but in general, we will have  $\ell \leq m$ .

149 2. Which are the  $\ell$  model variables (or linear combinations of model variables) that can be constrained by the measure-  
150 ments?

151 It turns out the the answer to these questions are found by performing a singular value decomposition of the Jacobian of the  
152 observation operator — see Eq. (C6) in the appendix. The right singular vectors can be used to construct a transformation of the  
153 model state vector — see Eq. (D16) in the appendix. The transformed vector components fall into two categories, namely, the  
154 signal-related components, which can be constrained by the measurements, and the noise-related components, which cannot  
155 be constrained by the measurements. From the singular values we can compute the number of signal degrees of freedom, i.e.,  
156 the number of model variables that can be constrained to better than observation error. We can further compute how much a set  
157 of measurements reduces the Shannon-entropy of the model state. This is a quantitative measure for the information content of  
158 the measurements.

159 Readers who are unfamiliar with these concepts are urged to read the brief introduction in the appendix. A more complete  
160 discussion of information aspects and inverse methods for atmospheric sounding can be found in Rodgers (2000).

## 161 2.5 Numerical assimilation experiments

162 We use the results of this analysis to modify our 3DVAR program. More specifically, we implement weak constraints into  
163 the 3DVAR program such that only the signal-related (transformed) model variables are allowed to be adjusted in the data-  
164 analysis procedure, while the noise-related components are not altered. We study the performance of the 3DVAR system by  
165 performing a numerical experiment. To this end, we first perform a reference run by driving the MATCH model with analysed  
166 meteorological data. These reference results are taken as the “true” chemical state of the atmosphere. We apply the optics  
167 model to the model output to generate synthetic “observations”. Next we run the MATCH model again, this time driven with  
168 48 hour-forecast meteorological data. The results are taken as a proxy for a background model-estimate that is impaired by  
169 uncertainties. Finally, we perform a 3DVAR-analysis of the “observations” and the background estimate in an attempt to restore  
170 the reference results. In this numerical experiment we have perfect knowledge of the true state, and we assume that our optics  
171 model is nearly perfect, thus providing nearly perfect observations. (We assumed an observation error standard deviation of  
172 10 %). The only factor that may prevent us from fully restoring the reference state is a lack of information in the observed  
173 parameters on the chemical composition of the aerosol particles. Thus, comparison of the retrieval and reference results gives  
174 us an indication of how strongly different model variables can be controlled by the information contained in the observations.



## 175 3 Results

### 176 3.1 Analysis of the information content of aerosol optical parameters

177 To be specific, we consider the set of parameters  $\{k_{\text{ext}}(\lambda_1), k_{\text{ext}}(\lambda_2), k_{\text{ext}}(\lambda_3), \beta_{\text{sca}}(\lambda_1), \beta_{\text{sca}}(\lambda_2), \beta_{\text{sca}}(\lambda_3)\}$ , where  $k_{\text{ext}}$  and  
178  $\beta_{\text{sca}}$  denote the extinction and backscattering coefficients, respectively, and the wavelengths  $\lambda_1 = 1064$  nm,  $\lambda_2 = 532$  nm, and  
179  $\lambda_3 = 355$  nm denote the first three Nd:YAG harmonics. Out of this six-parameter set we pick different subsets and analyse the  
180 singular values of the corresponding observation operators. From those we compute the number of signal degrees of freedom  
181 as well as the change in Shannon-entropy for each subset of measurements. The results are listed in Table 1, which shows a  
182 number of interesting findings:

- 183 1. When we increase the number of wavelengths from one to two, then the number of signal degrees of freedom  $N_s$  shows  
184 a corresponding increase from 1 to around 1.95–2.00. The change in entropy  $H$  indicates a similar trend; it increases  
185 from around 7 for a single wavelength to up to 12 for two wavelengths. (Compare, e.g., cases 1., 2., and 3. to cases 4.,  
186 5., and 6.) Hence we almost double the information contained in the measurements.
- 187 2. When we increase the number of wavelengths further from two to three, then  $N_s$  only increases from around 2.0 to  
188 around 2.7 (compare cases 4., 5., and 6. to case 7, or case 9. to case 10.) This is also reflected in  $H$ ; it only increases  
189 from around 12 to 13. This indicates that in our particular case there is little extra information to be gained by extending  
190 the number of spectral measurements beyond 2–3 wavelengths.
- 191 3. Supplementing extinction with backscattering measurements results in a significant increase in  $N_s$  and  $H$ . This can be  
192 seen by comparing, e.g., cases 5. and 11. By adding  $\beta_{\text{sca}}$  observations to  $k_{\text{ext}}$  observations the number of signal degrees  
193 of freedom increases from 2 to 3.9, so it almost doubles, while  $H$  increases from 12 to 18.
- 194 4. Case 12 clearly reveals the limitations of extending the set of observed parameters;  $N_s$  is only 4.6, significantly lower  
195 than the number of observed parameters,  $m=6$ .
- 196 5. An inspection of  $w_i$  shows that the singular values often display quite a dramatic decrease from the largest to the smallest  
197 value. By contrast, the contribution  $N_s^i$  to the total number of signal degrees of freedom decreases more gently. This fact  
198 is interesting in relation to the choice of the covariance matrix of the weak constrains — see Eqs. (D18), (D19) and  
199 (D21) in the appendix. In view of our findings here, we conclude that (D18) would yield a very sharp transition from  
200 unconstrained to constrained model variables, Eq. (D19) would give a smooth transition, and Eq. (D21) would give a  
201 moderately sharp transition.

202 We performed a sensitivity study on how the observation error affects the information content in the analysis. It is important  
203 to understand that the observation error is not the same as the measurement error. The latter contributes to the former, but the  
204 observation error contains also other sources of error. For instance, if we deal with morphologically complex particles, but  
205 our lack of knowledge forces us to make assumptions and invoke approximations about the particle shapes, then this source



**Table 1.** Signal degrees of freedom  $N_s$  and change in entropy  $H$  for the lowest model layer (closest to the surface) and different subsets of  $\{k_{\text{ext}}(\lambda_1), k_{\text{ext}}(\lambda_2), k_{\text{ext}}(\lambda_3), \beta_{\text{sca}}(\lambda_1), \beta_{\text{sca}}(\lambda_2), \beta_{\text{sca}}(\lambda_3)\}$ , where  $k_{\text{ext}}$  denotes the extinction coefficient,  $\beta_{\text{sca}}$  represents the backscattering coefficient, and the wavelengths  $\lambda_1 = 1064$  nm,  $\lambda_2 = 532$  nm, and  $\lambda_3 = 355$  nm denote the first three Nd:YAG harmonics. Also shown are the singular values  $w_i$  and their contributions  $N_s^i$  and  $H_i$  to  $N_s$  and  $H$ , respectively. The results have been obtained by assuming an observation standard deviation of 10 %.

No.	Parameters	Wavelengths	$i$	$w_i$	$N_s^i$	$H_i$	$N_s$	$H$
1.	$k_{\text{ext}}$	$\lambda_3$	1	163	1.00	7.35	1.00	7.35
2.	$k_{\text{ext}}$	$\lambda_2$	1	172	1.00	7.43	1.00	7.43
3.	$k_{\text{ext}}$	$\lambda_1$	1	194	1.00	7.52	1.00	7.52
4.	$k_{\text{ext}}$	$\lambda_2, \lambda_3$	1	237	1.00	7.89	1.95	10.1
			2	4.57	0.95	2.23		
5.	$k_{\text{ext}}$	$\lambda_1, \lambda_3$	1	253	1.00	7.98	2.00	11.9
			2	15.0	1.00	3.91		
6.	$k_{\text{ext}}$	$\lambda_1, \lambda_2$	1	259	1.00	8.02	1.99	11.5
			2	10.9	0.99	3.45		
7.	$k_{\text{ext}}$	$\lambda_1, \lambda_2, \lambda_3$	1	305	1.00	8.25	2.72	13.2
			2	15.8	1.00	3.98		
			3	1.64	0.73	0.94		
8.	$\beta_{\text{sca}}$	$\lambda_1$	1	226	1.00	7.82	1.00	7.82
9.	$\beta_{\text{sca}}$	$\lambda_1, \lambda_3$	1	292	1.00	8.19	2.00	12.3
			2	17.6	1.00	4.14		
10.	$\beta_{\text{sca}}$	$\lambda_1, \lambda_2, \lambda_3$	1	354	1.00	8.47	2.65	13.4
			2	18.2	1.00	4.18		
			3	1.36	0.65	0.76		
11.	$k_{\text{ext}}, \beta_{\text{sca}}$	$\lambda_1, \lambda_3$	1	386	1.00	8.59	3.89	17.9
			2	32.8	1.00	5.04		
			3	4.80	0.96	2.29		
			4	3.66	0.93	1.92		
12.	$k_{\text{ext}}, \beta_{\text{sca}}$	$\lambda_1, \lambda_2, \lambda_3$	1	467	1.00	8.87	4.61	19.4
			2	37.8	1.00	5.24		
			3	5.54	0.97	2.49		
			4	4.18	0.95	2.10		
			5	0.95	0.48	0.47		
			6	0.53	0.22	0.18		



**Table 2.** Signal degrees of freedom  $N_s$  and change in entropy  $H$  as a function of observation standard deviation, taken from the first model layer and case 12 in table 1.

Obs. std. dev. (%)	$N_s$	$H$
1	5.96	37.6
5	5.29	24.3
10	4.61	19.3
50	3.00	10.5
100	2.33	7.8

206 of error contributes to the observation error. The same is the case if we lack information about the particles' size distribution.  
207 Such assumptions also enter into our relatively simple optics model, so our previous assumption of an observation standard  
208 deviation of 10 % represents, most likely, a highly idealised case.<sup>3</sup>

209 To get an idea about the significance of the observation error on the amount of information we can extract from measure-  
210 ments, we consider case 12 in table 1, and we varied the observation standard deviation from 1% to 100%. Table 2 shows how  
211 the total entropy and signal degrees of freedom vary with the observation standard deviation. The larger the standard deviation,  
212 the less information can be obtained from the observations. Both the total entropy  $H$  and the signal degrees of freedom  $N_s$   
213 decrease with increasing standard deviation. For a standard deviation of 100 %, we only have two signal degrees of freedom  
214 contained in the six observed optical parameters. This demonstrates two important things.

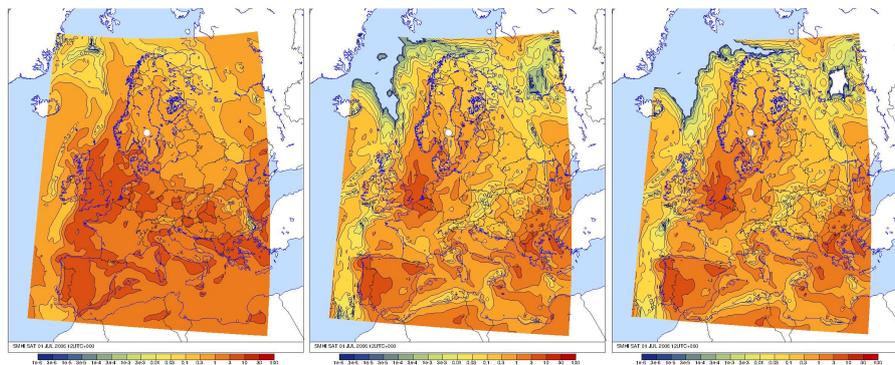
215 1. It is essential to develop accurate and realistic aerosol optics models. The most accurate measurements may intrinsically  
216 contain a wealth of information on aerosol properties. But we can only make use of this information to the extent that our  
217 observation operator is able to accurately describe the relation between the physical and chemical particle characteristics  
218 and their optical properties.

219 2. It is equally essential to accurately estimate the contribution of the uncertainties in the aerosol optics model to the  
220 observation error. If we underestimate this error, we will rely too much on the measurements than we should, thus  
221 assimilating noise. If we overestimate this error, we will waste information contained in the observations.

### 222 3.2 Numerical inverse-modelling experiment

223 We integrated the findings of 3.1 into our 3DVAR program by constraining the algorithm to varying only the signal-related  
224 model variables. We employed the weak-constraint approach described in the appendix. More specifically, the constraints are  
225 formulated by use of the ansatz given in Eq. (D21), where we set the constant  $c'' = \min\{w_k, 0.1\}$ , and where  $w_k$  is the smallest  
226 singular value of the scaled observation operator — see Eq. (C6).

<sup>3</sup>A more realistic optics model, such as the one investigated in Andersson and Kahnert (2016) would help to reduce the observation standard deviation. For future studies, such a model should be linearised and investigated in a similar way.



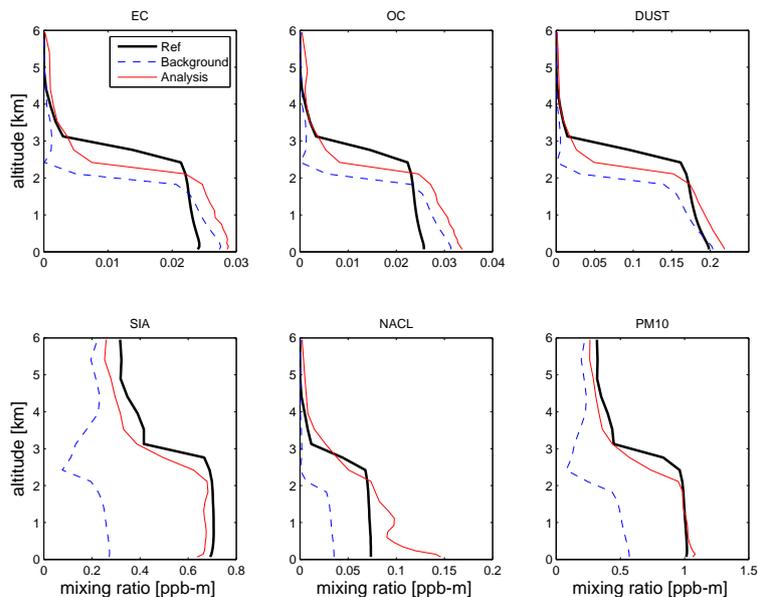
**Figure 1.** Ammonium sulphate mixing ratio over Europe. Left: reference field, centre: background field, right: 3DVAR analysis. The observation site is indicated by the white circle. Note the nonlinear colour scale!

227 To illustrate the method we conduct a numerical experiment as described in Sect. 2.5. We perform a 3DVAR analysis of the  
228 background field by assimilating four different vertical profiles of optical properties, namely, the backscattering coefficient,  
229  $\beta_{\text{bak}}$ , and the extinction coefficient,  $k_{\text{ext}}$ , each at a wavelength of 355 nm and 1064 nm. Thus in our case the number of  
230 singular values in each vertical layer is  $k = 4$ . We assume an idealised situation in which the observation standard deviation  
231 is only 10 %. As we see in Table 1 (row 11), the number of singular values larger than unity is  $\ell = k = 4$ , and the number of  
232 signal degrees of freedom is  $N_s = 3.9$ . So we have as many signal degrees of freedom as we have measurements.

233 As an example, Fig. 1 shows the ammonium sulphate mixing ratio in the lowest model layer (closest to the surface) computed  
234 for the reference run (left), the background estimate (centre), and the 3DVAR analysis (right)<sup>4</sup>. Clearly, the background field  
235 underestimates the reference field in most areas. We picked a location in Central Sweden (60°N, 15°E) to compute backscat-  
236 tering and extinction profiles from the reference results, which were then 3DVAR-analysed in conjunction with the background  
237 field. The analysis (right) restores the reference mixing ratios at and near the observation site. So, at least for ammonium  
238 sulphate mixing ratios in the lowest model layer we seem to obtain a satisfying solution to the inverse modelling problem.

239 A closer inspection of the analysis performance is given in Fig. 2. Each panel shows vertical profiles of mixing ratios at the  
240 observation site. We compare the analysis results (red solid line) to both the background estimate (blue dashed line) and the  
241 reference results (black solid line). The reference results of the secondary inorganic aerosol species (SIA, which is the sum  
242 of all sulphate, nitrate, and ammonium mixing ratios) are almost completely restored by the 3DVAR analysis at all altitudes.  
243 For elemental carbon (EC), organic carbon (OC), dust, and, even more so, sodium chloride (NaCl) the analysis overestimates  
244 the reference results at altitudes between 0–2 km, while above 2 km the reference results are at least partially restored by the  
245 analysis. When we compare the different scales on the x-axes, we see that SIA makes the dominant contribution to the aerosol  
246 mixing ratio. Accordingly, the total aerosol mass mixing ratio ( $\text{PM}_{10}$ ) is almost equally well restored by the analysis as the  
247 SIA mixing ratio.

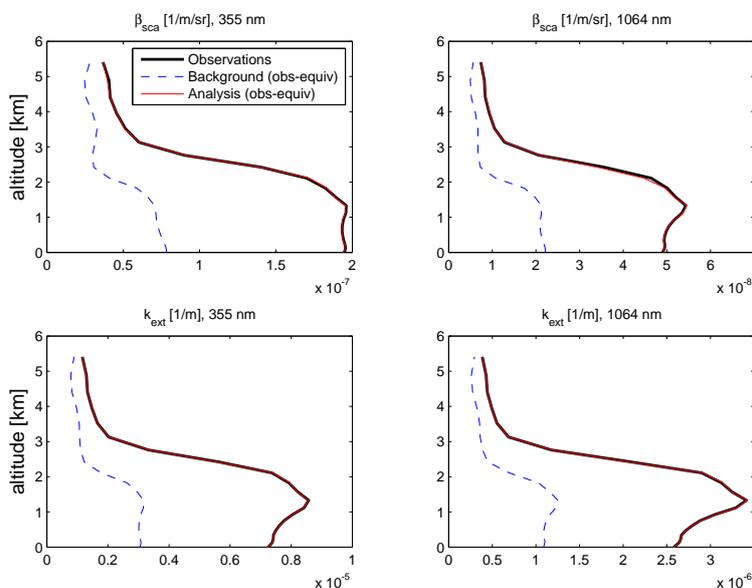
<sup>4</sup>The approach for generating a reference and background model-run has been explained in Sect. 2.5



**Figure 2.** Vertical profiles of elemental carbon (EC), organic carbon (OC), dust, secondary inorganic aerosols (SIA), sea salt, and total aerosol mass mixing ratio (PM<sub>10</sub>). Each panel shows the reference results (black solid line), background estimate (dashed blue line), and the 3DVAR analysis (red solid line).

248 Figure 3 shows the observations (black solid line) as well as the observation-equivalents of the background estimate (blue  
249 dashed line) and the 3DVAR analysis (red solid line) for all four observations, namely,  $\beta_{\text{bak}}$  at 355 nm wavelength (top left),  
250  $\beta_{\text{bak}}$  at 1064 nm (top right),  $k_{\text{ext}}$  at 355 nm (bottom left), and  $k_{\text{ext}}$  at 1064 nm (bottom right). We learn from this figure that the  
251 analysis follows the observations faithfully. The reason for this is that we assumed that the observations were highly accurate  
252 with an error standard deviation of only 10 %.

253 We have seen that the analysis provides a reasonable, but, as expected, not a perfect answer to the inverse problem. We  
254 have further seen that at (and near) the observation site it relies more on the observations than on the background estimate.  
255 However, the previous figures tell us little about the effect of the constraints we introduced. To learn about that we need to  
256 inspect the analysis in the abstract space of the transformed model variables  $\delta x'$  given in Eq. (C16). Figure 4 shows vertical  
257 profiles of all 20 variables  $\delta x'_i$ . The error variance within which each of these variables is allowed to vary in the analysis is  
258 given by the diagonal elements of the matrix  $\mathbf{B}_G$  in Eq. (D21). The first four of these are the singular values given in Table 1  
259 (row 11). The remaining 16 variances are set to 0.1. Thus the first element  $\delta x'_1$  has by far the largest freedom to be adjusted  
260 by the 3DVAR algorithm. The error variance of the second element,  $\delta x'_2$ , is smaller by roughly one order of magnitude. The  
261 error variances of  $\delta x'_3$  and  $\delta x'_4$  are of comparable magnitude, and each one is about one order of magnitude smaller than that  
262 of  $\delta x'_2$ . Finally, the error variances of the remaining 16 elements are about one order of magnitude smaller than that of  $\delta x'_4$ .  
263 Accordingly, the first element,  $\delta x'_1$ , is the one that deviates most strongly from zero. The elements  $\delta x'_2$ ,  $\delta x'_3$ , and  $\delta x'_4$  are varied



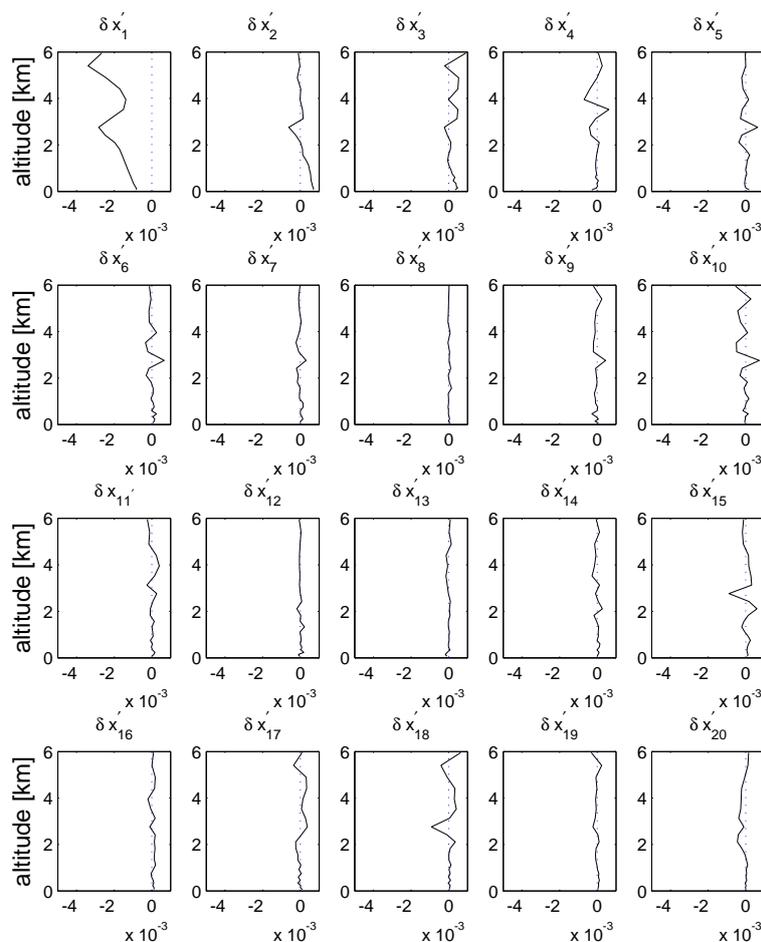
**Figure 3.** Observations (black solid line), and observation-equivalents of the background estimate (dashed blue line) and of the 3DVAR analysis (red solid line). The optical parameters and wavelengths are indicated above each panel.

264 much less in comparison. If we had imposed strong constraints, then the remaining elements would be exactly zero. However,  
265 our weak-constraint formulation allows even the other elements to deviate from zero within relatively tight limits. But several  
266 of them are, in fact, very close to zero, notably the elements  $\delta x'_i$  for  $i = 7-9, 11-14, 16,$  and  $19$ .

#### 267 4 Summary and conclusions

268 We have quantified the information content of extinction and backscattering measurements with regard to the chemical compo-  
269 sition of aerosol particles. This has been done by determining the singular values of the observation operator, by computing the  
270 number of signal degrees of freedom, and by calculating the change in Shannon-entropy caused by taking measurements. We  
271 first assumed a relatively low observation standard deviation of 10 %. In that case, when adding measurements of  $\beta_{\text{bak}}$  to mea-  
272 surements of  $k_{\text{ext}}$ , the information content nearly doubles. The same is true when we increase the number of wavelengths from  
273 a single wavelength to two wavelengths. However, when we further increase the number of wavelengths from two to three, then  
274 the gain in information is rather modest; there appeared to be little use in increasing the number of optical wavelengths beyond  
275 three. When the full set of six observations (both optical parameters at three wavelengths) is considered, then the number of  
276 signal degrees of freedom is 4.6. Thus we can constrain, at most, 4–5 model variables with this set of observations.

277 These conclusions depend, to be sure, on the assumed observation standard deviation. We therefore performed a sensitiv-  
278 ity study where we investigated how the observation standard deviation affects the information content. We observed a rather



**Figure 4.** Vertical profiles of the transformed model variables  $\delta x'$ .

279 dramatic decrease in both the entropy and signal degrees of freedom with increasing observation standard deviation. Note  
280 that not only the measurement error, but also the uncertainties in the aerosol optics model contribute to the observation er-  
281 ror. This highlights the importance of developing accurate aerosol optics models and of obtaining an accurate estimate of  
282 the observation error, especially of the uncertainty in the aerosol optics model. This is a prerequisite for extracting as much  
283 information as possible from the measurements, while avoiding to extract noise rather than signal. More often than not, com-  
284 putational limitations and lack of knowledge force us to introduce simplifying assumptions about the particles' morphologies.  
285 However, we know that aerosol optical properties can be highly sensitive to the shape (Mishchenko et al. (1997); Kahnert  
286 (2004)), small-scale surface roughness (Kahnert et al., 2012b), inhomogeneity (Mishchenko et al., 2014; Kahnert, 2015), ag-



287 gregation (Fuller and Mackowski, 2000; Liu and Mishchenko, 2007; Kahnert and Devasthale, 2011), irregularity (Muinonen,  
288 2000; Bi et al., 2010), porosity (Vilaplana et al., 2006; Lindqvist et al., 2011; Kylling et al., 2014), and combinations thereof  
289 (Lindqvist et al., 2009; Kahnert et al., 2013; Lindqvist et al., 2014). We need to know how much these sources of uncertainty  
290 contribute to the observation standard deviation. One way of estimating this is to compare aerosol optical properties computed  
291 with simple shape models to either measurements or to computations based on more realistic particle shape models — see  
292 Kahnert et al. (2016) for a recent review and a more detailed discussion.

293 We exploited our analysis of the information content and the number of signal degrees of freedom by formulating weak  
294 constraints in a 3DVAR algorithm. More specifically, we transformed the model variables into a new basis in which the com-  
295 ponents of the state vector can be divided into signal-related and noise-related components. We then added weak constraints  
296 to the assimilation algorithm in such a way that only the signal-related transformed model variables are varied by the 3DVAR  
297 analysis. Numerical experiments showed that the 3DVAR algorithm provided a reasonable solution to the inverse problem;  
298 when mapped into observation space, the analysis result closely reproduces the measurements. It also appeared that among the  
299 original model variables, secondary inorganic aerosol components were most faithfully retrieved by the inverse modelling solu-  
300 tion. Most importantly, it was demonstrated that the 3DVAR analysis follows, indeed, the imposed constraints; the transformed  
301 model variables are adjusted within certain limits according to how strongly they relate to the signal degrees of freedom.

302 The results presented here suggest further questions that should be addressed in future studies. We have performed this  
303 investigation with a mass transport model, thus focusing on the information content of optical measurements on the chemical  
304 composition of aerosols. When we include aerosol microphysical processes, then the model delivers the aerosols' size distribu-  
305 tion, as well as their size-resolved chemical composition. This makes the problem quite different from the one we investigated  
306 here. First, the dimension of the model space is considerably larger for an aerosol microphysics transport model. Constraining  
307 such a model with limited information from measurements becomes even more challenging than in the case of a mass transport  
308 model. On the other hand, an aerosol microphysics model delivers information on the particles size distribution and mixing  
309 state. Therefore, this would require us to make fewer assumptions in the aerosol optics model, which may reduce the obser-  
310 vation error. The present study should be extended to investigate the information contained in extinction and backscattering  
311 measurements for simultaneously constraining the chemical composition and the size of aerosol particles.

312 Another important, and often highly underrated issue concerns the choice of the aerosol optics model. In the present study  
313 we employed a simple homogeneous-sphere model in which all chemical components were assumed to be externally mixed.  
314 There is little one can put forward in defence of this model other than pure convenience. In this model the observation operator  
315 is linear, which is a prerequisite for much of the theoretical foundations of this study — see the appendix for details. However,  
316 it has been demonstrated that drastically simplifying assumptions, such as the external-mixture approximation, can give model  
317 results for aerosol optical properties that differ substantially from those obtained with more realistic nonlinear optics models  
318 (Andersson and Kahnert, 2016). It would therefore be important to extend the present study to include more accurate and  
319 realistic optics models. A first step could be to analyse the degree of nonlinearity of optics models that account for internal  
320 mixing of different aerosol species. If they turn out to be only mildly nonlinear, then one can linearise them and work with



321 the Jacobian of the nonlinear observation operator. Otherwise the theoretical methods employed in this paper would have to be  
322 extended in order to accommodate nonlinear observation operators.

### 323 **Appendix A: Inverse problems**

324 Suppose we have a system described by a set of variables  $x_1, \dots, x_n$ , summarised in a vector  $\mathbf{x}$ . Suppose also that we have an  
325 operator  $\hat{H} : \mathbb{R}^n \rightarrow \mathbb{R}^m$ ,  $\mathbf{x} \mapsto \mathbf{y} = \hat{H}(\mathbf{x})$  that allows us to compute a set of variables  $y_1, \dots, y_m$ , summarised in a vector  $\mathbf{y}$ . To  
326 take a specific example, we may think of  $\mathbf{x}$  as a vector of mixing ratios of chemical aerosol species,  $\mathbf{y}$  as a set of aerosol optical  
327 properties, and  $\hat{H}$  as an aerosol optics model. We consider the following two problems:

- 328 1. **Direct problem:** Given  $\mathbf{x}$  and  $\hat{H}$ , calculate  $\mathbf{y} = \hat{H}(\mathbf{x})$ .
- 329 2. **Inverse problem:** Given  $\mathbf{y}$  and  $\hat{H}$ , solve  $\mathbf{y} = \hat{H}(\mathbf{x})$  for  $\mathbf{x}$ .

330 A pair of such problems is inverse *to each other*; it is, therefore, somewhat arbitrary which problem we choose to call the  
331 direct problem, and which one we call the inverse problem. However, one of the problems is usually *well-posed*, while the  
332 other one is *ill-posed*. Such is also the case in aerosol optics modelling. It is customary to call the well-posed problem the  
333 *direct problem*, and the ill-posed one the *inverse problem*.

334 An equation  $\mathbf{y} = \hat{H}(\mathbf{x})$  is called *well-posed* if it has the following properties:

- 335 1. **Existence:** For every  $\mathbf{y} \in \mathbb{R}^m$ , there is at least one  $\mathbf{x} \in \mathbb{R}^n$  for which  $\mathbf{y} = \hat{H}(\mathbf{x})$ .
- 336 2. **Uniqueness:** For every  $\mathbf{y} \in \mathbb{R}^m$ , there is at most one  $\mathbf{x} \in \mathbb{R}^n$  for which  $\mathbf{y} = \hat{H}(\mathbf{x})$ .
- 337 3. **Stability:** The solution  $\mathbf{x}$  depends continuously on  $\mathbf{y}$ .

338 If any of these properties is not fulfilled, then the problem is called *ill-posed*.

### 339 **Appendix B: 3-dimensional variational data assimilation**

340 Data assimilation is usually employed for constraining models by use of measurements, but it can also be used to solve inverse  
341 problems. Here we focus on one specific data assimilation method known as 3-dimensional variational data assimilation, or  
342 3DVAR.

343 In a CTM we discretise the geographic domain of interest into a 3-dimensional grid. In each grid cell, the aerosol particles  
344 are characterised by the mass concentrations of each chemical component in the aerosol phase, such as sulphate, nitrate,  
345 ammonium, mineral dust, black carbon, organic carbon, and sea salt. Suppose we summarise all these mass concentrations  
346 from all grid cells into one large vector  $\mathbf{x} \in \mathbb{R}^n$ . The model provides us with a first guess of the atmospheric aerosol state,  
347 known as a *background estimate*<sup>5</sup>  $\mathbf{x}_b$ . Suppose also that we have  $m$  observations, which we summarise in a vector  $\mathbf{y} \in \mathbb{R}^m$ . We

<sup>5</sup>In the remote sensing and inverse modelling community, the background estimate is more commonly referred to as the *a priori* estimate.



348 further have an observation operator  $\hat{H} : \mathbb{R}^n \rightarrow \mathbb{R}^m$ ,  $\mathbf{x} \mapsto \hat{H}(\mathbf{x})$  that maps the state vector  $\mathbf{x}$  from model space to observation  
 349 space<sup>6</sup>. We further denote by  $\mathbf{x}_t$  the true state of the atmosphere, by  $\boldsymbol{\epsilon}_b = \mathbf{x}_t - \mathbf{x}_b$  the error of the background estimate, and by  
 350  $\boldsymbol{\epsilon}_o = \hat{H}(\mathbf{x}_t) - \mathbf{y}$  the observation error.<sup>7</sup> If the background errors are not correlated with the observation errors, then their joint  
 351 probability distribution becomes separable, i.e.

$$352 \quad P(\boldsymbol{\epsilon}_b, \boldsymbol{\epsilon}_o) = P_b(\boldsymbol{\epsilon}_b)P_o(\boldsymbol{\epsilon}_o). \quad (\text{B1})$$

353 The true state of the atmosphere is, of course, unknown. Therefore, our definition of the errors and their probability distri-  
 354 bution is only of conceptual use, but not of any practical value. However, we can reinterpret the probability distributions by  
 355 replacing  $\boldsymbol{\epsilon}_b$  in the argument of  $P_b$  with  $\mathbf{x} - \mathbf{x}_b$ , and by replacing  $\boldsymbol{\epsilon}_o$  in the argument of  $P_o$  with  $\hat{H}(\mathbf{x}) - \mathbf{y}$ . We further assume  
 356 that both the background and the observation errors are normally distributed. Thus we may write

$$357 \quad P_b(\mathbf{x}) = (2\pi |\mathbf{B}|)^{-1/2} \exp\left(-\frac{1}{2}(\mathbf{x} - \mathbf{x}_b)^T \cdot \mathbf{B}^{-1} \cdot (\mathbf{x} - \mathbf{x}_b)\right) \quad (\text{B2})$$

$$358 \quad P_o(\mathbf{x}) = (2\pi |\mathbf{R}|)^{-1/2} \exp\left(-\frac{1}{2}(\hat{H}(\mathbf{x}) - \mathbf{y})^T \cdot \mathbf{R}^{-1} \cdot (\hat{H}(\mathbf{x}) - \mathbf{y})\right). \quad (\text{B3})$$

359 Here  $\mathbf{B}$  and  $\mathbf{R}$  denote the covariance matrices of the background and observation errors, respectively, and  $|\cdot|$  denotes the  
 360 matrix determinant. In this form,  $P_b(\mathbf{x})$  represents the probability that the atmospheric aerosol particles are found in state  $\mathbf{x}$ ,  
 361 given a background estimate  $\mathbf{x}_b$  with error covariance matrix  $\mathbf{B}$ . Similarly,  $P_o(\mathbf{x})$  is the probability that the system is found in  
 362 state  $\mathbf{x}$ , given measurements  $\mathbf{y}$  with error covariances  $\mathbf{R}$ .<sup>8</sup>

363 Equations (B1)–(B3) can be summarised in the form

$$364 \quad P(\mathbf{x}) = \frac{1}{2\pi(|\mathbf{B}| \cdot |\mathbf{R}|)^{-1/2}} \exp(-J(\mathbf{x})) \quad (\text{B4})$$

$$365 \quad J(\mathbf{x}) = \frac{1}{2} \left[ (\mathbf{x} - \mathbf{x}_b)^T \cdot \mathbf{B}^{-1} \cdot (\mathbf{x} - \mathbf{x}_b) + (\hat{H}(\mathbf{x}) - \mathbf{y})^T \cdot \mathbf{R}^{-1} \cdot (\hat{H}(\mathbf{x}) - \mathbf{y}) \right], \quad (\text{B5})$$

366

367 where  $J$  is suggestively called the costfunction, since it can be interpreted as a measure for how “costly” it is for a state  $\mathbf{x}$  to  
 368 simultaneously deviate from the background estimate and the measurements within the permitted error bounds. The deviations  
 369 are weighted with the inverse error covariance matrices. For instance, this means that for measurements with a small error  
 370 variance, a deviation  $\hat{H}(\mathbf{x}) - \mathbf{y}$  becomes “more costly”.

371 We are interested in the most probable aerosol state of the atmosphere, i.e., in that state  $\mathbf{x}_a$  for which the probability  
 372 distribution attains its maximum. This is obviously the case when the argument of the exponential in Eq. (B4) assumes a

<sup>6</sup>The optics model  $\hat{H}$  usually has to invoke assumptions about physical aerosol properties that are relevant for the optical properties, but not provided by the CTM output, e.g. assumptions about the morphology of the particles. If the CTM is a simple mass-transport model without aerosol microphysics, then it is also necessary to invoke assumptions about the size distribution of the aerosols.

<sup>7</sup>The observation error must not be confused with the measurement error. The latter contributes to the former, but the observation error contains also other sources of error. For instance, if we deal with morphologically complex particles, but our lack of knowledge forces us to make assumptions and invoke approximations about the particle shapes, then this source of error contributes to the observation error. The same is the case if we lack information about the particles’ size distribution.

<sup>8</sup>The observation errors are often uncorrelated. In such case the matrix  $\mathbf{R}$  is diagonal, where the diagonal elements are the observation error variances.



373 minimum. Thus we seek to minimise the costfunction  $J$ . The variational method is based on computing the gradient of the  
 374 costfunction,  $\nabla J$ , and to use this in a descent algorithm to iteratively search for the minimum of  $J$ .

375 In practice it is common to introduce the variable  $\delta\mathbf{x} = \mathbf{x} - \mathbf{x}_b$ , and use the first-order Taylor expansion of the observation  
 376 operator,

$$377 \hat{H}(\mathbf{x}) = \hat{H}(\mathbf{x}_b) + \mathbf{H} \cdot \delta\mathbf{x}, \quad (\text{B6})$$

378 where the  $(m \times n)$ -matrix  $\mathbf{H}$  denotes the Jacobian of  $\hat{H}$  at  $\mathbf{x} = \mathbf{x}_b$ . If  $\hat{H}$  is only mildly non-linear, and if the components of  $\delta\mathbf{x}$   
 379 are sufficiently small, then we can substitute this first-order approximation into Eq. (B5), which yields

$$380 J = J_b + J_o \quad (\text{B7})$$

$$381 J_b(\delta\mathbf{x}) = \frac{1}{2} \delta\mathbf{x}^T \cdot \mathbf{B}^{-1} \cdot \delta\mathbf{x} \quad (\text{B8})$$

$$382 J_o(\delta\mathbf{x}) = \frac{1}{2} \left( \hat{H}(\mathbf{x}_b) + \mathbf{H} \cdot \delta\mathbf{x} - \mathbf{y} \right)^T \cdot \mathbf{R}^{-1} \cdot \left( \hat{H}(\mathbf{x}_b) + \mathbf{H} \cdot \delta\mathbf{x} - \mathbf{y} \right) \quad (\text{B9})$$

383 The components of the vector  $\delta\mathbf{x}$  are the *control variables* that are iteratively varied by the algorithm until the minimum of the  
 384 costfunction is found.

385 The solution to the equation  $\nabla J = \mathbf{0}$  is a solution to the inverse problem; we input the observations  $\mathbf{y}$  into the algorithm, and  
 386 as output we obtain a result in model space that is consistent with the measurements (within the given error bounds). What if  
 387 the measurements contain insufficient information about the state  $\mathbf{x}$ ? The algorithm will still provide an answer to the inverse  
 388 problem, but the missing information will be supplemented by the background estimate  $\mathbf{x}_b$ . The weighting of the two pieces of  
 389 information,  $\mathbf{x}_b$  and  $\mathbf{y}$ , is controlled by the respective error covariance matrices. Thus data assimilation is a statistical approach,  
 390 which can be expected to give good results *on average*, but not in every single time-step of the model run. This can become  
 391 highly problematic if we only have very few observations, i.e.,  $m \ll n$ . If we allow all model variables to be freely adjusted by  
 392 the assimilation algorithm in such a severely under-constrained case, then the algorithm may just assimilate noise rather than  
 393 signal, resulting in unreasonable solutions to the inverse problem (e.g. Kahnert (2009)). To avoid such problems, one needs to  
 394 systematically analyse the information content of the observations and constrain the assimilation algorithm to only operate on  
 395 the signal degrees of freedom.

### 396 Appendix C: Information content of measurements

397 Our ultimate goal is to formulate the data assimilation problem in such a way that the information contained in the mea-  
 398 surements is fully exploited, but not over-used. To this end, we first need to know how many independent quantities can be  
 399 determined from a specific set of measurements. We investigate this question by borrowing ideas from retrieval and information  
 400 theory — see Rodgers (2000) for more detailed explanations.

401 The main idea is to compare the variances of the model variables to those of the observations. Only those model variables  
 402 whose variance is larger than those of the observations can be constrained by measurements. However, to actually make such



403 a comparison is rather tricky. The first problem is that one cannot readily compare error covariance *matrices*. The second  
 404 problem is that model variables and measurements are in different spaces. We first address the second problem.

405 When we account for observation errors  $\epsilon_o$ , then the basic relation between model variables and observations is, to first order

$$406 \mathbf{y} = \hat{H}(\mathbf{x}_b) + \mathbf{H} \cdot \delta \mathbf{x} + \epsilon_o. \quad (\text{C1})$$

407 The error covariance matrices are given by the expectation values  $\mathbf{B} = \langle \delta \mathbf{x} \cdot \delta \mathbf{x}^T \rangle$ , and  $\mathbf{R} = \langle \epsilon_o \cdot \epsilon_o^T \rangle$ , where the dot denotes a  
 408 dyadic product. The covariance matrix of  $\delta \mathbf{y} = \mathbf{y} - \hat{H}(\mathbf{x}_b)$  is given by  $\langle \delta \mathbf{y} \cdot \delta \mathbf{y}^T \rangle = \mathbf{H} \cdot \mathbf{B} \cdot \mathbf{H}^T + \mathbf{R}$ , where we assumed that  
 409 background and observation errors are uncorrelated. This last equation suggests that we can compare model and observation  
 410 errors in the same space by transforming the background error covariance matrix from the space of  $(n \times n)$  matrices to the  
 411 space of  $(m \times m)$  matrices viz.  $\mathbf{H} \cdot \mathbf{B} \cdot \mathbf{H}^T$ .

412 To address the first problem, we diagonalise the covariance matrices by making the following change of variables

$$413 \delta \tilde{\mathbf{x}} = \mathbf{B}^{-1/2} \cdot \delta \mathbf{x} \quad (\text{C2})$$

$$414 \delta \tilde{\mathbf{y}} = \mathbf{R}^{-1/2} \cdot (\mathbf{y} - \hat{H}(\mathbf{x}_b)) \quad (\text{C3})$$

$$415 \tilde{\mathbf{H}} = \mathbf{R}^{-1/2} \cdot \mathbf{H} \cdot \mathbf{B}^{1/2}. \quad (\text{C4})$$

416 Here  $\mathbf{B}^{1/2}$  denotes the positive square root<sup>9</sup> of the matrix  $\mathbf{B}$ , and  $\mathbf{B}^{-1/2}$  denotes its inverse. In the new basis, the costfunction  
 417 in (B7)–(B9) becomes

$$418 J = \frac{1}{2} \delta \tilde{\mathbf{x}}^T \cdot \delta \tilde{\mathbf{x}} + \frac{1}{2} \left( \tilde{\mathbf{H}} \cdot \delta \tilde{\mathbf{x}} - \delta \tilde{\mathbf{y}} \right)^T \cdot \left( \tilde{\mathbf{H}} \cdot \delta \tilde{\mathbf{x}} - \delta \tilde{\mathbf{y}} \right). \quad (\text{C5})$$

419 The covariance matrices are now unit matrices. This can also be seen by considering the transformed errors, e.g.  $\tilde{\epsilon}_o = \mathbf{R}^{-1/2} \cdot \epsilon_o$   
 420 and computing  $\langle \tilde{\epsilon}_o \cdot \tilde{\epsilon}_o^T \rangle = \mathbf{R}^{-1/2} \cdot \langle \epsilon_o \cdot \epsilon_o^T \rangle \cdot \mathbf{R}^{-1/2} = \mathbf{1}$ , since  $\langle \epsilon_o \cdot \epsilon_o^T \rangle = \mathbf{R}$ . Similarly, we find  $\langle \delta \tilde{\mathbf{x}} \cdot \delta \tilde{\mathbf{x}}^T \rangle = \mathbf{1}$ . The covariance matrix  
 421 of the transformed measurement vector  $\delta \tilde{\mathbf{y}}$  is given by  $\langle \delta \tilde{\mathbf{y}} \cdot \delta \tilde{\mathbf{y}}^T \rangle = \tilde{\mathbf{H}} \cdot \tilde{\mathbf{H}}^T + \mathbf{1}$ . The first term is the model error covariance term  
 422 transformed into observation space, while the second term (the unit matrix) is the diagonalised observation error covariance  
 423 matrix.

424 We are still not in a position to make a meaningful comparison of model and observation errors, since the first term,  $\tilde{\mathbf{H}} \cdot \tilde{\mathbf{H}}^T$ ,  
 425 is still not diagonal. To make it so we need to perform one more transformation. To this end, we consider the singular value  
 426 decomposition of the matrix  $\tilde{\mathbf{H}}$ ,

$$427 \tilde{\mathbf{H}} = \mathbf{R}^{-1/2} \cdot \mathbf{H} \cdot \mathbf{B}^{1/2} = \mathbf{V}_L \cdot \mathbf{W} \cdot \mathbf{V}_R^T. \quad (\text{C6})$$

428 Here  $\tilde{\mathbf{H}}$  is a  $(m \times n)$ -matrix, the matrix of the left-singular vectors  $\mathbf{V}_L$  is a  $(m \times m)$ -matrix, the matrix  $\mathbf{V}_R$  containing the  
 429 right-singular vectors is a  $(n \times n)$ -matrix, and the  $(m \times n)$ -matrix  $\mathbf{W}$  consists of two blocks. If  $m < n$ , then the left block of  $\mathbf{W}$   
 430 is a  $(m \times m)$ -diagonal matrix containing the  $m$  singular values  $w_1, \dots, w_m$  on the diagonal; the right block is a  $(m \times (n - m))$ -  
 431 nullmatrix. Similarly, if  $m > n$ , then the upper block of  $\mathbf{W}$  is a  $(n \times n)$ -diagonal matrix containing the  $n$  singular values on the  
 432 diagonal, while the lower block is a  $((m - n) \times n)$ -nullmatrix.

<sup>9</sup>A matrix  $\mathbf{A}$  is called a square root of a matrix  $\mathbf{B}$  if  $\mathbf{A} \cdot \mathbf{A} = \mathbf{B}$ . The *positive* square root of  $\mathbf{B}$ , which is denoted by  $\mathbf{B}^{1/2}$ , has the property  $\mathbf{x}^T \cdot \mathbf{B}^{1/2} \cdot \mathbf{x} \geq 0$  for all  $\mathbf{x}$ . If  $\mathbf{B}$  is itself positive and symmetric, as is the case for covariance matrices, then the positive square root exists and is unique.



433 We now make another change of variables:

$$434 \quad \delta \mathbf{x}' = \mathbf{V}_R^T \cdot \delta \tilde{\mathbf{x}} \quad (\text{C7})$$

$$435 \quad \delta \mathbf{y}' = \mathbf{V}_L^T \cdot \delta \tilde{\mathbf{y}} \quad (\text{C8})$$

$$436 \quad \mathbf{H}' = \mathbf{V}_L^T \cdot \tilde{\mathbf{H}} \cdot \mathbf{V}_R. \quad (\text{C9})$$

437 The matrices  $\mathbf{V}_L$  and  $\mathbf{V}_R$  are orthogonal, i.e.,  $\mathbf{V}_L^T \cdot \mathbf{V}_L = \mathbf{1}$ , and similarly for  $\mathbf{V}_R$ . Thus, substitution of (C7)–(C9) into (C5)  
 438 yields

$$439 \quad J = \frac{1}{2} \delta \mathbf{x}'^T \cdot \delta \mathbf{x}' + \frac{1}{2} (\mathbf{H}' \cdot \delta \mathbf{x}' - \delta \mathbf{y}')^T \cdot (\mathbf{H}' \cdot \delta \mathbf{x}' - \delta \mathbf{y}'). \quad (\text{C10})$$

440 Evidently, the transformation given in (C7)–(C9) preserves the diagonality of the background and observation error covariance  
 441 matrices. What about the covariance matrix  $\langle \delta \mathbf{y}' \cdot \delta \mathbf{y}'^T \rangle$  in the new basis? Using  $\epsilon'_o = \mathbf{V}_L^T \cdot \tilde{\epsilon}_o = \mathbf{V}_L^T \cdot \mathbf{R}^{-1/2} \cdot \epsilon_o$ , as well as Eqs.  
 442 (C1), (C2)–(C4), and (C7)–(C9), we obtain  $\langle \delta \mathbf{y}' \cdot \delta \mathbf{y}'^T \rangle = \mathbf{H}' \cdot \mathbf{H}'^T + \mathbf{1}$ . The contribution of the background error covariances  
 443 in this coordinate system is  $\mathbf{H}' \cdot \mathbf{H}'^T$ , which is a diagonal matrix. This becomes clear from Eqs. (C6) and (C9), which yields

$$444 \quad \mathbf{H}' \cdot \mathbf{H}'^T = \mathbf{W} \cdot \mathbf{W}^T, \quad (\text{C11})$$

445 which is a  $(m \times m)$  diagonal matrix. Thus in this coordinate system we can readily compare the diagonal elements of the  
 446 transformed background error covariance matrix  $\mathbf{H}' \cdot \mathbf{H}'^T$  to the diagonal (unit) elements of the observation error covariance  
 447 matrix  $\mathbf{1}$ . Roughly, those singular values  $w_i$  on the diagonal of  $\mathbf{W}$  that are larger than unity correspond to model variables  $\delta x'_i$   
 448 that can be controlled by the measurements. Those singular values smaller than unity correspond to model variables that are  
 449 only related to noise.

450 In the above discussion we relied on plausibility arguments. We mention that there are more systematic ways of approaching  
 451 the problem. Here we merely state some key results without going into details. The interested reader is referred to Rodgers  
 452 (2000). However, in all approaches the main quantities of interest are always the singular values of the matrix  $\mathbf{R}^{-1/2} \cdot \mathbf{H} \cdot \mathbf{B}^{1/2}$ .

453 One can compute the number of signal degrees of freedom  $N_s$  from the expectation value of  $J_b$  in Eq. (B8). The result can  
 454 be expressed in terms of the singular values  $w_i$  of the transformed observation operator in Eq. (C6):

$$455 \quad N_s = \sum_i w_i^2 / (1 + w_i^2). \quad (\text{C12})$$

456 Another approach is based on information theory. Given a system described by a probability distribution function  $P(x)$ , one  
 457 defines the Shannon-entropy

$$458 \quad S(P) = - \int P(x) \ln \left( \frac{P(x)}{P_0(x)} \right) dx, \quad (\text{C13})$$

459 where  $P_0$  is a normalisation factor needed to make the argument of the logarithm dimensionless. A decrease in entropy ex-  
 460 presses an increase in our knowledge of the system. For instance, if we initially describe the system by  $P_i(x)$ , and, after taking  
 461 measurements, by  $P_f(x)$ , then the measurement process has changed the entropy by an amount

$$462 \quad H = S(P_i) - S(P_f). \quad (\text{C14})$$



463 In our case, we assume that all errors are normally distributed. In that case, one can show that

$$464 \quad H = \frac{1}{2} \sum_i \ln(1 + w_i^2). \quad (\text{C15})$$

465  $H$  can be interpreted as a measure for the information content of a set of measurements.

466 Our findings so far suggest a general strategy for how to optimise the amount of information that we can extract from  
 467 measurements. First, we need to compute the singular value decomposition in Eq. (C6), as well as the transformation given in  
 468 (C2) and (C7), which we can summarise as

$$469 \quad \delta \mathbf{x}' = \mathbf{V}_R^T \cdot \mathbf{B}^{-1/2} \cdot \delta \mathbf{x}. \quad (\text{C16})$$

470 Then we want to formulate the minimisation of the costfunction in such a way that only those components of  $\delta \mathbf{x}'$  are adjusted  
 471 by the assimilation algorithm that correspond to the largest singular values of the matrix  $\mathbf{W}$  in (C6). All other elements of  
 472  $\delta \mathbf{x}'$  should be left alone. In other words, we want to constrain the minimisation of the costfunction to the subspace of the  
 473 signal degrees of freedom of the state vector. Thus, in order to implement this idea, we first need to discuss how to incorporate  
 474 constraints into the theory.

#### 475 **Appendix D: Minimisation of the costfunction with constraints**

476 In the minimisation of the costfunction all elements of the control vector  $\delta \mathbf{x}$  are independently adjusted until the minimum  
 477 of  $J$  is found. This may not be a prudent approach if the information contained in the observations is insufficient to constrain  
 478 all model variables. In such case one should introduce constraints that reduce the number of independent control variables.  
 479 However, this needs to be done in a clever way; the goal is to neither under-use the measurements (thus wasting available  
 480 information), nor to over-use them (thus assimilating noise).

481 For reasons we will explain later we formulate the constraints as weak conditions. However, for didactic reasons as well as  
 482 for the sake of completeness, we will also mention how to formulate constraints as strong conditions.

#### 483 **D1 Minimisation of the costfunction with strong constraints**

484 Given  $k$  constraints in the form  $g_i(\delta \mathbf{x})=0$ ,  $i = 1, \dots, k$ , the most general way of finding the minimum of  $J(\delta \mathbf{x})$  under the  
 485 constraints  $g_i$  is the method of Lagrange multipliers. More specifically, one introduces  $k$  Lagrange multipliers  $\lambda_1, \dots, \lambda_k$  and  
 486 defines the function

$$487 \quad L(\delta x_1, \dots, \delta x_n, \lambda_1, \dots, \lambda_k) = J(\delta x_1, \dots, \delta x_n) + \sum_{i=1}^k \lambda_i g_i(\delta x_1, \dots, \delta x_n); \quad (\text{D1})$$

488 then one solves the minimisation problem

$$489 \quad \nabla L(\delta x_1, \dots, \delta x_n, \lambda_1, \dots, \lambda_k) = \mathbf{0}, \quad (\text{D2})$$



490 where  $\nabla = \nabla_{\delta x_1, \dots, \delta x_n, \lambda_1, \dots, \lambda_k}$  is now a  $(n+k)$ -dimensional gradient operator. Note that in this general formulation of the  
 491 problem the constraints can even be nonlinear. We are specifically interested in linear constraints, which can be expressed in  
 492 the form  $\mathbf{G} \cdot \delta \mathbf{x} = \mathbf{0}$ . Then the constrained minimisation problem becomes

$$493 \quad L(\delta \mathbf{x}, \boldsymbol{\lambda}) = J(\delta \mathbf{x}) + \boldsymbol{\lambda}^T \cdot \mathbf{G} \cdot \delta \mathbf{x} \quad (\text{D3})$$

$$494 \quad \nabla_{\delta \mathbf{x}, \boldsymbol{\lambda}} L(\delta \mathbf{x}, \boldsymbol{\lambda}) = \begin{pmatrix} \nabla_{\delta \mathbf{x}} J(\delta \mathbf{x}) + \boldsymbol{\lambda}^T \cdot \mathbf{G} \\ \mathbf{G} \cdot \delta \mathbf{x} \end{pmatrix} = \mathbf{0}. \quad (\text{D4})$$

495 Compared to the unconstrained minimisation problem, the introduction of  $k$  constraints has increased the dimension of the  
 496 problem from  $n$  to  $n+k$ . Naively, one may have expected that the dimension would, on the contrary, be reduced to  $n-k$ . This  
 497 is indeed the case if the constraints are linear, and if the function  $J$  is quadratic, as is the case in Eqs. (B7)–(B9). To see this,  
 498 let us first write those equations more concisely in the form

$$499 \quad J = \frac{1}{2} (\delta \mathbf{x}^T \cdot \mathbf{Q}_1 \cdot \delta \mathbf{x} + \mathbf{Q}_2^T \cdot \delta \mathbf{x} + \delta \mathbf{x}^T \cdot \mathbf{Q}_2 + \mathbf{Q}_3) \quad (\text{D5})$$

$$500 \quad \mathbf{Q}_1 = \mathbf{B}^{-1} + \mathbf{H}^T \cdot \mathbf{R}^{-1} \cdot \mathbf{H} \quad (\text{D6})$$

$$501 \quad \mathbf{Q}_2 = \mathbf{H}^T \cdot \mathbf{R}^{-1} \cdot (\hat{H}(\mathbf{x}_b) - \mathbf{y}) \quad (\text{D7})$$

$$502 \quad \mathbf{Q}_3 = (\hat{H}(\mathbf{x}_b) - \mathbf{y})^T \cdot \mathbf{R}^{-1} \cdot (\hat{H}(\mathbf{x}_b) - \mathbf{y}). \quad (\text{D8})$$

503 (Note that the covariance matrices and their inverses are symmetric, i.e.,  $\mathbf{R}^T = \mathbf{R}$ , etc.) The unconstrained minimisation  
 504 problem requires us to solve the equation  $\nabla J = \mathbf{Q}_1 \cdot \delta \mathbf{x} + \mathbf{Q}_2 = \mathbf{0}$ . Now we want to minimise the costfunction subject to the  
 505 the linear constraints

$$506 \quad \mathbf{G} \cdot \delta \mathbf{x} = \mathbf{0}, \quad (\text{D9})$$

507 where  $\mathbf{G}$  is a  $(k \times n)$ -matrix,  $\delta \mathbf{x}$  is an  $n$ -vector, and  $\mathbf{0}$  is the null-vector in  $\mathbb{R}^k$ . Let us denote the kernel<sup>10</sup> of  $\mathbf{G}$  by  $\ker(\mathbf{G})$ . Let  
 508 further  $\mathbf{z}_1, \dots, \mathbf{z}_{n-k}$  denote a basis of  $\ker(\mathbf{G})$ . We define the  $(n \times (n-k))$ -matrix

$$509 \quad \mathbf{Z} = \begin{pmatrix} \mathbf{z}_1 & \cdots & \mathbf{z}_{n-k} \end{pmatrix} \quad (\text{D10})$$

510 the column vectors of which are just the basis vectors of  $\ker(\mathbf{G})$ . Obviously,  $\mathbf{G} \cdot \mathbf{Z} = \mathbf{0}$ , where  $\mathbf{0}$  denotes the  $((k \times (n-k))$ -  
 511 nullmatrix. If  $\delta \mathbf{x}$  is a vector in  $\mathbb{R}^n$  for which there exists a vector  $\boldsymbol{\xi} \in \mathbb{R}^{n-k}$  such that  $\mathbf{Z} \cdot \boldsymbol{\xi} = \delta \mathbf{x}$ , then we automatically have  
 512  $\mathbf{G} \cdot \delta \mathbf{x} = \mathbf{0}$ , i.e.,  $\delta \mathbf{x}$  satisfies the linear constraints. Thus we can formulate the constrained minimisation problem by substitution  
 513 of  $\delta \mathbf{x} = \mathbf{Z} \cdot \boldsymbol{\xi}$  into Eq. (D5), which yields

$$514 \quad J = \frac{1}{2} (\boldsymbol{\xi}^T \cdot \mathbf{Z}^T \cdot \mathbf{Q}_1 \cdot \mathbf{Z} \cdot \boldsymbol{\xi} + \mathbf{Q}_2^T \cdot \mathbf{Z} \cdot \boldsymbol{\xi} + \boldsymbol{\xi}^T \cdot \mathbf{Z}^T \cdot \mathbf{Q}_2 + \mathbf{Q}_3) \quad (\text{D11})$$

$$515 \quad \mathbf{0} = \nabla J = \mathbf{Z}^T \cdot \mathbf{Q}_1 \cdot \mathbf{Z} \cdot \boldsymbol{\xi} + \mathbf{Z}^T \cdot \mathbf{Q}_2. \quad (\text{D12})$$

516 Thus we have reduced the  $(n+k)$ -dimensional constrained minimisation problem given in Eq. (D4) to a problem consisting of  
 517 the following two steps.

<sup>10</sup>The *kernel* or *nullspace* of a matrix is the set of all vectors  $\mathbf{z}$  such that  $\mathbf{G} \cdot \mathbf{z} = \mathbf{0}$ . The kernel is a subspace of the full vector space  $\mathbb{R}^n$  with  $\dim \ker(\mathbf{G}) = n - k$ .



518 1. Determine a basis of the nullspace  $\ker(\mathbf{G})$ ; this yields the matrix  $\mathbf{Z}$ .

519 2. Solve the unconstrained  $(n - k)$ -dimensional optimisation problem given in Eq. (D12). From the  $(n - k)$ -vector  $\boldsymbol{\xi}$  that  
520 minimises the costfunction in (D11), we then obtain the solution  $\delta\mathbf{x} = \mathbf{Z} \cdot \boldsymbol{\xi}$  that minimises the costfunction in (D5)  
521 subject to the constraint (D9).

## 522 D2 Minimisation of the costfunction with weak constraints

523 In the approach described in the previous section the solution satisfies the constraints exactly. Therefore, this approach is known  
524 as the minimisation of the costfunction with *strong constraints*. In the *weak-constraint* approach the constraints only need to  
525 be satisfied within specified error bounds.

526 The formulation of the weak-constraint approach is conceptually quite simple. One incorporates the constraints by adding  
527 an extra term to the costfunction (B7), i.e.

$$528 \quad J = J_b + J_o + J_G \quad (\text{D13})$$

$$529 \quad J_G = \frac{1}{2} \delta\mathbf{x}^T \cdot \mathbf{G}^T \cdot \mathbf{B}_G^{-1} \cdot \mathbf{G} \cdot \delta\mathbf{x}, \quad (\text{D14})$$

530 which also gives an extra term in the gradient of the costfunction,

$$531 \quad \nabla J_G = \mathbf{G}^T \cdot \mathbf{B}_G^{-1} \cdot \mathbf{G} \cdot \delta\mathbf{x}. \quad (\text{D15})$$

532 We will assume that the matrix  $\mathbf{B}_G = \text{diag}(\sigma_1^G, \dots, \sigma_k^G)$  is diagonal, where  $k$  is the number of constraints. The “error variances”  
533  $\sigma_i^G$  along the diagonal of  $\mathbf{B}_G$  allow us to fine-tune the influence of each constraint on the solution. If  $\sigma_i^G$  is small, then the  $i$ th  
534 constraint is relatively strong, and vice versa. The choice of these variances is a matter of experimenting and tuning. Typically,  
535 if the  $\sigma_i^G$  are made too large, then there is a risk that the minimisation algorithm ignores the constraints all together. In that case  
536 the solution will be very similar to the unconstrained solution. On the other hand, if the  $\sigma_i^G$  are made too small, then  $J_G$  can  
537 make the dominant contribution to  $J$ . In that case, there is a risk that the minimisation routine largely ignores the observations  
538 and returns a solution that lies quite close to the background estimate.

## 539 D3 Constraints designed for making optimum use of the information contained in the observations

540 We now want to incorporate the results of Section C into the variational data assimilation method. More specifically, we want  
541 to formulate weak constraints, Eq. (D14), based on the singular values of the observation operator in Eq. (C6). To this end,  
542 we make the change of variables given in Eq. (C16). We assume, without loss of generality, that the first  $\ell$  singular values are  
543 greater than unity. Thus we only want to use the corresponding components  $\delta x'_1, \dots, \delta x'_\ell$  as independent control variables in the  
544 3DVAR algorithm, while the remaining components remain unchanged, at least approximately within specified error bounds.



545 If we were to formulate this requirement as a strong constraint, as in Eq. (D9), then it would take the form

$$546 \quad \delta \mathbf{x}' = \mathbf{V}_R^T \cdot \mathbf{B}^{-1/2} \cdot \delta \mathbf{x} = \begin{pmatrix} \delta x'_1 \\ \vdots \\ \delta x'_\ell \\ 0 \\ \vdots \\ 0 \end{pmatrix}. \quad (\text{D16})$$

547 Thus the matrix expressing the constraints is given by  $\mathbf{G} = \mathbf{V}_R^T \cdot \mathbf{B}^{-1/2}$ , which is a  $(n \times n)$  matrix.

548 The weak constraint approach is, arguably, more suitable in our case. We have, in the preceding text, frequently used the  
 549 terms *signal degrees of freedom* and *noise degrees of freedom*. Although it was conceptually useful to make this distinction,  
 550 it is important to stress that there is no sharp boundary between the two. Rather, there is a smooth transition from singular  
 551 values  $w_1 > w_2 > \dots > w_\ell \geq 1$  to singular values  $1 > w_{\ell+1} > w_{\ell+2} > \dots > w_k$  ( $k = \min\{n, m\}$ ). For this reason we choose  
 552 to formulate the constraints as weak constraints. This allows us to make a smooth transition from free to constrained control  
 553 variables, where the transition from one regime to the other can be controlled by the singular values.

554 In order to apply the weak-constraint approach, we need to substitute the constraint-matrix  $\mathbf{G} = \mathbf{V}_R^T \cdot \mathbf{B}^{-1/2}$  into Eq. (D14),  
 555 which yields

$$556 \quad J_G = \frac{1}{2} \delta \mathbf{x}^T \cdot \mathbf{B}^{-1/2} \cdot \mathbf{V}_R \cdot \mathbf{B}_G^{-1} \cdot \mathbf{V}_R^T \cdot \mathbf{B}^{-1/2} \cdot \delta \mathbf{x}, \quad (\text{D17})$$

557 where  $\mathbf{B}_G$  is a  $(n \times n)$  matrix. We want to set up this matrix in such a way that we obtain a smooth transition from freely  
 558 adaptable control variables  $\delta x'_1, \dots, \delta x'_\ell$  to increasingly constrained variables  $\delta x'_{\ell+1}, \dots, \delta x'_k, \dots, \delta x_n$ . One possible choice of  
 559 the matrix  $\mathbf{B}_G$ , which is suggested by Eq. (C11), would be

$$560 \quad \mathbf{B}_G = \sigma_G \text{diag}(w_1^2, w_2^2, \dots, w_\ell^2, \dots, w_k^2, c, \dots, c), \quad (\text{D18})$$

561 where  $\sigma_G$  is a free scaling factor, and where the last  $n - k$  diagonal elements are equal to a constant  $c$  chosen to be much  
 562 smaller than  $w_k^2$ . Another possible choice would be

$$563 \quad \mathbf{B}_G = \sigma_G \cdot \text{diag}(\lambda_1, \lambda_2, \dots, \lambda_\ell, \dots, \lambda_k, c', \dots, c'), \quad (\text{D19})$$

$$564 \quad \lambda_i = w_i^2 / (1 + w_i^2), \quad (\text{D20})$$

565 where  $c' \ll \lambda_k$ . This ansatz is suggested by Eq. (C12), i.e., each of the elements  $\delta x'_1, \dots, \delta x'_k$  is weighted with its corresponding  
 566 contribution to the number of signal degrees of freedom. It turns out that Eq. (D18) gives a relatively sharp transition from  
 567 unconstrained to constrained model variables, while Eq. (D19) gives a very gentle transition. Another ansatz that lies in between  
 568 these two extremes would be

$$569 \quad \mathbf{B}_G = \sigma_G \text{diag}(w_1, w_2, \dots, w_\ell, \dots, w_k, c'', \dots, c''), \quad (\text{D21})$$



570 where  $c'' \ll w_k$ .

571 Despite the mathematical foundation of this approach we are left with some room for experimentation in the formulation of  
572 the matrix  $\mathbf{B}_G$ . It is a matter of experience to test different approaches and select the one that proves to be most suited.

## 573 Appendix E: Practical aspects of the implementation

574 One of the main practical problems is the dimension  $n$  of the model space. The grid-size is typically on the order  $N_x \times N_y \times$   
575  $N_z \sim 100 \times 100 \times 10$ , and the number of aerosol components is on the order of  $N_c \sim 10$ –100. Hence the dimension of the  
576 model space is  $n \sim 10^6$ – $10^7$ . In our case, the matrix  $\tilde{\mathbf{H}}$  in (C6) is a  $(m \times n)$  matrix. To numerically perform a singular value  
577 decomposition of such a large matrix would be a formidable task.

578 In variational data assimilation we encounter a similar problem in the inversion of the matrix  $\mathbf{B}$ . In our 3DVAR code this  
579 problem is alleviated by using a so-called spectral formulation. The idea is to make a Fourier-transformation in the horizontal  
580 coordinates and to assume that all horizontal error correlations are homogeneous and isotropic. Under these assumptions  
581 one obtains one background error covariance matrix for each horizontal wavenumber; each of these matrices has dimension  
582  $N_z \times N_c \sim 10^3$ – $10^4$ . This can further be reduced to about  $10^2$  by making a reduced eigenvalue diagonalisation. The details are  
583 explained in Kahnert (2008).

584 In our case we are primarily interested in constraining the chemical components. Therefore, the best solution for us is  
585 to simply restrict ourselves to the chemical subspace. To this end, we select a grid-point<sup>11</sup>  $(i, j, l)$  and consider the reduced  
586 background error covariance matrix  $\mathbf{B}^0$  with components  $B_{k;k'}^0 = B_{i,j,l,k;i,j,l,k'}$ , where  $k, k' = 1, \dots, N_c$ . Similarly, we consider  
587 reduced matrices  $\mathbf{R}^0$ ,  $\mathbf{H}^0$ , and  $\tilde{\mathbf{H}}^0 = (\mathbf{R}^0)^{-1/2} \cdot \mathbf{H}^0 \cdot \mathbf{B}^0$ , and we numerically compute the singular value decomposition of  
588 the latter. From this we obtain, for each level  $l$ , a constraint matrix  $(\mathbf{V}_R^0)^T \cdot (\mathbf{B}^0)^{-1/2}$  analogous to the one in Eq. (D16), but  
589 reduced to the chemical subspace. We then simply apply these chemical constraints throughout the horizontal domain (again,  
590 assuming horizontal homogeneity).

591 Another aspect concerns the positive square root of the background error covariance matrix, which appears in essential parts  
592 of the theory, namely, in Eqs. (C6) and (D16). In theoretical developments it is, arguably, didactically expedient to work with  
593 the matrix  $\mathbf{B}^{1/2}$ . But in practice there are numerically more efficient formulations. One such approach is discussed in Kahnert  
594 (2008) in the context of a spectral formulation of the variational method. In our present problem we employ the Cholesky  
595 decomposition<sup>12</sup> of the B-matrix,

$$596 \mathbf{B} = \mathbf{C}_u^T \cdot \mathbf{C}_u, \quad (\text{E1})$$

597 where  $\mathbf{C}_u$  is an upper triangular matrix. Thus the actual algorithm we used for formulating the constrained minimisation of  
598 the costfunction is obtained by replacing in the preceding formulas all incidences of the matrix  $\mathbf{B}^{1/2}$  with the matrix  $\mathbf{C}_u^T$  (and,  
599 similarly, by replacing the inverse matrix  $\mathbf{B}^{-1/2}$  by the inverse of the Cholesky factor,  $\mathbf{C}_u^{-T}$ ).

<sup>11</sup>Since the error correlations are assumed to be homogeneous in space any point in the horizontal direction will do.

<sup>12</sup>The Cholesky decomposition is, essentially, a special case of a LU-decomposition, which applies to symmetric real (or Hermitian complex), positive definite matrices.



600 *Author contributions.* MK worked with the theoretical developments and and numerical implementation, EA performed the testing of the  
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