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**Bromoform and  
dibromomethane in  
the tropics**

R. Hossaini et al.

# Bromoform and dibromomethane in the tropics: a 3-D model study of chemistry and transport

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Abstract

We have developed a detailed chemical scheme for the degradation of the short-lived source gases bromoform ( $\text{CHBr}_3$ ) and dibromomethane ( $\text{CH}_2\text{Br}_2$ ) and implemented it in the TOMCAT/SLIMCAT three-dimensional (3-D) chemical transport model (CTM).

The CTM has been used to predict the distribution of the two source gases (SGs) and 11 of their organic product gases (PGs). These first global calculations of the organic PGs show that their abundance is small. The longest lived organic PGs are  $\text{CBr}_2\text{O}$  and  $\text{CHBrO}$ , but their peak tropospheric abundance relative to the surface vmr of the SGs is less than 5%. We calculate their mean local tropospheric lifetimes in the tropics to be  $\sim 7$  and  $\sim 2$  days (due to photolysis), respectively. Therefore, the assumption in previous modelling studies that SG degradation leads immediately to inorganic bromine seems reasonable.

We have compared observed tropical SG profiles from a number of aircraft campaigns with various model experiments. In the tropical tropopause layer (TTL) we find that the CTM run using  $p$  levels and vertical winds from analysed divergence overestimates the abundance of  $\text{CH}_2\text{Br}_2$ , and to a lesser extent  $\text{CHBr}_3$ , although the data is sparse and comparisons are not conclusive. Better agreement in the TTL is obtained in the run using  $\theta$  levels and vertical motion from diabatic heating rates. Trajectory estimates of residence times in the two model versions confirm the more realistic transport in the  $\theta$ -level version. In the  $p$ -level model even when we switch off convection we still find significant amounts of the SGs considered may reach the cold point; the stratospheric source gas injection is only reduced by  $\sim 16\%$  for  $\text{CHBr}_3$  and  $\sim 2\%$  for  $\text{CH}_2\text{Br}_2$  without convection.

Overall, the relative importance of the SG pathway and the PG pathway for transport of bromine to the stratosphere has been assessed. Assuming a 10-day washout lifetime of  $\text{Br}_y$  we find the delivery of total Br from  $\text{CHBr}_3$  to be 0.72 pptv with  $\sim 53\%$  of this coming from SGI. Similarly, for  $\text{CH}_2\text{Br}_2$  we find a total Br value of 1.69 pptv with  $\sim 94\%$  coming from SGI. We infer that these species contribute  $\sim 2.4$  pptv of inorganic

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



bromine to the lower stratosphere with SGI being the dominant pathway.

## 1 Introduction

Bromine-containing very short-lived species (VSLS) are known to provide an additional supply of inorganic bromine ( $\text{Br}_y$ ) to the stratosphere. Emissions of such species are predominately of natural oceanic origin and have been shown to exhibit large variability, particularly in tropical coastal regions, that harbour substantial macroalgae (e.g. Quack and Wallace, 2003; Carpenter et al., 2009). It is also at tropical latitudes that deep convection allows the rapid ascent of such species from the marine boundary layer to the tropical tropopause layer (TTL). The rate of transport of these species, to and through, the TTL is currently poorly understood.

Current estimates of the contribution of inorganic bromine derived from VSLS ( $\text{Br}_y^{\text{VSLS}}$ ) to the stratospheric  $\text{Br}_y$  budget range from 3 to 8 ppt (Law and Sturges et al., 2007), with a more recent value of approximately 5 pptv derived by Dorf et al. (2008). Quantifying this additional source of bromine is important due its role in catalytic ozone depletion in the stratosphere. In addition, it is expected that reactive  $\text{Br}_y^{\text{VSLS}}$  will impact tropospheric composition (e.g. Von Glasow et al., 2004).

Two distinct pathways having been identified leading to the arrival of  $\text{Br}_y^{\text{VSLS}}$  in the stratosphere; namely source gas injection (SGI) and product gas injection (PGI) (e.g. Ko and Poulet et al., 2003). SGI refers to the transport of a source gas (SG, e.g. bromoform,  $\text{CHBr}_3$ ) to the stratosphere, where upon degradation will provide an in-situ source of  $\text{Br}_y$ . In contrast, PGI is the cross-tropopause transport of bromo-organic intermediates (e.g.  $\text{CBr}_2\text{O}$ ) and also inorganic products (e.g.  $\text{HBr}$ ,  $\text{BrO}$ ,  $\text{Br}$ ), produced from SG degradation in the troposphere. The efficiency of both SGI and PGI depends largely upon the photochemical loss of source gases (mainly via reaction with  $\text{OH}$  or photolysis) and rate of removal of degradation products (via wet deposition) versus the timescale for troposphere-stratosphere transport. For this reason, it is of interest to investigate the impact of convection on this additional source of bromine,  $\text{Br}_y^{\text{VSLS}}$ , in

### Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



model experiments.

Previous model work has concentrated on quantifying  $\text{Br}_y^{\text{VSLs}}$  from the more abundant very short-lived source gases such as  $\text{CHBr}_3$ . From a 2D model study, Dvortsov et al. (1999) concluded that  $\text{CHBr}_3$  contributes around 1 pptv additional  $\text{Br}_y$  to the lower stratosphere (LS). Similarly, Nielsen and Douglass (2001) also derived a value of 1 pptv from 3-D simulations with ~50% of this from SGI. In both of these studies the lifetime of  $\text{Br}_y$  following source gas degradation was assumed to be 10 days. More recently, Sinnhuber and Folkins (2006) used a 2-D mechanistic model of the tropical atmosphere to estimate that  $\text{CHBr}_3$  contributes between 0.8–2.1 pptv bromine via both SGI and PGI to the lower stratosphere with assumed  $\text{Br}_y$  lifetimes of 10–100 days. From this study, the contribution from SGI was approximately 0.5 pptv. The most recent model study, Kerkweg et al. (2008), confirmed earlier suggestions that  $\text{CHBr}_3$  contributes “substantial amounts” of  $\text{Br}_y$  to the lower stratosphere and that  $\text{Br}_y^{\text{VSLs}}$  should not be neglected in stratospheric modelling.

The impact of  $\text{Br}_y^{\text{VSLs}}$  in the stratosphere has been studied with multi-annual simulations by Feng et al. (2007) using the SLIMCAT 3-D chemical transport model (CTM). Salawitch et al. (2005) also performed a similar 2-D model study. Both studies report an ~10 DU decrease in the ozone column with an additional 6 pptv of  $\text{Br}_y^{\text{VSLs}}$  in the lower stratosphere. The impact of this additional bromine depends on the aerosol loading; the bromine causes a larger decrease in ozone when the aerosol loading is high and ClO is elevated.

The model studies discussed above have not directly considered the bromo-organic products (i.e. product gases, PGs) formed following source gas degradation. This omission is addressed in this study which evaluates the contribution of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  to the stratospheric Br budget, along with the relative contribution of SGI and PGI. Furthermore, results include novel estimates of the major and minor PGs formed following  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  removal. Vertical source gas profiles are compared to measurements made during several aircraft campaigns in the tropical troposphere and near-tropopause regions. To date, there have been no measurements of organic PGs

**Bromoform and dibromomethane in the tropics**

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



in the tropical atmosphere. We provide the first model estimates of the local lifetimes and abundances of these species in the tropical atmosphere.

Section 2 describes the derived chemical scheme for degradation of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$ . Kinetic and mechanistic assumptions are also discussed. Section 3 contains a description of the basic CTM setup, along with details of sensitivity runs carried out. Section 4 presents the model results. Conclusions and recommendations for future research are discussed in Sect. 5.

## 2 Chemistry scheme

Here we outline our chemical scheme to describe the tropospheric degradation of bromoform and dibromomethane. Reasonable mechanistic and kinetic assumptions have been made and are discussed. Kinetic data either is taken from Sander et al. (2006) (hereafter “JPL”) or the Leeds Master Chemical Mechanism (hereafter “MCM”, see <http://mcm.leeds.ac.uk/MCM/>). A summary of reactions and kinetic data used within the scheme is given in Tables 1 and 2 for  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$ , respectively.

### 2.1 Bromoform

The degradation of  $\text{CHBr}_3$  has been examined in previous theoretical studies (e.g. McGivern et al., 2002, 2004). Its local tropospheric lifetime is  $\sim 26$  days with photolysis being the dominant loss process (e.g. Ko and Poulet et al., 2003). Our bromoform scheme considers 7 organic species:  $\text{CHBr}_3$ ,  $\text{CBr}_3\text{O}_2$ ,  $\text{CHBr}_2\text{O}_2$ ,  $\text{CBr}_3\text{OOH}$ ,  $\text{CHBr}_2\text{OOH}$ ,  $\text{CBr}_2\text{O}$  and  $\text{CHBrO}$  and is summarised in Fig. 1. The following subsections discuss the details of this scheme.

#### 2.1.1 Removal of $\text{CHBr}_3$ source gas

The scheme assumes removal of  $\text{CHBr}_3$  occurs via reaction with OH/Cl radicals and also, more rapidly, by photolysis (Reactions R1–R3). The rates of reaction with OH

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

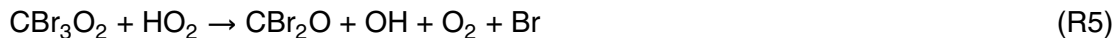


and Cl ( $k_1$ ,  $k_2$ ) are calculated using the JPL recommended temperature-dependent expressions. The rate of photolysis ( $j_3$ ) is calculated using JPL absorption cross section data along with a parameterisation for their temperature-dependence (Moortgat et al., 1993). The quantum yield for Br atoms following Reaction (R3) is assumed to be unity. It is also assumed that the immediate products of Reactions (R1–R3) ( $\text{CBr}_3$  and  $\text{CHBr}_2$ ) will be rapidly oxidised under tropospheric conditions.



### 2.1.2 Removal of peroxy species

The two peroxy radicals formed in Reactions (R1–R3) are assumed to be removed via reaction with NO and  $\text{HO}_2$  (Reactions R4–R9). Self-reaction of these species is deemed slow and is therefore not considered here. The  $\text{CBr}_3\text{O}_2 + \text{NO}$  Reaction (R4) is assumed to produce  $\text{CBr}_2\text{O}$ , an expected major product of bromoform degradation (Ko and Poulet et al., 2003). Excited intermediates, such as  $\text{CBr}_3\text{OONO}^*$  (not considered here), are expected to fragment rapidly to form  $\text{CBr}_3\text{O}$ , which itself would undergo a rapid decomposition to  $\text{CBr}_2\text{O}$  (e.g. McGivern et al., 2002). The rate constant for Reaction (R4) ( $k_4$ ) is calculated using the recommended JPL expression for the analogous species  $\text{CCl}_3\text{O}_2$ . Similarly, rate constants for the  $\text{CBr}_3\text{O}_2 + \text{HO}_2$  reactions ( $k_5$ ,  $k_6$ ) are assumed equal (i.e. equal branching ratio of products) and taken from the MCM. These reactions produce  $\text{CBr}_2\text{O}$  and the minor hydroperoxide product  $\text{CBr}_3\text{OOH}$ , respectively.



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

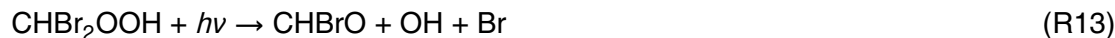


The reaction of  $\text{CHBr}_2\text{O}_2 + \text{NO}$  (Reaction R7) is assumed to produce a second major product of bromoform degradation, namely formyl bromide ( $\text{CHBrO}$ ). As for R4 it is likely that R7 would proceed via an excited intermediate (not considered here as sufficiently short-lived) such as  $\text{CHBr}_2\text{OOH}^*$ . The rate constant for this reaction ( $k_7$ ) is assumed equal to the analogous species  $\text{CHCl}_2\text{O}_2$  and taken from the MCM. This is also the case for the  $\text{CHBr}_2\text{O}_2 + \text{HO}_2$  reactions (Reactions R8, R9) which produce  $\text{CHBrO}$  and the minor hydroperoxide,  $\text{CHBr}_2\text{OOH}$  respectively.



### 2.1.3 Removal of minor end products

Removal of the two hydroperoxide species produced in Reactions (R6) and (R9) is assumed to be achieved via reaction with OH (Reactions R10, R12) and also by photolysis (Reactions R11, R13). Rate constants for the OH reactions are assumed equal to that of the analogous chlorine-containing species,  $\text{CCl}_3\text{OOH}$  and  $\text{CHCl}_2\text{OOH}$ , from the MCM. The photolysis rates are calculated using the absorption cross sections of methylhydroperoxide ( $\text{CH}_3\text{OOH}$ ). Reaction of these hydroperoxides with OH leads to the reformation of the respective peroxy radical (initially formed in Reactions R1 and R3).



### 2.1.4 Removal of major end products and Br<sub>y</sub>

Removal of the major products of bromoform degradation, CBr<sub>2</sub>O and CHBrO, is assumed to occur via photolysis (Reactions R14, R15). For CBr<sub>2</sub>O+hν we assume a yield of two Br atoms, with other photolysis pathways, such as that leading to HBr production, not considered. Similarly, for CHBrO+hν it is assumed the quantum yield for Br atoms is unity. Photolysis rates for both reactions are calculated using the recommended JPL cross section data.



In the current scheme all inorganic bromine species produced are grouped together as Br<sub>y</sub> without any further partitioning. Depending on the model run (see Sect. 3.2), Br<sub>y</sub> is removed in the troposphere by washout given a specified assumed lifetime.



## 2.2 Dibromomethane

The degradation of CH<sub>2</sub>Br<sub>2</sub> has also been examined in previous theoretical studies (e.g. McGivern et al., 2002, 2004). The local tropospheric lifetime is quoted as ~120 days with reaction with OH being the dominant loss process (Ko and Poulet et al., 2003). Our scheme considers six organic species, CH<sub>2</sub>Br<sub>2</sub>, CHBr<sub>2</sub>O<sub>2</sub>, CH<sub>2</sub>BrO<sub>2</sub>, CHBrO, CHBr<sub>2</sub>OOH and CH<sub>2</sub>BrOOH. The major products of CH<sub>2</sub>Br<sub>2</sub> degradation are expected to be CHBrO and Br<sub>y</sub> with CHBr<sub>2</sub>OOH being a minor product.

### 2.2.1 Removal of CH<sub>2</sub>Br<sub>2</sub> source gas

Our scheme assumes removal of CH<sub>2</sub>Br<sub>2</sub> is achieved via reaction with OH/Cl radicals and also, less rapidly, by photolysis (Reactions R17–19). The rates of reaction

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



with OH and Cl ( $k_{17}$ ,  $k_{18}$ ) are calculated using the JPL temperature-dependent expressions. The rate of photolysis ( $j_{19}$ ) is calculated using JPL absorption cross section data at 295–298 K. The quantum yield for Br atoms is assumed to be unity. As for the CHBr<sub>3</sub> scheme, it is assumed that following H abstraction and photolysis, the immediate products of source gas degradation (CHBr<sub>2</sub>, CH<sub>2</sub>Br) will be rapidly oxidised under tropospheric conditions forming associated peroxy radicals.



### 2.2.2 Removal of peroxy species

The two peroxy radicals formed in Reactions (R17–R19) are assumed to be removed via reaction with NO and HO<sub>2</sub>. Loss of CHBr<sub>2</sub>O<sub>2</sub> via these Reactions (R20–22) is treated as that described in Sect. 2.1.2 for the bromoform scheme.



The CH<sub>2</sub>BrO<sub>2</sub>+NO Reaction (R23) is assumed to produce NO<sub>2</sub> and Br<sub>y</sub>. The rate constant for this reaction ( $k_{23}$ ) is calculated using the recommended JPL expression. For reactions with HO<sub>2</sub> (Reactions R24, R25), rate constants are taken from the MCM. The products of these reactions are Br<sub>y</sub> and the hydroperoxide CH<sub>2</sub>BrOOH, respectively.



### 2.2.3 Removal of end products

Removal of CHBrO (Reaction R26), produced in Reactions (R20) and (R21), is achieved via photolysis as discussed in Sect. 2.1.3. Similarly removal of CHBr<sub>2</sub>OOH (Reaction R27, R28) produced in Reactions (R22) is analogous to removal in Reactions (R12) and (R13) (i.e. by reaction with OH and by photolysis respectively). Removal of CH<sub>2</sub>BrOOH, produced in Reaction (R25), has yet to be considered and is also assumed to be removed via reaction with OH (Reaction R29) and also by photolysis (Reaction R30). The rate constant for the OH reaction ( $k_{29}$ ) is taken from the MCM and the reaction products are expected to be the peroxy species CH<sub>2</sub>BrO<sub>2</sub> and water. The photolysis rate ( $j_{30}$ ) is calculated using the cross sections of CH<sub>3</sub>OOH.



## 3 Model and experiments

### 3.1 TOMCAT/SLIMCAT 3-D CTM

TOMCAT is an off-line 3-D CTM described in detail by Chipperfield (2006). The model has performed well in previous tropospheric studies and has been shown to simulate key chemistry and transport reasonably (e.g. Arnold et al., 2005). The model uses the Prather (1986) conservation of second-order moments advection scheme, a parameterisation of convection (Stockwell and Chipperfield, 1998) and also a parameterisation of boundary layer mixing (Holtslag and Boville, 1993). The CTM in TOMCAT mode uses a hybrid  $\sigma$ - $p$  vertical coordinate and diagnoses the large-scale vertical motion from divergence. The CTM has an option (“SLIMCAT”) for running with isentropic ( $\theta$ ) levels in the upper troposphere and stratosphere with the vertical motion calculated from heating rates. The SLIMCAT model only considers transport by large-scale advection; there

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



is no parameterisation of convection and boundary layer mixing. The CTM includes a scheme to calculate trajectories (e.g. Monge-Sanz et al., 2007).

### 3.2 Simulations

In all simulations described here (see Table 3) the resolution of the model was  $5.6^\circ \times 5.6^\circ$  with 38 (or 31) vertical levels ( $\sim 1$  km deep in mid troposphere) extending from the surface to  $\sim 35$  km. The model was forced using the European Centre for Medium-Range Weather Forecasts (ECMWF) 6-hourly analyses. The model was initialised on 1/1/2006 and run for 2 years. Year 1 was treated as model spin-up and year 2 output (2007) was saved every 3.75 days for analysis.

For the “base run” (run *B*), the TOMCAT model included specified oxidant fields along with the  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  degradation schemes described in Sect. 2. Monthly mean diurnal mean fields of the concentration of fixed species (e.g. OH, NO,  $\text{HO}_2$ ) were read from a previous TOMCAT full chemistry run for 2005. The background concentration of atomic chlorine, which was not calculated in the tropospheric full chemistry run, was set to  $1 \times 10^4$  molecules  $\text{cm}^{-3}$ . The model chemical scheme used a climatological tropical ozone profile for photolysis calculations (Chipperfield, 1999)

Figure 3 shows example tropical zonal mean profiles of temperature and the primary oxidant OH. The mixing ratio of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  source gases were fixed uniformly in space and time at 1.2 pptv in the bottom two layers of the model in the tropical regions ( $\pm 20^\circ$ ). This value seems reasonable for a typical surface mixing ratio of these SGs and furthermore provides the best fit to observed profiles (see Sect. 4). In this study we are interested in the relative mixing ratios of bromine species in the TTL compared to the surface and we do not need to introduce the complication of specifying uncertain emission fluxes. All other advected tracers were initialised at zero at the start of the simulation. The lifetime of  $\text{Br}_y$  was assumed to be infinite for run *B*. A number of sensitivity runs were also performed. Run  $S_{\text{NOCONV}}$  differed from run *B* in that model convection was switched off (note, mixing in the PBL remained switched on). Runs  $S_{10}$ ,  $S_{20}$  and  $S_{40}$  differed from run *B* in that the lifetime of  $\text{Br}_y$  was set to 10, 20

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Bromoform and dibromomethane in the tropics**

R. Hossaini et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

and 40 days below the cold-point tropopause (CPT), respectively. Run  $S_{2OH}$  differed from the base in that model  $[OH]$  was doubled. Run  $S_{L31}$  differed in that the model employed coarser vertical resolution (factor of 2) above  $\sim 300$  hPa and therefore had only 31 levels. Finally, a SLIMCAT simulation (run  $S_{SLIMCAT}$ ) was also carried out in which artificial mixing in the tropical troposphere was assumed by fixing both  $CHBr_3$  and  $CH_2Br_2$  SGs in the lower 8 levels of the  $\sigma - \theta$  model (surface to  $\sim 10$  km).

## 4 Results and discussion

### 4.1 Source gas injection

Figure 4 shows the mean modelled loss rates (due to reaction with OH and photolysis) and the local photochemical lifetimes for  $CHBr_3$  and  $CH_2Br_2$  in the tropics. Results from this analysis are also summarised in Table 4. For  $CHBr_3$ , the dominant loss process is photolysis. The calculated local lifetime of  $CHBr_3$  ( $\tau_{local}$ ) ranges between  $\sim 25$ – $30$  days in the TTL and has a surface value of  $\sim 15$  days. This is generally consistent with previous model calculations (e.g. Warwick et al., 2006; Sinnhuber and Folkins, 2006). For  $CH_2Br_2$ , in the mid-troposphere loss is dominated by reaction with OH with photolysis being slow. At the CPT, the two loss channels are roughly equal. The local lifetime of  $CH_2Br_2$  ranges from  $\sim 50$  days at the surface to a maximum of  $\sim 520$  days in the TTL. This is somewhat large given the working definition of a VSLS as a species whose lifetime is less than 6 months (Law and Sturges et al., 2007).

Figure 5 shows the tropical zonal mean profiles for the source gases  $CHBr_3$  and  $CH_2Br_2$  from the base run  $B$  and sensitivity runs  $S_{NOCONV}$  (no convection),  $S_{2OH}$  ( $2 \times [OH]$ ) and  $S_{SLIMCAT}$ . The location of the CPT and the approximate base of the TTL is shown for reference. Note that we define the base of the TTL as the level of maximum convective outflow (approximately 12 km) and the top of the TTL as the cold-point (e.g. Law and Sturges et al., 2007). The results here show that with the full TOMCAT model transport (run  $B$ ) the mean  $CHBr_3$  mixing ratio at the CPT

(~17 km) is ~0.126 pptv resulting in an SGI contribution of ~0.38 pptv of Br<sub>y</sub> to the lower stratosphere. This is in general agreement with the SGI value of 0.5 pptv predicted by Sinnhuber and Folkins (2006). Similarly for CH<sub>2</sub>Br<sub>2</sub>, run *B* predicts a mean of 0.797 pptv of SG at the tropopause, resulting in a value for SGI of ~1.59 pptv; i.e. over 4× that of bromoform. Without convection (run *S*<sub>NOCOCONV</sub>) the SGI values reduce to 0.32 pptv and 1.57 pptv, i.e. ~84% and ~98% of run *B* values for CHBr<sub>3</sub> and CH<sub>2</sub>Br<sub>2</sub>, respectively. These results suggest that SGI via both species is not overly sensitive to model parameterised convection, particularly at the CPT. For CHBr<sub>3</sub>, this apparently contradicts the findings of Nielsen and Douglass (2001) who report a treatment of convection is required in their model simulations in order for CHBr<sub>3</sub> to reach the tropical lower stratosphere. Similarly, Warwick et al. (2006) report from a 3-D model study CHBr<sub>3</sub> to be highly dependent on convection in the tropical upper troposphere. Without further details on the experiments performed in these other studies we cannot comment further on the differences. In our experiments, although we switch off convection we still include the parameterisation of mixing in the PBL which causes mixing of surface-emitted tracers in the bottom few km. If we also switch off PBL mixing then we see CHBr<sub>3</sub> especially largely confined to the lowest model level (~100 m) and a large decrease in UT values. It may be that the studies of Nielsen and Douglass (2001) and Warwick et al. (2006) also included transport due to this process in their definition of convection. Given that turbulent mixing in the PBL can transport tracers to the lower free troposphere, then the lifetimes of CHBr<sub>3</sub> and CH<sub>2</sub>Br<sub>2</sub> would indicate that resolved vertical advection by the analysed winds could still cause some transport to the UT.

There is uncertainty in the modelled OH profile in the UT and so a sensitivity run was performed to investigate the impact of a large (×2) change in [OH] on the modelled SG profiles. From the run *S*<sub>2OH</sub> profile in Fig. 5 it is clear the abundance of CH<sub>2</sub>Br<sub>2</sub> and associated SGI is more sensitive to changes in OH concentration than that of CHBr<sub>3</sub>. This is expected given the dominance of the CH<sub>2</sub>Br<sub>2</sub>+OH reaction over photolysis, relative to that of CHBr<sub>3</sub>+OH. For the SLIMCAT run *S*<sub>SLIMCAT</sub>, SG profiles have been scaled to approximately mimic that of the base run in the mid-troposphere. It can be

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**Bromoform and dibromomethane in the tropics**R. Hossaini et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

seen that SLIMCAT predicts a lower abundance of both SGs in the TTL and near-tropopause regions.

TOMCAT model runs with ECMWF winds have been reported previously to exhibit too rapid vertical motion in the lower stratosphere region (e.g. Chipperfield, 2006; Monge-Sanz et al., 2007). The problem is more obvious with ERA-40 winds than that with the more recent ECMWF datasets and is mainly related to the noisy analysed wind fields. This is known to affect all CTMs using wind velocities or divergence to obtain the vertical motion (e.g. Scheele et al., 2005; Wohltman and Rex, 2008). SLIMCAT runs are not affected by the same problem as, in this case, above 350 K vertical motion is computed from diagnosed heating rates. For this reason, the spurious vertical transport present in TOMCAT runs due to analysis noise is eliminated from SLIMCAT runs. In addition, SLIMCAT uses isentropic levels in the stratosphere which helps to separate vertical and horizontal motion and has also proven to provide more realistic transport in the LS than TOMCAT (e.g. Chipperfield, 2006).

Krueger et al. (2008) performed a Lagrangian model study in the TTL region using ECMWF operational winds and found that the use of the ECMWF vertical wind field resulted in significantly faster motion than the use of computed heating rates. A study by Wolthman and Rex (2008) with ECMWF winds has also shown improvements in the vertical velocities when obtained from diagnosed heating rates with respect to the vertical velocity field from the analyses (although we do not use the vertical velocity from the analyses).

The trajectory calculation inside the TOMCAT/SLIMCAT CTM has been used to estimate the mean tracer vertical transport in the TTL in runs  $B$  and  $S_{\text{SLIMCAT}}$ . Trajectories were initialised at 80 hPa (run  $B$ ) and 380 K (run  $S_{\text{SLIMCAT}}$ ) and advected backwards in time using the vertical winds from the analysed divergence field and diagnosed heating rates respectively. The mean vertical motion, calculated from the trajectory displacement, was 0.64 mm/s in run  $B$  and 0.324 mm/s in run  $S_{\text{SLIMCAT}}$ . Analysis of HCFC and HFC data from the WB-57 indicates a transit time from 360 K to 380 K of about 3–4 months (E. Atlas, personal communication, 2009). Analysis of CO<sub>2</sub> data (Park et

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



al., 2009) shows a similar transit time, which corresponds to a range in vertical transport, between the lower boundary of the TTL and the tropopause, of 0.5–0.14 mm/s. Krueger et al. (2008) find residence times in the winter TTL (2001/02) of 36 days using operational ECMWF with diagnosed heating rates, and only 20 days when using the corresponding ECMWF vertical winds. Our calculations show a residence time in the 360–380 K region of 20 days based on the TOMCAT run and 52 days based on the SLIMCAT run for the period November–December 2005. Therefore, the vertical motion in the  $\theta$ -coordinate model run is more realistic, which leads to overall better agreement with the observations (see below). The residence times in the TTL calculated here are within the range, 20–80 days, as published (WMO, 2007).

Overall, given the predicted mixing ratios of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  in the near-tropopause region, it seems the latter species may be deemed more significant. Wamsley et al. (1994) reported  $\text{CH}_2\text{Br}_2$  to have an atmospheric lifetime long enough to reach the stratosphere and  $\text{CHBr}_3$  (with shorter lifetime) to contribute negligible amounts to stratospheric Br. The results here are also consistent with Laube et al. (2008) who, based on observations, deduced  $\text{CH}_2\text{Br}_2$  to be the “dominant” very short-lived SG. Their results, along with Schaufli et al. (1998), find  $\text{CH}_2\text{Br}_2$  to be present up to  $\sim 18.5$  km ( $\sim 0.15$  pptv). The results here confirm  $\text{CH}_2\text{Br}_2$  at this level with run *B* predicting  $\sim 0.5$  pptv and run  $S_{\text{SLIMCAT}}$   $\sim 0.16$  pptv. The latter seems more reasonable given the observed values discussed above, along with the overestimation of SG in the near-tropopause region by TOMCAT (see below).

Figure 6 shows the modelled profiles of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  versus tropospheric observations from the PEM TROPICS-B (e.g. Colman et al., 2001; [http://www-gte.larc.nasa.gov/pem/pemtb\\_obj.htm](http://www-gte.larc.nasa.gov/pem/pemtb_obj.htm)) and INTEX-B (<http://www.espo.nasa.gov/intex-b/>) campaigns. The model profiles shown here are averaged over the same spatial domain and for the same months (but for 2007) as the observations. This shows that the model profiles from run *B*, constrained with a surface mixing ratio of 1.2 pptv for both species, fit the observations in the mid troposphere well. Particularly, for  $\text{CH}_2\text{Br}_2$ , the model is able to reproduce observed mixing ratios and profile shape. Note that the difference

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



between runs  $B$  and  $S_{\text{NOCONV}}$  here is small, showing that modelled convection is only playing a small role in this region.

Figure 7 shows the modelled profiles of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  against tropical observations which extend into the TTL. These are the 2007 NASA TC-4 (<http://www.espo.nasa.gov/tc4/>), the NASA PRE-AVE ([http://espoarchive.nasa.gov/archive/arcs/pre\\_ave/](http://espoarchive.nasa.gov/archive/arcs/pre_ave/)) campaign and the 2006 NASA CR-AVE (<http://www.espo.nasa.gov/ave-costarica2/>) campaigns. For TC-4 DC8 flights targetted recent convective outflow while only a single WB57 flight targetted convection. Again the model profiles are averaged over the same spatial domain and for the same months as the observations. For  $\text{CHBr}_3$ , the base model  $B$  performs reasonably well against the observations in the lower troposphere and near the tropopause. The model is able to reproduce the gradient seen from the surface to  $\sim 500$  hPa and the modelled profile lies within the min-max variability of observations at most levels. Of importance is the model's ability to simulate  $\text{CHBr}_3$  in the near-tropopause region. A number of previous model studies have significantly overestimated  $\text{CHBr}_3$  in this region when compared with observations (e.g. Warwick et al. 2006; Nielsen and Douglass, 2001). It can be seen that TOMCAT performs well in this regard. However, it seems that TOMCAT may not capture convection effectively shown by the lack of a signature "C-shape" in the profile. Results from TC4, which targetted active convective outflow during most flights, may not be representative of the region as a whole. Similarly for  $\text{CH}_2\text{Br}_2$ , the modelled profile seems reasonable against observations in the lower troposphere. There is less variability seen here than for  $\text{CHBr}_3$ , due to the longer lifetime of  $\text{CH}_2\text{Br}_2$ . However, in the near-tropopause region, TOMCAT overestimates the abundance of source gas. This will be in part due to the too fast modelled vertical transport through the TTL region in run  $B$  (see above). Note that the overestimation of the SGs in the TTL in TOMCAT is not due to the model vertical resolution. The SG profiles are nearly identical in run  $B$  and  $S_{L31}$  which has the higher vertical resolution (not shown). The vertical transport is controlled by the vertical winds and not by numerical diffusion.

From Fig. 7 it is apparent that run  $S_{\text{SLIMCAT}}$ , with slower vertical transport in the TTL,

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Bromoform and dibromomethane in the tropics**

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

reproduces observed  $\text{CH}_2\text{Br}_2$  in the near-tropopause region fairly well and better than the TOMCAT runs. Similarly for PRE-AVE flights, the run *B* profile fits  $\text{CHBr}_3$  observations well in the important upper TTL and tropopause regions. A larger difference is seen here between runs with and without convection along with more of a “C-shape” in the modelled profile. This is due to the model being sampled in the months January and February when convection is stronger. The  $\theta$ -level model in this case seems to underestimate observed  $\text{CHBr}_3$  in the TTL, although there are few observations at this level. For  $\text{CH}_2\text{Br}_2$ , both runs *B* and  $S_{\text{SLIMCAT}}$  seem reasonable in the near-tropopause region, with the former perhaps fitting the observations better. Unlike comparisons with TC4 and CR-AVE data, run *B* does not show a significant overestimation of  $\text{CH}_2\text{Br}_2$ . This could be explained by variation in the strength of deep convective uplift between campaigns. Finally, for CR-AVE data, modelled  $\text{CHBr}_3$  is overestimated in runs *B* and  $S_{\text{NOCONV}}$  in the TTL and near-tropopause regions. In these regions run  $S_{\text{SLIMCAT}}$  performs well. This is the also the case for  $\text{CH}_2\text{Br}_2$ , where the TOMCAT overestimation is greater.

Figure 8 shows a tracer-tracer plot of  $\text{CHBr}_3$  vs.  $\text{CH}_2\text{Br}_2$  in the 350–80 hPa region from runs *B*,  $S_{\text{NOCONV}}$ ,  $S_{2\text{OH}}$  and  $S_{\text{SLIMCAT}}$ . Mean observations from the TC-4 data set (also in this region, see Fig. 7a and b) are included on this figure along with a power line of best fit of all datasets. The origins of the model lines (high SG mixing ratios) are arbitrary and the plot tests the ability of the the different model runs to fit both SG profiles simultaneously. The model lines here indicate that with the current model setup (chemistry and transport), both tracers cannot be simulated correctly at the same time. The  $S_{\text{SLIMCAT}}$  run performs the best as can be seen in the gradient relative to that of the observations. Run  $S_{\text{SLIMCAT}}$  has the slowest transport through the TTL and allows more loss of  $\text{CH}_2\text{Br}_2$  relative to  $\text{CHBr}_3$  in the lower TTL where loss by OH is faster (Fig. 4).

## 4.2 Product gas injection

Figure 9 shows the annual tropical zonal mean abundance of the product gases arising from  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  degradation. For bromoform the major degradation products are  $\text{CBr}_2\text{O}$  and  $\text{CHBrO}$  with  $\text{CBr}_3\text{OOH}$  and  $\text{CHBr}_2\text{OOH}$  being minor products. This is consistent with the suggestions of Ko and Poulet et al. (2003). The mixing ratios of these species in the TTL are very low ( $<0.03$  pptv) for the assumed SG surface vmrs. From this we infer that the contribution of these species to PGI and thus total Br from  $\text{CHBr}_3$  is negligible. The mixing ratios of the peroxy radicals in the scheme,  $\text{CBr}_3\text{O}_2$  and  $\text{CHBr}_2\text{O}_2$ , were found to be near zero throughout the profile (not shown).

Similarly, for dibromomethane the model predicts the major degradation product to be  $\text{CHBrO}$  and a minor product to be  $\text{CHBr}_2\text{OOH}$ . Again, this is consistent with the suggestions of Ko and Poulet et al. (2003) and the mixing ratios of these species are also near-zero throughout the profile. As for  $\text{CHBr}_3$ , it is apparent that the contribution of organic products arising from  $\text{CH}_2\text{Br}_2$  SG degradation to PGI and thus total bromine, is negligible. Furthermore, although poorly quantified, organic PGs would be expected to be removed from the atmosphere by washout processes due to their solubility. The model work described in this paper has not accounted for this and thus the extremely low near-tropopause mixing ratios reported here for  $\text{CBr}_2\text{O}$ ,  $\text{CHBrO}$  and other PGs could indeed be overestimates. The results here suggest that PGI is dominated by the transport of inorganic products.

Figure 10 shows the calculated loss rates due to photolysis and the resultant lifetime of  $\text{CBr}_2\text{O}$  and  $\text{CHBrO}$  in the tropics. We find  $\text{CBr}_2\text{O}$  to have a lifetime of  $\sim 7$  days and  $\text{CHBrO}$  of  $\sim 1$  day due to photolysis. Overall, these results show that the assumption made in previous model studies of instantaneous conversion between organic bromine product gases and  $\text{Br}_y$  following SG degradation seems reasonable.

### Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



### 4.3 Total bromine

Figures 11 and 12 show the modelled contribution of SGI, PGI and total Br delivered from  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  degradation, respectively. Results are shown for sensitivity runs  $S_{10}$ ,  $S_{20}$  and  $S_{40}$  (i.e.  $\text{Br}_y$  lifetime of 10, 20 and 40 days). The base of the TTL and CPT are included for reference. Values of these fields are summarised in Table 5 and are quoted at the location of the CPT (i.e. the approximate contribution to the lower stratosphere). All results here are calculated as an annual zonal mean in the tropics. Assuming a  $\text{Br}_y$  lifetime of 10 days, our model predicts that  $\text{CHBr}_3$  contributes  $\sim 0.72$  pptv additional bromine to the lower stratosphere. We find the fraction of this value delivered via SGI and PGI to be approximately equal. This is in general agreement with the work of Dvortsov et al. (1999) and Nielsen and Douglass (2001) who report similar values of  $\sim 1$  pptv. These studies also find the delivery via SGI and inorganic PGI to be approximately equal. Similarly, the results here are in good agreement with the work of Sinnhuber and Folkins (2006) who report total bromine from  $\text{CHBr}_3$  reaching the cold point to be 0.8 pptv for the same assumed 10-day  $\text{Br}_y$  lifetime. In addition, they report a PGI value of 0.3 pptv which is consistent with the 0.35 pptv reported here.

From  $\text{CH}_2\text{Br}_2$  and with the same 10 day  $\text{Br}_y$  lifetime, we find a delivery of  $\sim 1.69$  pptv of bromine to the lower stratosphere, with  $\sim 94\%$  from SGI. In this case, the contribution from PGI is small. Despite only two Br atoms per molecule (as opposed to 3 for  $\text{CHBr}_3$ ), the dominance of the SGI pathway is due to the longer local lifetime of  $\text{CH}_2\text{Br}_2$ , allowing more SG to reach the upper troposphere. The results reported here for  $\text{CH}_2\text{Br}_2$  may constitute an upper limit given the observed overestimation of SG in the near-tropopause region in TOMCAT. Furthermore, we infer a total Br contribution from both SGs to be  $\sim 2.4$  pptv to the lower stratosphere. Increasing the assumed  $\text{Br}_y$  lifetime below the tropopause to 20 and 40 days raises this value to  $\sim 2.9$  and  $\sim 3.6$  pptv. The mean lifetime of  $\text{Br}_y$  in the troposphere and in particular the TTL region is, however, uncertain at present.

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## 5 Conclusions

We have performed a 3-D model study using the TOMCAT/SLIMCAT CTM in order to quantify, first, the contribution of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  to the stratospheric bromine budget and, second, the relative magnitude of SGI and PGI. A detailed chemical scheme describing the tropospheric degradation of both source gases, along with simplified product gas chemistry, has been developed.

We have thus provided novel estimates of the organic products gases arising from  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  degradation. The major degradation products have been found to be  $\text{CBr}_2\text{O}$  and  $\text{CHBrO}$  whose local lifetimes are calculated at  $\sim 7$  and  $\sim 2$  days respectively. We find their contribution to total bromine negligible and thus infer that assumption of instantaneous production of  $\text{Br}_y$  following  $\text{CHBr}_3/\text{CH}_2\text{Br}_2$  degradation in model studies is reasonable. It is likely that this assumption will also be valid for other short-lived source gases (e.g.  $\text{CHBr}_2\text{Cl}$ ,  $\text{CHBrCl}_2$ ) whose degradation products are comparably short-lived. However, attempt at measurements of species such as  $\text{CBr}_2\text{O}$  and  $\text{CHBrO}$  in the troposphere and TTL would certainly be useful.

The TOMCAT/SLIMCAT CTM has been shown to perform reasonably against observations of these SGs in the tropical troposphere-lower stratosphere region. The  $\sigma$ - $\theta$  level model (SLIMCAT) tends to agree better than the  $\sigma$ - $p$  level model (TOMCAT) due to the more realistic vertical transport in the TTL. The SLIMCAT run also gives the best simultaneous relative comparisons of the profiles of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$ . We find a treatment of convection is not required in our TOMCAT simulations to transport significant quantities of SG to the TTL and lower stratosphere. In future work the convective transport parameterisation will be included in SLIMCAT to investigate the impact of this against the background of slower resolved advection.

The results presented here have shown  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  together could contribute around 2.4 pptv of Br to the lower stratosphere when a  $\text{Br}_y$  lifetime of 10 days is assumed along with mean surface mixing ratios of 1.2 pptv for both source gases. Assuming the  $\text{Br}_y^{\text{VLSL}}$  value of 5 pptv inferred by Dorf et al. (2008), then a shortfall of  $\sim 2.6$

### Bromoform and dibromomethane in the tropics

R. Hossaini et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



pptv remains. This may, in part, be supplied from Br-containing SGs such as CH<sub>2</sub>BrCl, CHBr<sub>2</sub>Cl and CHBrCl<sub>2</sub> (local lifetimes of 150, 69 and 78 days respectively, Law and Sturges et al., 2007). Furthermore, although poorly quantified, it is possible “additional bromine source gases” which remain unknown may contribute (Laube et al., 2008).

5 Our results are generally in agreement with previous model work in that the contribution of bromine-containing VSLs may supply a significant amount of Br<sub>y</sub> to the lower stratosphere. Therefore, in future stratospheric simulations it will be important to take this into account.

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---

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

---

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



---

**Bromoform and dibromomethane in the tropics**R. Hossaini et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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**Bromoform and dibromomethane in the tropics**R. Hossaini et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

---

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

**Table 1.** Summary of bromoform chemistry scheme. **a** Photolysis rates calculated from (where available) recommended absorption cross section data. **b** Paramaterisation for the temperature-dependence of SG absorption cross sections included. **c** Br<sub>y</sub> first-order loss due to washout adjusted to give lifetimes of 10, 20, 40 and ∞ days. **d** Rate constant estimated from reactions of analogous chlorine-containing species or generalised expression.

	Reaction	Rate/cm <sup>3</sup> molecule <sup>-1</sup> s <sup>-1</sup>	Comment	Ref.
(R1)	CHBr <sub>3</sub> +OH	$k(T)=1.35E-12.exp(-600/T)$	–	JPL
(R2)	CHBr <sub>3</sub> +Cl	$k(T)=4.85E-12.exp(-850/T)$	–	JPL
(R3)	CHBr <sub>3</sub> + <i>hν</i>	<b>a, b</b>	Assume quantum yield of 1 for Br	JPL
(R4)	CBr <sub>3</sub> O <sub>2</sub> +NO	$k(T)=7.30E-12.exp(-270/T)$	<b>d</b>	JPL
(R5)	CBr <sub>3</sub> O <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	<b>d</b>	MCM
(R6)	CBr <sub>3</sub> O <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	As (R5) and equal partitioning of products	MCM
(R7)	CHBr <sub>2</sub> O <sub>2</sub> +NO	$k(T)=2.54E-12.exp(360/T)$	<b>d</b>	MCM
(R8)	CHBr <sub>2</sub> O <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	<b>d</b>	MCM
(R9)	CHBr <sub>2</sub> O <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	As (R8) and equal partitioning of products	MCM
(R10)	CBr <sub>3</sub> OOH +OH	$k(T)=1.90E-12.exp(190/T)$	<b>d</b>	MCM
(R11)	CBr <sub>3</sub> OOH + <i>hν</i>	<b>a</b>	Assume cross sections of CH <sub>3</sub> OOH	JPL
(R12)	CHBr <sub>2</sub> OOH+OH	$k(T)=1.90E-12.exp(190/T)$	<b>d</b>	MCM
(R13)	CHBr <sub>2</sub> OOH+ <i>hν</i>	<b>a</b>	Assume cross sections of CH <sub>3</sub> OOH	JPL
(R14)	CBr <sub>2</sub> O+ <i>hν</i>	<b>a</b>	Assume quantum yield of 2 for Br	JPL
(R15)	CHBrO+ <i>hν</i>	<b>a</b>	Assume quantum yield of 1 for Br	JPL
(R16)	Br <sub>y</sub> washout	<b>c</b>	Variable in model – see text	

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

**Table 2.** As Table 1 but for dibromomethane.

Reaction	Rate/cm <sup>3</sup> molecule <sup>-1</sup> s <sup>-1</sup>	Comment	Ref.
(R17) CH <sub>2</sub> Br <sub>2</sub> +OH	$k(T)=2.00E-12.exp(-840/T)$	–	JPL
(R18) CH <sub>2</sub> Br <sub>2</sub> +Cl	$k(T)=6.30E-12.exp(-800/T)$	–	JPL
(R19) CH <sub>2</sub> Br <sub>2</sub> + <i>hν</i>	<b>a</b>	Assume quantum yield of 1 for Br	JPL
(R20) CHBr <sub>2</sub> O <sub>2</sub> +NO	$k(T)=2.54E-12.exp(360/T)$	As Table 1 (Reaction R7)	
(R21) CHBr <sub>2</sub> O <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	As Table 1 (Reaction R8)	
(R22) CHBr <sub>2</sub> O <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	As Table 1 (Reaction R9)	
(R23) CH <sub>2</sub> BrO <sub>2</sub> +NO	$k(T)=4.00E-13.exp(300/T)$	<b>d</b>	MCM
(R24) CH <sub>2</sub> BrO <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	<b>d</b>	MCM
(R25) CH <sub>2</sub> BrO <sub>2</sub> +HO <sub>2</sub>	$k(T)=3.30E-13.exp(820/T)$	As (Reaction R24) and equal partitioning of products	MCM
(R26) CHBr <sub>2</sub> OOH+OH	$k(T)=1.90E-12.exp(190/T)$	As Table 1 (Reaction R12)	
(R27) CHBr <sub>2</sub> OOH+ <i>hν</i>	<b>a</b>	As Table 1 (Reaction R13)	
(R28) CH <sub>2</sub> BrOOH+OH	$k(T)=1.90E-12.exp(190/T)$	<b>d</b>	MCM
(R29) CH <sub>2</sub> BrOOH+ <i>hν</i>	<b>a</b>	Assume cross sections of CH <sub>3</sub> OOH	JPL
(R30) CHBrO+ <i>hν</i>	<b>a</b>	As Table 1 (Reaction R15)	
(R31) Br <sub>y</sub> washout	<b>c</b>	As Table 1 (Reaction R16)	

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

**Table 3.** CTM experiments performed.

Run	Convection	Br <sub>y</sub> lifetime	Vertical levels	[OH]
<i>B</i>	yes	Infinite	38	1×
<i>S</i> <sub>NOCOCONV</sub>	no	Infinite	38	1×
<i>S</i> <sub>10</sub>	yes	10 days below cold-point	38	1×
<i>S</i> <sub>20</sub>	yes	20 days below cold-point	38	1×
<i>S</i> <sub>40</sub>	yes	40 days below cold-point	38	1×
<i>S</i> <sub>2OH</sub>	yes	Infinite	38	2×
<i>S</i> <sub>L31</sub>	yes	Infinite	31	1×
<i>S</i> <sub>SLIMCAT</sub>	no	Infinite	38 ( $\sigma$ - $\theta$ )	1×

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

**Table 4.** Mean modelled lifetimes (days) of  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  in the tropics in different regions from the surface to cold-point ( $\sim 17$  km,  $\sim 95$  hPa). Lifetime with respect to photolysis ( $\tau_{hv}$ ), OH ( $\tau_{\text{OH}}$ ) and overall local lifetime ( $\tau_{\text{local}}$ ).

Source Gas	Pressure/hPa	$\tau_{hv}$ /days	$\tau_{\text{OH}}$ /days	$\tau_{\text{local}}$ /days
$\text{CHBr}_3$	1000	28	35	16
$\text{CHBr}_3$	1000–500	28	42	17
$\text{CHBr}_3$	500–200	25	132	21
$\text{CHBr}_3$	200–95	33	292	29
$\text{CH}_2\text{Br}_2$	1000	21004	53	52
$\text{CH}_2\text{Br}_2$	1000–500	20495	66	65
$\text{CH}_2\text{Br}_2$	500–200	13231	243	237
$\text{CH}_2\text{Br}_2$	200–95	2535	655	453

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

**Table 5.** Summary of modelled total Br, SGI and PGI for  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  with  $\text{Br}_y$  lifetimes of 10, 20, 40 and  $\infty$  days below the cold-point tropopause.

Source Gas	$\text{Br}_y$ Lifetime	SGI/ppt	PGI/ppt	Total Br/ppt	%SGI
$\text{CHBr}_3$	10 days	0.377	0.345	0.722	53
$\text{CH}_2\text{Br}_2$	10 days	1.594	0.099	1.693	94
<b>Total</b>	10 days	1.971	0.434	<b>2.415</b>	81
$\text{CHBr}_3$	20 days	0.377	0.732	1.109	34
$\text{CH}_2\text{Br}_2$	20 days	1.594	0.194	1.788	89
<b>Total</b>	20 days	1.971	0.92	<b>2.891</b>	68
$\text{CHBr}_3$	40 days	0.377	1.323	1.700	22
$\text{CH}_2\text{Br}_2$	40 days	1.594	0.334	1.928	82
<b>Total</b>	40 days	1.971	1.657	<b>3.628</b>	54
$\text{CHBr}_3$	$\infty$	0.377	3.168	~3.6	11
$\text{CH}_2\text{Br}_2$	$\infty$	1.594	0.763	~2.4	66
<b>Total</b>	$\infty$	1.971	3.931	<b>~6.0</b>	33

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

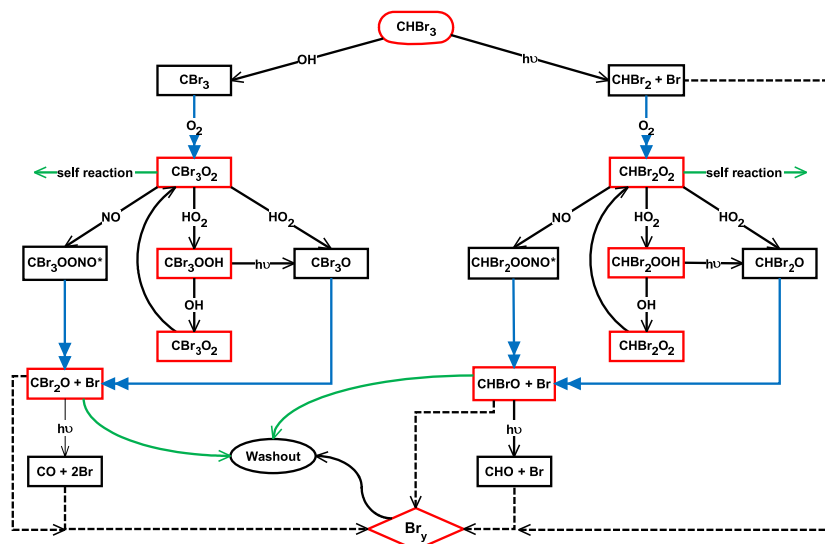
Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.



**Fig. 1.** Scheme for the tropospheric degradation of  $\text{CHBr}_3$ . The red boxes denote the 9 organic species (1 SG+6 PGs) considered directly in the scheme, i.e. species whose production and loss are both calculated explicitly (see text). Black boxes denote species which are sufficiently short-lived for exclusion from the scheme. Blue lines therefore denote rapid reactions, dashed lines indicate the production of  $\text{Br}_y$  and green denotes reactions not considered in the present scheme.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Bromoform and dibromomethane in the tropics

R. Hossaini et al.

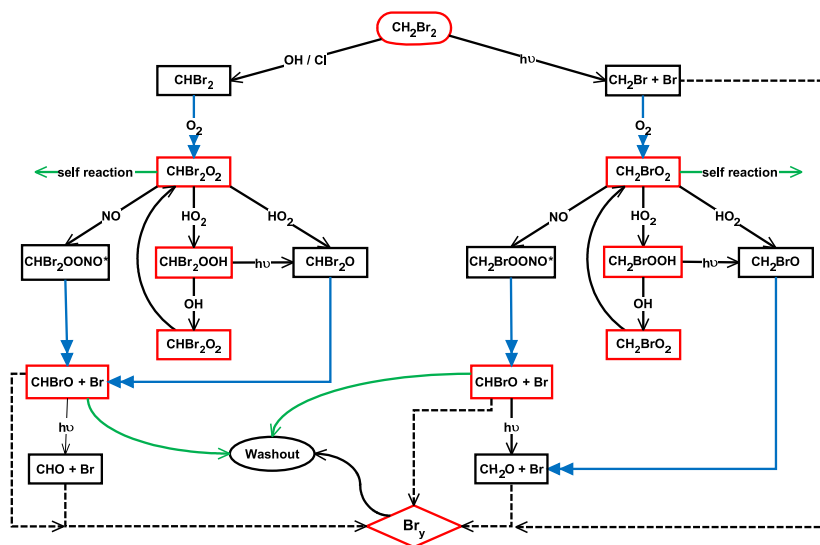


Fig. 2. As Fig. 1 but for  $\text{CH}_2\text{Br}_2$ .

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

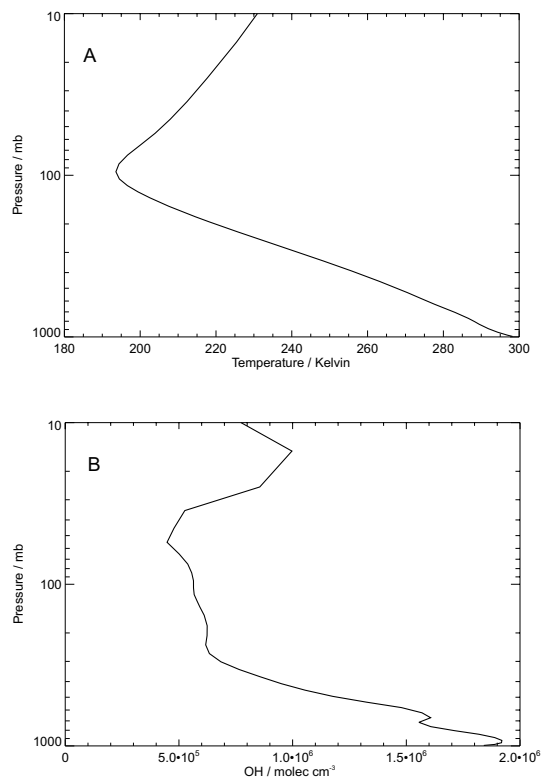
Printer-friendly Version

Interactive Discussion



**Bromoform and dibromomethane in the tropics**

R. Hossaini et al.

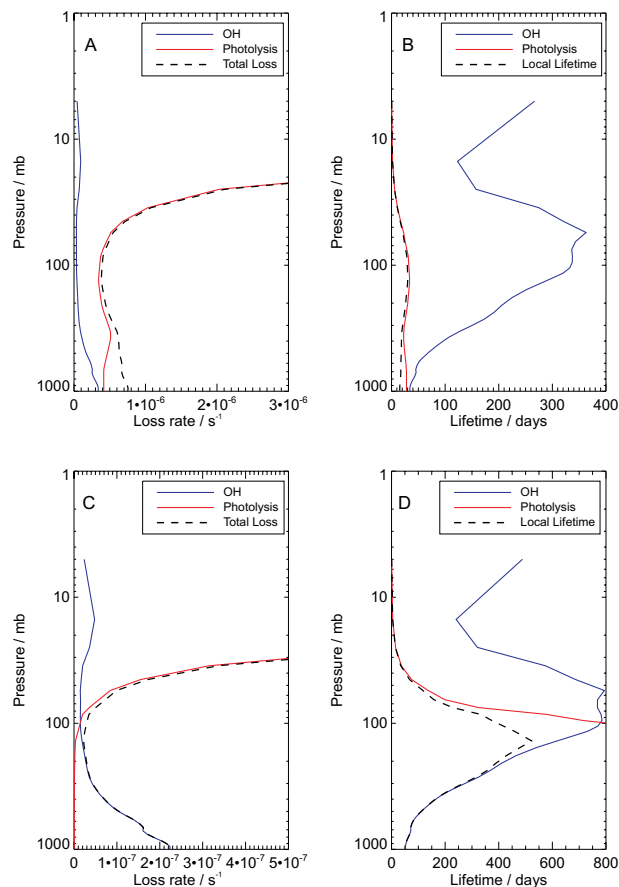


**Fig. 3.** Tropical ( $\pm 20^\circ$  latitude) annual mean zonal mean profiles of **(a)** temperature (K) from model run *B* (i.e. ECMWF analyses) and **(b)** specified (24-h mean) OH concentration profile used in CTM experiments.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Bromofrom and dibromomethane in the tropics

R. Hossaini et al.

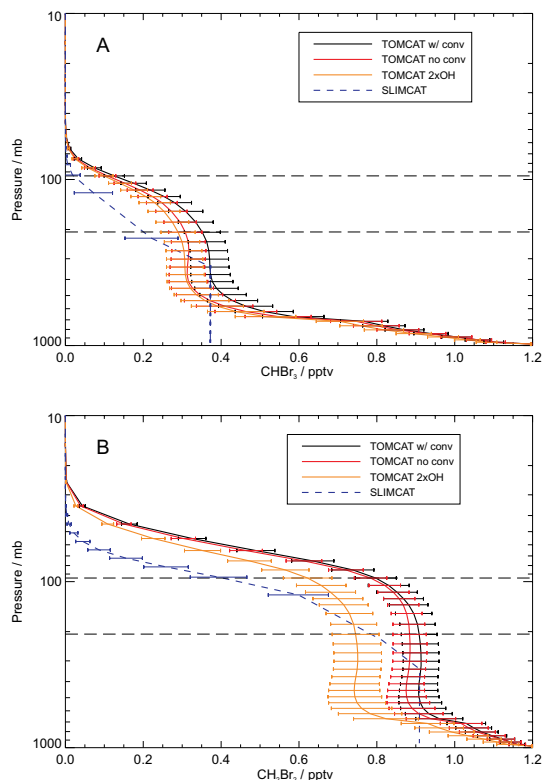


**Fig. 4.** Tropical ( $\pm 20^\circ$  latitude) annual mean zonal mean profiles of **(a)**  $\text{CHBr}_3$  loss rate ( $\text{s}^{-1}$ ) due to photolysis (red line), reaction with OH (blue line), and the overall total (black dashed line), and **(b)**  $\text{CHBr}_3$  lifetime (days) due to the same processes from the TOMCAT model runs. Panels **(c)** and **(d)** are the equivalent panels for  $\text{CH}_2\text{Br}_2$ .

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Bromofrom and dibromomethane in the tropics

R. Hossaini et al.



**Fig. 5.** Tropical ( $\pm 20^\circ$  latitude) annual mean zonal mean modelled profiles of **(a)**  $\text{CHBr}_3$  and **(b)**  $\text{CH}_2\text{Br}_2$ . Black line denotes output from the run *B* (i.e. with convection), red line denotes output from sensitivity run  $S_{\text{NOCONV}}$  (no convection), orange line from run  $S_{2\text{OH}}$  and dashed blue line from run  $S_{\text{SLIMCAT}}$ . Variability in time and space shown with  $\pm 1$  standard deviation. The location of the cold-point tropopause (CPT) and the base of the TTL (black dashed lines) are shown for reference.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

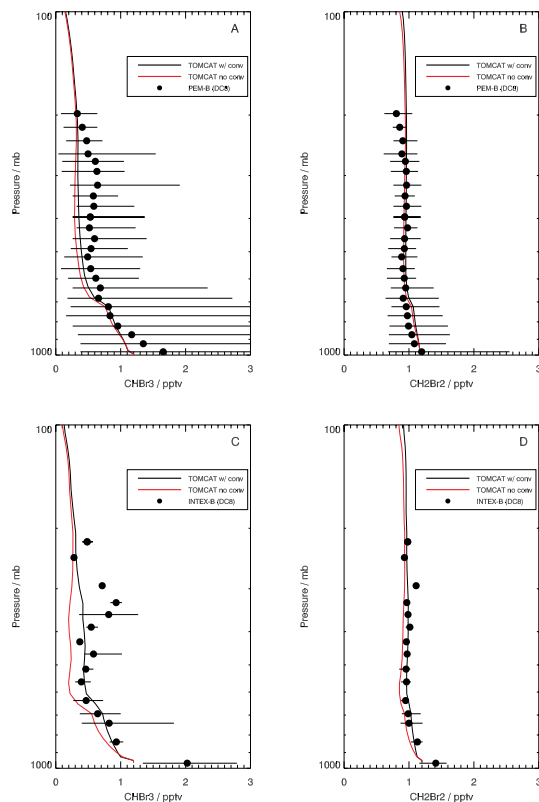
Printer-friendly Version

Interactive Discussion



## Bromofrom and dibromomethane in the tropics

R. Hossaini et al.

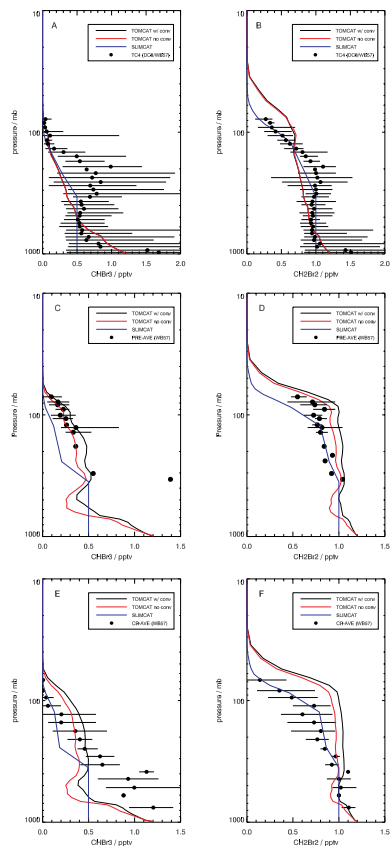


**Fig. 6.** Comparison of observed DC-8 profiles and model results for **(a)**  $\text{CHBr}_3$  and **(b)**  $\text{CH}_2\text{Br}_2$  from the PEM TROPICS-B mission (March-April 1999). The observations have been averaged vertically in  $\sim 0.5$  km bins for tropical latitudes and between  $84^\circ$  and  $211^\circ$  W longitude. The model (runs *B* and  $S_{\text{NOCONV}}$ ) is sampled over the same spatial domain and over the same months (for 2007). **(c)** and **(d)** are similar plots for the NASA INTEX-B campaign (March 2006). The observations have been averaged at tropical latitudes and between  $95^\circ$  and  $99^\circ$  W longitude. Horizontal lines on all observations denote min-max variability.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


## Bromoform and dibromomethane in the tropics

R. Hossaini et al.



**Fig. 7.** Comparison of observed DC-8 and WB57 profiles and model results for (a)  $\text{CHBr}_3$  and  $\text{CH}_2\text{Br}_2$  (b) from the NASA TC4 campaign (July–August 2007). (c) and (d) are similar comparisons for the PRE-AVE campaign (January–February 2004). (e) and (f) are similar comparisons for the CR-AVE campaign. (January–February 2006). The TC-4 observations have been averaged in tropical latitudes and between  $70^\circ$  and  $106^\circ$  W longitude. The PRE-AVE observations are averaged in tropical latitudes and between  $81^\circ$  and  $85^\circ$  W longitude. Similarly, the CR-AVE observations have been averaged at tropical latitudes and between  $79^\circ$  and  $86^\circ$  W longitude.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

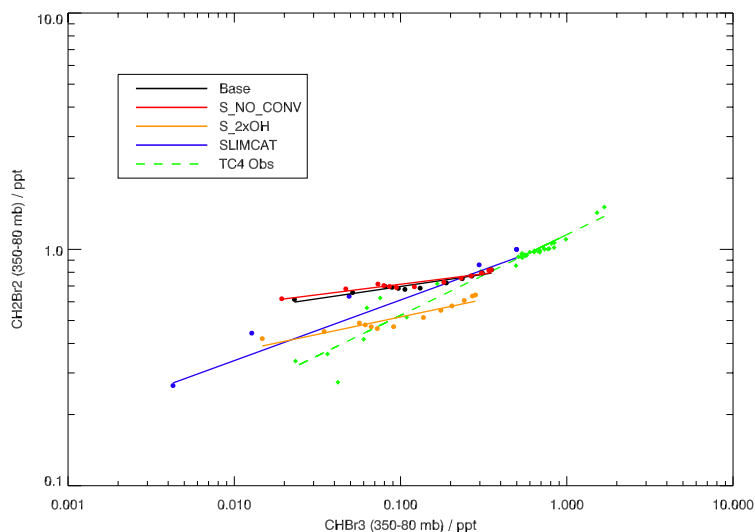
Printer-friendly Version

Interactive Discussion



**Bromoform and dibromomethane in the tropics**

R. Hossaini et al.

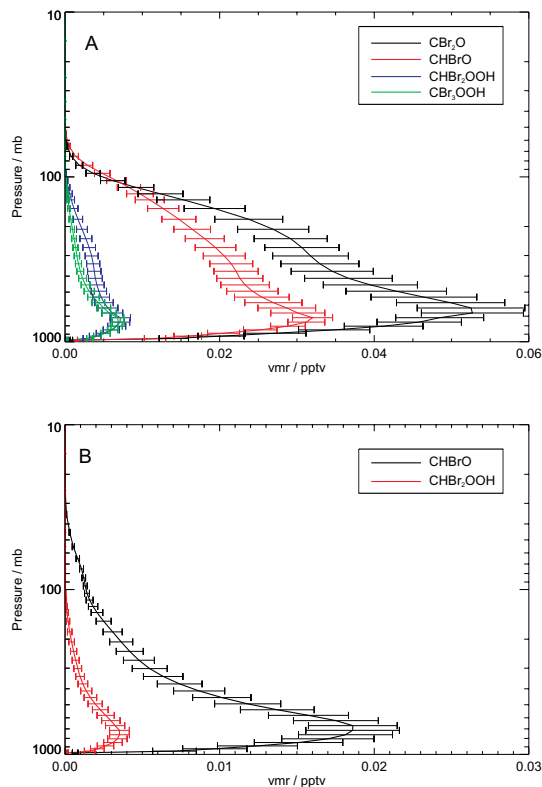


**Fig. 8.** Correlation plot of observed  $\text{CH}_2\text{Br}_2$  versus observed  $\text{CHBr}_3$  between 350 and 80 hPa from the TC-4 campaign (Fig. 7a). Also shown are model results from runs  $B$ ,  $S_{2\text{OH}}$ , and  $S_{\text{SLIMCAT}}$  in the same region. Power lines of best fit are included on all data sets of the form,  $Y=(aX)^b$ .

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**Bromofrom and dibromomethane in the tropics**

R. Hossaini et al.

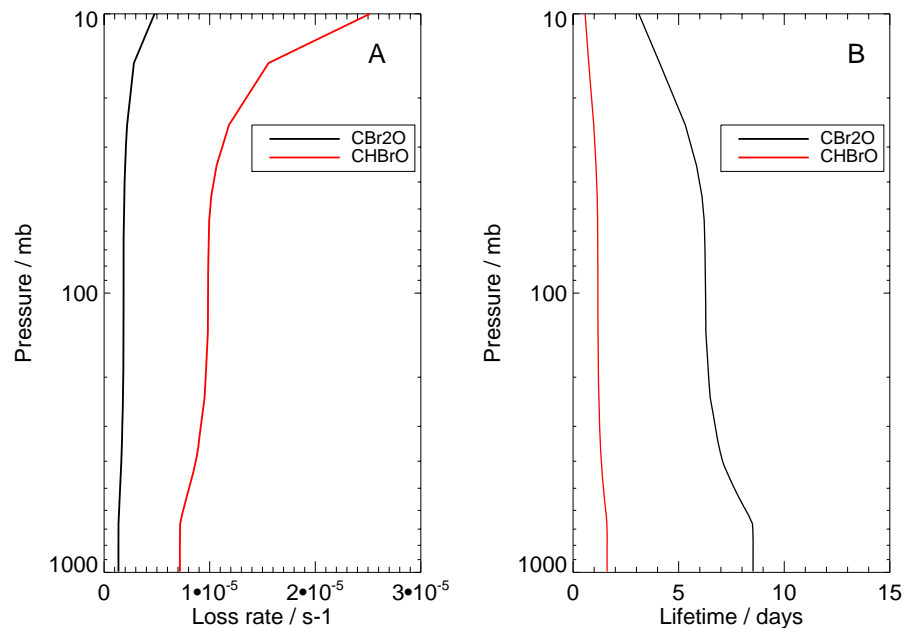


**Fig. 9.** Modelled tropical ( $\pm 20^\circ$  latitude) annual mean zonal mean profiles of the principal organic PGs arising from (a)  $\text{CHBr}_3$  and (b)  $\text{CH}_2\text{Br}_2$  degradation in the tropics. The assumed surface mixing ratio of both SGs is 1 pptv. The output is from run B with variability in time and space shown with  $\pm 1$  standard deviation.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**Bromofrom and dibromomethane in the tropics**

R. Hossaini et al.

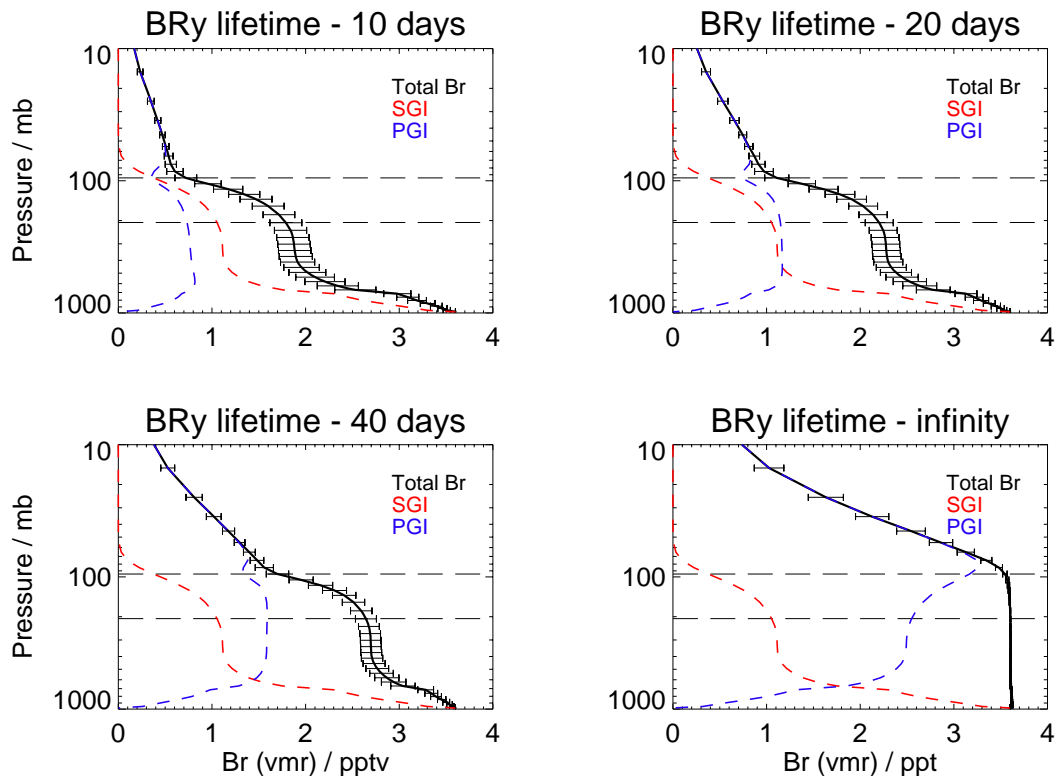


**Fig. 10.** Modelled tropical ( $\pm 20^\circ$  latitude) annual mean zonal mean profiles of **(a)** the loss rate (solely due to photolysis) and **(b)** the lifetime of  $\text{CBr}_2\text{O}$  (black line) and  $\text{CHBrO}$  (red line).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Bromoform and dibromomethane in the tropics

R. Hossaini et al.



**Fig. 11.** Modelled tropical ( $\pm 20^\circ$  latitude) annual mean zonal mean profiles of SGI, PGI and total Br arising from  $\text{CHBr}_3$  degradation in the tropics from runs (a)  $B$ , (b)  $S_{10}$ , (c)  $S_{20}$  and (d)  $S_{40}$ . These runs assume a  $\text{CHBr}_3$  surface mixing ratio of 1.2 pptv. The location of the cold-point tropopause (CPT) and base of the TTL (black dash lines) are shown for reference.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

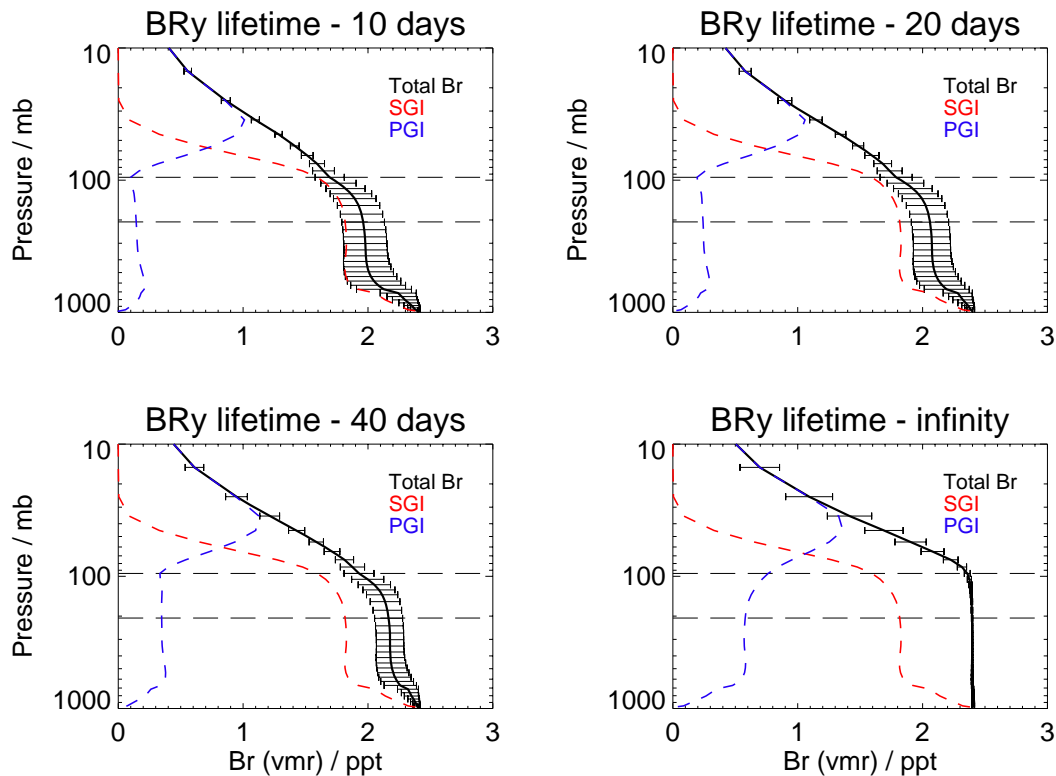
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Interactive Discussion



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**Fig. 12.** As Fig. 11 but for  $\text{CH}_2\text{Br}_2$ . The assumed surface mixing ratio of  $\text{CH}_2\text{Br}_2$  is 1.2 pptv.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)