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4D-Var wind-ozone coupling

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On the extraction of wind information from the assimilation of ozone profiles in Météo–France 4D-Var operational NWP suite

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Abstract

By applying four-dimensional variational data-assimilation (4D-Var) to a combined ozone and dynamics Numerical Weather Prediction model (NWP), ozone observations generate wind increments through the ozone-dynamics coupling. The dynamical impact of Aura/MLS satellite ozone profiles is investigated using Météo–France operational *ARPEGE* NWP 4D-Var assimilation system for a period of 3 months. A data-assimilation procedure has been designed and run on 6-h windows. The procedure includes: (1) 4D-Var assimilating both ozone and operational NWP standard observations, (2) *ARPEGE* transporting ozone as a passive-tracer, (3) *MOCAGE*, the Météo–France chemistry and transport model re-initializing the *ARPEGE* ozone background at the beginning time of the assimilation window. The Degrees of Freedom for Signal diagnostics show that the MLS data covering the 68.1–31.6 hPa vertical pressure range are the most informative and their information content is nearly of the same order as tropospheric humidity-sensitive radiances. Furthermore, with the help of error variance reduction diagnostics, the ozone contribution to the reduction of the horizontal divergence background-error variance is shown to be better than tropospheric humidity-sensitive radiances. Moreover, by using observation minus forecast statistics, it is found that the ozone assimilation reduces the wind bias in the lower stratosphere.

1 Introduction

Over the years, data assimilation schemes have evolved into very sophisticated systems, such as four-dimensional variational 4D-Var (Rabier et al., 2000), which is operational in the numerical weather prediction (NWP) system of Météo–France, *ARPEGE*. The scheme combines a large variety of both space- and surface-based meteorological observations with background information of the atmospheric state. During the 4D-Var process, the evolution of the linearized forecast model and its adjoint act as an additional constraint (Andersson et al., 1994).

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Ozone is a radiatively important atmospheric trace gas and 4D-Var is adapted to produce four-dimensional ozone analyses consistent with atmospheric dynamics (Eskes et al., 2005; Geer et al., 2006; Lahoz et al., 2007). In return, the assimilation of ozone can have a beneficial impact on the dynamics forecast modeling. Ozone influences the model temperature through the radiation parametrization, often referred to as “ozone/radiation interaction” (Cariolle and Morcrette, 2006). Moreover, an accurate ozone knowledge may improve the use of satellite radiances sensitive to ozone such as long-wave channels of both the High Infrared Resolution Sounder HIRS (e.g., Derber and Wu, 1998) and the Advanced Microwave Sounder AMSU-B (e.g., John and Buehler, 2004). In addition to these potential benefits, 4D-Var ozone assimilation can improve the wind field through dynamics-ozone coupling. This takes advantage of the fact that the photochemical lifetime of ozone is relatively long in the upper troposphere-lower stratosphere (UTLS) region where vertical and horizontal motions associated with synoptic scales systems lead to similar ozone and isentropic potential vorticity (IPV) anomalies (Reed, 1950; Danielsen, 1968; Davis, 1999; Semane et al., 2002). This could bring a valuable correction of model uncertainties in the UTLS region. Daley (1995) highlighted the feasibility of inferring dynamical information on wind fields from the assimilation of chemical constituent observations, given sufficiently dense, frequent and accurate measurements. Riishøjgaard (1996) demonstrated the use of ozone measurements to reconstruct the flow field in a barotropic vorticity equation model. The study of Holm et al. (1999) showed in detail the wind-ozone coupling in a 4D-Var system. Peuch et al. (2000) demonstrated the dynamical impact of total ozone column observations in Observing System Simulation Experiments (OSSEs) using the *ARPEGE* 4D-Var suite. All studies agree that two prerequisites for producing a beneficial interaction between the ozone and wind fields within 4D-Var are the use of high quality observations and background ozone fields. These have proven to be very strong constraints and, up to now, few results obtained with actual ozone observations have been presented. In this work, the dynamical impact of ozone 4D-Var assimilation is thus investigated within a framework in which *ARPEGE* is run during a period of 3

months from 23 January to 22 April 2006 with the following input:

- Frequent and accurate UTLS ozone profiles from the Microwave Limb Sounder (MLS) aboard the Aura satellite with a vertical resolution of 2.7 km in the UTLS and a horizontal resolution of 165 km (Waters et al., 2006).
- Ozone background provided by *MOCAGE*, a comprehensive chemistry and transport model (CTM) of Météo-France (Peuch et al., 1999). This background is consistent with dynamical fields of ARPEGE as MOCAGE is driven by ARPEGE dynamics.

On the one hand, the information content of MLS data is examined by using the Degrees of Freedom for Signal diagnostics. In addition, with the help of the error variance reduction diagnostic of Desroziers et al. (2005), the reduction of the initial background error variance of wind fields due to MLS ozone data assimilation is described. The results are compared to humidity-sensitive radiances of HIRS and AMSU-B sounders. On the other hand, the observation minus forecast (OMF) statistics are used to examine how the wind fields are affected by the assimilation of MLS data. The paper is outlined as follows: Sect. 2 presents the MLS ozone assimilation procedure, Sect. 3 describes the dynamical impact of ozone data, and Sect. 4 summarizes the results and discusses their implication.

2 MLS ozone assimilation procedure

Thanks to the 4D-Var assimilation process, ozone observations affect not only the analysis of the ozone field itself, but also the analysis of the wind field through the adjoint of the ozone advection model. According to Peuch et al. (2000), ozone observations will improve the wind field only if they are sufficiently accurate and numerous, if the linearized ozone advection model is appropriate for simulating time-space evolution over the assimilation window and finally if the background is consistent with the observations. The 4D-Var experiment carried out in this study attempts to nearly meet these

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requirements through the use of MLS observations and ozone background derived from the *MOCAGE* CTM at the beginning of each 4D-Var *ARPEGE* data-assimilation cycle.

2.1 MLS measurements

5 The Aura satellite was launched on 15 July 2004 and placed into a near-polar Earth orbit at ~ 705 km. It orbits the Earth around 14 times per day. The MLS instrument aboard Aura uses the microwave limb sounding technique to measure chemical constituents between the upper troposphere and the lower mesosphere. It is an advanced version of the MLS instrument on the Upper Atmosphere Research Satellite (UARS). This is
10 because it provides more dense spatial coverage with its daily 3500 profiles, between 82° N and 82° S, while UARS MLS could only produce up to 1300 every day. In addition, Aura MLS ozone profiles are retrieved with a horizontal resolution of 165 km and a typical vertical resolution of about 2.7 km in the stratosphere (Froidevaux et al., 2006). Although Aura MLS observations contain useful information on ozone concentrations
15 between 316 hPa and 0.1 hPa, only the data between 215.4 and 0.46 hPa are recommended for scientific research. Precision is around 20–50 ppbv and 0.1–0.2 ppmv in the vertical ranges 215.4–22 hPa and 21.5–0.46 hPa, respectively (Waters et al., 2006). In this study, the first publicly available Aura MLS dataset, version 1.5, of ozone profiles from 23 January to 22 April 2006, is used. As high biases exist at 215.4 hPa for
20 equatorial latitudes, data for this level covering the 30° N– 30° S latitude range are discarded. The observational error in the used data-assimilation experiment is prescribed directly from the precision estimate supplied with the Aura MLS data; no bias correction is applied.

2.2 *MOCAGE* model

25 *MOCAGE* (*MOD*èle de *Chimie Atmosphérique à Grande Echelle*) is a three-dimensional chemistry transport model of the troposphere and stratosphere. It pro-

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vides a number of optional configurations with varying domain geometries and res-
olutions, as well as chemical and physical parametrization packages. It is used for
operational chemical weather prediction (e.g., Dufour et al., 2004; Hollingsworth et al.,
2008; Rouil et al., 2008) and chemical data assimilation research (e.g., Semane et al.,
2007; Massart et al., 2007; El Amraoui et al., 2008a; El Amraoui et al., 2008b). The
transport scheme of MOCAGE is semi-lagrangian (Rasch and Williamson, 1990) and
the chemical one used in this study is the comprehensive scheme RACMOBUS, which
combines the stratospheric scheme REPROBUS (Lefèvre et al., 1994) and the tropo-
spheric scheme RACM (Stockwell et al., 1997). *MOCAGE* is forced by external wind
and temperature fields from the operational meteorological model of Météo-France,
ARPEGE (Courtier et al., 1991) and is used with its global horizontal resolution of 2°
both in latitude and longitude, including 47 hybrid levels from the surface up to 5 hPa.
The vertical resolution is about 800 m in the vicinity of the tropopause and in the lower
stratosphere.

2.3 Assimilation methodology

The notation in this paper will follow Ide et al. (1997) as closely as possible. Météo-
France uses the 4D-Var scheme as part of its ARPEGE NWP system (Rabier et al.,
2000). This scheme minimizes the following function $J(\delta\mathbf{x})$:

$$J(\delta\mathbf{x}) = J_b(\delta\mathbf{x}) + J_o(\delta\mathbf{x}) + J_c(\delta\mathbf{x}) \quad (1)$$

where the increment vector $\delta\mathbf{x}$ is the difference between the model state \mathbf{x} and the
background state \mathbf{x}_b . The $J_b(\delta\mathbf{x})$ term in Eq. (1) refers to the background cost function

$$J_b(\delta\mathbf{x}) = \delta\mathbf{x}^T \mathbf{B}^{-1} \delta\mathbf{x} \quad (2)$$

and the $J_o(\delta\mathbf{x})$ term refers to the observation cost function

$$J_o(\delta\mathbf{x}) = (\mathbf{d} - \mathbf{H}(\delta\mathbf{x}))^T \mathbf{R}^{-1} (\mathbf{d} - \mathbf{H}(\delta\mathbf{x})) \quad (3)$$

where $\mathbf{d} = \mathbf{y}^\circ - \mathbf{H}(\mathbf{x}_b)$ is the departure between the observation vector \mathbf{y}° and its model equivalent in observation space $\mathbf{H}(\mathbf{x}_b)$. The operator \mathbf{H} is a generalized interpolator (including model integration to observation time over the assimilation window) from the model grid to the observation location and \mathbf{H} represents its tangent-linear. The $J_c(\delta\mathbf{x})$ in Eq. (1) is a penalty term controlling gravity waves. In Eq. (2) and Eq. (3), \mathbf{B} and \mathbf{R} respectively represent the background and observation error covariance matrices. In order to assimilate MLS observations, the following changes in the operational version of the 4D-Var *ARPEGE* system have been applied. First, the ozone variable is incorporated in the model state vector \mathbf{x} . Hereafter, the ozone part of \mathbf{x} is denoted \mathbf{x}_{O_3} . The time-space evolution of \mathbf{x}_{O_3} is simulated by a passive-tracer advection equation using a semi-lagrangian transport scheme. Second, MLS ozone observations are included in Eq. (3). Third, in contrast to the other meteorological variables for which the background \mathbf{x}_b of Eq. (2) is provided by the *ARPEGE* 3-h forecast run, the ozone part of \mathbf{x}_b (hereafter represented by $\mathbf{x}_{b_{\text{O}_3}}$) is provided by *MOCAGE* at the beginning time of each 6-h assimilation window. Figure 1 depicts a vertical cross section from 90°N to 90°S of the normalized differences in the zonal mean ozone fields from *MOCAGE* (comprehensive chemistry) and *ARPEGE* (where ozone is only advected) after 6, 12, 18 and 24 h of simulation. It clearly shows that, after 6 h, ozone fields computed by *ARPEGE* remain very close to the one computed with the dedicated *MOCAGE* CTM; discrepancies grow after 12–24 h of simulation, both due to a less appropriate representation of transport and to the lack of chemical source and sink terms in the *ARPEGE* model. Finally, in the operationally used version of *ARPEGE*, the part of background-error covariance matrix \mathbf{B} for dynamical meteorological variables is estimated using Analysis Ensemble Method (Houtekamer et al., 1996). In this study, the part of the \mathbf{B} related to \mathbf{x}_{O_3} (hereafter represented by \mathbf{B}_{O_3}) is computed statistically using the NMC method (Parrish and Derber, 1992) with 12/36-h forecast differences. The 12/36-h forecasts are both constructed by integrating the passive-tracer advection equation from an ozone initial condition given by *MOCAGE*. It must be noticed, here, that the constructed \mathbf{B}_{O_3} matrix corresponds more to the ozone short-term forecasts errors of *ARPEGE* rather

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than to the MOCAGE ones, but both errors should be similar for the following reasons: (1) the photochemical lifetime of ozone is relatively long in the UTLS region and (2) MOCAGE uses a dynamical forcing provided by ARPEGE. Further, this way of \mathbf{B}_{O_3} matrix construction has the advantage to provide it on the ARPEGE grid directly. Figure 2a, b, and c respectively depict pressure, ozone average vertical correlations and ozone horizontal correlation scale as a function of model levels in the UTLS. Figure 2b shows that the vertical correlation length of $x_{b_{\text{O}_3}}$ errors is small in the stratosphere and increases towards the tropopause. Furthermore, ozone is assimilated in a univariate manner: no correlation is specified between ozone and wind (relative vorticity) in the formulation of the \mathbf{B} matrix. The univariate ozone background errors were chosen because they minimize the feedback effects of ozone on other meteorological variables (Dethof and Holm, 2004). Therefore, ozone observations affect the analysis of the meteorological fields only through the coupling of trace gas concentration and air transport. No ozone/radiation interaction is included: the ozone climatology of Fortuin and Langematz (1995) is used in the radiation scheme. Observation operator calculations within satellite radiances assimilation (e.g., ozone-sensitive channel 9 of HIRS and channel 18 of AMSU-B) also use ozone climatology instead of the new introduced model ozone.

2.4 Experiment setup

In order to investigate the impact of the assimilation of MLS ozone profiles on ARPEGE wind fields, two assimilation experiments have been run at a regular low resolution of T107 truncation on 41 levels from the ground up to 1 hPa. The operational version was at a resolution of T358 (stretched grid) on 41 levels in 2006; the current (2008) operational version is at a resolution of T538 on 60 levels. In the first experiment (denoted CTL), MLS data are not assimilated. In the second experiment (denoted MLS), MLS ozone profiles, between 215.4 and 1 hPa in the extratropics and between 147 and 1 hPa in the latitudes range (30°S – 30°N), are assimilated. The MLS and CTL experiments are run over a period of 3 months from 23 January to 22 April 2006. Other

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assimilated observations in the CTL and MLS experiments are based on the operational data including surface observations, radiosondes, aircraft reports, and wind profilers, as well as satellite observations such as winds from the geostationary satellites, MODIS winds, QuikSCAT winds, and radiance data from both HIRS and AMSU (Unit-A and Unit-B) sounders aboard the National Oceanic and Atmospheric Administration polar-orbiting satellites (NOAA-15, NOAA-16 and NOAA-17). 4D-Var assimilates observations within ± 3 h of the analysis time (00:00, 06:00, 12:00 and 18:00 UTC). This is in line with the current operational set-up in *ARPEGE*, as this study aims at being in a fully realistic NWP context. The flow of 4D-Var procedure for the case of 12:00 UTC analysis time is given in Fig. 3.

3 Ozone data dynamical impact study

3.1 OMA, OMF and DFS diagnostics

Figure 4 displays $100 \times \frac{OMA}{y^o}$ (solid line with squares) and $100 \times \frac{OMF}{y^o}$ (dashed line with circles) both averaged in time from 23 January to 21 February 2006 (00:00 and 12:00 UTC analysis times) and over the following latitude bands (a) 30° N– 90° N, (b) 30° N– 30° S, and (c) 30° S– 90° S. Here, OMA and OMF respectively refer to Observation minus Analysis and Observation minus First-guess for ozone. This shows that when MLS data are assimilated, the departures are reduced and the analysis agrees better with the MLS data in comparison to the background. For instance, the assimilation of MLS profiles corrects the background overestimation of ozone in the Northern Hemisphere, by driving the analyses to a value slightly close to MLS observations owing to the negative correction induced by the analysis increment. This overestimation feature of the background has been already diagnosed by El Amraoui et al. (2008a) during winter months inside the vortex. The reverse process occurs in the tropics and the Southern Hemisphere at pressure levels greater than 46.4 hPa, where there is a substantial correction of the model underestimation of ozone.

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A traditional way of estimating data impact in a assimilation system is to perform Observing System Experiments (OSEs). It consists of removing one particular dataset over a long assimilation period. One then evaluates the forecast scores, compared to a reference assimilation using the complete set of observations (Bouttier and Kelly, 2001). This procedure provides the impact of observations on the forecasts and not on the analyses themselves (Desroziers et al., 2005). In the perspective of diagnosing the direct impact of observations on the analyses, other diagnostics such as Degree of Freedom for Signal and variance reduction were developed. This study is based on these two diagnostics. The Degrees of Freedom for Signal (*DFS*) provide a measure of the gain in information brought by the observations (Rodgers, 2000; Fisher, 2003). In a linear framework, *DFS* is algebraically defined by the trace of the **HK** matrix

$$DFS = Tr(\mathbf{HK}) \quad (4)$$

where **K** refers to the Kalman gain matrix, which only depends on the specified statistics of the assimilating system

$$\mathbf{K} = \mathbf{BH}^T(\mathbf{HBH}^T + \mathbf{R})^{-1} \quad (5)$$

Like in Chapnik et al. (2006), the *DFS* is approximated here by using the method of Desroziers and Ivanov (2001). Indeed, one normal ozone analysis \mathbf{x}^a and a perturbed one \mathbf{x}^{a*} , performed with perturbed observations \mathbf{y}^{o*} , are used to estimate the partial Degrees of Freedom for Signal associated with a particular subset of observations (called *DFS_i*) through the following expression:

$$DFS_i = \delta \mathbf{y}_i^{oT} \mathbf{R}_i^{-1} \mathbf{\Pi}_i \mathbf{H} \delta \mathbf{x}^a (\delta \mathbf{y}^o) \quad (6)$$

where the subscript *i* refers to a specific subset and \mathbf{R}_i presents its observation error covariance matrix. $\mathbf{\Pi}_i$ is the projection operator that allows us to pass from the complete dataset used in the assimilation to the subset *i*. $\delta \mathbf{y}_i^o = \mathbf{\Pi}_i \delta \mathbf{y}^o$ is the projection of the vector of perturbations $\delta \mathbf{y}^o$ onto the subset *i*,

$$\delta \mathbf{y}^o = \mathbf{y}^{o*} - \mathbf{y}^o \quad (7)$$

and $\delta x^a(\delta y^o)$ is the perturbation on the analysis produced by a perturbation on the complete assimilated dataset:

$$\delta x^a(\delta y^o) = x^{a*} - x^a \quad (8)$$

The partial Degrees of Freedom for Signal associated with a particular subset of observations can be determined if the associated error characteristics of these observations are not correlated to the rest of the observation errors in the specified \mathbf{R} matrix. In fact, it is assumed that the complete set of observations can be split between observation subsets with independent errors, then matrix \mathbf{R} is block-diagonal. The complete assimilated dataset over a 6-h assimilation window is perturbed one-time and the Eq. (6) is evaluated for each assimilation cycle for a one-month period.

Figure 5a, b and c respectively show the observation number, $100 \times \frac{\sigma_{oi}}{y_i^o}$ and DFS_i , all averaged in time from 23 January to 21 February 2006 (00:00 and 12:00 UTC analysis times) and for the following specific subsets: MLS ozone data on specific pressure levels and humidity-sensitive channel radiances from the HIRS sounder on NOAA-16 (channels 11 and 12) and the AMSU-B sounders on NOAA-16 and NOAA-17 (channels 18, 19 and 20). σ_{oi} refers to the observation-error of the subset i . One can see that the MLS data covering the 68.1–31.6 hPa vertical pressure range are the most informative and their information content is nearly of the same order as humidity-sensitive channel radiances. The fact that the first pressure 100–215.4 hPa range does not provide significantly more information in comparison to the pressure 68.1–31.6 hPa range, is probably due to a combination of factors. First, the variances of observation error for the first range are significantly greater than those of the second one. Second, the assumed horizontal correlation of background errors introduces a degree of redundancy into the information provided by the observation within the scale length of the correlation. This latter is smaller for the second pressure range in comparison to the first one (see Fig. 2c).

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3.2 Error variance reduction diagnostics

The error variance reduction method introduced by Desroziers et al. (2005) diagnoses the direct impact of a specific observation subset on a specific analysis field. It provides, hence, useful information on the use of observations in an operational analysis. It is used here in order to objectively diagnose the impact of ozone data on the wind analysis in terms of horizontal divergence and relative vorticity, in comparison to humidity-sensitive channel radiances from both HIRS and AMSU-B sounders. The error variance reduction, for horizontal divergence or relative vorticity, is estimated by using the following expression:

$$r_i = \delta y_i^{oT} \mathbf{R}_i^{-1} \Pi_i \mathbf{HBL}^T \mathbf{L} \delta x^a (\delta y^o) \quad (9)$$

where \mathbf{L} is a mapping, which corresponds here to a projection operator onto the two selected model variables that are directly linked to the wind field: horizontal divergence and relative vorticity. DFS_i can be derived from the r_i expression by considering the special case of $\mathbf{L} = \mathbf{B}^{-1/2}$.

Figure 6 presents r_i for relative vorticity (a) and horizontal divergence (b) averaged over the same period and for the same data as those used for the DFS_i diagnostics. It clearly shows that the contribution of MLS ozone dataset to the reduction in error between the background error and the analysis error of wind fields is larger for horizontal divergence field in comparison to HIRS and AMSU-B channel contributions. The vertical distribution of error variance reductions brought by ozone data for both horizontal divergence and relative vorticity is not uniform as a function of pressure. The lower pressure levels of MLS in the UTLS (100–215.4 hPa) has the most important influence on relative vorticity. This was expected as there is a strong correlation between ozone and potential vorticity in the vicinity of the tropopause level. The difference between DFS_i and r_i at these levels could be the fact that the assumed horizontal correlation of background errors introduces a degree of redundancy into the information provided by the observation within the scale length of the correlation leading to a smaller DFS_i in

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comparison to r_j . Besides, the relative vorticity and horizontal divergence fields are not influenced by ozone in a similar way. This is likely a manifestation of the multivariate analysis structures implied by the balance operator (based on the geostrophic linear balance between mass and wind). Further, the large error variance reductions brought by humidity radiances on relative vorticity compared to those on horizontal divergence, can be explained by the fact that these radiances (e.g., HIRS channels) are rather sensitive to temperature (but to a lesser extent when compared to humidity). This gives them the possibility to affect the relative vorticity not only via the 4D-var wind-humidity coupling, but also implicitly through the balance operator owing to the temperature-relative vorticity correlations. Note that humidity-sensitive radiances can also influence the analysis of temperature field and thereby its background-error variance reduction, through the model physics given the fact that temperature analysis response is dependent on the “drying” or “wetting” analysis effect of these radiances. Model dynamics are able to efficiently extract information on an unobserved component of the flow (the wind) from information on both humidity and ozone data. Intrinsically linked to the four-dimensional nature of the assimilation, the wind-ozone and wind-humidity couplings generate wind increments both from ozone and humidity-sensitive measurements leading to the reduction of the wind background error variance. As it has been elegantly introduced by Andersson et al. (1994), the process could be described as retrieval/assimilation of “water-vapor winds” from humidity-sensitive radiances and of “ozone winds” from MLS observations.

3.3 OMF wind statistics

The OMF residuals produced by a data assimilation system provide a convenient metric of evaluating global analyses. Here, OMF statistics from *ARPEGE* are used to examine how wind assimilation output and their associated OMF bias and standard deviation are affected by the additional assimilation of MLS ozone data. Specifically, *ARPEGE* wind fields (zonal and meridional components) from the first-guess, 24-h and 48-h forecasts are compared against radiosondes for experiments with and without MLS

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ozone data. OMF are averaged in time from 23 January to 22 April 2006 (for 00:00 and 12:00 UTC) and over the globe. The number of observations per pressure level, used in the OMF statistics, is given in Table 1. Figure 7 depicts vertical profiles of OMF bias for the CTL (solid line) versus the MLS experiment (dashed line) for both zonal (left hand-side panels) and meridional (right hand-side panels) wind components. The top, middle and bottom panels correspond to the first guess, 24-h and 48-h forecasts, respectively. In comparison to the CTL, the MLS experiment has smaller biases in the lower stratosphere for both zonal and meridional winds. The middle and bottom panels of Fig. 7 confirm that information added in the MLS experiment is successfully retained in the 24/48-h forecasts in the lower stratosphere. OMF standard deviations, not shown here, are very close for both experiments. Overall, the addition of MLS ozone data improves the agreement between the wind fields and the verifying radiosondes data over the globe. The observation minus first guess biases (top panels of Fig. 7) show that the wind response signal due to the MLS data assimilation is nearly vertically uniform between 20 and 100 hPa for the zonal component. It has, however, a pronounced positive peak between 30 and 50 hPa for the meridional wind component. This response signal difference between the two components of the wind has also been pointed out by Riishøjgaard (1996). In fact, the direction of the flow with respect to the gradient of the tracer concentration is a key parameter that governs the wind response to the tracer data assimilation. Indeed, when the flow is aligned with the gradient of the tracer concentration, the tracer assimilation affects considerably the wind during the adjoint calculation (this can be easily inferred from the simplified Euler-Lagrange equations used by Holm et al. 1999). In the top right-hand side panel of Fig. 7, the peak in the bias reduction of the wind meridional component corresponds to the altitudes where the meridional component of ozone gradient is large; this is typically true for the middle and high latitudes. To summarize, the MLS ozone data provide a beneficial signal in the ARPEGE wind fields. Even though the ozone data dynamical impact is not very large, it is not neutral and mostly positive over a meteorological data sparse region (e.g., the lower stratosphere). It must be noticed that the dynamical impact of the MLS

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data depends very much on how the information contained in the data is extracted and incorporated into the initial condition. Yet, a good MLS dataset, which is unbiased with respect to the background, could probably improve the wind analysis even more. Note that, in this work, no bias correction has been applied in the first approach.

4 Conclusions

In this work, MLS ozone profiles have been assimilated in the ARPEGE 4D-Var together with operational observations. This was carried out to determine to what extent the wind-ozone coupling, within the forecast model and its adjoint, could provide wind increments in response to remotely-sensed MLS ozone information. In the designed MLS assimilation experiment the ozone is re-initialized from the *MOCAGE* CTM every 6 h, however, the meteorological fields continue their temporal evolution through the assimilation cycles. A data impact study with 4D-Var analysis was conducted from 23 January to 22 April 2006. The Degrees of Freedom for Signal diagnostics showed that the MLS data covering the 68.1–31.6 hPa vertical pressure range are the most informative and their information content is nearly of the same order as humidity-sensitive radiances. Furthermore, the error variance reduction diagnostics showed that MLS ozone observations positively contribute to the reduction of the global initial background error variance of the wind fields. Moreover, they bring additional information and in terms of horizontal divergence, this contribution is larger than that given by HIRS and AMSU-B. Further, the statistics on observation minus forecast produced by *ARPEGE* allowed a description of how the ozone observations affect wind fields. Indeed, it was found that the major impact of ozone on wind fields, diagnosed within the MLS data assimilation experiment, consists of a slight improvement in the wind fields in the lower stratosphere where meteorological data are sparse. In addition, the forecasts issued after assimilation of MLS ozone data are found to be closer to the wind observations than the forecasts issued without MLS. The main conclusion is that it is possible to improve the wind simulation of *ARPEGE* by adding MLS ozone data to the assimilation sys-

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tem. The improvement of lower stratospheric wind forecasts inferred from MLS data demonstrated the potential benefit for including ozone assimilation in an operational framework. In ongoing work, the impact of dynamical forcing resulted from *ARPEGE* (from the MLS experiment) on the *MOCAGE* CTM, is under examination. In fact, the new meteorological fields produced by the MLS experiment are used as new dynamical forcing for *MOCAGE*. The impact of this new forcing on *MOCAGE* in terms of ozone evolution will be assessed in comparison to a control run of *MOCAGE* using the standard dynamical forcing derived from the *ARPEGE* control experiment without MLS assimilation. As the MLS data are not available in near-real time, their assimilation could not be implemented in an operational weather forecasting. However, one could envision the assessment, as done in this work, of the derived benefit from assimilating ozone data (IASI and GOME2 instruments) of Metop satellite measurements launched on 19 October 2006, and which are available in near-real time to operational centres.

Acknowledgements. The authors would like to thank Jean Pailleux from Météo-France for his useful suggestions. This research is supported by Météo-France and the Centre National d'Etudes Spatiales (CNES). This work was also supported in the context of the EU-funded Framework V project ASSET (contract EVK2-CT-2002-00137). MLS data were downloaded from the NASA Goddard Space Flight Center.

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Table 1. Number of radiosonde observations (N. Obs) per pressure level in hPa, used in OMF wind statistics.

Pressure (hPa)	N. Obs.	Pressure (hPa)	N. Obs	Pressure (hPa)	N. Obs	Pressure (hPa)	N. Obs
10	137 905	70	150 144	250	196 573	600	108 615
20	165 619	100	206 210	300	199 299	700	188 118
30	170 677	150	228 913	400	190 044	850	140 236
50	159 185	200	214 988	500	186 961	925	103 808

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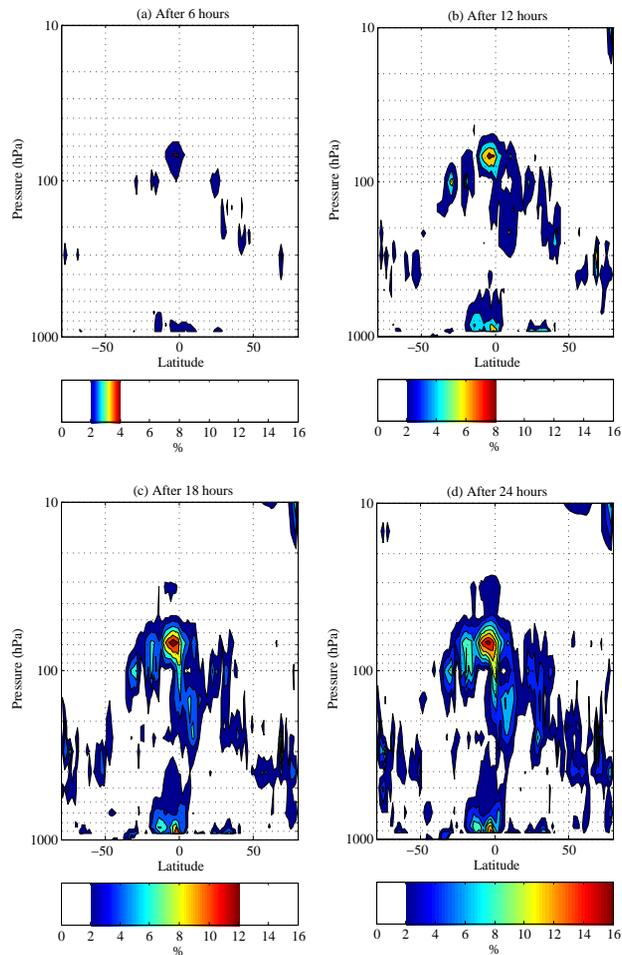


Fig. 1. MOCAGE and ARPEGE zonal mean difference percentage, $100 \times \frac{ARPEGE - MOCAGE}{MOCAGE}$, after 6 (a), 12 (b), 18 (c) and 24 h (d) of simulation.

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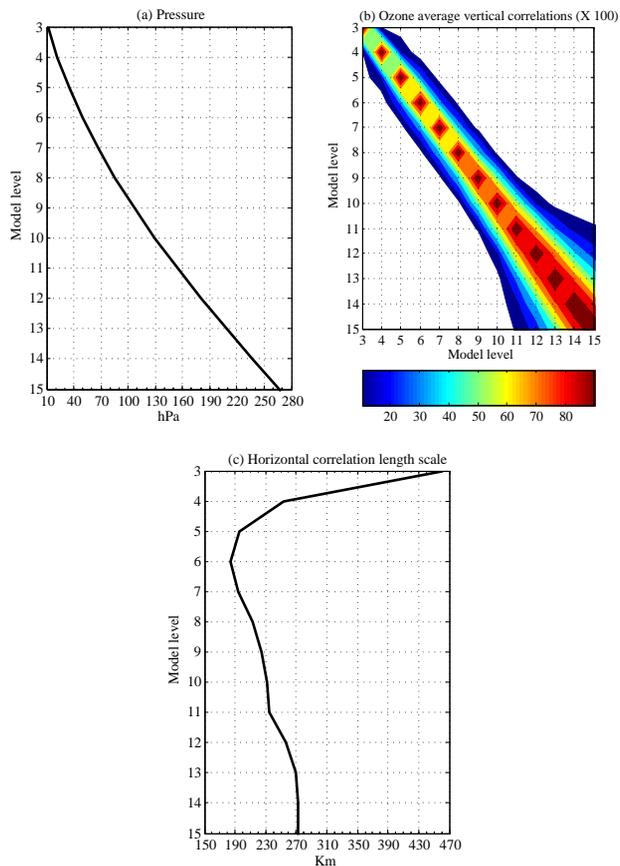


Fig. 2. (a) Pressure as a function of model levels in the UTLS, (b) Ozone average vertical correlations, and (c) Ozone horizontal correlation scale as a function of model levels in the UTLS.

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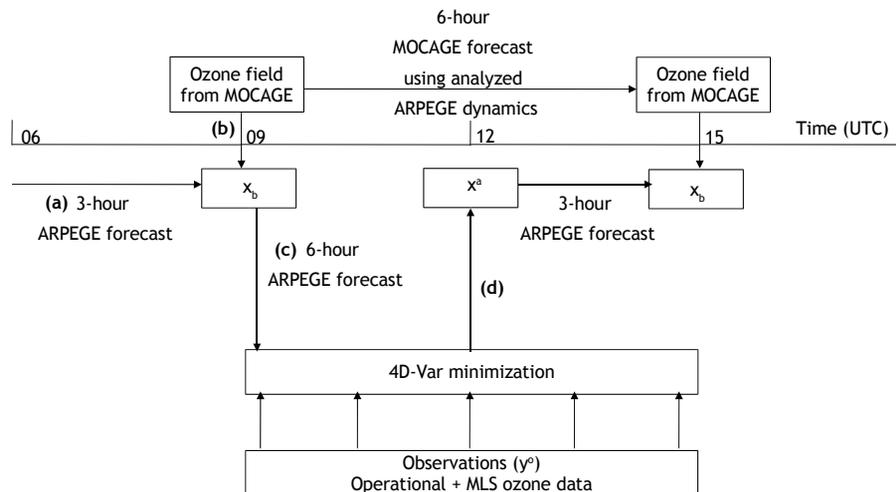


Fig. 3. Flow of 4D-Var procedure for the case of 12:00 UTC analysis time. **(a)** 3-h forecast from previous ARPEGE analysis is used as background \mathbf{x}_b , **(b)** The ozone background field is provided by MOCAGE, **(c)** 6-h forecast from \mathbf{x}_b to perform the innovation vector, $\mathbf{d} = \mathbf{y}^o - \mathbf{H}(\mathbf{x}_b)$, over a 6-h assimilation window (09:00 UTC–15:00 UTC), and **(d)** After the 4D-Var minimization, the increment analysis is added to the first guess field (valid at 12:00 UTC) to obtain the analysis \mathbf{x}^a at 12:00 UTC.

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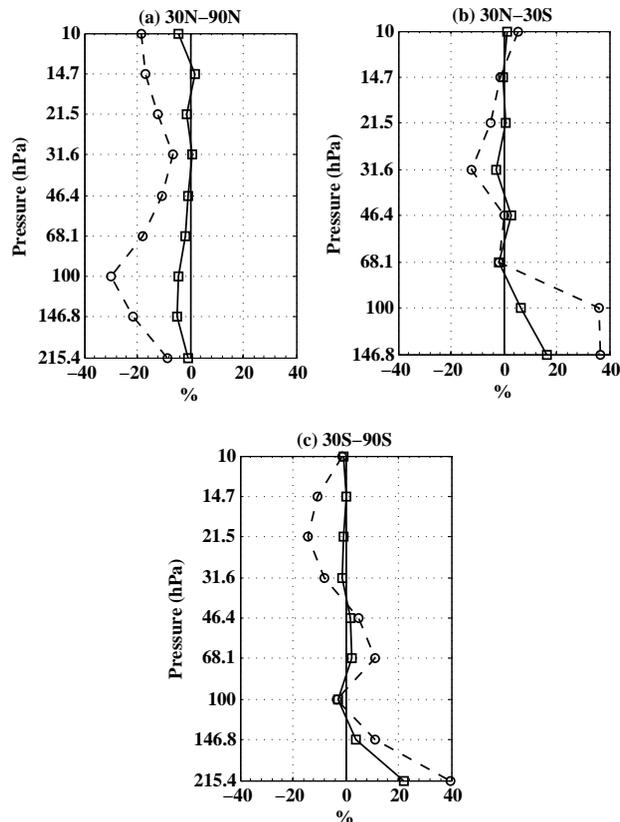


Fig. 4. $100 \times \frac{OMA}{y_o}$ (solid line with squares) and $100 \times \frac{OMF}{y_o}$ (dashed line with circles), both averaged in time from 23 January to 21 February 2006 (00:00 and 12:00 UTC analysis times) and over the following latitude bands **(a)** 30° N–90° N, **(b)** 30° N–30° S, and **(c)** 30° S–90° S.

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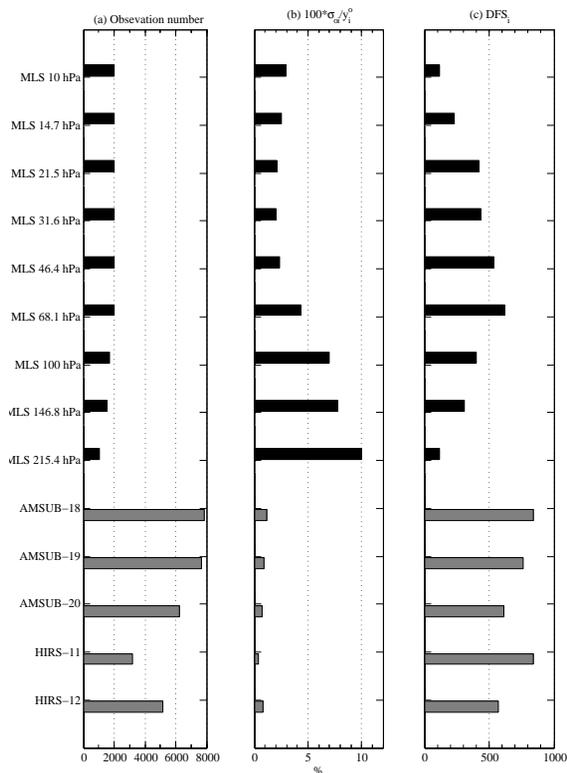


Fig. 5. (a) Observation number averaged in time from 23 January to 21 February 2006 (00:00 and 12:00 UTC analysis times) and over the globe for the following specific subsets: MLS ozone data on specific pressure levels and humidity-sensitive channel radiances from the HIRS sounder on NOAA–16 (channels 11 and 12) and the AMSU-B sounders on NOAA–16 and NOAA–17 (channels 18, 19 and 20), (b) $100 \times \frac{\sigma_{oi}}{y_{oi}^2}$ averaged over the globe and over the same period and for the same data as before, (c) DFS_i averaged over the same period and for the same data as before.

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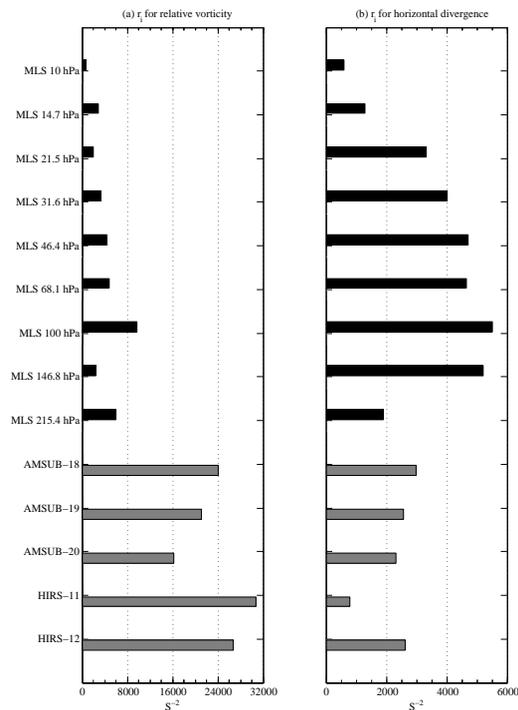


Fig. 6. r_i mean for relative vorticity **(a)** and horizontal divergence **(b)** brought by the following specific subsets: MLS ozone data on specific pressure levels and humidity-sensitive channel radiances from the HIRS sounder on NOAA-16 (channels 11 and 12) and the AMSU-B sounders on NOAA-16 and NOAA-17 (channels 18, 19 and 20). The mean is obtained by averaging the daily r_i (00:00 and 12:00 UTC analysis times) of a one-month period from 23 January 2006 to 21 February 2006.

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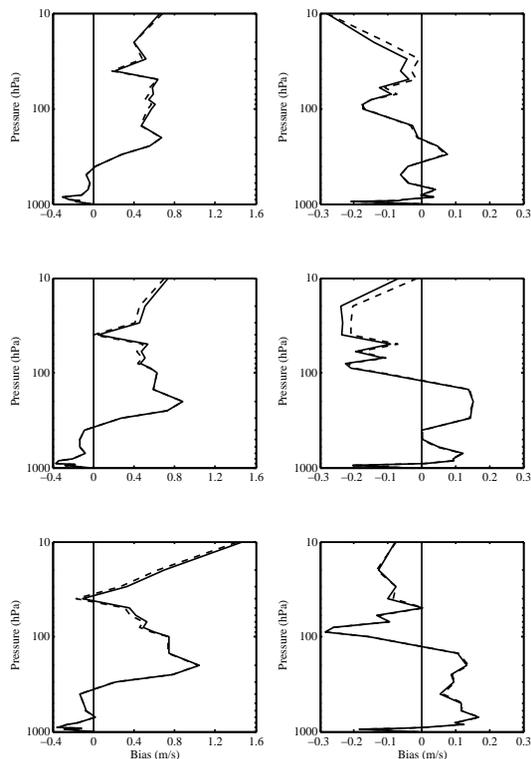


Fig. 7. Comparison of *ARPEGE* winds against radiosondes. Observation minus forecast residuals (OMF) are globally and time averaged from 23 January to 22 April 2006 for the CTL experiment (solid line) and the MLS experiment (dashed line) for both zonal (left hand-side panels) and meridional (right hand-side panels) wind components. The top, middle and bottom panels correspond to the first guess, 24-h and 48-h forecasts, respectively.

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