

# Global temperature estimates in the troposphere and stratosphere: a validation study of COSMIC/FORMOSAT-3 measurements

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## Abstract

This paper mainly focuses on the validation of temperature estimates derived with the newly launched Constellation Observing System for Meteorology Ionosphere and Climate (COSMIC)/Formosa Satellite 3 (FORMOSAT-3) system. The analysis is based on the radio occultation (RO) data sample collected during the first year observation from April 2006 to April 2007. For the validation, we have used the operational stratospheric analyses (models) including the National Centers for Environmental Prediction-Reanalysis (NCEP-Reanalysis), the Japanese 25-year Reanalysis (JRA-25), and the United Kingdom Met Office (MetO) data sets. Comparisons done in different formats reveal excellent agreement between the COSMIC and model outputs. Spatially, the largest deviations are noted in the polar latitudes, and height-wise, the tropical tropopause region noted the maximum differences. However, these differences are only 2–4 K. We found that among the three models the NCEP data sets have the best resemblance with the COSMIC measurements. We also have done comparison of specific humidity and refractivity profiles with other measurements/models. Specific humidity profiles show comparatively large differences at altitudes below 5 km. Refractivity profiles derived by the COSMIC and other datasets show very good agreement.

## 1 Introduction

Temperatures are important in many aspects of atmospheric research and many studies have investigated the accuracy of derived temperature profiles (e.g., Hertzog et al., 2003; Pommerau et al., 2002; Knudsen et al., 2002). Knowledge of the upper tropospheric and lower stratospheric temperatures is important for understanding the structure and dynamics of the region, and is related to the issues connected to global climate change and the stratosphere-troposphere exchanges (Holton et al., 1995; Baray et al., 1998; Burris et al., 1998; Steinbrecht et al., 1998). A wide variety of observational techniques have been used to measure temperature in the troposphere and

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lower stratosphere and their variations in time and space. These techniques include balloon soundings (e.g., Kitamura and Hirota, 1989; Tsuda et al., 1991), rocketsonde measurements (e.g., Dewan et al., 1984; Hamilton, 1991; Eckermann et al., 1995), and lidar studies (e.g., Wilson et al., 1991; Hauchecorne and Chanin, 1980; Chanin and Hauchecorne, 1991; Whiteway and Carswell, 1994; LeBlanc et al., 1998; Sivakumar et al., 2003) etc. Most of these observations are mainly over land areas of the Northern Hemisphere (NH) and the period of observations has been limited. This raises the problem of insufficient data for establishing climatology for a global scale, despite the good results from many of the ground-based and space-borne instruments (Steiner and Kirchengast, 2000). Recently Global Positioning Systems (GPS) Radio occultation (RO) observations started offering several important and unique features complementary to other methods of observing the Earth's atmosphere. The GPS-RO technique has the advantages of global coverage, high accuracy, high vertical resolution (less than 1 km), long-term stability, self-calibration and capability to operate in all-weather conditions (Kursinski et al., 1997; Wickert et al., 2001; Hajj et al., 2002).

The GPS-RO technique has been very useful in collecting the global high-resolution data sets of atmospheric refractivity, temperature, pressure, and water vapour profiles. Using the GPS/MET data, initial investigations of the temperature variations in the tropical tropopause region were performed by Nishida et al. (2000) and Randel et al. (2003). The CHAMP and SAC-C profiles occurring within 30 min and 200 km are compared and agreed to better than 0.86 K and to within 0.1 K in the mean between 5 and 15 km altitude, after removing the expected variability of the atmosphere (Hajj et al., 2004). Wang et al. (2004) compared more than 100 temperature profiles collected using the Michelson Interferometer for Passive Atmospheric Sounding (MIPAS) and the Challenging Mini-Satellite Payload (CHAMP) during 14 days in September/October 2002. They found good agreement in the upper troposphere and lower stratosphere.

The US/Taiwanese Constellation Observing System for Meteorology Ionosphere and Climate (COSMIC/FORMOSAT-3) satellite was successfully launched on 15 April 2006. This mission is expected to bring the use of the innovative use of GPS RO data for

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atmospheric and ionospheric research, for improving global weather forecasts and climate change related studies. A distinctive feature of the COSMIC mission, compared to previous RO missions, is tracking both setting and rising neutral atmospheric occultations in the lower troposphere in an open-loop (OL) mode (Schreiner et al., 2007). This is very important for detecting the moisture variation in the lower troposphere, which is very important for weather prediction and climate analysis.

In this paper, we use COSMIC and operational stratospheric analyses data sets (hereafter models) for the validation of upper troposphere and lower stratosphere temperature profiles. The model data sets include those implemented at the National Centers for Environmental Prediction-Reanalysis (NCEP-Reanalysis), the Japanese 25-year Reanalysis (JRA-25), and the UK Met Office (MetO). In the next section we provide some details of the data analysis. Results obtained from the present analysis in terms of comparison between COSMIC satellite and models are described in Sect. 3. We describe the global temperature behaviour noted by each measurement and their differences of temperatures with respect to the COSMIC, for summer and winter seasons for particular pressure levels. Comparison of the latitudinal temperature has been carried out for different pressure levels. We also made an attempt to compare the height profiles of specific humidity and refractivity derived by the COSMIC, model data sets, and TES-AURA and MLS-AURA satellite measurements. Finally, the conclusions of the present analysis are summarized in Sect. 4.

## 2 Data analysis

### 2.1 (a) COSMIC

GPS-RO measurements are active limb sounding measurements of the Earth's atmosphere. This technique has emerged as a promising approach for global atmospheric measurements. Under the assumption of geometric optics and local spherical symmetry of the atmosphere, the phase delay measurements can be directly inverted to yield

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the index of refraction profile with vertical resolution that varies from about 50 m in the lower troposphere to about 200–300 m in the lower stratosphere. As mentioned previously, the COSMIC satellite system (Rocken et al., 2000), a constellation of 6 LEO micro-satellites, was put into orbit on 15 April 2006. Since 21 April 2006, GPS-RO technique has been providing accurate and high vertical resolution global data sets of atmospheric parameters. In the present study, we used the COSMIC RO data collected during the one year period from 21 April 2006 to 30 April 2007.

Figure 1 shows the histogram of number of occultations (per day) made by COSMIC soundings that cover the period from 21 April 2006 to 30 April 2007. The total number of occultations is 425 797 during this period of observation. During the initial months the daily occultations were less than 400, and it has increased enormously, from August 2006, to about 1400–1600 soundings within the height interval 0–40 km. The maximum number of occultations (per day) is observed on 4 March 2007 and is 2493.

## 2.2 (b) NCEP-Reanalysis (NCEP)

The reanalysis data set was created through the cooperative efforts of the National Centers for Environmental Prediction (NCEP) and National Center for Atmospheric Research (NCAR) to produce relatively high-resolution global analyses of atmospheric fields over a long time period (Kalnay et al., 1996; Kistler et al., 2001). Data from many different sources (including rawinsondes, balloons, aircrafts, ships, surface stations and satellites) were put through a quality check, fed into an assimilation model that includes parameterizations for all major physical processes, and finally examined again for self consistency. The data sets contain daily averages of geopotential height, wind, and temperature on 17 pressure levels above the surface (from 1000 hPa to 10 hPa) with spatial resolution of  $2.5^{\circ} \times 2.5^{\circ}$ , with a resulting horizontal grid of  $144 \times 73$  points.

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## 2.3 (c) JRA-25 Reanalysis (JRA-25)

The Japanese 25-year reanalysis (JRA-25) is an up-to-date data set produced by the Japan Meteorological Agency (JMA) and the Central Research Institute of Electrical Power Industry (CRIEPI). The main components of the system are a spectral forecast model, various quality control (QC) processes, a three dimensional variational (3D-Var) data assimilation process, and a land surface model. JRA-25 reanalysis is a basic meteorological grid point data set with uniform resolution of 120 km in the horizontal and from the surface to about 50 km in the vertical. The output contains over 200 meteorological parameters globally with 23 vertical levels from 1000 hPa to 0.4 hPa and the resolution of  $2.5^{\circ} \times 2.5^{\circ}$  intervals in longitude/latitude wise. A 3D-Var data assimilation system is adopted for assimilating various kinds of satellite data effectively (Onogi et al., 2005, 2007).

## 2.4 (d) UK Met Observations (MetO)

A data assimilation system has been developed at the UK Meteorological Office (UKMO) to analyze the various observations available in the troposphere and stratosphere. The data assimilation system is based on the analysis correction scheme used at the UKMO for operational weather forecasting (Swinbank and O'Neill, 1994). The analysis consists of temperatures, wind components, and geopotential heights on a global grid of resolution  $2.5 \times 3.75^{\circ}$  (latitude/longitude) with 22 vertical pressure levels from 1000 to 0.32 hPa corresponds to the altitude range of 0–55 km. The aim of the assimilation project is to provide a self-consistent sequence of three-dimensional grid point analyses of the atmosphere, up to the lower mesosphere. The sequence of synoptic analyses is constructed from a heterogeneous mixture of observational data. Since November 2000, the MetO stratospheric analyses have been produced using a new three-dimensional variational data assimilation system (Lorenc et al., 2000).

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### 3 Results and discussion

In this section we focus on comparison of COSMIC temperature profiles with the NCEP, MetO and JRA-25 reanalysis data sets. Prior to the comparisons with models, for each RO profile a coinciding profile was extracted from the model values, i.e., by spatially interpolating the model values of nearest coordinates (time and locations) to the locations of the RO soundings.

In order to note the global coverage of COSMIC RO, an example of the global occultations for the NH summer season is shown in Fig. 2, where the total number of occultations is 49 302. Here the NH summer season represents the period covered from June to August. In Fig. 2 (right panel) we also show the histogram of number of occultation occurred at every  $10^\circ$  latitude range. It is evident in the figure that the data coverage is very good for the latitude sector  $80^\circ$  S– $80^\circ$  N. The coverage in the equatorial region is lesser than the midlatitude regions. The near Polar regions ( $80^\circ$ – $90^\circ$ ) are marked with very low coverage. The summer occultation coverage at the midlatitudes of the NH and SH is more than 3000 occultations, whereas the equatorial region shows a comparatively smaller number (little more than 1000). We have examined the winter coverage and found that the number of occultation is more than the summer occultations.

#### 3.1 (a) Mean temperature

First, we compare the mean temperature profiles retrieved from COSMIC RO measurements with models. Figure 3 shows the height profiles of zonal-mean temperatures for five different latitude bands during the one week period 10 April 2007 to 16 April 2007. Here we present the data sets of COSMIC, NCEP, JRA-25, and MetO. To determine the mean temperatures, all the available longitudinal data at individual height are sorted into latitude bins of a specified width. The selected latitudinal ranges cover the high-, mid- and low-latitude regions in both Northern and Southern Hemispheres, and they are  $50^\circ$  N– $80^\circ$  N,  $20^\circ$  N– $50^\circ$  N,  $20^\circ$  N– $20^\circ$  S,  $20^\circ$  S– $50^\circ$  S, and  $50^\circ$  S– $80^\circ$  S. In

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Fig. 3 in addition to the mean temperature profiles (left sub-panel) we also show the differences (right sub-panel) with respect to the COSMIC temperatures. In general, comparison of the COSMIC and model temperatures reveals very good agreement. There exists some difference at altitudes below 8 km, which is due to interference from water vapor existence and thus temperature retrieval is incomplete at those heights. Also the correction due to ionospheric residuals will create issues at altitudes above 45 km, and sometimes even from 35 km and above (Rocken et al., 1997; Syndergaard 2000). Hence, we have restricted our comparisons to the height range between 8 and 30 km. The mean differences between COSMIC and models at most of the individual heights vary between 1–2 K, and the mean difference averaged over the height region 8–30 km is noted as values between  $-0.75$  and  $0.9$  K for all the five latitude bands. Larger deviations are seen in the equatorial ( $20^\circ$  S– $20^\circ$  N) troposphere region, where the data sets exhibit differences of about 2.5–4 K. It is not clear whether these differences are due to larger uncertainties in the COSMIC data or based on the performed analyses. The enhanced gravity wave activity near the equator could also introduce larger differences. These differences are expected to be significantly reduced when large data sets are averaged. Using the CHAMP data, it has been shown that the increase of the tropical bias and root mean square (rms) deviation is related to gravity waves (Marquardt et al., 2002). Comparing the models, the NCEP model shows the least difference than all other models. The figure demonstrates positive differences in the troposphere and negative differences in stratosphere heights indicating warmer and colder temperatures by COSMIC. The warming is larger in the upper troposphere than in the lower troposphere is due to the moist adiabatic lapse rate decreasing with the increasing temperatures and lower stratosphere the cooling due to increased longwave emission is to a large  $^\circ$  offset by increased absorption of upwelling longwave radiation from the troposphere (Ringer and Healy, 2008). Schmidt et al. (2004) compared the CHAMP temperature data with nearby radiosonde observations and they found that temperature bias is less than 0.5 K in the upper troposphere and lower stratosphere region. The cold bias above 20 km is related to the bending angle smoothing and the

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interpolation procedure for the downward integration of the refractivity profile. Recently, Kishore et al. (2006) made statistical estimates of the bias between the CHAMP, SAC-C, and radiosonde temperature measurements obtained for different latitude sectors during the period from March 2002 to December 2003. Their results demonstrated that the mean bias in the troposphere is less than 0.7 K and it is comparatively larger in the stratosphere (between 0.9–1.5 K).

### 3.2 (b) Longitudinal – Latitudinal temperature variation

Here we would like to describe the spatial structure of the temperatures derived from COSMIC RO measurements and comparison with the models JRA-25, MetO, and NCEP during the NH summer (June, July, and August) of 2006. Figure 4a–d represents the temperature contours (at 100 hPa) derived by the COSMIC, JRA-25, MetO and NCEP data sets, respectively, during 2006 summer. The data used in this study have been obtained on standard pressure levels, and interpolated onto a  $5^{\circ} \times 5^{\circ}$  longitude-latitude grid. Describing the features, it can be seen that the temperature structure of the COSMIC and models looks similar. The northern Polar region temperature is close to 228 K ( $-45^{\circ}\text{C}$ ) whereas the temperature at southern Polar region is about 28 K lower (Note that the northern summer corresponds to the southern winter.). In southern midlatitudes ( $35^{\circ}\text{S}$ – $45^{\circ}\text{S}$ ) the maximum temperatures is about 216 K for COSMIC and model observations. The tropical minimum temperature is about 194 K ( $-79^{\circ}\text{C}$ ), which is observed over the eastern Pacific. Therefore the tropical tropopause is expected to occur at about 16 km. Hoinka (1998) reported that for all latitudes the temperature at 100 hPa level provides a first estimate of tropopause temperature.

The global temperature structure generated by the COSMIC and model values for the northern winter (December 2006, January and February 2007) at 100 hPa pressure level is shown in Fig. 5a–d. Looking at the temperature pattern at the Polar regions we can see that the Arctic temperature is about 206 K, whereas it is close to 228 K in the Antarctic region. In the tropics, temperature reaches about 192 K ( $-81^{\circ}\text{C}$ ), which is observed in the western Pacific. Newell et al. (1972) had reported that the largest region

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of lowest temperatures (colder than  $-82.5^{\circ}\text{C}$  in the seasonal mean) at 100 hPa occurs over the western Pacific Ocean in winter. Secondary regions of limited extent appear over Northern South America and Africa. Newell and Gould-Stewart (1981) presumed a significant water vapor transport from the troposphere into the stratosphere in these regions and it is most active in the west Pacific in winter and moves to the monsoon region for the NH summer. Both COSMIC and model noticed the NH maximum temperatures of  $\sim 224\text{ K}$  at  $50^{\circ}\text{ N}$ . Concluding the summer and winter temperature features observed in the contours, it can be stated that the structure generated by the COSMIC and models has encouraging similarity, which is further established with the temperature differences explained below.

The longitude-latitude structure of the temperature differences between COSMIC and model data sets at 100 hPa level are shown in Fig. 6 and Fig. 7, for summer and winter, respectively. In each figure, (a) corresponds to COSMIC–JRA-25, (b) denotes COSMIC–MetO, and (c) represents the COSMIC–NCEP data sets. Examining the Fig. 6 we can see that the difference between the COSMIC and other measurements is not very large. Maximum deviation of 2 K is observed in limited areas. Such deviations are generally observed in the high latitudes in the Southern Hemisphere (SH). It is evident that the COSMIC and the NCEP model values match better than the other two pairs. Note that the solid line denotes warmer COSMIC temperature and dotted line denotes colder COSMIC temperature. For a large latitude/longitude sector, the summer temperature difference is about 0.75–1.25 K. Schmidt et al. (2004) compared the CHAMP with ECMWF analyses data sets in the troposphere and lower stratosphere for the period of May 2001 and September 2002 and they mentioned that the average bias in the upper troposphere and lower stratosphere (250–30 hPa) is less than 1.5 K. Gobiet et al. (2005) made a seasonal latitude-height temperature structure of CHAMP and compared with ECMWF operational analysis using 2.5 year data sets. They noted mean temperature cold biases of  $-2.5\text{ K}$  (maximum) and warm biases up to  $+3.5\text{ K}$  in the Austral polar vortex region.

Figure 7 shows the global distribution of the winter temperature differences between

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COSMIC and models. It illustrates that in the NH the largest differences occur with cold biases of  $-4.5$  K with MetO and JRA-25. However, the NCEP data show differences of only about  $-1.5$  to  $-2.5$  K over the Arctic region. The corresponding differences observed in the SH are smaller, typically ranging between  $-1.5$  and  $-2.5$  K for all the models. The mean differences are relatively small between  $30^\circ$  S and  $30^\circ$  N at about  $1.0$ – $1.25$  K. Figure 7 further indicates that the COSMIC temperature observed to be warmer in the latitude range of  $30^\circ$  S– $30^\circ$  N and colder in the northern and southern polar regions. In the beginning of winter the polar vortex edge was located at the northern high latitudes especially above Europe, but moved significantly as the warming progressed and it might be one of the causes for increased temperature differences. From the summer and winter patterns shown in Figs. 6 and 7, it can be seen that the differences observed in NH winter (SH summer) are significantly larger than the NH summer (SH winter) and this reflects the SH winter circulation in the upper troposphere and lower stratosphere is less disturbed than in the NH winter. We note that the maxima of the seasonal mean differences between COSMIC and models at 100 hPa is larger during NH winter than summer season. The MIPAS and CHAMP comparison studies demonstrated a maximum bias of  $\sim 1$ – $1.5$  K at 30 km (Wang et al., 2004). Further, using the MIPAS and SAC-C data sets, Jiang et al. (2004) estimated the mean differences for 5–30 km and for two different seasons. Their estimate was less than  $\sim 2$  K.

It is appropriate to show the statistical estimates of the bias between the COSMIC–NCEP, COSMIC–MetO and COSMIC–JRA-25 measurements obtained for different latitude sectors for summer and winter seasons at 100 hPa and 50 hPa, and are shown in Tables 1 and 2. The tables depict that the winter differences are consistently larger than the summer for both pressure levels. The RMS deviations are in the range of  $1.5$ – $2.25$  K. The mean differences are larger at high latitudes in both summer and winter seasons for both pressure levels, especially for the MetO values.

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### 3.3 (c) Latitudinal temperature variations at three different pressure levels

The latitudinal temperature structure, at three different pressure levels in December 2006, observed by the models and satellite measurements is shown in Fig. 8a–c. The three chosen pressure levels are 100 hPa, 50 hPa and 20 hPa. In the right panel Fig. 8d–f we show the differences with respect to the COSMIC measurements for the corresponding pressure levels. From Fig. 8a–c it is evident that the overall latitudinal structure is similar in each pressure level. Looking at the differences (Fig. 8d–f), large temperature differences are observed at 100 hPa and less at 20 hPa level. The TES-AURA data exhibit cold biases up to  $\sim 5$  K in the southern latitude ( $50^\circ$  S) region and  $\sim 4$  K in the NH ( $70^\circ$  N) region. The MLS-AURA is illustrating cold bias for all the pressure levels and at almost all latitudes. The maximum difference observed in this case is  $\sim 4$  K in the upper troposphere and lower stratosphere. In the case of radiosonde temperatures, the differences are little larger at 100 hPa level where it exemplifies warm bias of  $\sim 4$  K over the Arctic and cold bias over the Antarctic ( $\sim 4$  K). In particular, the NCEP, NCEP 2, JRA-25 data agree well with COSMIC measurements and they show warm bias of  $1.5$  to  $2$  K and cold bias of  $1$  to  $2$  K. The MetO data sets are consistent with COSMIC in the latitude range of  $60^\circ$  S– $60^\circ$  N with a warm bias of  $2$ – $3.5$  K and cold bias of  $1$ – $2$  K. Overall, it can be concluded that the temperature differences are comparatively smaller over the latitude sector  $60^\circ$  N– $60^\circ$  S.

### 3.4 (d) Comparison of vertical profiles of temperature, specific humidity and refractivity

Comparison of vertical specific humidity, temperature and refractivity profiles derived from COSMIC, CHAMP, ECMWF, NCEP, Radiosonde and TES-AURA data sets are discussed here. The data sets correspond to September 2006 and for the equatorial region from ( $20^\circ$  S to  $20^\circ$  N). Figure 9a–c shows the height profiles of specific humidity, temperature and refractivity profiles. For estimating the specific humidity profile, we followed the method reported by Heise et al. (2006). They explained the method using

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the CHAMP data and compared with radiosonde and ECMWF data sets. Our results suggest that the COSMIC and other data sets have identical humidity values except for the altitudes lower than 4 km. The mean temperature profiles (Fig. 9b) derived from various data sets show excellent similarity. The only slight difference is noted in TES-AURA measurements at altitudes below the tropopause. Finally, the refractivity profiles shown in Fig. 9c also reveal similar values. We have conducted similar analysis for other months also and found impressive likeness among various measurements.

#### 4 Conclusions

By April 2007 nearly half a million COSMIC radio occultations have been collected high-quality temperature profiles in the troposphere and lower stratosphere. Aiming a validation of the COSMIC temperature profiles, at first we have compared the COSMIC temperature profiles averaged for five different latitude bands with model data sets such as NCEP, JRA-25 and MetO. For the troposphere and stratosphere heights (8–30 km) considered in the present study, we found good agreement among various data sets. The mean differences averaged globally and over height region between 8 and 30 km are estimated as  $0.55 \pm 0.015$  K and  $0.65 \pm 0.02$  K, respectively.

The global temperature structures at 100 hPa level revealed by the COSMIC and models, for the northern summer and winter seasons, again show very good similarities in most of the latitude/longitude sectors. Further, we have given attention to note the difference between the COSMIC measurements and the model values obtained for the northern summer and winter seasons. For the summer season, the maximum differences are relatively smaller at about 0.75–1.25 K for the 70° S–70° N latitude range and it increased to 1–2.5 K in the Polar region. Larger differences at the Polar latitudes are caused by the less RO coverage or lower vertical resolution of model values. Also the SH is less covered with regular meteorological observations. The MetO data show comparatively larger differences (2.0–2.5 K) in the equatorial latitudes.

During the northern winter season the bias is comparatively larger for the JRA-25 and

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MetO data sets. In the Arctic region the JRA-25 and MetO data sets show cold biases up to 4.5 K. However, the NCEP data show better resemblance with the COSMIC data. The observed differences can be attributed to the low vertical resolution of the model measurements or error introduced by the spherical symmetry assumption when making the COSMIC retrieval. It is also probably affected by the stratospheric circulation in the NH characterized by strong major warming or strong planetary wave activity. Model data sets may not take into account such effects of rapid synoptic developments.

Comparison of latitudinal distribution of zonal mean temperatures performed for the 100 hPa, 50 hPa, and 20 hPa levels mostly show encouraging agreements. Noticeable differences of 3–4 K are noted in the Polar regions. TES-AURA measurements depict comparatively larger differences than other measurements. Finally, comparison studies did for the specific humidity and refractivity profiles also reveals very good agreements between the COSMIC and other measurements. In the future we look forward to perform further investigations with additional COSMIC and other GPS RO measurements.

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**Table 1.** The temperature bias between the COSMIC and model data sets calculated for different latitude sectors during summer and winter seasons and at 100 hPa.

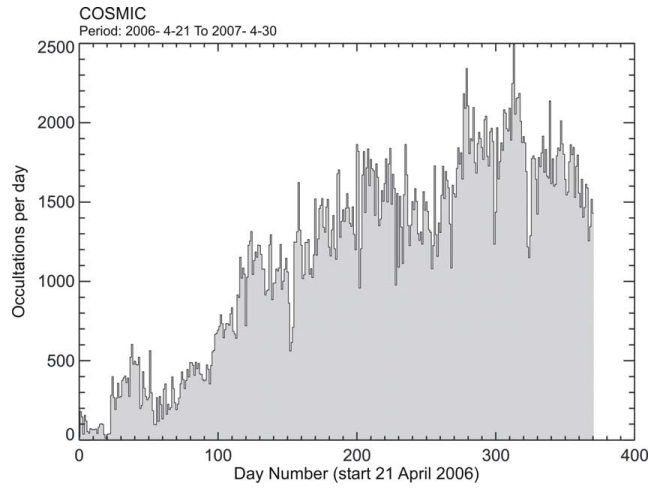
Latitude	Summer (100 hPa)			Winter (100 hPa)		
	COSMIC – NCEP (K)	COSMIC – MetO (K)	COSMIC – JRA25 (K)	COSMIC – NCEP (K)	COSMIC – MetO (K)	COSMIC – JRA25 (K)
50° N–80° N	-0.043 ± 0.69	0.993 ± 0.78	0.346 ± 0.46	-0.066 ± 0.43	-0.981 ± 1.03	-1.479 ± 1.23
20° N–50° N	-0.184 ± 0.45	0.752 ± 0.76	0.053 ± 0.74	-0.324 ± 0.78	-0.187 ± 0.56	-0.727 ± 1.45
20° N–20° S	0.296 ± 0.78	-0.531 ± 0.89	-0.602 ± 0.98	-0.245 ± 0.77	0.614 ± 0.88	0.140 ± 0.49
20° S–50° S	-0.478 ± 0.88	0.261 ± 0.55	-0.401 ± 0.67	-0.332 ± 0.72	0.350 ± 0.78	0.722 ± 1.15
50° S–80° S	-0.854 ± 1.23	1.921 ± 2.01	1.381 ± 1.44	-1.078 ± 1.56	-0.602 ± 1.49	-0.543 ± 1.47

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**Table 2.** Temperature bias between the COSMIC and model data sets calculated for different latitude sectors during summer and winter seasons and at 50 hPa.

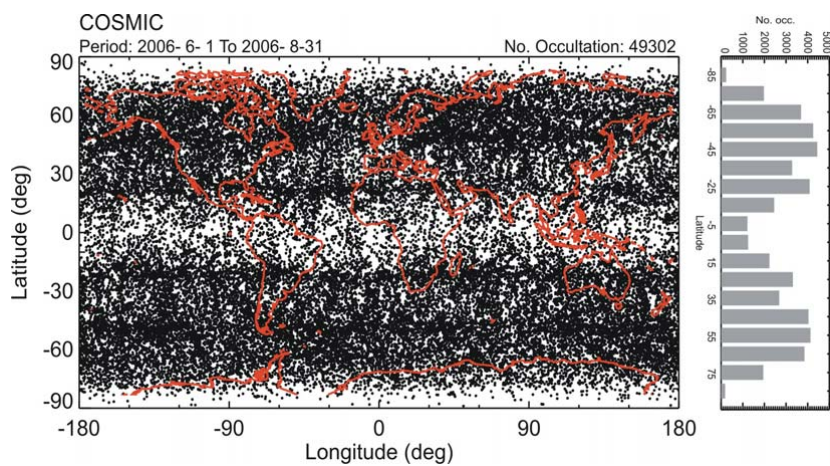
Latitude	Summer (50 hPa)			Winter (50 hPa)		
	COSMIC – NCEP (K)	COSMIC – MetO (K)	COSMIC – JRA25 (K)	COSMIC – NCEP (K)	COSMIC – MetO (K)	COSMIC – JRA25 (K)
50° N–80° N	-0.672 ± 1.29	0.562 ± 0.89	0.391 ± 0.66	-1.094 ± 2.13	-1.594 ± 2.43	-1.302 ± 1.97
20° N–50° N	-0.196 ± 0.58	-0.346 ± 0.75	-0.246 ± 0.59	-0.052 ± 0.89	-0.075 ± 1.16	0.093 ± 0.45
20° N–20° S	-0.227 ± 0.65	-0.213 ± 0.79	-0.313 ± 0.88	0.066 ± 0.99	0.087 ± 0.45	0.811 ± 1.39
20° S–50° S	-0.162 ± 0.66	-0.780 ± 1.05	-0.880 ± 1.47	-0.164 ± 0.82	-0.145 ± 0.74	0.235 ± 0.89
50° S–80° S	-0.279 ± 0.73	1.252 ± 1.62	1.452 ± 1.78	0.609 ± 1.19	1.775 ± 2.09	1.805 ± 2.26

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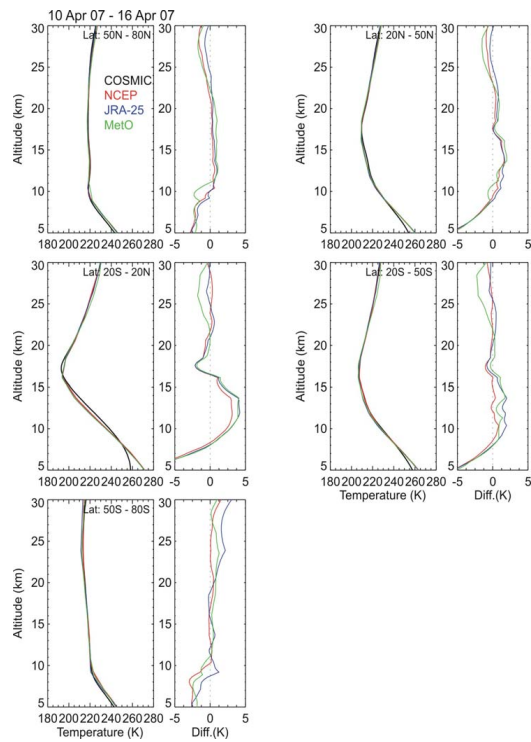
**Fig. 1.** Histogram of number of occultations (per day) for the period from 21 April 2006 to 30 April 2007.

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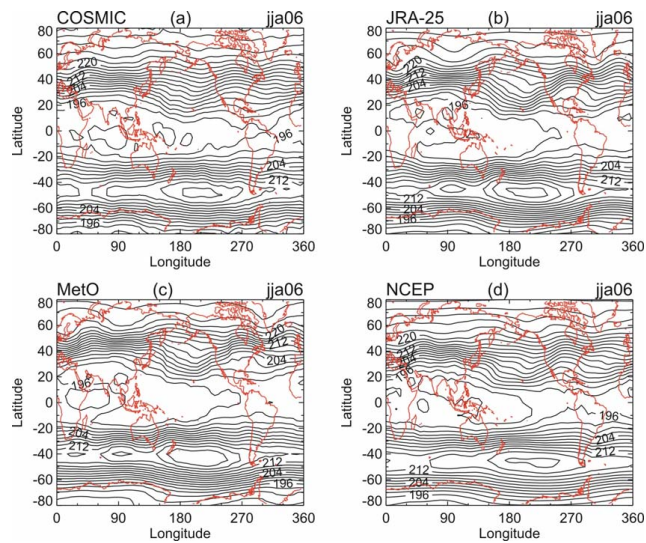
**Fig. 2.** Global coverage of COSMIC occultations for the Northern Hemisphere summer (1 June 2006–31 August 2006). Histogram of total number of occultations occurred in every 10° latitude range is shown in the right panel.

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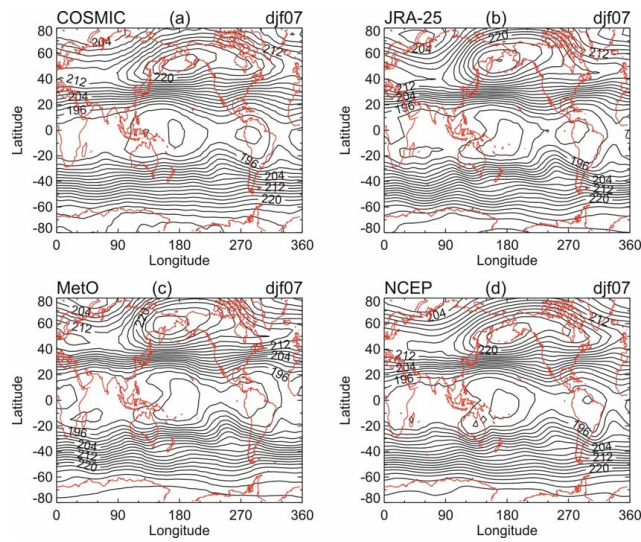
**Fig. 3.** Height profiles of the COSMIC and model temperatures derived for various latitude bands and for the period 10 April 2007–16 April 2007. In each right sub panel the difference in temperature with respect to COSMIC measurements and model values is plotted.

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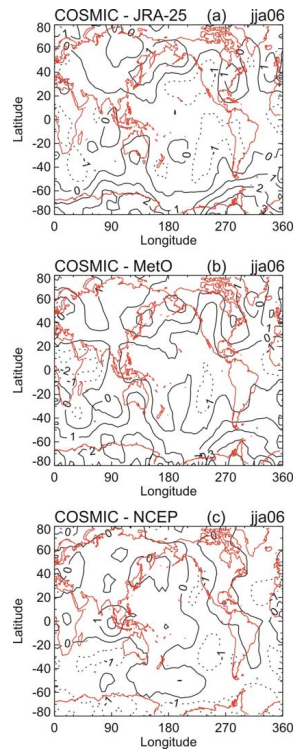
**Fig. 4.** Mean global temperatures at 100 hPa by (a) COSMIC, (b) JRA-25, (c) MetO and (d) NCEP data sets during Northern Hemisphere summer (June–August, 2006). The contour interval is 2 K.

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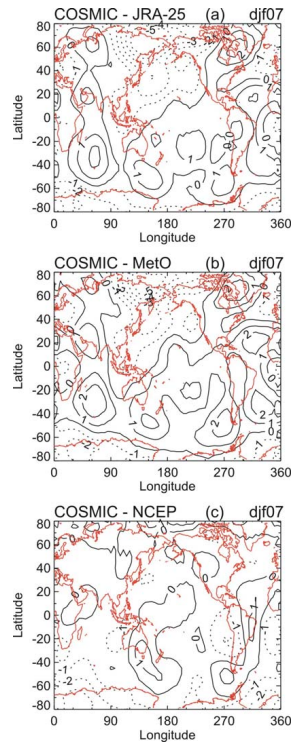
**Fig. 5.** Same as Fig. 4, but for the Northern Hemisphere winter (December 2006, January, February 2007).

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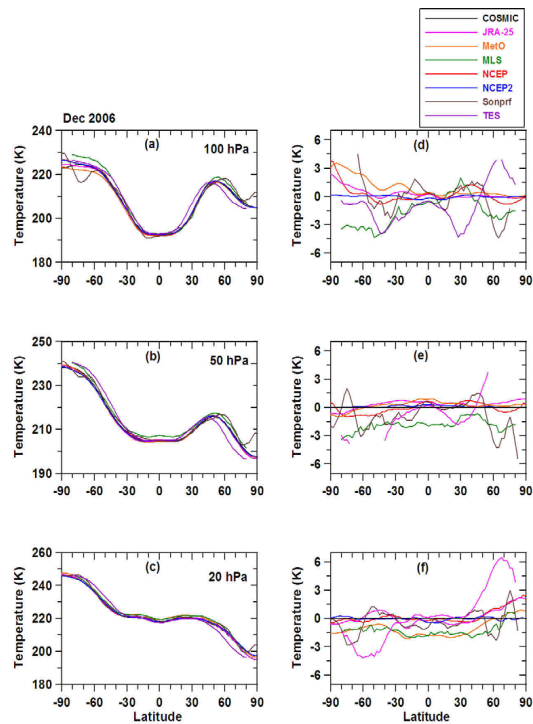
**Fig. 6.** Global zonal mean temperature differences (in Kelvin) between (a) COSMIC–JRA-25, (b) COSMIC–MetO, and (c) COSMIC–NCEP during Northern Hemisphere summer (June–August, 2006). The contour interval is 1 K.

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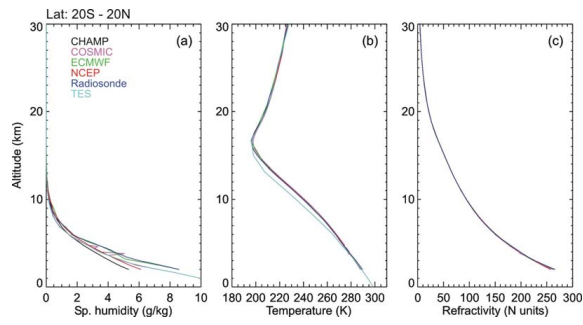
**Fig. 7.** Same as Fig. 6, but for Northern Hemisphere winter (December 2006, January, February 2007). The contour interval is 1 K.

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**Fig. 8.** Latitudinal distribution of mean zonal temperatures for (a) 100 hPa, (b) 50 hPa, and (c) 20 hPa. Various data sets shown in the figure correspond to December 2006. Distribution of differences with respect to COSMIC measurements and other data sets are given in the right panel (d–f).

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**Fig. 9.** Comparison of average vertical height profiles of (a) specific humidity, (b) temperature, and (c) refractivity profiles derived from COSMIC, CHAMP, ECMWF, NCEP, Radiosonde, and TES-AURA data sets for the latitude sector 20° S–20° N during September 2006.