

Lightning activity in Brazilian thunderstorms during TROCCINOX: implications for NO_x production

H. Huntrieser¹, U. Schumann¹, H. Schlager¹, H. Höller¹, A. Giez², H.-D. Betz³,
D. Brunner^{4,*}, C. Forster^{5,**}, O. Pinto Jr.⁶, and R. Calheiros⁷

¹Institut für Physik der Atmosphäre, Deutsches Zentrum für Luft- und Raumfahrt (DLR),
Oberpfaffenhofen, Germany

²Flugabteilung, Deutsches Zentrum für Luft- und Raumfahrt (DLR), Oberpfaffenhofen,
Germany

³Physics Department, University of Munich, Germany

⁴Institute for Atmospheric and Climate Science, ETH Zurich, Switzerland

⁵Norwegian Institute for Air Research (NILU), Atmosphere and Climate Change Department,
Kjeller, Norway

⁶National Institute for Space Research, INPE, Brazil

⁷Instituto de Pesquisas Meteorológicas – Universidade Estadual Paulista, IPMet/UNESP,
Bauru, Brazil

14813

*now at: Laboratory for Air Pollution and Environmental Technology, Empa, Swiss Federal
Laboratories for Materials Testing and Research, Dübendorf, Switzerland

**now at: Institut für Physik der Atmosphäre, Deutsches Zentrum für Luft- und Raumfahrt (DLR),
Oberpfaffenhofen, Germany

Received: 12 September 2007 – Accepted: 5 October 2007 – Published: 16 October 2007

Correspondence to: H. Huntrieser (heidi.huntrieser@dlr.de)

Abstract

During the TROCCINOX field experiment in January and February 2005, the contribution of lightning-induced nitrogen oxides (LNOx) from tropical and subtropical thunderstorms in Southern Brazil was investigated. Airborne trace gas measurements (NO, NO_y, CO and O₃) were performed up to 12.5 km with the German research aircraft Falcon. During anvil penetrations in selected tropical and subtropical thunderstorms of 4 and 18 February, NO_x mixing ratios were on average enhanced by 0.7–1.2 and 0.2–0.8 nmol mol⁻¹ totally, respectively. The relative contributions of boundary layer NO_x (BL-NO_x) and LNO_x to anvil-NO_x were derived from the NO_x-CO correlations. On average ~80–90% of the anvil-NO_x was attributed to LNO_x. A Lightning Location Network (LINET) was set up to monitor the local distribution of cloud-to-ground (CG) and intra-cloud (IC) radiation sources (here called “strokes”) and compared with lightning data from the operational Brazilian network RINDAT (Rede Integrada Nacional de Detecção de Descargas Atmosféricas). The horizontal LNO_x mass flux out of the anvil was determined from the mean LNO_x mixing ratio, the horizontal outflow velocity and the size of the vertical cross-section of the anvil, and related to the number of strokes contributing to LNO_x. The values of these parameters were derived from the airborne measurements, from lightning and radar observations, and from a trajectory analysis. The amount of LNO_x produced per LINET stroke depending on measured peak current was determined. The results were scaled up with the Lightning Imaging Sensor (LIS) flash rate (44 flashes s⁻¹) to obtain an estimate of the global LNO_x production rate. The final results gave ~1 and ~2–3 kg(N) per LIS flash based on measurements in three tropical and one subtropical Brazilian thunderstorms, respectively, suggesting that tropical flashes may be less productive than subtropical ones. The equivalent mean annual global LNO_x nitrogen mass production rate was estimated to be 1.6 and 3.1 Tg a⁻¹, respectively. By use of LINET observations in Germany in July 2005, a comparison with the lightning activity in mid-latitude thunderstorms was also performed. In general, the same frequency distribution of stroke peak currents as for tropical thunderstorms over

14815

Brazil was found. The different LNO_x production rates per stroke in tropical thunderstorms compared with subtropical and mid-latitude thunderstorms seem to be related to the different stroke lengths (inferred from comparison with laboratory data and observed lengths). In comparison, the impact of other lightning parameters as stroke peak current and stroke release height was assessed to be minor. The results from TROCCINOX suggest that the different vertical wind shear may be responsible for the different stroke lengths.

1 Introduction

A general introduction to the LNO_x topic and overviews of past and present measurements of LNO_x in thunderstorms are given in accompanying papers by Huntrieser et al. (2007) (HH07) and by Schumann and Huntrieser (2007) (SH07). Observations from local field experiments have been extrapolated to the global scale to estimate the average amount of LNO_x produced annually over the globe which is one crucial, yet highly uncertain, parameter in the global NO_x budget. A LNO_x nitrogen mass source strength between 2 and 20 Tg a⁻¹ has frequently been given in the literature in the past (WMO, 1995; Bradshaw et al., 2000). More recently, lower values between 1 and 14 Tg a⁻¹ have been reported based on estimates from airborne and satellite measurements (Huntrieser et al., 2002; Beirle et al., 2004; Ridley et al., 2004; Boersma et al., 2005; Beirle et al., 2006; Ott et al., 2007b). Furthermore, chemical transport models (CTMs) have been used to reduce the LNO_x range by a comparison of modelled NO_x concentrations, for different LNO_x source strengths and vertical distributions, with local field and satellite measurements. The results obtained with model fits indicate best-estimate values for the global LNO_x nitrogen mass between 2 and 8 Tg a⁻¹ (SH07).

Different methods have been used to estimate the amount of LNO_x based on ground-based, airborne and laboratory measurements, and theoretical calculations, as reviewed by SH07. Airborne NO_x measurements can be combined with lightning observations to estimate the amount of LNO_x produced per flash or per metre flash length.

14816

These numbers have been scaled up with the mean flash length and the annual global flash frequency. A large uncertainty in the estimate of LNO_x still results from the assumed NO production rates by CG and IC flashes (Martin et al., 2007). It has been suggested that most components of a discharge produce NO_x with varying, not determined efficiencies (Chameides, 1986; Coppens et al., 1998; Dye et al., 2000; Rahman et al., 2007). In addition, it has been pointed out that the different flash lengths for CG and IC flashes may play an important role in the LNO_x production rate (Defer et al., 2003).

Results from the European Lightning Nitrogen Oxides Experiment (EULINOX) and Stratosphere-Troposphere Experiment: Radiation, Aerosols, and Ozone (STERAO) (DeCaria et al., 2000; Fehr et al., 2004; DeCaria et al., 2005; Ridley et al., 2005; Ott et al., 2007b) indicate that IC flashes produce about as much NO per flash as CG flashes (IC/CG production ratio 0.5–2). In addition, laboratory results from Gallardo and Cooray (1996) and model simulations from Zhang et al. (2003) support that IC and CG flashes are similarly energetic. On the other hand, laboratory studies by Wang et al. (1998) showed that LNO_x depends less on energy and more on atmospheric pressure and the peak current of the flash. They concluded that “NO production per metre discharge length as a function of peak current appears to provide a more appropriate scaling factor for estimates of total global NO production”. The present study makes use of this finding by combining Wang et al. (1998) NO_x measurements for laboratory flashes with our NO_x and lightning peak current measurements from the field. First results were briefly presented in Huntrieser et al. (2006), indicating differences for tropical and subtropical thunderstorms in Brazil, which are discussed here in more detail. A further study is in preparation by Ott et al. (2007a¹). The authors find that the mean peak currents and the NO production amounts per flash in five different thunderstorms decrease with increasing latitude: the lowest value of NO pro-

¹ Ott, L. E., Pickering, K. E., DeCaria, A. J., Stenchikov, G. L., Lin, F.-F., Wang, D., Lang, S., and Tao, W.-K.: Production of lightning NO_x and its vertical distribution calculated from 3-D cloud scale chemical transport simulations, in preparation, *J. Geophys. Res.*, 2007a.

14817

duction (360 moles/flash) was found for a EULINOX storm (48° N) and the largest value of NO production (700 moles/flash) was found for a CRYSTAL-FACE storm (26° N). Recently, Barthe et al. (2007) incorporated the relationship between produced LNO_x per m laboratory spark and atmospheric pressure according to Wang et al. (1998) in their simulations with an explicit electrical scheme and a 3-D mesoscale model (Meso-NH).

Up to now, only a few airborne experiments have been conducted that are suitable to provide an estimate of the LNO_x production rate in the tropics (see SH07). In this paper we present measurements from the “Tropical Convection, Cirrus and Nitrogen Oxides Experiment” (TROCCINOX) carried out in the wet season in January and February 2005 in the State of São Paulo and its surroundings in southern Brazil (10° S to 28° S and 38° W to 55° W). Both tropical and subtropical thunderstorms were investigated, since the operation area was located along the South Atlantic convergence zone (SACZ) (HH07). The main questions of this study are: 1.) How much LNO_x is produced by these tropical and subtropical thunderstorms? 2.) What are the relative contributions from strokes with different peak currents? 3.) How large is the LNO_x production rate per stroke or flash? 4.) Is this LNO_x production rate different for tropical and subtropical thunderstorms? 5.) What are the possible reasons for the difference? 6.) Can the findings from TROCCINOX help to explain the large LNO_x productivity observed in Florida thunderstorms during CRYSTAL-FACE?

To answer these questions we analyse airborne measurements of NO, NO_y, CO, and O₃ mixing ratios, the J(NO₂) photolysis rate and meteorological parameters performed in the outflow of thunderstorms, trajectory analyses with the FLEXPART model and measurements from LINET, which was set up during TROCCINOX to monitor the local lightning distribution (Sect. 2). This system registers very low and low frequency (VLF/LF) radiation sources (here called “strokes”) from both CG and IC flashes. LINET data are compared with data from the operational Brazilian lightning detection network RINDAT (Sect. 3) and with LIS data (Sect. 4). Airborne NO_x and ground-based lightning measurements are combined to give an estimate of the amount of LNO_x produced per LINET stroke, and as a function of peak current according to Wang et al. (1998)

14818

(Sect. 4). From the ratio between LIS and LINET during one overpass, the amount of LNO_x per LIS flash is determined. Knowing the global and annual LIS flash rate (44 ± 5 flashes s^{-1} , Christian et al., 2003; Christian and Petersen, 2005), the equivalent annual global LNO_x production rate based on individual TROCCINOX thunderstorms is estimated (Sect. 4). The different LNO_x production rates estimated in tropical and subtropical thunderstorms are investigated through a comparison of LINET measurements, e.g. frequency distributions of stroke peak currents and mean peak currents (Sect. 5). In addition, the lightning properties are compared with those in mid-latitude thunderstorms over Germany, where the same lightning location network (LINET) was set up in July 2005. Airborne NO_x measurements over Germany are available from previous campaigns (Huntrieser et al., 1998, 2002), but not for July 2005 (Sect. 5). The results are discussed and summarised in Sects. 6–7. The present study is the first to our knowledge that investigates whether tropical, subtropical and mid-latitude thunderstorms have different potentials to produce LNO_x by combining lightning peak current measurements with airborne NO_x and meteorological measurements.

2 Data and model description

For general information on the TROCCINOX field experiment, see the papers by Schumann et al. (2004), HH07 and SH07. The following subsections describe the airborne data obtained mainly from the research aircraft Falcon of the Deutsches Zentrum für Luft- und Raumfahrt (DLR) as well as partly from the Russian M55 Geophysica aircraft (Sect. 2.1), lightning data from LINET, LIS and RINDAT (Sect. 2.2) and model simulations from FLEXPART (Sect. 2.3). In addition, we use data from two S-band Doppler radars in Bauru (22.4° S, 49.0° W) and in Presidente Prudente (22.1° S, 51.4° W) operated by the Instituto de Pesquisas Meteorológicas (IPMet). Two different radar reflectivity products are presented: surveillance Plan Position Indicator (PPI, range 450 km) and 3.5 km Constant Altitude PPI (CAPPI, range 240 km). The meteorological environment of tropical, subtropical and mid-latitude thunderstorms was characterised with

14819

analysis data (temperature, water vapour mixing ratio, pressure, wind velocity and direction) from the European Centre for Medium Range Weather Forecasts (ECMWF) with 3 h temporal resolution, 1° horizontal resolution and 60 vertical levels. The equivalent potential temperature is calculated as described in HH07. The separation of tropical and subtropical air masses is based on meteorological data, as already discussed for the two selected flights of 4 and 18 February 2005 in HH07. The 4 and 18 February flights were classified as tropical and subtropical, respectively.

2.1 Airborne instrumentation: Falcon and Geophysica

Airborne measurements up to 12.5 km were carried out with the Falcon, which was equipped with DLR instruments to measure NO, NO_y, O₃, CO and J(NO₂). The chemical instrumentation is the same as that used during several DLR field campaigns in the past (HH07). Position, altitude, temperature, humidity, pressure and the 3-dimensional wind vector (*u*, *v* and *w*) were measured with the standard Falcon meteorological measurement systems (Schumann et al., 1995). Wind and pressure were measured with a Rosemount flow angle sensor (model 858) at the Falcon's noseboom tip. The aerodynamic measurements were analysed according to an extensive in-flight calibration programme (Bögel and Baumann, 1991).

In addition, NO and CO measurements were obtained from the high-flying Geophysica aircraft (~20 km) (Stefanutti et al., 2004). The SIOUX instrument, developed and operated by the DLR, measures the NO mixing ratio (chemiluminescence technique) with a time resolution of 1 s, and an accuracy and precision of 10% and 5%, respectively. The CO-TDL instrument (cryogenic Tunable Diode Laser technique) operated by the Istituto Nazionale di Ottica Applicata/Consiglio Nazionale delle Ricerche (INOA/CNR), measures the CO mixing ratio with an averaging time of 5 s, the accuracy and precision being 5% and 2%, respectively.

All flight altitude values refer to pressure height and all times to UTC (Coordinated Universal Time) time (see also HH07).

14820

2.2 Lightning measurements: LINET, LIS and RINDAT

During the TROCCINOX field campaign from 21 January to 27 February 2005, the VLF/LF (5–300 kHz) lightning detection network LINET was used to monitor the local lightning distribution with high spatial resolution. LINET was set up by DLR in cooperation with IPMet. The network included six sensors from DLR to observe the area 19.5–24.5° S and 46.5–51.5° W (see Table 1). The average distance to the next closest sensor was ~80 km. For comparison of lightning characteristics, measurements from southern Germany in summer 2005 with 19 sensors (from both DLR and the University of Munich), monitoring the area 47–51° N and 5–14° E, were also included in this study (Table 1). The average distance to the next closest sensor was ~80 km in the outer region and ~20 km in the inner region. The basically similar features of the LINET arrays in Germany and Brazil allow comparison of the characteristics of thunderstorms systems in both regions (Schmidt et al., 2005).

The LINET system has been developed by the University of Munich and the sensor technology and measurement procedures have been described in detail by Betz et al. (2004), Schmidt et al. (2004, 2005), Betz et al. (2007) and Schmidt (2007). For an overview of system characteristics see SH07. LINET continuously measures the transient magnetic components of VLF/LF emissions from lightning discharges. These signals are emitted by certain components of the flashes, and therefore a direct comparison with published flash statistics (e.g., IC/CG ratio) is not possible. At the current stage, VLF sources are considered separately. A routine algorithm to combine them into flashes is under development. In Sect. 6a small set of strokes were combined manually to flash “components”, which indicate that LINET locates few VLF strokes per flash components (on average 3 and up to 9).

It is known that the amplitude of a measured electromagnetic signal is proportional to the peak current (Uman et al., 1975; Rakov et al., 1992; Cummins et al., 1998; Orville, 1999; Jerauld et al., 2005; Schulz et al., 2005). Thus, the peak current of LINET strokes is estimated from the VLF pulse amplitude. The registered amplitude depends on the

14821

distance between the VLF pulse and the measuring LINET sensor. The registered pulse is normalised by the reciprocal value of the distance between pulse source and sensor, and averaged over all sensors that registered the VLF pulse. Owing to refined antenna techniques, optimised waveform handling and a shorter sensor base line of <100 km, a high detection efficiency of low peak currents is possible. The detection efficiency, stroke-current dependent, is highest in the LINET centre area (2°×2°) and decreases rapidly down to 30% towards the periphery. Currents as low as ~1–2 kA can be detected by the system within the LINET centre area (periphery ~5 kA). In comparison, most other VLF/LF lightning networks report only strokes >5–10 kA (Cummins et al., 1998).

In addition to LINET data, spaceborne measurements from LIS on board the Tropical Rainfall Measurement Mission (TRMM) satellite (Christian et al., 1999; Thomas et al., 2000; Boccippio et al., 2002) were used to estimate the total regional flash density (sum of CG and IC flashes) over the TROCCINOX area. For an overview of system characteristics see SH07. Locally, the duration of a measurement is 90 s, which is long enough to estimate the flashing rate of most thunderstorms in the field of view during the passage (see <http://thunder.msfc.nasa.gov/lis/>). At noon the detection efficiency is 73±11% and at night 93±4% (Boccippio et al., 2002). Here we used LIS science products (total count of flashes) from the “LIS space time domain search” (see <http://thunder.nsstc.nasa.gov/lightning-cgi-bin/lis/LISSearch.pl>). A recent comparison between LIS and LINET data showed a good agreement between two systems that are based on completely different measurement techniques (Schmidt et al., 2005). Here LIS data for one overpass of 4 February 2005 were compared with LINET data (see Sect. 4.5). LNOx estimates per LINET stroke were scaled up with LIS observations to provide an estimate of the regional and global strength of the LNOx production.

LINET data were also compared with data from the operational Brazilian lightning detection network RINDAT (see <http://www.rindat.com.br/>). Like LINET, RINDAT operates in the VLF/LF range. The detection efficiency for strokes with peak currents above 10 kA is 80–90% and the location accuracy is 0.5–2.0 km (Pinto and Pinto, 2003). The

14822

RINDAT system mainly registers CG flashes. A recent comparison between LINET and RINDAT strokes indicates reasonable agreement for CG strokes when LINET peak currents are above 12 kA (Schmidt et al., 2005); see further comparisons in Sect. 3.

2.3 Transport modelling: FLEXPART

5 The distribution of LNO_x in the vicinity of thunderclouds was simulated with the Lagrangian particle dispersion model FLEXPART. General information on FLEXPART used for TROCCINOX is given in HH07. The model has mainly been used for studying long-range air pollution transport (e.g. Stohl et al., 2003a, b; Huntrieser et al., 2005), but also to investigate LNO_x transport (Stohl et al., 2003b; Beirle et al., 2006; Cooper
10 et al., 2006).

The LNO_x emissions used here as input for FLEXPART were based on lightning stroke data from the LINET system. LNO_x was released uniformly in the vertical between 5 km (freezing level with negative charge centre) and 13 km altitude (cloud top) at the accurate horizontal position of observed VLF sources. Because of the low resolution
15 of the ECMWF wind fields (0.5° horizontally) used as input for FLEXPART (HH07), the distribution of lightning sources is assumed to be uniform in the vertical. A hundred particles were released per stroke, carrying the mass of LNO_x produced (here set to 1 kg). The convection scheme, used in these FLEXPART applications, transports the particles upward into the anvil, from where they follow trajectories computed with the
20 ECMWF wind fields. No quantitative estimate of the amount of LNO_x is possible from these simulations; they can, however, be used to estimate the extension of the LNO_x field advected out of the anvil region.

3 Observations during the field experiment

An overview of the observations on the two selected TROCCINOX days, 4 and 18
25 February 2005 with thunderstorms in tropical and subtropical air masses, respectively,

14823

is given in HH07. Here we briefly focus on the performance of the lightning detection network LINET (Sect. 3.1) and on the representation of the Falcon measurements in the anvil outflow (Sect. 3.2).

3.1 Performance of LINET compared with RINDAT

5 To evaluate the performance of the LINET system in more detail, a comparison with the operational lightning detection network in Brazil (RINDAT) was carried out for 4 and 18 February 2005. Horizontal distributions of RINDAT and LINET strokes were compared for the LINET centre area on 4 February 2005 (Fig. 1a), and for the northern
LINET periphery area on 18 February 2005 (Fig. 1b), 00:00 UTC–24:00 UTC. Overall,
10 a general agreement was found, but with a slight shift of RINDAT strokes to the west compared with LINET strokes. In some areas the density of LINET strokes was much larger than of RINDAT strokes (probably because IC strokes and strokes with low peak are not registered by RINDAT). The correlations between LINET and RINDAT peak currents (absolute values) for 222 and 173 selected strokes of 4 February (21.5–22.5° S
15 and 48.5–49.5° W) and 18 February (19.4–20.0° S and 47.7–49.2° W) are shown in Figs. 2a and b, respectively. On 4 February only negative CG strokes were compared. On 18 February no separation between CG and IC strokes was possible for the LINET data, owing to the location of the selected strokes along the northern periphery area. LINET strokes were therefore compared with both positive and negative CG
20 strokes from RINDAT. About 10% of the selected LINET strokes were positive ones. The strokes shown in Fig. 2 were selected manually to represent peak current values over the entire current range. LINET peak currents above 13–14 kA are in general also detected by RINDAT, occasionally even LINET peak currents down to 7 kA. The slope (0.83) seen in Fig. 2a indicates that a 20 kA LINET stroke is on average registered as
25 14 kA by RINDAT. In Fig. 2b the slope is slightly lower (0.68) owing to the lower LINET detection efficiency along the northern periphery. In addition, the mean peak current is higher for LINET strokes (35 kA in Fig. 2a and 31 kA in Fig. 2b) than for RINDAT strokes (27 kA in Fig. 2a and 20 kA in Fig. 2b). The high correlation coefficient ($r^2 =$

14824

0.92 and 0.95, respectively) between the peak currents of the two systems indicates a good agreement in general. Lower RINDAT CG+ peak currents (<30 kA) are frequently registered as IC+ by LINET (~40%) and stronger RINDAT CG- peak currents (>100 kA) are frequently registered as IC- by LINET (~40%). This finding can be compared with results from EULINOX in Germany where flashes registered with a LPATS system (same technology as used for RINDAT) were compared with the French Office National d'Etudes et de Recherches Aéropatiales (ONERA) VHF interferometer measurements. Théry (2001) found that 61% of the positive LPATS flashes (those of low intensity) and 32% of the negative LPATS flashes were in fact IC flashes. A recent study by Pinto et al. (2007) also confirmed that a large percentage of the positive CG flashes registered by RINDAT over Brazil are in fact IC flashes. For the analysed dataset we found that weak positive RINDAT peak currents (<10 kA) are occasionally (<10%) registered as negative strokes by LINET.

3.2 NO_x in the anvil outflow derived from aircraft measurements

In HH07 it was briefly discussed whether the outflow altitude where LNO_x maximises was reached with the Falcon aircraft (important question for comparison with results from other field campaigns and for further calculations in Sect. 4). It was concluded that this altitude was reached with certainty on 18 February, but on 4 February the Falcon measured the largest mixing ratios in the uppermost flight levels so that larger mixing ratios at higher altitudes inside the anvil cannot be excluded. Therefore, for the latter day measurements from the high-flying Geophysica in the upper part of this thunderstorm were briefly analysed as discussed below.

The Falcon measurements in two of the anvils of 4 February (anvil 1a and 5a, listed in Table 2a and described in Sect. 4.1) can be compared with coincident measurements with the high-flying Geophysica. The Geophysica penetrated anvil 1a during ascent between 15.9 and 16.6 km (penetration at flight time: 67 070–67 298 s, at position: 21.3–21.5° S and 49.1–49.3° W) and anvil 5a during descent between 17.2 and 16.5 km (penetration at flight time: 66 569–66 696 s, at position: 21.8–21.9° S and

14825

48.5–48.7° W). The mean NO mixing ratios in anvil 1a and anvil 5a were 0.30 and 0.35 nmol mol⁻¹, respectively. The mean anvil-NO mixing ratio is the mean value of all NO 1s-values measured between the entrance and exit of the anvil (determined from the distinct increase and decrease in the NO mixing ratio). The mean mixing ratios measured by the Geophysica between ~16–17 km altitudes are distinctly lower than the NO mixing ratios measured by the Falcon at lower altitudes (10.6–10.7 km): 0.80 and 1.16 nmol mol⁻¹ in anvil 1a and anvil 5a, respectively. The Geophysica measurements in the anvils on 4 February indicate an increase in NO mixing ratios with decreasing altitude, opposite to the Falcon measurements. Hence the outflow level where NO mixing ratios maximise was likely to be located between the altitudes at which the Falcon and Geophysica penetrated the anvils. CO measurements from the Geophysica (personal communication P. Mazzinghi, INOA/CNR) can be used to determine this outflow level more precisely (on the assumption that LNO_x maximises where CO maximises). The vertical CO profile (ascent and descent in the vicinity of the selected anvils) shows enhanced mixing ratios mainly between ~10–14 km altitudes. The mixing ratios were rather constant throughout this layer, ~130–140 nmol mol⁻¹. The outflow level, where the CO mixing ratio maximises (132–138 nmol mol⁻¹), was located between ~12.0–12.5 km, about 1.5–2 km above the Falcon penetration. At the levels where the Falcon penetrated the anvils (10.6–10.7 km), however, the CO mixing ratio (132 nmol mol⁻¹) was similar to the lowest Geophysica CO mixing ratios in the outflow level. The Falcon data may therefore underestimate the mean NO mixing ratios in the selected anvils to a degree which cannot be quantified from the available dataset. These mean NO mixing ratios are needed for further calculations in the next section. Preliminary results from cloud-resolved modelling for the 4 February thunderstorms by Pickering et al. (2007) suggest that the anvil outflow NO maximum is located between 12 and 13 km, which supports our estimates derived from the vertical CO profile.

14826

4 Estimate of the LNO_x production rate per flash and per year

In this section the measurements in selected tropical and subtropical thunderstorms of 4 and 18 February 2005 are discussed in more detail. The spatial and temporal distributions of LINET strokes are presented (Sect. 4.1). The contribution from observed LINET strokes to measured anvil-NO_x mass and the resulting LINET stroke rates are estimated (Sect. 4.2). Furthermore, the contribution of BL-NO_x and LNO_x to measured anvil-NO_x is estimated (Sect. 4.3). The horizontal LNO_x mass flux rate out of the anvils is calculated by means of estimated LNO_x mixing ratios and horizontal outflow wind velocities from the flights combined with the size of the vertical cross-section of the anvils (Sect. 4.4). LNO_x nitrogen mass flux rates (g s^{-1}) and LINET stroke rates (strokes s^{-1}) are combined to estimate the production rate of LNO_x (in g of nitrogen mass or number of NO_x molecules) per LINET stroke and per LIS flash (Sect. 4.5). Finally, the annual global LNO_x nitrogen mass production rate is estimated (in Tg a^{-1}). Figure 3 gives an overview of these different steps described in detail in the following subsections, starting with the selection of a thunderstorm (TS) and ending with an estimate of the annual global LNO_x production rate $G(\text{LNO}_x)$.

4.1 Spatial and temporal LINET stroke distributions

The spatial distributions of LINET strokes of 4 and 18 February 2005 are shown in Figs. 4a and b, respectively. For the selected thunderstorms, strokes occurring before the penetrations by the Falcon are highlighted in colour. Superimposed is the Falcon track showing the successful, repeated penetrations of the subtropical thunderstorm system of 18 February, and the zigzag pattern between the tropical thunderstorms (labelled 1a, 5a and 2b) of 4 February. The time periods of the anvil penetrations are listed in Table 2a. The direction of the thunderstorm movement (red arrows in Fig. 4) is inferred from lightning data. The main wind direction in the anvil outflow (green arrows), as inferred from Falcon wind measurements, controls the transport of LNO_x out of the anvils. On 4 February the main wind direction in the flight level (influenced

14827

by the Bolivian High, see Fig. 5c in HH07) varied between north-east and south-east in vicinity of anvil 1a and 2b, and was from the south-west in the vicinity of anvil 5a. The NO_x mixing ratio along the flight track is also superimposed in Fig. 4. Elevated mixing ratios exceeding $0.6 \text{ nmol mol}^{-1}$ NO_x were frequently measured in the anvil outflow downstream of nearby lightning strokes.

The selected thunderstorms of 4 and 18 February occurred in the centre and at the northern border line of the LINET network, respectively. Because of a higher sensitivity in the network centre, the fraction of strokes with low currents ($<10 \text{ kA}$) was much higher on 4 February (87%) than on 18 February (45%). For the latter thunderstorm system no separation between IC and CG strokes was possible because of the large distance from the centre. For an adequate comparison of the stroke rates in these storms, it was necessary to restrict comparisons to higher stroke peak currents ($\geq 10 \text{ kA}$) which were observed with about the same detection efficiency, independently of their location within the LINET network. On 4 February strokes were widespread with some at the LINET periphery. LINET strokes were therefore compared with LIS flashes and RINDAT strokes to determine the detection efficiency of the LINET system relative to the other two systems. The change in detection efficiency for these selected LINET strokes towards the LINET periphery was only minor ($<10\%$) compared with the other two systems and not considered further.

The temporal distributions of LINET stroke rates in the selected thunderstorms for peak currents $\geq 10 \text{ kA}$ are presented in Fig. 5. The storms of 4 February were mainly in a mature stage during the aircraft passage. In comparison, the long-lived storm system of 18 February was in a decaying stage and probed long after the peak lightning activity (first lightning was registered already 6 h before the first penetration).

4.2 Contribution of LNO_x to anvil-NO_x and determining LINET stroke rates

For evaluation of the LNO_x production rate per stroke, it is necessary to estimate which of the LINET strokes during the storm lifetime contributed to the measured anvil-NO_x enhancement and its horizontal and vertical extension. This is a very difficult task

14828

which might be best performed by using cloud-scale modelling. First cloud-resolving model simulations have been performed for selected TROCCINOX storms of 4 February 2005 by Chaboureau et al. (2007) and by Pickering et al. (2007) but cloud-resolving simulations are presently not available for the thunderstorm system of 18 February 2005. Instead we make use of FLEXPART lightning tracer simulations, as explained in Sect. 2.3. The simulations follow lightning tracers from the horizontal LINET stroke distributions (Fig. 4b) using ECMWF wind fields. For the long-lived thunderstorm case of 18 February, ambient wind velocities were strong and LNO_x was advected far downwind. The ECMWF wind agrees well with Falcon measurements of wind velocity and direction, except in the core of the anvil penetrations (Fig. 6). As shown later in this section, comparison of the ECMWF wind fields and FLEXPART results with radar and airborne wind and anvil-NO_x observations, as indicated in Fig. 3, supports the validity of the FLEXPART simulations of the 18 February thunderstorm system, in spite of the coarse horizontal resolution (0.5°) of the ECMWF wind velocity fields used.

For the thunderstorm system of 18 February, a time sequence of FLEXPART lightning tracer simulations (Fig. 7) indicates a rather fast development of an elongated area with enhanced LNO_x downwind of the storm system, following the wind in the upper troposphere (UT). Tracer distributions for six different simulations are shown in this figure (output resolution: 30 min and 0.08 degrees, horizontal cross-sections at 10 km altitude corresponding to the flight level) considering transport of emissions from strokes in various time intervals. The simulated tracer distributions may be compared with the anvil-NO_x observations from the Falcon (Fig. 4b). Only the last four simulations (Fig. 7c–f) indicate distinctly enhanced LNO_x along the right anvil transect, as observed by the Falcon. Furthermore, mixing ratios in the left transects, closer to the core, were twice as high as in the right transects. Given the measured UT wind velocity of 15 to 20 m s⁻¹, it is clear that strokes that occurred between 19:00–19:30 UTC (along the left anvil transect, 49.7–49.8° W) or earlier do not contribute to the anvil-NO_x enhancement observed along the right anvil transect. The air with enhanced LNO_x is advected further downwind to the right in Figs. 7d–f. Only strokes after 19:30 and be-

14829

fore 20:55 UTC (when the storm decayed) (Fig. 7c) were therefore considered to have contributed to the observed anvil-NO_x. During this 85 min period about 130 strokes with peak currents ≥ 10 kA were detected, corresponding to a stroke rate of 0.025 strokes s⁻¹ (Table 2a).

For the 18 February 2005 thunderstorm system, the horizontal extension (Δx) of FLEXPART lightning tracer in Fig. 7c, perpendicular to the wind direction (see Fig. 4b), was estimated to be ~ 30 –35 km. This width agrees well with the extension of the flight path segment with enhanced NO_x observed during the single anvil transects (28–35 km); see the grey scale along the flight track in Fig. 4b and Table 2. This parameter (Δx) will be used to estimate the horizontal LNO_x mass flux out of the anvil in Sect. 4.4.

Finally, a radar image of the 18 February thunderstorm system (Bauru radar, elevation angle 0°), indicates a pronounced, elongated structure of the storm system (Fig. 8), similar to the FLEXPART result at 10 km altitude. The 18 February thunderstorm system is located in the upper, northern domain of the radar range, about 240 km from the radar site. Unfortunately, the radar information is sparse in this region and no more detailed data are available since the domain is out of the quantification range where volumetric data are collected.

For the thunderstorms of 4 February, no FLEXPART simulations were performed since the storms just developed ~ 1 h ago and this time was considered too short for realistic simulations. In addition, the ambient UT wind velocities were low (4–7 m s⁻¹) and LNO_x remained in the vicinity of the storms. Instead, as indicated in Fig. 3, the LNO_x production rate per stroke and the width Δx were estimated from a combination of horizontal LINET stroke distributions, radar images, Falcon wind and anvil-NO_x observations. The average altitude (arithmetical mean) of all IC strokes (Table 4a) in anvil 1a (10.0 km) and anvil 5a (11.6 km) was below or just above the flight level (10.6 km and 10.7 km, respectively; see Table 2a), indicating that the majority of LNO_x, produced by the observed strokes left the anvil at about the flight level.

It is assumed that all LINET strokes observed in the vicinity of these storms between storm initiation and Falcon penetration (coloured in Fig. 4a) contributed to the observed

14830

anvil-NO_x (Fig. 4a). (This is a working hypothesis with large uncertainties, which cannot be quantified without cloud-model simulations.) For comparison with the 18 February case, only the number of strokes with peak currents ≥ 10 kA is counted. In anvil 1a about 278 strokes were registered between 16:55 and 18:20 UTC, in anvil 5a about 130 strokes were registered between 18:05 and 18:45 UTC and in anvil 2b about 311 strokes were registered between 17:55 and 19:20 UTC; see Fig. 5, which corresponds to the following stroke rates: 0.055, 0.054 and 0.061 strokes s⁻¹, respectively (see Table 2a).

For each of the three anvil penetrations, the width (Δx) of the LNO_x plume perpendicular to the wind direction was estimated from the horizontal LINET stroke distribution, from the anvil-NO_x observations (Fig. 4a) and from the radar images at the time of the penetrations (shown only hourly in Fig. 9). The Δx values are ~ 35 , ~ 25 and ~ 45 km for anvils 1a, 5a and 2b, respectively (see Table 2a).

4.3 Contribution of BL-NO_x to anvil-NO_x

The boundary layer (BL) contribution ($\chi_{\text{BL-NO}_x}$) to the NO_x mixing ratio in the anvil ($\chi_{\text{Anvil-NO}_x}$) is derived from the correlation between NO_x and CO mixing ratios in the BL and in the anvil. It is assumed, that BL air is transported upwards rapidly within strong, well-developed updrafts with little ambient mixing and without chemical loss of NO_x and CO. Hence, about the same CO mixing ratio is observed in the main anvil outflow ($\chi_{\text{Anvil-CO}}$) as in the BL layer ($\chi_{\text{BL-CO}}$):

$$\chi_{\text{Anvil-CO}} = \chi_{\text{BL-CO}} \quad (1)$$

and LNO_x (χ_{LNO_x}) is the difference between anvil-NO_x and BL-NO_x:

$$\chi_{\text{LNO}_x} = \chi_{\text{Anvil-NO}_x} - \chi_{\text{BL-NO}_x} \quad (2)$$

These assumptions are supported by cloud-model simulations (Pickering et al., 1992; Thompson et al., 1997; Ott et al., 2007b) and airborne thunderstorm observations (Dickerson et al., 1987; Hauf et al., 1995; Huntrieser et al., 1998; Höller et al., 1999;

14831

Huntrieser et al., 2002; Lopez et al., 2006; Bertram et al., 2007; Koike et al., 2007). The ratio of NO_x to CO in the BL (<2 km) is conserved during the rapid upward transport into the anvil:

$$\chi_{\text{BL-NO}_x} / \chi_{\text{BL-CO}} = (\chi_{\text{Anvil-NO}_x} - \chi_{\text{LNO}_x}) / \chi_{\text{Anvil-CO}} \quad (3)$$

Vertical NO_x, CO, and O₃-profiles from the 18 February flight are shown in Fig. 10a. The CO mixing ratios, measured during the anvil penetrations, are in a similar range as those measured at ~ 2 km altitude (see red box), supporting the assumption of rapid upward transport from the top of the BL into the anvils. Unfortunately, no NO_x measurements are available below 3 km for this flight. Instead, NO_x measurements in the BL were only available for ten TROCCINOX “fair weather” flights without active thunderstorms (Fig. 1a in HH07). NO_x and CO data from all available flights in the BL (<2 km) were therefore used to estimate the average BL NO_x-CO correlation. It can be justified that this relationship is representative, since CO mixing ratios in the BL were in the same range both for “thunderstorm” and for “fair weather” flights. The BL data were sampled mainly during take-off and landing near the campaign base. Hence, it was assumed that these values are representative for the entire BL covered by the selected flights.

In Fig. 10b the correlation between measured NO_x and CO for the Falcon flight of 18 February is shown (black dots). Different types of air mass origin (Pacific, Amazon basin, anvil and background), as discussed in HH07, are marked. The measured NO_x mixing ratios were mainly below 0.2 nmol mol⁻¹, except during the anvil penetrations. The average NO_x-CO correlation in the BL for all TROCCINOX flights (data from Fig. 1 in HH07) is also shown in Fig. 10b (red-yellow dots). Average CO mixing ratios during the anvil penetrations of the 18 and 4 February flights were 95–105 and 105–115 nmol mol⁻¹, respectively. From the measured BL-CO (90–120 nmol mol⁻¹) and the correlation, the average BL-NO_x mixing ratio and its standard deviation (std) were estimated to be 0.11 ± 0.07 nmol mol⁻¹. For the anvil penetrations of 4 and 18 February listed in Table 2a, average LNO_x volume mixing ratios (χ_{LNO_x}) were determined by sub-

traction of the mean BL-NO_x contribution (0.11 nmol mol⁻¹) from the mean anvil-NO_x values. The mean values for anvil-NO_x range between 0.2–0.8 nmol mol⁻¹ in the subtropical thunderstorm of 18 February and between 0.7–1.2 nmol mol⁻¹ in the tropical thunderstorms of 4 February (Table 2a in HH07). As a result, χ_{LNO_x} values in the range
5 from 0.1 to 1.1 nmol mol⁻¹ were obtained, as listed in Table 2a in the present paper.

Overall, the contribution of BL-NO_x to anvil-NO_x in the selected thunderstorms of 4 February (anvil 1a, 5a, and 2b) and 18 February (only anvil penetrations I, III and V closest to the core considered here) was ~10–20%. This range is slightly lower than the average found in European thunderstorms with 25 to 40% (Huntrieser et al., 1998,
10 2002). In the investigated TROCCINOX thunderstorms, the contribution from LNO_x clearly dominated the anvil-NO_x budget with ~80–90%. This contribution is higher than observed during the TRACE-A experiment at the end of the dry (burning) season, where only 30–40% of anvil-NO_x was attributed to LNO_x (Pickering et al., 1996).

4.4 Estimate of the horizontal LNO_x mass flux

15 Cloud-model simulations indicate that most LNO_x produced in a thunderstorm is transported into the anvil (Skamarock et al., 2003). If the total LNO_x mass in the anvil region (dependent on the LNO_x mixing ratio and the volume covered by this LNO_x) and the total number of flashes in the thunderstorm that contributed to this LNO_x were known, the LNO_x production rate per flash could be estimated, assuming a constant
20 LNO_x production per flash. Up to now, however, no method exists which can determine the required parameters exactly. Model approaches have e.g. estimated the horizontal NO_x flux out of the anvil through a vertical control surface (Skamarock et al., 2003; Barthe et al., 2007). A combination of in situ aircraft observations and cloud-model
25 simulations was used to separate the outflow flux into a LNO_x flux and an environmental NO_x flux. This approach was originally introduced by Chameides et al. (1987) for airborne measurements in thunderstorms during GTE/CITE and has also been applied by us for measurements in LINOX and EULINOX thunderstorms (Huntrieser et

14833

al., 1998, 2002). Alternatively, the NO content in the thunderstorm is estimated from the product of airborne in situ measurements of NO at certain levels in the anvil and the estimated volume of the appropriate cloud segments (Ridley et al., 2004). The total volume is derived from the sum of the vertically staggered flight segments. The two
5 methods are described in more detail in SH07.

The TROCCINOX thunderstorm penetrations listed in Table 2a provide only snapshots of the conditions at a certain level of the cloud at a certain time. It is not known how representative these anvil penetrations are for the average anvil conditions (see also discussion in Sect. 3.2). These are, however, the only measurements that are
10 available. Time series of trace gas measurements (NO_x, CO, and O₃) during the penetrations listed in Table 2a have already been presented and discussed in HH07. On the 18 February flight, the anvil outflow from the selected thunderstorm system was successfully penetrated 6 times (Fig. 4b). In addition to the mentioned trace gases, NO_y was measured and mixing ratios during the 6 penetrations are shown together with the
15 vertical velocity (absolute values) in Fig. 11. The 1 s absolute velocity values mainly varied between 0.1 and 1.0 m s⁻¹ indicating that the measurements were carried out outside the core region of the thunderstorm cell, where far higher vertical velocities are to be expected. The highest NO_y mixing ratios were measured during the anvil penetrations with the strongest vertical velocities, which is closer to the core region
20 (~10–30 km) where most lightning occurs (penetration I, III and V). The closest penetration to the maximum anvil outflow level was penetration III, where the mean updraft velocity (0.8 m s⁻¹) was distinctly higher than the mean downdraft velocity (0.2 m s⁻¹), and the highest mean NO_y mixing ratio (1.1 nmol mol⁻¹) was measured. About 30 km further downwind (penetration II, IV, VI), the measurements indicate that a large part
25 of the outflow already mixed with the ambient air (similar mean updraft and downdraft velocities).

Moreover, on 4 February the selected thunderstorms were penetrated only once, but rather close to the core. Hence, too few repeated anvil penetrations and limited radar reflectivity data are available to apply the method introduced by Ridley et al. (2004). We

14834

therefore use a modified version of the method introduced by Chameides et al. (1987) and assume that the measurements during each anvil penetration (snapshots) are representative average anvil conditions. We consider the horizontal mass flux of LNO_x through a vertical control surface. The vertical surface dimensions can be estimated from the combination of e.g. airborne measurements and FLEXPART simulations as explained before in Sect. 4.2 and as indicated in Fig. 3. Repeated penetrations of the 18 February thunderstorm system indicated that Δz was >1.3 km (10.7–9.4 km, Table 2a). The entire vertical extent of the anvil outflow can be most clearly seen in vertical profiles of the CO mixing ratio measured by the high-flying Geophysica (personal communication, P. Mazzinghi, INOA/CNR). On 18 February the most distinct enhancement in the CO mixing ratio was observed between ~ 9 –12 km altitudes and Δz was set to ~ 3 km (see Table 2a). On 4 February the enhancement in CO was less clear owing to elevated background mixing ratios: enhanced mixing ratios were mainly observed between ~ 10 –14 km altitudes and Δz was set to ~ 4 km (see Table 2a).

The horizontal LNO_x mass flux F_{LNO_x} (in nitrogen mass per time, g s^{-1}) was calculated for each thunderstorm penetration listed in Table 2a according to:

$$F_{LNO_x} = \chi_{LNO_x} \cdot \frac{M_N}{M_{air}} \cdot \rho_a (V_a - V_s) \cdot \Delta x \cdot \Delta z \quad (4)$$

where χ_{LNO_x} is the mean NO_x volume mixing ratio produced by lightning (mol mol^{-1}), M_N and M_{air} are the molar masses of nitrogen (14 g mole^{-1}) and air (29 g mole^{-1}), respectively, ρ_a is the air density (g m^{-3}) calculated from measured temperature and pressure in the anvil, and $V_a - V_s$ is the difference between the wind vectors in the anvil outflow and at the steering level (see Table 2b). The last term $\Delta x \cdot \Delta z$ is the area (m^2) of the vertical cross-section perpendicular to the wind direction in the anvil outflow. In general, the wind at the steering level (~ 700 hPa) determines the mean motion of a thunderstorm cell (Keenan and Carbone, 1992), but this parameter is not available from the airborne measurements. Instead, horizontal LINET stroke distributions, as shown in Fig. 4, were plotted with a higher temporal resolution (10 min) and the storm

14835

motion (V_s) was determined from the temporal stroke evolution. The parameters in Eq. 4, except Δx (Sect. 4.2) and Δz , were calculated directly from Falcon measurements by averaging the measured data over the time period when the thundercloud was penetrated (between entry and exit of anvil), see Table 2a.

F_{LNO_x} values were calculated for the selected thunderstorms by insertion of the parameters listed in Table 2a into Eq. 4, which give nitrogen mass flux values between 48 and 178 g s^{-1} (Table 2a). The flux values for subtropical thunderstorms (only anvil I, II, and V considered) and tropical thunderstorms are within a similar range. The flux values in Table 2a can be divided by the molar mass for nitrogen and the area of the vertical cross-section ($\Delta x \cdot \Delta z$) to estimate the flux in the unit $\text{mol m}^{-2} \text{ s}^{-1}$. The range of these fluxes, 3.3 – $7.1 \times 10^{-8} \text{ mol m}^{-2} \text{ s}^{-1}$, is well comparable to nitrogen mass flux values simulated by Barth et al. (2007) who ran different cloud-scale models (range 2.7 – $13.0 \times 10^{-8} \text{ mol m}^{-2} \text{ s}^{-1}$) and to Barthe et al. (2007), who simulated $6 \times 10^{-8} \text{ mol m}^{-2} \text{ s}^{-1}$ on average in the anvil outflow of a STERAO storm.

The parameters listed in Table 2a have large uncertainties. The relative maximal error of the F_{LNO_x} estimate was therefore calculated. The uncertainty for χ_{LNO_x} is given by the standard deviation (on average $\sim 50\%$ of the mean value); for $V_a - V_s$ the standard deviations listed in Table 2b indicate an uncertainty of up to $\sim 50\%$; for Δx the uncertainty was ~ 5 – 10 km corresponding to $\sim 40\%$; and for Δz the vertical anvil extension on 4 February varied between 3.5 – 6 km and on 18 February between 2 – 4 km indicating an uncertainty up to $\sim 50\%$. Summing up these uncertainties, the relative maximal error of the F_{LNO_x} estimate is $\sim 190\%$.

4.5 Estimate of the LNO_x production rate per stroke and per year

For the estimate of the LNO_x production rate P_{LNO_x} (nitrogen mass per stroke, in g stroke^{-1}), the horizontal LNO_x mass flux F_{LNO_x} (g s^{-1}) is divided by the LINET stroke

rate R_{LINET} (strokes s^{-1}):

$$P_{\text{LNOx}} = \frac{F_{\text{LNOx}}}{R_{\text{LINET}}} \quad (5)$$

P_{LNOx} estimates for the selected anvil penetrations resulted in values between 1.9 and 5.6 kg stroke $^{-1}$, see Table 2a. Unfortunately, the dataset in Table 2a is very sparse. Nevertheless, mean P_{LNOx} values for three tropical and one subtropical thunderstorms (only anvil penetrations I, III and V considered) are estimated to 2.4 and 4.5 kg stroke $^{-1}$, respectively, which corresponds to 4.8×10^{25} and 9.0×10^{25} molecules NO stroke $^{-1}$. These results suggest that a subtropical thunderstorm may produce more LNOx per LINET stroke than a tropical thunderstorm (factor ~ 2). Possible reasons for this difference will be discussed in Sects. 5 and 6.

For comparison with other published results, the P_{LNOx} estimates per LINET stroke were scaled to P_{LNOx} estimates per LIS flash. During the TROCCINOX field period from 21 January to 27 February 2005, only one overpass of 4 February at 21:23:45–21:25:21 UTC provided a sufficient large set of coincident LINET and LIS measurements. On this day, lightning activity in the LINET centre area (21.5–22.5° S and 48.5–49.5° W) and close-by (covering totally 20.0–23.0° S and 48.5–50.5° W) was suitable for comparison. Overall 82 LIS flashes and 481 LINET strokes were registered in the studied area during the ~ 90 s measurement. LINET strokes with peak currents down to at least 4 kA (absolute value) were sensed by LIS. For the selected time period, Fig. 12 shows the horizontal distributions of all available LINET strokes (black dots) and LIS flashes (red dots) for the area where most lightning occurred (21.4–22.4° S and 48.5–50.0° W) together with radar reflectivity (grey). For the P_{LNOx} estimate, only stronger LINET strokes with peak currents ≥ 10 kA are considered (in total 41 strokes, yellow dots) as mentioned before in Sect. 4.1. The LIS detection efficiency at night (0.93) was taken into account (21:23–21:25 UTC = 19:23–19:25 Brazilian Summer Time). This implies a LINET/LIS ratio of about $(41/82) \times 0.93 = 0.5$, considering only LINET strokes with peak currents ≥ 10 kA.

14837

By means of this ratio, the mean values for P_{LNOx} per LIS flash for tropical and subtropical thunderstorms (only anvil penetrations I, III, and V considered) are 1.2 and 2.2 (range 0.9–2.8) kg, respectively, corresponding to 2.4 and 4.5 (range 1.9–5.6) $\times 10^{25}$ molecules NO. These estimates for TROCCINOX are well within the range of more recent estimates. From a review of previous investigations, SH07 derive a best-estimate of 3.5 (range 0.5–10) kg of nitrogen per flash.

The estimates for P_{LNOx} per LIS flash were multiplied with the number of LIS flashes occurring globally, 44 flashes s^{-1} . If the selected tropical and subtropical TROCCINOX thunderstorms were representative for the globe, the implied mean annual global LNOx production rate G_{LNOx} would be ~ 1.6 and 3.1 Tg a^{-1} , respectively (factor ~ 2 difference). These values are close to previous best estimates for mid-latitude thunderstorms over Europe, 3–4 Tg a^{-1} (Huntrieser et al., 1998, 2002). The individual estimates for the single thunderstorm penetrations listed in Table 2a, however, range from 1.3 to 3.9 Tg a^{-1} , indicating a wide range of values and large uncertainties depending on where (horizontally and vertically) the anvil was penetrated.

Finally, the relative maximal errors of the P_{LNOx} and G_{LNOx} estimates (Table 2a) were calculated. The uncertainty for R_{LINET} was estimated from the standard deviations of the time series of the LINET stroke rates (Fig. 5). The standard deviations varied between 50–90% of the mean values. From the estimates for $F_{\text{LNOx}} \sim 190\%$ and $R_{\text{LINET}} \sim 90\%$, the relative maximal error of the P_{LNOx} estimate for LINET strokes was $\sim 280\%$. For the P_{LNOx} estimate for LIS flashes, it was assumed that the uncertainty in the conversion of LINET strokes (≥ 10 kA) to LIS flashes was $\sim 30\%$ (depending on which LIS detection efficiency was used: day or night). This gives a relative maximal error of $\sim 310\%$. For the G_{LNOx} estimate, the uncertainty in the global LIS flash rate was given with $\sim 10\%$, which gives a final relative maximal error of $\sim 320\%$. Given this relative maximum error, the final range for the G_{LNOx} values listed in Table 2a is between 0.4 and 12 Tg a^{-1} . This range is comparable to other ranges given for G_{LNOx} in previous publications (see Sect. 1 and SH07).

14838

5 Possible explanations for different LNOx production rates in tropical, subtropical and mid-latitude thunderstorms

The results in the previous section lead us to hypothesise that tropical thunderstorms over Brazil may produce less LNOx per stroke than subtropical thunderstorms. In this section we investigate whether these differences in the LNOx production rate may be related to differences in the stroke peak currents (Sect. 5.1), stroke lengths (Sect. 5.1) or stroke release heights (Sect. 5.2) (relationships investigated by Wang et al., 1998, in the laboratory). Furthermore, mean stroke peak currents observed by LINET are compared for several tropical and mid-latitude thunderstorms and for one subtropical thunderstorm (Sects. 5.3–5.4).

5.1 LNOx production rate as a function of stroke peak current

In this subsection we combine the result of laboratory measurements by Wang et al. (1998) with our field measurements to determine the LNOx production rate as a function of peak current. Wang et al. (1998) determined the NO production rate per unit laboratory spark, whereas our analysis provides the NO production rate per LINET stroke in the field. We assume that both follow the same dependency on peak current as given by Wang et al. (1998). Hence, both differ only by a constant factor, which has the dimension of laboratory spark length per LINET stroke. This factor will be determined below.

The relationship between the peak current and NO produced per spark as found by Wang et al. (1998) from measurements in the laboratory (at 1.01×10^5 Pa) is given by:

$$n_{\text{NOlab}}(I) = a + b \times I + c \times I^2 \quad (6)$$

where $n_{\text{NOlab}}(I)$ is the NO production normalised to 1 m spark length (10^{21} molecules NO m^{-1}), $a=0.14$, $b=0.026$, and $c=0.0025$ and I is the peak current of the spark (kA). The number of NO molecules can be converted to the mass of nitrogen according to:

$$M_{\text{NOlab}}(I) = M_W \times n_{\text{NOlab}}(I) \quad (7)$$

14839

where $M_{\text{NOlab}}(I)$ is the nitrogen production per to 1 m spark length (10^{-3} g m^{-1}) and M_W is a constant (molecular weight of N, unit g molecule $^{-1}$). According to Eq. (7) a laboratory spark with 10 kA would produce 0.015 g nitrogen m^{-1} .

Furthermore, Eq. (7) was multiplied with a constant factor C_{LINET} (m laboratory spark per LINET stroke) to convert the production per laboratory sparks and metre to the total number of LINET strokes. Here we assume that all LINET strokes in one specific anvil have the same length, independent of peak current and flash component, since no further information is available from our dataset. The factor C_{LINET} was estimated from Eq. (8). Values from Table 3 for anvil 1a (lower half of table, here only strokes ≥ 10 kA considered) were inserted in Eq. (8); the mean LNOx mass flux, F_{LNOx} , (120 g s^{-1}), and the LINET stroke rate (strokes s^{-1}), $R_{\text{LINET}}(I)$, for a given peak current I :

$$F_{\text{LNOx}} = C_{\text{LINET}} \int_{I=10}^{\infty} R_{\text{LINET}}(I) \times M_{\text{NOlab}}(I) dI \quad (8)$$

where $R_{\text{LINET}}(I)$ was estimated from Eq. (9):

$$R_{\text{LINET}}(I) = \frac{1}{t} \int_{t=0}^t N_{\text{LINET}}(I, t) dt \quad (9)$$

and $N_{\text{LINET}}(I, t)$ is the number of LINET strokes for a given peak current and time.

The value of R_{LINET} in Eq. (9), integrated over all peak currents (here ≥ 10 kA), is 0.055 strokes s^{-1} (Table 3 lower half, anvil 1a). The total integral in Eq. (8) (peak currents from anvil 1a inserted) is estimated to 2.71×10^{-3} strokes s^{-1} g m^{-1} laboratory spark. The corresponding LINET factor, C_{LINET} , was estimated to 44.4×10^3 ($=120/(2.71 \times 10^{-3})$) m laboratory spark stroke $^{-1}$ (see Table 3). We expect a LINET stroke in the atmosphere to be shorter than this calculated length (44 km), but probably broader than a laboratory spark. One metre LINET stroke is therefore probably more efficient in producing LNOx than a laboratory spark.

14840

Applied to LINET strokes, Eq. (7) changes to:

$$M_{\text{NO}_{\text{linet}}}(I) = C_{\text{LINET}} \times M_{\text{NO}_{\text{lab}}}(I) \quad (10)$$

where $M_{\text{NO}_{\text{linet}}}(I)$ is the nitrogen mass production per LINET stroke (g stroke^{-1}). A LINET stroke with a peak current of 10 kA (anvil 1a) would produce 0.7 kg nitrogen according to Eq. (10). This value is rather large because peak currents <10 kA were neglected (Table 3, lower half) and the total LNOx mass was distributed only over strokes ≥ 10 kA in Eq. (8).

The same calculations were performed for other selected thunderstorm penetrations of 4 and 18 February 2005, as shown in Figs. 13a–d and as listed in Table 3. In Table 3 every thunderstorm penetration is listed twice. For the first calculation (upper half in the table) all registered strokes were considered (peak currents down to 2 and 6 kA, depending on the detection efficiency in that area). For a comparison between the 4 and 18 February selected penetrations, however, only strokes with peak currents ≥ 10 kA were considered, as listed in the lower half of the table.

In Figs. 13a–d the frequency distributions of LINET strokes (grey bars) per 1 kA peak current interval are shown for the selected thunderstorms of 4 and 18 February 2005. The stroke frequency rapidly decreases with increasing peak current. Superimposed are the laboratory results by Wang et al. (1998) concerning the NO dependency on peak current modified for LINET strokes (blue dashed line) according to Table 3 (different C_{LINET} values considered), and in addition the estimated total amount of nitrogen mass produced per 1 kA LINET stroke interval for the selected thunderstorm systems (red line). The total mass estimates are listed in Table 3.

In Fig. 13e–f the same type of calculations were performed for datasets with a larger number of LINET strokes to point out more clearly the differences between the stroke peak current frequency distributions of 4 and 18 February. All strokes (≥ 10 kA) in the LINET centre area on 4 February between 00:00 and 24:00 UTC were considered in Fig. 13e. In Fig. 13f the same calculations were performed for the selected thunderstorm system of 18 February for all strokes (≥ 10 kA) between 14:00 UTC and

14841

21:00 UTC (see Fig. 4b). The integral over all peak currents gives a total nitrogen mass of 8.8×10^3 kg produced by 4359 tropical strokes on 4 February (Fig. 13e), and a larger value of 11.2×10^3 kg produced by 2034 subtropical strokes on 18 February (Fig. 13f). This example also suggests that a subtropical stroke may produce a larger amount of nitrogen mass than a tropical stroke (here by a factor 2.7), mainly owing to differences in the stroke length (160 and 44 km, respectively). The higher production rate of LNOx by subtropical strokes was not caused by the stroke peak currents, since this frequency distribution was shifted to lower peak currents on 18 February (mean 31 kA, calculated from data in Fig. 2) compared with 4 February (mean 35 kA), as indicated in Figs. 13e–f.

From the stroke frequency distributions of 4 February it was estimated that strokes with peak currents ≥ 5 kA (only 30% of all strokes) produce the bulk amount (70%) of the total nitrogen mass. This result indicates that the numerous weak strokes with peak currents <5 kA are less important for the LNOx production. In Fig. 14 the same stroke frequency distribution separated, however, into IC and CG strokes, indicates that these weak strokes are mainly IC strokes. Furthermore, Figs. 13a–b indicates that there was a large fraction of these strokes with low peak currents in anvil 1a compared with anvil 5a. Yet, the high stroke rate in anvil 1a (0.252 s^{-1}) produces a similar mean LNOx mass flux, F_{LNOx} , value as in anvil 5a with a much lower stroke rate (0.167 s^{-1}) (Table 3, upper half). (The calculations for F_{LNOx} are based on similar penetration levels: 10.6 and 10.7 km.) Furthermore, the calculated stroke length was slightly shorter in anvil 5a (~ 25 km) compared with anvil 1a (~ 30 km), and the IC stroke release height (Table 4a) was slightly higher in anvil 5a (11.6 km) compared with anvil 1a (10.0 km). This result indicates that the lower stroke rate, shorter stroke length and higher stroke release height (see Sect. 5.2) in anvil 5a cannot explain the similar F_{LNOx} values determined for anvil 1a and 5a. Only if the higher stroke peak currents in anvil 5a (mean 12 kA, Table 4a) compared with anvil 1a (mean 8 kA) are considered, these may give an explanation in this case.

14842

5.2 LNOx production rate as a function of atmospheric pressure

A further explanation for the different LNOx production rates of tropical and subtropical strokes in the selected Brazilian thunderstorms may be related to the release height of the strokes. Laboratory measurements by Wang et al. (1998) indicate that the LNOx production rate increases with increasing atmospheric pressure:

$$n_{\text{NOlab}}(\rho) = a + b \times \rho \quad (11)$$

where $n_{\text{NOlab}}(\rho)$ is the NO production normalised to 1 m spark length (10^{21} molecules NO m^{-1}), $a=0.34$, and $b=1.30$, and ρ is the pressure (10^5 Pa). A laboratory spark at 1000 hPa (ground level) would produce 0.038 gN m^{-1} ; at 500 hPa (300 hPa) about 0.023 (0.017) gN m^{-1} would be produced. The average height of IC strokes at mid-latitudes (Table 4a) is ~ 10.0 km (270 hPa) and in the tropics ~ 10.5 km (250 hPa) (Table 4a). The calculated difference in LNOx production rate (factor 1.1) between these two altitudes (0.016 and 0.015 gN m^{-1} , respectively) is only minor and cannot explain the distinctly higher LNOx production rate of subtropical strokes. Only if we make the unrealistic assumption that all subtropical strokes are CG strokes (mean release height ~ 700 hPa) and all tropical strokes are IC strokes (mean release height 250 hPa) can a factor of ~ 2 ($=0.029/0.015$) difference be achieved.

In the last subsection it was concluded that weak strokes with peak currents < 5 kA are less important for the LNOx production. The majority of strokes with peak currents < 5 kA are IC strokes according to the frequency distributions of IC and CG strokes in Fig. 14. These IC strokes are released in the UT at low pressure. Taking this further relationship into account (decreasing LNOx production rate with decreasing pressure), we find that the large number of very weak strokes with peak currents < 5 kA only have a minor contribution to the LNOx budget.

14843

5.3 Comparison of mean stroke peak currents in several tropical and one subtropical Brazilian thunderstorms

The results in the previous subsections lead us to hypothesise that the different stroke lengths (calculated) may mainly contribute to the different LNOx production rates determined for several tropical and one subtropical Brazilian thunderstorms. The contribution from the different stroke peak current frequency distributions was found to be minor, but maybe important to explain differences between single tropical thunderstorms as mentioned in Sect. 5.1. In this subsection, values of the mean peak current (also separated for CG and IC strokes) are analysed in detail for a larger number of tropical Brazilian thunderstorms in the period with available LINET measurements (21 January–27 February 2005) to investigate the differences between a number of tropical thunderstorms and the subtropical thunderstorm of 18 February.

For an equivalent comparison only strokes in the centre of the LINET detection network (from 21.5° S to 22.5° S and 48.5° W to 49.5° W, 00:00–24:00 UTC) were considered to avoid changes in detection efficiency and in the IC/CG VLF source ratio towards the border line. Selected days with a large number of LINET strokes suitable for statistical calculations are listed in Table 4a (4b) for peak currents ≥ 1 kA (≥ 10 kA).

The selected days were classified according to HH07 into different categories: tropical, transition tropical-subtropical and subtropical cases, by use of the meteorological parameters listed in Table 4c (daily mean values) and as indicated in Fig. 15 (3 h values). In Table 4c the equivalent potential temperature (Θ_e) at 850 and 500 hPa, and the wind velocity and direction at 200 hPa are listed for the selected LINET days in Table 4a–b. As suggested in HH07, Θ_e in tropical air masses exceeded 345 K at 850 hPa and 332 K at 500 hPa and the UT wind velocity was in general low ~ 5 – 10 m s^{-1} , influenced by the Bolivian High.

The 4 February 2005 was selected as a case representative for tropical thunderstorm activity in general. In the selected region $\sim 36\,000$ strokes were registered during the whole day (Table 4a). As expected for tropical thunderstorms, the fraction of IC strokes

14844

dominates over CG strokes and amounts to at least 57%. The \geq symbol indicates that the fraction of IC strokes may be even larger. The 3D procedure applied to discriminate between IC and CG strokes categorises some strokes as uncertain. In most cases, this stroke is a CG stroke (and was here defined as CG), but for unfavourable positions of the measuring network stations it cannot be excluded that it is an IC stroke. Note that we do not deal here with flash counts but with flash component (stroke) counts. In Table 4a the fraction of positive and negative IC and CG strokes for 4 February (all data) was determined to $\geq 35\%$ (IC+), $\geq 22\%$ (IC-), $< 24\%$ (CG-) and $< 19\%$ (CG+). The overall mean peak current (magnitude) was 6 kA. The mean peak currents for the different types of strokes (as mentioned above) were +4, -5, -10 and +5 kA, respectively. The ratio of positive to negative strokes was 1.2. The mean height of IC strokes was 9.7 km. Furthermore, the last column in Table 4a indicates that the fraction of peak currents ≥ 10 kA was 12%.

These results from the 4 February tropical thunderstorms can be compared with other tropical thunderstorms (see Table 4a–b) and the 18 February subtropical thunderstorm. In Table 4a a high mean peak current of 13 kA is given for the 18 February which, however, is not comparable to the rest of the data in Table 4a, since the storm was located along the northern periphery of the LINET network (detection efficiency lower). For an equivalent comparison, only strokes ≥ 10 kA, as listed in Table 4b, were considered and an area along the northern periphery of the LINET detection network ($19.4\text{--}20.0^\circ$ S and $47.7\text{--}49.2^\circ$ W, see Fig. 4b), where the subtropical thunderstorm of 18 February 2005 developed, was selected. The calculated mean peak currents for this area indicate a slightly lower mean peak value, 17 kA, for the subtropical thunderstorm of 18 February compared with the mean peak value for tropical thunderstorms of 4 February for the same area, 20 kA (same value as found for the LINET centre area listed in Table 4b, indicating that the detection efficiency for higher peak currents is about the same in the LINET centre and along the northern periphery, as also discussed in Sect. 4.1).

Overall, the mean peak currents in different tropical thunderstorms of 4 February

14845

(and other tropical thunderstorms listed in Table 4b) were highly variable between 18 and 26 kA (probably depending on thunderstorm intensity). In the next section these values are compared with mean peak currents in mid-latitude thunderstorms over Germany to investigate if any major differences exist.

5.4 Comparison of mean stroke peak currents in several tropical and mid-latitude thunderstorms

The LINET network was also operated in southern Germany in summer 2005 (Sect. 2.2). LINET measurements covered an area reaching from 47° N to 51° N and from 5° E to 14° E. The 29 July 2005 was one of the days in summer 2005 with the highest lightning activity over Germany. In the LINET area $\sim 500\,000$ strokes were registered during the whole day. In Fig. 16a–b the cloud distribution over Europe on this day is shown together with the horizontal distribution of LINET strokes over southern Germany and the positions of the LINET sensors, respectively.

For further estimates with LINET data, only data in the LINET centre region were considered (here 48.5° N to 49.5° N and 11° E to 12° E), as mentioned previously. Other days during the German field campaign in June and July 2005 with a high LINET stroke activity in this area were 29 June and 4, 10 and 15 July. The stroke statistics from these days are listed in Table 4a–b (for peak currents ≥ 1 kA and ≥ 10 kA, respectively) (lower half of table) and can be compared with LINET stroke statistics from Brazil (upper half of table). On 29 June and 29 July strokes in two intense, isolated fast-moving thunderstorms (TS) with high flash rates were considered. On 4, 10 and 15 July all strokes in the LINET centre area were considered. The mean peak currents on the latter days (Table 4a) were 5–6 kA, comparable to the lowest values observed in tropical Brazilian thunderstorms with 5–12 kA (probably lower over Germany owing to the slightly higher detection efficiency in the LINET centre because of a more dense network of LINET sensors). For the same reasons, the fraction of peak currents ≥ 10 kA in these German thunderstorms with 10–15% was comparable to the lowest values in tropical Brazilian thunderstorms with 8–34%. If only strokes with peak currents ≥ 10 kA are

14846

considered, however, (Table 4b), the stroke statistics results from the 4, 10 and 15 July German thunderstorms (mean peak current 19–20 kA, ratio of peak currents ≥ 10 kA is 30–32%) are within the range for tropical Brazilian thunderstorms (18–26 kA, 24–50%).

On 29 June and 29 July 2005 two strong, isolated fast moving thunderstorm systems ahead of cold fronts over Germany with a high stroke frequency were selected, as listed in Table 4a–b and as marked in Fig. 16b for 29 July. In Table 4a the mean peak currents (≥ 1 kA) in these storms were higher with 9 kA compared with 4, 10 and 15 July, and also the ratio of peak currents ≥ 10 kA was higher with 25 and 29%, more similar to the upper values in Brazilian thunderstorms. As expected for mid-latitude thunderstorms (e.g. Prentice and Mackerras, 1977), the fraction of IC strokes ≥ 19 –46% was on average less than in tropical Brazilian thunderstorms and the CG fraction dominated with < 54 –81% (all mid-latitude cases in Table 4a considered). Furthermore, the mean height of IC strokes was ~ 0.5 km lower over Germany compared with Brazil.

To investigate how the different configurations and numbers of sensors over Germany and Brazil influenced the results, an additional dataset was analysed. For all selected German days, a dataset with a reduced number of LINET sensors was created. In the original LINET dataset from 29 July between 5 and 8 sensors (average 7) were considered for the stroke determination in the selected thunderstorm system (indicated in Fig. 16b). In the reduced LINET dataset between 4 and 6 sensors were used (average 5). The number and position of these sensors were selected to be as similar as possible to the configuration over Brazil (on 4 February four sensors) for an equivalent comparison. The dataset of 29 July with a reduced number of sensors registers fewer strokes and a slightly enhanced mean peak current of 11 kA compared with the original dataset (Table 4a, last line). If only strokes ≥ 10 kA are considered (Table 4b, last line), however, no differences between the original dataset and the dataset with a reduced number of sensors were found, which is also the case for the rest of the selected German days (not shown). LINET measurements over Brazil and Germany are therefore well comparable (except for the lowest peak currents), in agreement with the findings by Schmidt et al. (2005).

14847

In Fig. 17a frequency distributions of LINET strokes (in the LINET centre area) as a function of peak current for the two different datasets on 29 July are given in colour (for comparison all curves were normalised). The general agreement between the two curves in red and blue is good, except that the reduced dataset (lower detection efficiency) registers fewer strokes with low peak currents. The distributions show that strokes with a peak current of 2–3 kA were most frequently measured. For comparison, the frequency distribution in tropical thunderstorms over Brazil of 4 February 2005 is also shown (in black). Distinctly more strokes with low peak currents (< 7 kA) and fewer high peak currents (≥ 7 kA) were registered in these thunderstorms compared with the selected intense Germany thunderstorm system of 29 July. Next, a larger dataset is used to investigate this observed difference in peak current distributions for selected German and Brazilian thunderstorms.

In Fig. 17b frequency distributions of LINET strokes (all strokes in the LINET centre area) as a function of peak current for different datasets given in colour are shown (for comparison all curves were normalised). Five different datasets are compared: 1.) tropical Brazil (78 745 strokes) including 230105, 040205 and 250205 (days with a large number of strokes, as listed in Table 4a–b), 2.) tropical Brazil (36 207 strokes) including only 040205, 3.) mid-latitude Germany (58 543 strokes) including 290605 (here all data), 040705, 100705, 150705 and 290705 (here all data), 4.) mid-latitude Germany (6738 strokes) including only 290605 (selected intense thunderstorm), 5.) mid-latitude Germany (3444 strokes) including only 290705 (selected intense thunderstorm). This comparison shows no distinct differences in general between the stroke frequency distributions over Germany and over Brazil (dark blue and black lines). Again, this result confirms that the two LINET configurations are comparable.

For an estimate of the LNO_x mass produced by the observed strokes, the stroke frequency distributions in Fig. 17a were multiplied with the modified Wang et al., 1998 relationship (Eq. 10), as described in Sect. 5.1. In Fig. 17c the estimated amount of NO produced per 1 kA stroke interval for the dataset in Fig. 17a is shown, with different values for the calculated stroke lengths C_{LINET} (from Table 3). The amount of produced

14848

NO in Fig. 17c is a factor ~ 8 higher for the selected intense mid-latitude German thunderstorm system of 29 July 2005 (both distributions) compared with tropical Brazilian thunderstorms of 4 February 2005 in general (based on normalised distributions). In this example, the wide range of calculated stroke lengths (126 and 30 km, respectively) and stroke peak currents inserted (mean 10 and 6 kA, respectively) indicates that a large difference in the LNOx production rate may result in selected cases.

In the tropical Brazilian thunderstorms of 4 February 2005 the majority of the strokes $\sim 90\%$ have peak currents < 10 kA. Figure 17c indicates that these strokes produce only $\sim 50\%$ of the total LNOx amount. In addition, $\sim 70\%$ of all strokes have peak currents < 6 kA, but these produce only 30% of the total LNOx amount. In comparison, in the mid-latitude German intense thunderstorm system of 29 July 2005, about half of the strokes ($\sim 50\%$) have peak currents < 6 kA, but these strokes produce only $\sim 10\%$ of the LNOx amount. Only a small fraction of all strokes $\sim 10\%$ have peak currents > 20 kA, which produce $\sim 60\%$ of the total LNOx amount. These examples again indicate that the LNOx amount produced by a thunderstorm mainly depends on the number of strokes with peak currents > 5 kA.

6 Discussion

The results in the previous section indicate that the amount of nitrogen produced by lightning in a thunderstorm is not well correlated with the number of strokes only (Table 3). We suggest, in accordance with Wang et al. (1998) and Barthe et al. (2007), that other lightning parameters such as stroke length, peak current and release height also necessarily have to be taken into account. The amount of nitrogen produced per m flash and kA for a given pressure level, see parameter S in Table 3, would be more appropriate for comparison, but all of these parameters are in general not available from operational lightning detection networks. Up to now, all these parameters have not been taken into account concurrently, which may explain the wide range of values for the LNOx production rate per flash given in the literature (SH07). Furthermore, for

14849

a more equivalent comparison it would also be desirable to always relate the estimates from a specific lightning detection network to global lightning detection systems such as LIS.

From the analyses of TROCCINOX data, which suggest a higher LNOx production rate for a subtropical thunderstorm compared with several tropical thunderstorms, the stroke length was found to be the most important parameter responsible for the differences. The calculated stroke length was on average a factor ~ 3 longer in the investigated subtropical thunderstorm (Table 3). We suggest that the reason for the longer strokes may be related to the enhanced vertical shear in the horizontal wind in this thunderstorm compared with tropical thunderstorms. In the subtropical case, the vertical wind velocity difference between the pressure levels 200 hPa and 700 hPa ($V_a - V_s$) is 2–3 times larger than in the tropical thunderstorm (Table 2a–b, see also Fig. 14 in HH07). The elevated vertical wind shear may distribute charged particles in the subtropical cloud over longer horizontal distances generating longer flashes. Recent simulations by Barthe and Pinty (2007) of an ideal supercellular storm case, using a 3-D mesoscale model with an explicit lightning flash scheme, give some hints in this direction. Furthermore, conceptual models of the electrical structure in a mesoscale convective system (MCS) by Stolzenburg et al. (1994) and in a supercell by Wiens et al. (2005) (both storm types related to elevated wind shear) indicate that the charged regions in the upper part of the cloud may stretch far away from the convective region with precipitation. A lateral displacement of upper level charge because of vertical wind shear was first suggested by Pierce (1955), Brook et al. (1982) and Hill (1988). Later, observations by Rutledge and MacGorman (1988) of MCS and by Engholm et al. (1990) of summer as well as winter storms confirmed a tilted deformation of the charge centre by the vertical wind shear. More recently, these findings have been discussed by Gilmore and Wicker (2002) and Carey and Buffalo (2007).

Unfortunately, for our dataset the actual stroke length cannot be determined from the available lightning data. From the RINDAT data it is possible, however, to compare the length between different VLF sources along a flash (“flash component”). VLF sources

14850

within <1 s were grouped to a flash component. In Fig. 18a–b the horizontal locations of the flash components are shown for two selected cases. For the 18 February subtropical thunderstorm (strokes considered from 14:00 UTC–21:00 UTC, Fig. 4b) a clear pattern with longer flash components mainly between NW and SE results, influenced by the strong vertical wind shear (Fig. 18a). In contrast, for the 4 February tropical thunderstorms (17:00 UTC–18:00 UTC, Fig. 4a) the horizontal distribution of these flash components shows no preferred direction (Fig. 18b), owing to the lower vertical wind shear. In addition, the flash components are on average shorter compared with the subtropical thunderstorm. The average length of the selected 224 flash components on 4 February was 1.6 km compared with 3.1 km on 18 February based on the 173 flash components, a factor 1.9 difference in the lengths.

The thunderstorm observations during TROCCINOX can also be compared with observations in Florida thunderstorms during CRYSTAL-FACE. Ridley et al. (2004) observed that NO mixing ratios in fresh (not heavily polluted) anvils were distinctly higher than observed anywhere else up to then in fresh anvils. Average anvil-NO mixing ratios varied between $\sim 1\text{--}4$ nmol mol $^{-1}$, compared with $0.2\text{--}1.6$ nmol mol $^{-1}$ NO $_x$ on average in Brazilian anvils during TROCCINOX (HH07) and 1.3 ± 0.7 nmol mol $^{-1}$ NO $_x$ in more polluted thunderstorms during EULINOX (Huntrieser et al., 2002). For two thunderstorms during CRYSTAL-FACE, Ridley et al. (2004) determined $(3.3\text{--}6.6) \times 10^{25}$ and $(17\text{--}23) \times 10^{25}$ molecules NO per flash (LIS) for a moderate size and a large storm, respectively. Only the values for the moderate size storm are within the range of the mean value for a subtropical thunderstorm during TROCCINOX: 4.5×10^{25} molecules NO per flash (LIS), respectively.

The reason for these extreme NO production rates in Florida anvils may be related to a combination of the lightning parameters mentioned above. First, the high flash rate in these storms may play an important role. The high flash density over Florida (>9 flashes km $^{-2}$ a $^{-1}$, Orville et al., 2002), however, is comparable to that of Brazil (mean 6–8 flashes and up to 10–15 flashes km $^{-2}$ a $^{-1}$ based on National Lightning Detection Network (NLDN) and RINDAT, respectively) (Pinto and Pinto, 2003). Fur-

14851

thermore, it has been reported that the mean peak current of CG flashes is shifted to higher values in Florida thunderstorms compared with other regions in the United States: see mean peak values for different regions listed in Table 5. The NLDN detection efficiency for flashes with peak currents >5 kA is 80–90% since the upgrade of the system in 1994–1995 (Cummins et al., 1998) and 90–95% since the upgrade in 2002–2003 (Cummins et al., 2006). Values based on LINET data are also listed in Table 5 for comparison, which indicate a low mean value in the investigated subtropical thunderstorm, but just within the range of mean LINET values for tropical and mid-latitude thunderstorms. In the latter two storm types, the range of the mean peak currents is comparable, but the highest values were observed in tropical thunderstorms. The results presented in Table 5 indicate a latitudinal variation of mean peak current for NLDN, but not for LINET. Finally, the higher wind velocities in the UT over Florida (15 and 20 m s $^{-1}$, Garrett et al., 2005) compared with tropical Brazil indicates that the stroke length may be longer in Florida thunderstorms. We therefore conclude that several lightning parameters, such as flash density, mean peak current and possibly flash length (related to UT wind velocity), achieve especially high values in Florida thunderstorms, which may be responsible for the high anvil-NO mixing ratios and LNO $_x$ production rates per stroke observed. Other authors also mentioned that the most severe thunderstorms on earth occur in the central-eastern United States owing to the unique combination of steep lapse rates, moist boundary layer inflow and substantial wind shear (Doswell, 2001; Del Genio et al., 2007).

Furthermore, in our study we find that the LNO $_x$ production rate per stroke may be a factor ~ 2 (up to 8) larger for subtropical and mid-latitude strokes than for tropical strokes (Table 2a–b, Fig. 17c). We therefore suggest that different estimates for the amount of LNO $_x$ produced per stroke depending on region are needed for more accurate global LNO $_x$ estimates, taking stroke peak currents, stroke release height and stroke lengths (related to the vertical wind shear between upper and lower troposphere) into account. In agreement, recent simulations with the GEOS-CHEM model (Hudman et al., 2007) reveal that upper tropospheric NO $_x$ mixing ratios observed over the southern United

14852

States during ICARTT (July–August 2004) can only be reproduced if the LNO_x yield in the model is increased by a factor 4 relative to the Ott et al. (2007b) value estimated from EULINOX and other mid-latitude and subtropical storms. For the same model and period Martin et al. (2006) increased northern mid-latitude lightning NO_x emissions by a factor 4 to fit aircraft observations while tropical lightning NO_x emissions remained unchanged. Furthermore, preliminary cloud-model results from Pickering et al. (2007) for the 4 February 2005 case suggest that IC flashes in tropical thunderstorms in Brazil produce less LNO_x (factor 1.6) compared with previous cloud-model results based on 5 mid-latitude and one subtropical thunderstorms. From our analyses of TROCCINOX data, we suggest that mainly the different vertical wind shear in these regions (as it impacts the stroke length) may explain why different LNO_x production rates per stroke are needed in the models.

7 Summary and conclusions

During the TROCCINOX field experiment in southern Brazil in the wet season of 2005, lightning-produced NO_x (LNO_x) in and around tropical and subtropical thunderstorms was investigated in large detail with the DLR Falcon and partly with the high-flying M55 Geophysica aircraft. On two selected days (4 and 18 February), enhanced NO_x mixing ratios in the range of 0.2 to 1.2 nmol mol⁻¹ (width 25–45 km) on average were observed during anvil penetrations at 9–11 km altitudes. The main part of this NO_x enhancement, 80–90%, was attributed to LNO_x while the contribution from the boundary layer (BL) was in general less important (10–20%). Overall, NO_x mixing ratios in the anvil outflow region were comparable in magnitude to those measured at mid-latitudes during other campaigns (SH07), but distinctly lower than observed in fresh anvils during CRYSTAL-FACE over Florida, ~1–4 nmol mol⁻¹ NO (Ridley et al., 2004).

A lightning detection network, LINET, was set up to monitor the local stroke distribution (here VLF sources), which generally agreed well with observations from the operational lightning detection network in Brazil, RINDAT, and with observations from

14853

LIS. The horizontal LNO_x mass flux in TROCCINOX thunderstorms was estimated from the airborne measurements and trajectory simulations, and combined with LINET stroke rates. The average LNO_x production rate per stroke and per LIS flash were estimated. The final results gave ~1 and ~2–3 kg per LIS flash for three tropical and one subtropical Brazilian thunderstorms, respectively, suggesting a higher LNO_x production rate in the latter storm type (factor ~2). The estimated LNO_x production rates were multiplied with the number of LIS flashes occurring globally, 44 flashes s⁻¹. From these values, the mean annual global LNO_x production rate was estimated to 1.6 and 3.1 Tg a⁻¹, respectively, from the two storm types mentioned above. These values are well within the range of more recent estimates in SH07. The spread of the results for different thunderstorms penetrations (1.3–3.9 Tg a⁻¹), however, indicates a large variability. This variety mainly results from the extrapolation from single measurements in the anvil outflow, which depends on the aircraft position relative to the thunderstorm core. The large uncertainty (~320% relative max. error) in the estimates gives a final range of 0.4–12 Tg a⁻¹.

The analyses from TROCCINOX indicate that the amount of nitrogen produced by lightning in a thunderstorm is not well correlated with the number of strokes only. It is suggested, from laboratory findings by Wang et al. (1998), that other lightning parameters such as stroke length, peak current and release height also necessarily have to be taken into account. Wang et al. (1998) observed a positive correlation between NO production rates per m laboratory spark and stroke peak currents and ambient pressure, which was applied to the present study. According to the first relationship, data from TROCCINOX indicate that the total amount of nitrogen produced by a thunderstorm mainly depends on the number of strokes with peak currents >5 kA. The analyses from TROCCINOX further suggest that the reason for the higher LNO_x production rate in the subtropical thunderstorm of 18 February (factor ~2), compared with tropical thunderstorms of 4 February, may be related to the different stroke lengths (factor ~3 longer calculated for the subtropical storm) and is not related to differences in the mean stroke peak current (slightly lower in the subtropical thunderstorm compared

14854

with the tropical thunderstorms). Furthermore, the mean IC stroke height was ~ 0.5 km lower in the subtropical thunderstorm, though this difference may cause only a minor difference in the LNO_x production rate. From these results it was suggested that the longer stroke length in the subtropical thunderstorm may be related to the higher vertical wind shear between 200 and 700 hPa (factor $\sim 2-3$), spreading charged particles in the anvil over larger distances. We therefore suggest that the vertical wind shear may be an important parameter influencing the amount of LNO_x produced per flash, which has not been taken into account up to now. For future model studies with CTMs, we recommend the incorporation of global meteorological fields of the vertical wind shear to account for the different regional LNO_x production rates per stroke. Furthermore, for future field campaigns we suggest using lightning location systems that determine the flash length more precisely (e.g. the French ONERA VHF interferometric mapper or the New Mexico Tech Lightning Mapping Array, LMA).

In addition, a comparison between the lightning activity in TROCCINOX thunderstorms over Brazil and in mid-latitude thunderstorms over Germany was carried out, as measurements with the LINET system were available for both areas in 2005. The stroke frequency distribution as a function of peak current and the range of mean peak currents were in general similar in Brazilian tropical thunderstorms and German mid-latitude thunderstorms. For a selected case, however, it was estimated that the LNO_x production rate per LINET stroke may be distinctly higher in intense German mid-latitude thunderstorms compared with tropical Brazilian thunderstorms in general (up to factor ~ 8), owing to the longer stroke length (related to higher vertical wind shear) and the higher mean peak current in these storms.

Finally, the results from Brazilian thunderstorms during TROCCINOX were compared with Florida thunderstorms during CRYSTAL-FACE, where especially high anvil-NO mixing ratios and LNO_x production rates per flash have been observed. Results from the available literature indicate that several lightning parameters, such as flash density, mean peak current and stroke length (related to UT wind velocity), all achieve especially high values in Florida thunderstorms compared with other regions, which we

14855

suggest may be responsible for the high NO mixing ratios observed in Florida storms.

We conclude that one advantage of the present study is the use of a lightning location system where the whole spectra of peak currents down to ~ 1 kA can be detected. In addition, CG and IC strokes can be separated, and measurements with the same system were performed in tropical, subtropical and mid-latitude regions. With this system new insights into the LNO_x potential for different types of strokes are possible, but it has to be pointed out that the present study is based on a very limited dataset containing few, random airborne penetrations of single subtropical and tropical thunderstorm systems. Unfortunately, it is not known how representative these measurements are. The main uncertainties originate from the estimate of mean anvil-NO_x mixing ratios during few anvil penetrations, and from the limited altitude range covered by the aircraft. Also the attribution of the set of observed stroke events to the LNO_x increase is very uncertain. Furthermore, the data were obtained only from a small set of thunderstorm systems over Germany and Brazil. Finally, the scaling between LINET and LIS observations is uncertain because only one short time segment was available to compare both systems directly. Owing to these restrictions the uncertainties in the presented results may be large and the stated generalisations ought to be used with caution. We are aware of the very limited dataset presented in this paper, but with this paper we would like to give a hint in which direction further research concerning measurements of LNO_x in the field could take. Recently during AMMA (August 2006), the anvil outflow from African thunderstorms was probed with several aircraft. In addition, recent measurements in tropical thunderstorms over Australia (Darwin) from the SCOUT campaign in December 2005 can be investigated in the same manner as presented in this paper. During both field experiments lightning measurements were carried out with the LINET system.

Even though the majority of global lightning occurs in the tropics (Christian et al., 2003), our study suggests that subtropical and mid-latitude thunderstorms may contribute with an essential fraction to global LNO_x, since we observe a tendency that the stroke length (related to the vertical wind shear) is longer in these storms. As a

14856

result, more LNO_x may be produced per stroke in these storms with elevated vertical wind shear compared with tropical low-shear thunderstorms as observed over Brazil. Finally, we would like to point out that the estimated amount of LNO_x produced per flash, as given in many studies (SH07), largely depends on the type of lightning detection system used and therefore these estimates should be compared more carefully in future. Furthermore, the relationship between the used lightning system and LIS must be known if the values are scaled up with the annual global LIS flash rate (44 s⁻¹), as demonstrated in this paper.

Acknowledgements. TROCCINOX was partially funded by the Commission of the European Community under the contract EVK2-CT-2001-00122 and by the DLR (Deutsches Zentrum für Luft- und Raumfahrt) and other TROCCINOX partners. TROCCINOX was performed as a coordinated action of European and Brazilian research institutes and agencies together with the Brazilian project Tropical Convection and Cirrus Experiment Brazil (TroCCiBras), see detailed acknowledgements in HH07. Furthermore, we are especially grateful to A. Minikin for his contribution to flight planning and his work as mission scientist on some of the flights, to A. Roiger, M. Lichtenstern and P. Stock (DLR-Oberpfaffenhofen) for instrument preparation, to C. Schwierz (ETH-Zurich) for meteorological support during the field campaign and post-campaign meteorological analyses and to A. Stohl (NILU-Kjeller) for post-campaign meteorological analyses and helpful comments on draft manuscripts. We thank the Falcon pilots (M. Grossrubatscher, D. Günther, M. Hinterwaldner and R. Welsler), engineers and scientists of the flight department of DLR for the excellent support during the campaigns and the Geophysica team (especially P. Mazzinghi, F. D'Amato and S. Viciani, INOA/CNR, for the CO measurements). Further, T. Fehr (DLR), the Brazilian IPMet team and K. Schmidt, B. Fuchs and W. P. Oettinger from the LINET team (University of Munich) are greatly acknowledged for setting up and running the LINET system in Brazil and Germany. We thank P. S. Borges (IPMet, São Paulo State University/UNESP, Bauru) for preparing the radar images from southern Brazil. We express our gratitude to the lightning team at MSFC-NASA for the access to the LIS data. The European Centre for Medium Weather Forecasts (ECMWF) and the German Weather Service (DWD) are acknowledged for permitting access to the ECMWF archives. Finally, we are especially grateful to E. Defer (LERMA, Observatoire de Paris), who critically reviewed draft manuscripts and gave insightful comments and helpful suggestions for the data analysis.

14857

References

- Barth, M. C., Kim, S.-W., Wang, C., et al.: Cloud-scale model intercomparison of chemical constituent transport in deep convection, *Atmos. Chem. Phys. Discuss.*, 7, 8035–8085, 2007, <http://www.atmos-chem-phys-discuss.net/7/8035/2007/>.
- Barthe, C. and Pinty, J.-P.: Simulation of a supercellular storm using a three-dimensional mesoscale model with an explicit lightning flash scheme, *J. Geophys. Res.*, 112, D06210, doi:10.1029/2006JD007484, 2007.
- Barthe, C., Pinty, J.-P., and Mari, C.: Lightning-produced NO_x in an explicit electrical scheme tested in a Stratosphere-Troposphere Experiment: Radiation, Aerosols, and Ozone case study, *J. Geophys. Res.*, 112, D04302, doi:10.1029/2006JD007402, 2007.
- Beirle, S., Platt, U., Wenig, M., and Wagner, T.: NO_x production by lightning estimated with GOME, *Adv. Space Res.*, 34, 793–797, 2004.
- Beirle, S., Spichtinger, N., Stohl, A., Cummins, K. L., Turner, T., Boccippio, D., Cooper, O. R., Wenig, M., Grzegorski, M., Platt, U., and Wagner, T.: Estimating the NO_x produced by lightning from GOME and NLDN data: a case study in the Gulf of Mexico, *Atmos. Chem. Phys.*, 6, 1075–1089, 2006.
- Bertram, T. H., Perring, A. E., Wooldridge, P. J., et al.: Direct measurements of the convective recycling of the upper troposphere, *Science*, 315, 816–820, 2007.
- Betz, H.-D., Schmidt, K., Oettinger, W. P., and Wirz, M.: Lightning detection with 3D-discrimination of intracloud and cloud-to-ground discharges, *Geophys. Res. Lett.*, 31, L11108, doi:10.1029/2004GL019821, 2004.
- Betz, H.-D., Schmidt, K., Fuchs, B., Oettinger, W. P., and Höller, H.: Cloud lightning: Detection and utilization for total lightning measured in the VLF/LF regime, *J. Lightning Res.*, 2, 1–17, <http://www.jorl.org>, 2007.
- Boccippio, D. J., Koshak, W. J., and Blakeslee, R. J.: Performance assessment of the tropical transient detector and lightning imaging sensor. Part I: Predicted diurnal variability, *J. Atmos. Oceanic Technol.*, 19, 1318–1332, 2002.
- Boersma, K. F., Eskes, H. J., Meijer, E. W., and Kelder, H. M.: Estimates of lightning NO_x production from GOME satellite observations, *Atmos. Chem. Phys.*, 5, 2311–2331, 2005.
- Bögel, W. and Baumann, R.: Test and calibration of the DLR Falcon wind measuring system by maneuvers, *J. Atmos. Oceanic Technol.*, 8, 5–18, 1991.
- Bradshaw, J., Davis, D., Grodzinsky, G., Smyth, S., Newell, R., Sandholm, S., and Liu, S.:

14858

- Observed distributions of nitrogen oxides in the remote free troposphere from the NASA global tropospheric experiment programs, *Rev. Geophys.*, 38, 61–116, 2000.
- Brook, M., Nakano, M., Krehbiel, P., and Takeuti, T.: The electrical structure of the Hokuriku winter thunderstorms, *J. Geophys. Res.*, 87, 1207–1215, 1982.
- 5 Carey, L. D. and Buffalo, K. M.: Environmental control of cloud-to-ground lightning polarity in severe storms, *Mon. Wea. Rev.*, 135, 1327–1353, 2007.
- Chaboureaud, J.-P., Cammas, J.-P., Duron, J., Mascart, P. J., Sitnikov, N., and Voessing, H.-J.: A numerical study of tropical cross-tropopause transport by convective overshoots, *Atmos. Chem. Phys.*, 7, 1731–1740, 2007, <http://www.atmos-chem-phys.net/7/1731/2007/>.
- 10 Chameides, W. L.: The role of lightning in the chemistry of the atmosphere, in *The Earth's Electrical Environment*, pp. 70–77, Nat. Acad. Press, Washington, D.C., 1986.
- Chameides, W. L., Davis, D. D., Bradshaw, J., Rodgers, M., Sandholm, S., and Bai, D. B.: An estimate of the NO_x production rate in electrified clouds based on NO observations from the GTE/CITE 1 fall 1983 field operation, *J. Geophys. Res.*, 92, 2153–2156, 1987.
- 15 Chowdhuri, P., Anderson, J. G., Chisholm W. A., et al.: Parameters of lightning strokes: A Review, *IEEE Transactions on Power Delivery*, Vol. 20, No. 1, 346–358, 2005.
- Christian, H. J., Blakeslee, R. J., Goodman, S. J., et al.: The Lightning Imaging Sensor, *Proceedings of the 11th International Conference on Atmospheric Electricity*, Guntersville, Alabama, June 7–11, 1999, pp. 746–749, 1999.
- 20 Christian, H. J., Blakeslee, R. J., Boccippio, D. J., et al.: Global frequency and distribution of lightning as observed from space by the Optical Transient Detector, *J. Geophys. Res.*, 108, 4005, doi: 10.1029/2002JD002347, 2003.
- Christian, H. J. and Petersen, W.: Global lightning activity, *Conference on Meteorological Applications of Lightning Data*, 85th AMS Annual Meeting, San Diego, CA, 10–12 January, 2005.
- 25 Cooper, O. R., Stohl, A., Trainer, M., et al.: Large upper tropospheric ozone enhancements above mid-latitude North America during summer: In situ evidence from the IONS and MOZAIC ozone measurement network, *J. Geophys. Res.*, 111, D24S05, doi:10.1029/2006JD007306, 2006.
- 30 Coppens, F., Berton, R., Bondiou-Clergerie, A., and Gallimberti, I.: Theoretical estimate of NO_x production in lightning corona, *J. Geophys. Res.*, 103, 10 769–10 785, 1998.
- Cummins, K. L., Murphy, M. J., Bardo, E. A., Hiscox, W. L., Pyle, R. B., and Pifer, A. E.: A combined TOA/MDF technology upgrade of the U.S. National Lightning Detection Network,

14859

- J. Geophys. Res.*, 103, 9035–9044, doi:10.1029/98JD00153, 1998.
- Cummins, K., Cramer, J. A., Biagi, C. J., Krider, E. P., Jerauld, J., Uman, M. A., and Rakov, V. A.: The U.S. national lightning detection network: post-upgrade status, *The 86th AMS Annual Meeting, 2nd Conference on Meteorological Applications of Lightning Data*, Paper 6.1, 28 January–3 February 2006, Atlanta, GA, 2006.
- 5 DeCaria, A. J., Pickering, K. E., Stenchikov, G. L., Scala, J. R., Stith, J. L., Dye, J. E., Ridley, B. A., and Laroche, P.: A cloud-scale model study of lightning-generated NO_x in an individual thunderstorm during STERAO-A, *J. Geophys. Res.*, 105, 11 601–11 616, 2000.
- DeCaria, A. J., Pickering, K. E., Stenchikov, G. L., and Ott, L. E.: Lightning-generated NO_x and its impact on tropospheric ozone production: A three-dimensional modelling study of a Stratosphere-Troposphere Experiment: Radiation, Aerosols, and Ozone (STERAO-A) thunderstorm, *J. Geophys. Res.*, 110, D14303, doi:10.1029/2004JD005556, 2005.
- Defer, E., Laroche, P., Dye, J. E., and Skamarock, W.: Use of total lightning lengths to estimate NO_x production in a Colorado thunderstorm, paper presented at 12th International Conference on Atmospheric Electricity, 9–13 June, Int. Comm. on Atmos. Electr., Versailles, France, 2003.
- Del Genio, A. D., Yao, M.-S., and Jonas, J.: Will moist convection be stronger in a warmer climate?, *Geophys. Res. Lett.*, 34, L16703, doi:10.1029/2007GL030525, 2007.
- Dickerson, R. R., Huffman, G. J., Luke, W. T., et al.: Thunderstorms: An important mechanism in the transport of air pollutants, *Science*, 235, 460–465, 1987.
- 20 Doswell, C. A., III: *Severe Convective Storms*, Meteorol. Monogr. Ser., Vol. 28(50), 570 pp., Am. Meteorol. Soc., Boston, Mass., USA, 2001.
- Dye, J. E., Ridley, B. A., Skamarock, W., et al.: An overview of the Stratospheric-Tropospheric Experiment: Radiation, aerosols, and ozone (STERAO)-Deep convection experiment with results for the July 10, 1996 storm, *J. Geophys. Res.*, 105, 10 023–10 045, 2000.
- 25 Engholm, C. D., Williams, E. R., and Dole R. M.: Meteorological and electrical conditions associated with positive cloud-to-ground lightning, *Mon. Wea. Rev.*, 118, 470–487, 1990.
- Fehr, T., Höller, H., and Huntrieser, H.: Model study on production and transport of lightning-produced NO_x in an EULINOX supercell storm, *J. Geophys. Res.*, 109, D09102, doi:10.1029/2003JD003935, 2004.
- 30 Gallardo, L. and Cooray, V.: Could cloud-to-cloud discharges be as effective as cloud-to-ground discharges in producing NO_x?, *Tellus*, 48B, 641–651, 1996.
- Garrett, T. J., Navarro, B. C., Twohy, C. H., et al.: Evolution of a Florida cirrus anvil, *J. Atmos.*

14860

- Sci., 62, 2352–2372, 2005.
- Gilmore, M. S. and Wicker, L. J.: Influences of the local environment on supercell cloud-to-ground lightning, radar characteristics, and severe weather on 2 June 1995, *Mon. Wea. Rev.*, 130, 2349–2372, 2002.
- 5 Hauf, T., Schulte, P., Alheit, R., and Schlager, H.: Rapid vertical trace gas transport by an isolated mid-latitude thunderstorm, *J. Geophys. Res.*, 100, 22 957–22 970, 1995.
- Hill, R. D.: Interpretation of bipole patterns in a mesoscale storm, *Geophys. Res. Lett.*, 23, 643–645, 1988.
- Höller, H., Finke, U., Huntrieser, H., Hagen, M., and Feigl, C.: Lightning produced NO_x (LINOX) - Experimental design and case study results, *J. Geophys. Res.*, 104, 13 911–13 922, 1999.
- 10 Hudman, R. C., Jacob, D. J., Turquety, S., et al.: Surface and lightning sources of nitrogen oxides in the United States: Magnitudes, chemical evolution, and outflow, *J. Geophys. Res.*, 112, D12S05, doi:10.1029/2006JD007912, 2007.
- Huntrieser, H., Schlager, H., Feigl, C., and Höller, H.: Transport and production of NO_x in electrified thunderstorms: Survey of previous studies and new observations at mid-latitudes. *J. Geophys. Res.*, 103, 28 247–28 264, 1998.
- 15 Huntrieser, H., Feigl, C., Schlager, H., Schröder, F., Gerbig, C., van Velthoven, P., Flatøy, F., Théry, C., Petzold, A., Höller, H., and Schumann, U.: Airborne measurements of NO_x, tracer species and small particles during the European Lightning Nitrogen Oxides Experiment, *J. Geophys. Res.*, 107 (D11), 4113, doi: 10.1029/2000JD000209, ACH 5-1 - ACH 5-24, 2002.
- 20 Huntrieser, H., Heland, J., Schlager, H., et al.: Intercontinental air pollution transport from North America to Europe: Experimental evidence from airborne measurements and surface observations, *J. Geophys. Res.*, 110, D01305, doi:10.1029/2004JD005045, 2005.
- Huntrieser, H., Schlager, H., Höller, H., Schumann, U., Betz, H.-D., Boccippio, D., Brunner, D., Forster, C., and Stohl, A.: Lightning-produced NO_x in tropical, subtropical and mid-latitude thunderstorms: New insights from airborne and lightning observations, European Geosciences Union, General Assembly 2006, Vienna, Austria, 2–7 April 2006, Oral Paper Nr. EGU06-A-03286, *Geophys. Res. Abstracts*, Vol. 8, 2006.
- 25 Huntrieser, H., Schlager, H., Roiger, A., Schumann, U., Höller, H., Kurz, C., Brunner, D., Schwierz, C., Richter, A., and Stohl, A.: Lightning-produced NO_x over Brazil during TROC-CINOX: airborne measurements in tropical and subtropical thunderstorms and the importance of mesoscale convective systems, *Atmos. Chem. Phys.*, 7, 2987–3013, 2007, <http://www.atmos-chem-phys.net/7/2987/2007/>.
- 30

14861

- Jerauld, J., Rakov, V. A., Uman, M. A., Rambo, K. J., Jordan, D. M., Cummins, K. L., and Cramer, J. A.: An evaluation of the performance characteristics of the U.S. National Lightning Detection Network in Florida using rocket-triggered lightning, *J. Geophys. Res.*, 110, D19106, doi:10.1029/2005JD005924, 2005.
- 5 Keenan, T. D. and Carbone, R. E.: A preliminary morphology of precipitation systems in tropical northern Australia, *Q. J. R. Meteorol. Soc.*, 118, 283–326, 1992.
- Koike, M., Kondo, Y., Kita, K., et al.: Measurements of reactive nitrogen produced by tropical thunderstorms during BIBLE-C, *J. Geophys. Res.*, 112, D18304, doi:10.1029/2006JD008193, 2007.
- 10 Langford, A. O., Portmann, R. W., Daniel, J. S., Miller, H. L., and Solomon, S.: Spectroscopic measurements of NO₂ in a Colorado thunderstorm: Determination of the mean production by cloud-to-ground lightning flashes, *J. Geophys. Res.*, 109, D11304, doi: 10.1029/2003JD004158, 2004.
- Lopez, J. P., Fridlind, A. M., Jost, H.-J., et al.: CO signatures in subtropical convective clouds and anvils during CRYSTAL-FACE: An analysis of convective transport and entrainment using observations and a cloud-resolving model, *J. Geophys. Res.*, 111, D09305, doi:10.1029/2005JD006104, 2006.
- 15 Lyons, W. A., Uliasz, M., and Nelson, T. E.: Large peak current cloud-to-ground lightning flashes during the summer months in the contiguous United States, *Mon. Wea. Rev.*, 126, 2217–2233, 1998.
- 20 Martin, R. V., Sioris, C. E., Chance, K., et al.: Evaluation of space-based constraints on global nitrogen oxide emissions with regional aircraft measurements over and downwind of eastern North America, *J. Geophys. Res.*, 111, D15308, doi:10.1029/2005JD006680, 2006.
- Martin, R. V., Sauvage, B., Folkens, I., et al.: Space-based constraints on the production of nitric oxide by lightning, *J. Geophys. Res.*, 112, D09309, doi:10.1029/2006JD007831, 2007.
- 25 Orville, R. E.: Peak-current variations of lightning return strokes as a function of latitude, *Nature*, 343, 149–151, 1990.
- Orville, R. E.: Comments on “Large peak current cloud-to-ground lightning flashes during the summer months in the contiguous United States”, *Mon. Wea. Rev.*, 127, 1937–1938, 1999.
- 30 Orville, R. E. and Huffines, G. R.: Lightning ground flash measurements over the contiguous United States: 1995–1997, *Mon. Wea. Rev.*, 127, 2693–2703, 1999.
- Orville, R. E., Huffines, G. R., Burrows, W. R., Holle, R. L., and Cummins, K. L.: The North American lightning detection network (NALDN) – First results: 1998–2000, *Mon. Wea. Rev.*,

14862

130, 2098–2109, 2002.

- Ott, L. E., Pickering, K. E., Stenchikov, G. L., Huntrieser, H., and Schumann, U.: Effects of lightning NO_x production during the July 21 European Lightning Nitrogen Oxides Project storm studied with a three-dimensional cloud-scale chemical transport model, *J. Geophys. Res.*, 112, D05307, doi:10.1029/2006JD007365, 2007b.
- Petersen, W. A. and Rutledge, S. A.: Some characteristics of cloud-to-ground lightning in tropical northern Australia, *J. Geophys. Res.*, 97, 11 553–11 560, 1992.
- Petersen, W. A., Christian, H. J., and Rutledge, S. A.: TRMM observations of the global relationship between ice water content and lightning, *Geophys. Res. Lett.*, 32, L14819, doi:10.1029/2005GL023236, 2005.
- Pickering, K. E., Thompson, A. M., Scala, J. R., Tao, W.-K., Dickerson, R. R., and Simpson, J.: Free tropospheric ozone production following entrainment of urban plumes into deep convection, *J. Geophys. Res.*, 97, 17 985–18 000, 1992.
- Pickering, K. E., Thompson, A. M., Wang, Y., et al.: Convective transport of biomass burning emissions over Brazil during TRACE A, *J. Geophys. Res.*, 101, 23 993–24 012, 1996.
- Pickering, K. E., Huntemann, T., Ott, L., Barth, M., Huntrieser, H., Schlager, H., Schumann, U., Vaughan, G., and Volz-Thomas, A.: Cloud-resolved simulations of lightning- NO_x in observed tropical thunderstorms, European Geosciences Union, General Assembly 2007, Vienna, Austria, 15–20 April 2007, Oral Paper Nr. EGU2007-A-11013, *Geophys. Res. Abstr.*, Vol. 9, 2007.
- Pierce, E. T.: The development of lightning discharges, *Quart. J. Roy. Meteor. Soc.*, 81, 229–240, 1955.
- Pinto, I. R. C. A. and Pinto, O., Jr.: Cloud-to-ground lightning distributions in Brazil, *J. Atmos. Solar-Terrestrial Physics*, 65, 733–737, 2003.
- Pinto, O., Jr., Pinto, I. R. C. A., and Naccarato, K. P.: Maximum cloud-to-ground lightning flash densities observed by lightning location systems in the tropical region: A review, *Atmos. Res.*, 84, 189–200, 2007.
- Price, C., Penner, J., and Prather, M.: NO_x from lightning: 1. Global distribution based on lightning physics, *J. Geophys. Res.*, 102, 5929–5941, 1997.
- Prentice, S. A., and Mackerras, D.: The ratio of cloud to cloud-ground lightning flashes in thunderstorms, *J. Appl. Meteorol.*, 16, 545–550, 1977.
- Rahman, M., Corray, V., Rakov, V. A., Uman, M. A., Liyanage, P., DeCarlo B. A., Jerauld, J., and Olsen III, R. C.: Measurements of NO_x produced by rocket-triggered lightning, *Geophys. Res. Lett.*, 34, L03816, doi:10.1029/2006GL027956, 2007.

14863

- Rakov, V. A., Thottapillil, R., and Uman, M. A.: On the empirical formula of Willett et al. relating lightning return-stroke peak current and peak electric field, *J. Geophys. Res.*, 97, 11 527–11 533, 1992.
- Ridley, B. A., Dye, J. E., Walega, J. G., Zheng, J., Grahek, F. E., and Rison, W.: On the production of active nitrogen by thunderstorms over New Mexico, *J. Geophys. Res.*, 101, 20 985–21 005, 1996.
- Ridley, B. A., Ott, L., Pickering, K., et al.: Florida thunderstorms: A faucet of reactive nitrogen to the upper troposphere, *J. Geophys. Res.*, 109, D17305, doi:10.1029/2004JD004769, 2004.
- Ridley, B. A., Pickering, K. E., and Dye, J. E.: Comments on the parameterization of lightning-produced NO in global chemistry-transport models, *Atmos. Environ.*, 39, 6184–6187, 2005.
- Rutledge, S. A. and MacGorman, D. R.: Cloud-to-ground lightning activity in the 10–11 June 1985 mesoscale convective system observed during the Oklahoma-Kansas PRE-STORM project, *Mon. Wea. Rev.*, 116, 1393–1408, 1988.
- Schmidt, K., Betz, H.-D., Oettinger, W. P., Wirz, M., and Diendorfer, G.: A new lightning detection network in southern Germany, 27th International Conference on Lightning Protection (ICLP), September 2004, Avignon, France, 2004.
- Schmidt, K., Betz, H.-D., Oettinger, W. P., Wirz, M., Pinto Jr., O., Naccarato, K. P., Höller, H., Fehr, T., and Held, G.: A comparative analysis of lightning data during the EU-Brazil TROCCINOX / TroCCiBras campaign, VIII International Symposium on Lightning Protection (SIPDA), 21–25 November 2005, São Paulo, Brazil, 2005.
- Schmidt, K.: Ortung und Analyse von Blitzen durch Registrierung von VLF-Atmospherics innerhalb eines Messnetzes, Ph.D. thesis, Ludwig-Maximilians-Universität, Munich, Germany, 2007.
- Schulz, W., Cummins, K., Diendorfer, G., and Dorninger, M.: Cloud-to-ground lightning in Austria: A 10-year study using data from a lightning location system, *J. Geophys. Res.*, 110, D09101, doi:10.1029/2004JD005332, 2005.
- Schumann, U., Konopka, P., Baumann, R., Busen, R., Gerz, T., Schlager, H., Schulte P., and Volkert, H.: Estimate of diffusion parameters of aircraft exhaust plumes near the tropopause from nitric oxide and turbulence measurements, *J. Geophys. Res.*, 100, 14 147–14 162, 1995.
- Schumann, U., Huntrieser, H., Schlager, H., Bugliaro, L., Gatzert, C., and Hoeller, H.: Nitrogen Oxides from thunderstorms- Results from experiments over Europe and the Continental

14864

- Tropics, paper presented at Deutsch-Österreichisch-Schweizerische Meteorologen-Tagung (DACH), Deutsche Meteorologische Gesellschaft, Karlsruhe, Germany, 7–10 September, 2004.
- 5 Schumann, U. and Huntrieser, H.: The global lightning-induced nitrogen oxides source rate, *Atmos. Chem. Phys.*, 7, 3823–3907, 2007, <http://direct.sref.org/1680-7324/acp/2007-7-3823>.
- Skamarock, W. C., Dye, J. E., Defer, E., Barth, M. C., Stith, J. L., Ridley, B. A., and Baumann, K.: Observational- and modelling-based budget of lightning-produced NO_x in a continental thunderstorm, *J. Geophys. Res.*, 108, 4305, doi:10.1029/2002JD002163, 2003.
- 10 Stefanutti, L., MacKenzie, A. R., Santacesaria, V., et al.: The APE-THESEO Tropical Campaign: An overview, *J. Atmos. Chem.*, 48, 1–33, 2004.
- Stohl, A., Forster, C., Eckhardt, S., Huntrieser, H., Heland, J., Schlager, H., Aufmhoff, H., Arnold, F., and Cooper, O.: A backward modelling study of intercontinental pollution transport using aircraft measurements, *J. Geophys. Res.*, 108, 4370, doi: 10.1029/2002JD002862, 2003a.
- 15 Stohl, A., Huntrieser, H., Richter, A., Beirle, S., Cooper, O., Eckhardt, S., Forster, C., James, P., Spichtinger, N., Wenig, M., Wagner, T., Burrows, J., and Platt, U.: Rapid intercontinental air pollution transport associated with a meteorological bomb, *Atmos. Chem. Phys.*, 3, 969–985, 2003b.
- Stolzenburg, M., Marshall, T. C., Rust, W. D., et al.: Horizontal distribution of electrical and meteorological conditions across the stratiform region of a mesoscale convective system, *Mon. Wea. Rev.*, 122, 1777–1797, 1994.
- Théry, C.: Evaluation of LPATS data using VHF interferometric observations of lightning flashes during the Eulinox experiment, *Atmos. Res.*, 56, 397–409, 2001.
- 20 Thomas, R. J., Krehbiel, P. R., Rison, W., Hamlin, T., Boccippio, D. J., Goodman, S. J., and Christian, H. J.: Comparison of ground-based 3-dimensional lightning mapping observations with satellite-based LIS observations in Oklahoma, *Geophys. Res. Lett.*, 27, 1703–1706, 2000.
- Thompson, A. M., Tao, W.-K., Pickering, K. E., Scala, J. R., and Simpson, J.: Tropical deep convection and ozone formation, *Bull. Am. Meteorol. Soc.*, 78, 1043–1054, 1997.
- 30 Uman, M. A., McLain, D. K., and Krider, E. P.: The electromagnetic radiation from a finite antenna, *Am. J. Phys.*, 43, 33–38, 1975.
- Wacker, R. S. and Orville, R. E.: Changes in measured lightning flash count and return stroke peak current after the 1994 U.S. National Lightning Detection Network upgrade: 1. Observa-

14865

- tion, *J. Geophys. Res.*, 104, 2151–2157, 1999.
- Wang, Y., DeSilva, A. W., and Goldenbaum, G. C.: Nitric oxide production by simulated lightning: Dependence on current, energy, and pressure, *J. Geophys. Res.*, 103, 19 149–19 159, 1998.
- 5 Wiens, K. C., Rutledge, S. A., and Tessendorf, S. A.: The 29 June 2000 supercell observed during STEPS. Part II: Lightning and charge structure, *J. Atmos. Sci.*, 62, 4151–4177, 2005.
- WMO: Scientific Assessment of Ozone Depletion: 1994, World Meteorological Organization Global Ozone Research and Monitoring Project, Report No. 37, 1995.
- Zhang, X. J., Helsdon, J. H., and Farley, R. D.: Numerical modeling of lightning-produced NO_x using an explicit lightning scheme: 1. Two-dimensional simulations as a “proof of concept”, *J. Geophys. Res.*, 108, 4579, doi:10.1029/2002JD003224, 2003.
- 10

Table 1. Positions of LINET sensors in Brazil (Jan–Feb 2005) and Germany (July 2005).

Country/ Station Name	Latitude	Longitude
Brazil	° S	° W
Marilia	22.235	49.965
Novo Horizonte	21.466	49.226
Bauru	22.358	49.027
Quirinhos	22.951	49.896
Araquara	21.813	48.199
Botucatu	22.848	48.432
Germany	° N	° E
Ravensburg	47.801	09.696
Regensburg	49.043	12.103
Weissenburg	49.019	10.960
Berchtesgaden	47.634	13.001
Lahr	48.365	07.828
Karlsruhe	49.093	08.426
Basel	47.561	07.969
Bamberg	49.880	10.914
Weiden	49.667	12.184
Stegen	48.076	11.139
Passau	48.572	13.424
Garching	48.269	11.674
Peissenberg	47.801	11.010
Geretsried	47.870	11.476
Buchloe	48.037	10.728
Stadtbergen	48.349	10.850
Lichtenau	47.881	11.080
Lagerlechfeld	48.181	10.840
Oberpfaffenhofen	48.087	11.280

14867

Table 2a. Estimates of horizontal LNOx mass flux F_{LNOx} , LINET stroke rate R_{LINET} , LNOx production rate per LINET stroke and per LIS flash P_{LNOx} , and global LNOx production rate per year G_{LNOx} .

Flight and Anvil Penetration / tropical (t) or subtropical (s)	Entry and Exit Time (UTC), s	Pressure Altitude, km	Mean, χ_{LNOx} , nmol mol ⁻¹	$ V_x - V_a ^1$, m s ⁻¹	ρ_a , kg m ⁻³	Δx , km	Δz , km	F_{LNOx}^2 , g(N) s ⁻¹	R_{LINET}^3 , (LINET strokes)s ⁻¹	P_{LNOx} , g(N) (LINET stroke) ⁻¹	P_{LNOx} , g(N) (LIS flash) ⁻¹	G_{LNOx} , Tg(N) a ⁻¹
040205_1a (t)	66 199–66 433	10.6	0.76	6.5	0.36	35	4	120	0.055	2205	1103	1.5
040205_5a (t)	67 682–67 833	10.7	1.10	5.9	0.36	25	4	113	0.054	2082	1041	1.4
040205_2b (t)	69 905–70 169	10.1	0.57	9.2	0.39	45	4	178	0.061	2914	1457	2.0
mean tropical ⁴										2400	1200	1.6
180205b.I (s)	74 056–74 209	10.6	0.42	17.7	0.36	28	3	109	0.025	4258	2129	3.0
180205b.II (s)	74 453–74 623	10.7	0.18	20.0	0.36	33	3	62	0.025	2430	1215	1.7
180205b.III (s)	75 013–75 186	10.1	0.65	12.2	0.39	32	3	143	0.025	5623	2811	3.9
180205b.IV (s)	75 601–75 761	10.1	0.21	20.0	0.39	30	3	71	0.025	2792	1396	1.9
180205b.V (s)	76 102–76 280	9.4	0.39	11.9	0.41	33	3	91	0.025	3568	1784	2.5
180205b.VI (s)	76 584–76 757	9.4	0.13	17.7	0.41	35	3	48	0.025	1876	938	1.3
mean subtropical ⁴										4483	2241	3.1
relative max. error			~50%	~50%		~40%	~50%	~190%	~90%	~280%	~310%	~320%

¹ Horizontal anvil outflow velocity, calculated from values in Table 2b.² The horizontal LNOx mass flux out of the anvil, see Eq. (4).³ Only LINET strokes with peak currents ≥ 10 kA were considered for an equivalent comparison between 040205 (strokes mainly inside the LINET centre) and 180205b (strokes along LINET periphery).⁴ The mean value for the tropical anvil penetrations 1a, 5a and 2b of 4 February 2005 is given. The mean value for the subtropical anvil penetrations I, III, and V of 18 February 2005 (penetrations closest to the maximum anvil outflow) is given.

14868

Table 2b. Measured wind velocity and direction in the anvil outflow and at the steering level¹.

Flight and Anvil Penetration	Measured Wind Direction in Anvil Outflow $d_a, ^\circ$	Measured Wind Velocity in Anvil Outflow $V_a, \text{m s}^{-1}$	Wind Direction at Steering Level ¹ $d_s, ^\circ$	Wind Velocity at Steering Level ¹ $V_s, \text{m s}^{-1}$
040205_1a	71±31	5.2±1.7	160	3.9
040205_5a	182±50	3.9±2.7	280	3.9
040205_2b	107±23	6.9±2.0	350	3.7
180205b_I	283±12	16.6±4.1	185	4.3
180205b_II	277±4	19.4±2.3	185	4.3
180205b_III	274±11	11.5±5.3	185	4.3
180205b_IV	279±5	19.2±1.3	185	4.3
180205b_V	268±10	11.6±5.7	185	4.3
180205b_VI	278±5	16.9±1.4	185	4.3

¹ The wind at the steering level ~3 km (~700 hPa) determines the mean motion of a thunderstorm cell.

14869

Table 3. Lightning-produced NO mass per LINET stroke P_{LNOx} , for tropical and subtropical thunderstorms considering different stroke peak currents.

Flight and Anvil Penetration/ tropical (t) or subtropical (s)	Registered/ Considered Stroke Peak Currents ¹ , kA	Duration, of Stroke Activity ² , min	Number of LINET Strokes ²	R_{LINET} , (LINET strokes) s^{-1}	F(LNOx) ³ , g(N) s^{-1}	Integral ⁴ , $\text{strokes s}^{-1} \times 10^{-2} \text{g m}^{-1}$ laboratory spark	C_{LINET} ⁵ , m laboratory spark stroke^{-1}	Total Nitrogen Mass Produced ⁶ , $\times 10^3 \text{g}$	P_{LNOx} ⁷ , $\text{g(N) (LINET stroke)}^{-1}$	S, $\text{g(N) (kA)}^{-1} (\times 10^3)$ laboratory spark ⁻¹	Pressure, hPa
040205_1a (t)	≥ 2	85	1287	0.252	120	4.04	29732	613	476	2.0	240
040205_5a (t)	≥ 2	40	400	0.167	113	4.52	24946	271	677	2.3	235
040205_2b (t)	≥ 2	85	1492	0.293	178	4.31	41239	906	607	1.8	260
180205b_III (s)	≥ 6	85	236	0.046	143	1.13	126393	731	3097	1.9	260
040205_1a (t)	≥ 10	85	278	0.055	120	2.71	44390	613	2205	2.8	240
040205_5a (t)	≥ 10	40	130	0.054	113	3.72	30285	271	2082	2.6	235
040205_2b (t)	≥ 10	85	311	0.061	178	2.48	71516	906	2914	2.1	260
180205b_III (s)	≥ 10	85	130	0.025	143	0.89	160324	731	5623	2.1	260

¹ On 4 February 2005 strokes with peak currents down to 2 kA were registered (mainly within the LINET centre). On 18 February 2005 only strokes with peak currents ≥ 6 kA were registered (mainly along the LINET periphery). Thus, on 18 February the detection efficiency for low peak currents was lower than on 4 February and the stroke characteristics are not comparable (upper half of the table). For a more equivalent comparison between the 4 and 18 February only strokes with peak currents ≥ 10 kA were considered (lower half of the table).

² On 4 February 2005 the registered/considered strokes until penetration were active from 16:55 to 18:20 UTC within anvil 1a (85 min), from 17:55 to 19:20 UTC within anvil 2b (85 min) and from 18:05 to 18:45 UTC within anvil 5a (40 min) On 18 February 2005 the registered/considered strokes until storm decay were active between 19:30 and 20:55 UTC (85 min).

³ The horizontal LNOx mass flux out of the anvil (see Eq. 4). Values from Table 2a are given.

⁴ For every 1 kA, the total number of strokes N_{LINET} , within the duration of the stroke activity, are summed up and divided by the duration of the stroke activity which gives $R_{\text{LINET}}(l)$ (see Eq. 9). $R_{\text{LINET}}(l)$ is then multiplied with the Wang et al. 1998 laboratory relationship $M_{\text{NOlab}}(l)$ (see below⁵ and Eq. 7) and summed up over all 1 kA intervals (part of Eq. 8 and Eq. 9): $\int_{l=1}^{\infty} R_{\text{LINET}}(l) \times M_{\text{NOlab}}(l) dl$ where $R_{\text{LINET}}(l) = \frac{1}{T} \int_{t=0}^T N_{\text{LINET}}(l, t) dt$

⁵ Conversion of the Wang et al. 1998 laboratory relationship to LINET strokes (see Eq. 8): $C_{\text{LINET}} = F_{\text{LNOx}} / \int_{l=1}^{\infty} R_{\text{LINET}}(l) \times M_{\text{NOlab}}(l) dl$

⁶ The total nitrogen mass produced by the thunderstorm within the duration of the stroke activity. For every 1 kA interval and the duration of the stroke activity, $M_{\text{NOlab}}(l)$ (Eq. 7) modified for LINET (Eq. 10) is multiplied with the total number of strokes $N_{\text{LINET}}(l)$ and summed up over all peak currents.

⁷ Nitrogen mass produced per considered stroke.

14870

Table 4a. LINET statistics of positive and negative CG and IC stroke fractions (here VLF pulses) and mean peak currents estimated for strokes with peak currents ≥ 1 kA.

Date	Type of Thunderstorms/ Area ^{1,2}	Number of Strokes ³	Mean Peak Current, kA	CG Strokes (Fraction), %	IC Strokes (Fraction), %	Height IC Strokes, km	Ratio Positive/Negative Strokes	Mean Peak Current (Fraction) for CG-, kA (%)	Mean Peak Current (Fraction) for CG+, kA (%)	Mean Peak Current (Fraction) for IC-, kA (%)	Mean Peak Current (Fraction) for IC+, kA (%)	Peak Current ≥ 10 kA (Fraction), %
Tropical (Brazil)												
230105	LINET centre	11324	6	<44	≥ 56	10.3 \pm 2.9	1.0	-9 (23)	+5 (21)	-5 (28)	+4 (28)	10
240105	LINET centre	419	6	37	63	10.5 \pm 3.3	1.1	-11 (21)	+5 (16)	-5 (26)	+4 (37)	11
250105	LINET centre	848	7	47	53	9.6 \pm 2.8	1.6	-11 (21)	+5 (26)	-8 (17)	+5 (36)	15
040205	LINET centre	36234	6	43	57	9.7 \pm 3.1	1.2	-10 (24)	+5 (19)	-5 (22)	+4 (35)	12
040205	LINET centre: anvil_1a*	1278	8	24	76	10.0 \pm 3.4	0.7	-15 (20)	+5 (4)	-8 (40)	+5 (36)	22
040205	LINET centre: anvil_5a**	439	12	88	12	11.6 \pm 3.2	0.7	-17 (55)	+7 (33)	-4 (5)	+4 (7)	34
040205	outside LINET centre: anvil_2b***	1466	8	(55)	(45)	(14.0 \pm 2.5)	0.7	-11 (34)	+6 (21)	-7 (25)	+6 (20)	21
250205	LINET centre	31221	5	42	58	10.3 \pm 3.6	1.5	-7 (19)	+5 (23)	-5 (20)	+4 (38)	8
Transition Trop.-Subtrop. (Brazil)												
290105	LINET centre	419	6	40	60	8.6 \pm 3.0	1.3	-12 (25)	+5 (15)	-6 (19)	+4 (41)	16
050205	LINET centre	1608	6	66	34	9.1 \pm 3.3	1.8	-9 (25)	+6 (41)	-5 (11)	+5 (23)	13
190205	LINET centre	17228	5	47	53	9.6 \pm 3.3	1.6	-9 (21)	+5 (26)	-5 (17)	+4 (36)	9
Subtropical (Brazil)												
180205	LINET periphery****	3368	(13)	-	-	-	0.5	-	-	-	-	(57)
180205	LINET periphery: anvil I-VI*****	236	(13)	-	-	-	0.6	-	-	-	-	(55)
Mid-latitude or Subtr. (Germany)												
290605	LINET centre: isolated TS	4232	9	54	46	9.3 \pm 2.4	0.6	-11 (35)	+6 (19)	-11 (29)	+5 (17)	25
040705	LINET centre	6337	5	72	28	10.8 \pm 3.3	0.8	-6 (39)	+4 (33)	-6 (16)	+4 (12)	10
100705	LINET centre	15174	6	65	35	9.6 \pm 3.3	0.9	-7 (36)	+4 (29)	-6 (17)	+4 (18)	12
150705	LINET centre	8607	6	81	19	10.2 \pm 3.3	0.8	-8 (47)	+4 (34)	-5 (10)	+3 (9)	15
290705	LINET centre: isolated TS	2254	9	76	24	9.1 \pm 3.4	0.5	-12 (52)	+6 (24)	-8 (14)	+4 (10)	29
290705 ⁴	LINET centre: isolated TS, red.	1761	11	81	19	9.0 \pm 3.5	0.4	-13 (58)	+7 (23)	-11 (12)	+6 (7)	37

¹ The LINET centre area over Brazil covers 21.5–22.5° S and 48.5–49.5° W (area with highest detection efficiency).

² The LINET centre area over Germany covers 48.5–49.5° N and 11.0–12.0° E (area with highest detection efficiency).

³ For statistical reasons only days with at least 400 strokes (≥ 1 kA) in the LINET centre area were considered and strokes that were defined as IC or CG strokes (undefined strokes were neglected). The numbers given are the total number of strokes registered between 00:00 and 24:00 UTC or for selected anvils.

⁴ Reduced dataset (sensor configuration similar to Brazilian configuration).

*Anvil_1a (21.2–21.7° S and 48.9–49.2° W) is located at the edge of the LINET centre area and partly outside.

**Anvil_5a (21.7–21.9° S and 48.4–48.7° W) is located mainly inside the LINET centre area and comparable to other estimates.

***Anvil_2b (21.0–21.5° S and 49.7–50.1° W) is located just outside the LINET centre area where the fraction of IC strokes in general decreases, so estimates for this anvil penetration (especially IC height) are not directly comparable to the other estimates.

****Subtropical thunderstorm system (19.4–20.0° S and 47.7–49.2° W, 14:00–21:00 UTC) is located along the LINET periphery (detection efficiency decreases) and therefore not well comparable to other estimates in this table (where peak currents ≥ 1 kA are considered).

*****Anvil I-VI (19.3–19.8° S and 48.9–49.2° W, 19:30–21:00 UTC) is located along the LINET periphery (detection efficiency decreases) and therefore not well comparable to other estimates in this table (where peak currents ≥ 1 kA are considered).

14871

Table 4b. Same as Table 4a, but for strokes with peak currents ≥ 10 kA.

Date	Type of Thunderstorms/ Area ^{1,2}	Number of Strokes ³	Mean Peak Current, kA	CG Strokes (Fraction), %	IC Strokes (Fraction), %	Height IC Strokes, km	Ratio Positive/Negative Strokes	Mean Peak Current (Fraction) for CG-, kA (%)	Mean Peak Current (Fraction) for CG+, kA (%)	Mean Peak Current (Fraction) for IC-, kA (%)	Mean Peak Current (Fraction) for IC+, kA (%)	Peak Current ≥ 10 kA (Fraction), %
Tropical (Brazil)												
230105	LINET centre	1144	19	<70	> 30	10.8 \pm 2.9	0.3	-23 (57)	+15 (13)	-13 (18)	+16 (12)	30
240105	LINET centre	47	19	77	23	10.0 \pm 3.7	0.3	-20 (66)	+13 (11)	-20 (11)	+13 (12)	28
250105	LINET centre	130	21	61	39	9.3 \pm 2.3	0.6	-25 (42)	+16 (19)	-26 (19)	+14 (20)	31
040205	LINET centre	4379	20	73	27	9.5 \pm 3.2	0.3	-22 (62)	+19 (11)	-17 (13)	+16 (14)	34
040205	LINET centre: anvil_1a*	129	18	65	35	10.2 \pm 2.7	0.2	-20 (62)	+16 (3)	-13 (24)	+16 (11)	26
040205	LINET centre: anvil_5a**	123	26	98	2	(8.5 \pm 0.8)	0.1	-27 (86)	+19 (12)	-14 (1)	+14 (1)	50
040205	outside LINET centre: anvil_2b***	308	19	70	30	(14.2 \pm 2.4)	0.3	-22 (58)	+14 (12)	-16 (19)	+16 (11)	28
250205	LINET centre	2406	18	65	35	11.0 \pm 4.0	0.4	-20 (50)	+17 (15)	-16 (20)	+14 (15)	24
Transition Trop.-Subtrop. (Brazil)												
290105	LINET centre	68	18	80	20	8.3 \pm 3.5	0.1	-20 (74)	+13 (6)	-13 (16)	+12 (4)	31
050205	LINET centre	212	22	80	20	8.0 \pm 2.6	0.9	-23 (45)	+24 (35)	-16 (8)	+17 (12)	33
190205	LINET centre	1499	18	77	23	10.2 \pm 3.9	0.4	-19 (62)	+15 (15)	-16 (12)	+15 (11)	25
Subtropical (Brazil)												
180205	LINET periphery****	1914	17	-	-	-	0.4	-	-	-	-	21
180205	LINET periphery: anvil I-VI*****	130	17	-	-	-	0.4	-	-	-	-	22
Mid-latitude or Subtr. (Germany)												
290605	LINET centre: isolated TS	1065	23	56	44	9.2 \pm 2.3	0.2	-24 (48)	+24 (8)	-24 (38)	+16 (6)	44
040705	LINET centre	607	20	71	29	9.7 \pm 3.0	0.2	-19 (56)	+26 (15)	-18 (24)	+23 (5)	30
100705	LINET centre	1845	20	71	29	9.3 \pm 3.4	0.2	-21 (60)	+17 (11)	-20 (23)	+19 (6)	32
150705	LINET centre	1253	19	90	10	10.3 \pm 3.3	0.1	-20 (80)	+19 (10)	-19 (9)	+16 (1)	31
290705	LINET centre: isolated TS	659	21	87	13	7.4 \pm 3.4	0.1	-22 (76)	+17 (11)	-24 (11)	+18 (2)	43
290705 ⁴	LINET centre: isolated TS, red.	659	21	87	13	7.4 \pm 3.4	0.1	-22 (76)	+17 (11)	-24 (11)	+18 (2)	43

¹ The LINET centre area over Brazil covers 21.5–22.5° S and 48.5–49.5° W (area with highest detection efficiency).

² The LINET centre area over Germany covers 48.5–49.5° N and 11.0–12.0° E (area with highest detection efficiency).

³ For statistical reasons only days with at least 40 strokes (≥ 10 kA) in the LINET centre area were considered and strokes that were defined as IC or CG strokes (undefined strokes were neglected). The numbers given are the total number of strokes registered between 00:00 UTC and 24:00 UTC or for selected anvils.

⁴ Reduced dataset (sensor configuration similar to Brazilian configuration).

*Anvil_1a (21.2–21.7° S and 48.9–49.2° W) is located at the edge of the LINET centre area and partly outside.

**Anvil_5a (21.7–21.9° S and 48.4–48.7° W) is located mainly inside the LINET centre area and comparable to other estimates.

***Anvil_2b (21.0–21.5° S and 49.7–50.1° W) is located just outside the LINET centre area where the fraction of IC strokes in general decreases, so estimates for this anvil penetration (especially IC height) are not directly comparable to the other estimates.

****Subtropical thunderstorm system (19.4–20.0° S and 47.7–49.2° W, 14:00–21:00 UTC) is located along the LINET periphery (detection efficiency decreases) but can be compared with the other estimates in this table, since only peak currents ≥ 10 kA are considered here.

*****Anvil I-VI (19.3–19.8° S and 48.9–49.2° W, 19:30–21:00 UTC) is located along the LINET periphery (detection efficiency decreases) but can be compared with the other estimates in this table, since only peak currents ≥ 10 kA are considered here.

14872

Table 4c. Equivalent potential temperature (Θ_e) and wind velocity and direction (calculated from ECMWF analyses) in the LINET centre area¹ for selected days with LINET strokes as listed in Table 4a–b.

Date	Mean and Std ² Θ_e at 850 hPa, K	Mean and Std Θ_e at 500 hPa, K	Mean and Std Wind Velocity at 200 hPa, m s ⁻¹	Wind Direction at 200 hPa
Tropical (Brazil)				
230105	349±2	341±1	11±3	SE
240105	349±2	341±1	8±3	SE–SW
250105	349±2	343±1	6±1	SW–NW
040205	344±3	340±1	5±2	SE–SW
250205	342±4	337±1	17±2	SW
Transition Tropical-Subtropical (Brazil)				
290105	344±2	342±2	16±1	W
050205	343±2	339±3	7±2	SW
190205	347±3	330±2	34±2	W
Subtropical (Brazil)				
180205	343±4	330±1	27±1	W
Mid-latitude (Germany)				
290605	331±3	326±2	21±2	NW–SW
040705	323±6	323±1	14±2	W–SW
100705	321±3	319±1	8±2	NE
150705	323±3	320±1	20±7	NW
Subtropical (Germany)				
290705	340±5	325±2	19±3	SW

¹ The LINET centre area over Brazil covers 21.5–22.5° S and 48.5–49.5° W, and over Germany 48.5–49.5° N and 11.0–12.0° E.

² Std = Standard deviation.

14873

Table 5. Overview over mean peak currents for negative (and partly for all¹) CG flashes measured in different regions (mainly United States = U.S.) indicating a latitudinal gradient as suggested by Orville (1990) and Orville et al. (2002). Higher mean peak currents were estimated in the older studies owing to a lower detection efficiency of low peak currents. The U.S. NLDN was upgraded in 1994–1995 (Cummins et al., 1998) and in 2002–2003 (Cummins et al., 2006).

Reference (chronological)	Lightning Detection System	Mean Peak Current CG- (SE U.S.), kA	Mean Peak Current CG- (contiguous U.S.), kA	Mean Peak Current CG- (N and Central U.S.), kA
Orville (1990)	NLDN (=National Lightning Detection Network): 1988	Florida: 40–45 ¹		New England: 25
Petersen and Rutledge (1992)	Northern Australia: 1989–1990 (magnetic direction-finding lightning sensors from U.S.)	Darwin in tropical northern Australia: 39 ¹ (similar as in Florida)		
Price et al. (1997)	NLDN: summer 1988	Florida and Midwest: 36		
Lyons et al. (1998)	NLDN: 1991–1995	Large peak currents >75: preferable in SE U.S.		
Wacker and Orville (1999)	NLDN: 1989–1993 (pre-upgrade)	Florida: 41	contiguous U.S.: 38	Kansas: 33
Wacker and Orville (1999)	NLDN: 1994 (network upgraded)	Florida: 36	contiguous U.S.: 34	Kansas: 30
Wacker and Orville (1999)	NLDN: 1995 (network upgraded)	Florida: 33	contiguous U.S.: 30	Kansas: 26
Orville and Huffines (1999)	NLDN: 1995–1997	continental coastal areas of the U.S.: >26 ²		mountainous region: 15–20 ²
DeCaria et al. (2000)	NLDN: 12 July 1996			Colorado: 15 ¹
Orville et al. (2002)	NALDN (=North American Lightning Detection Network, including Canada plus contiguous U.S.): 1998–2000	Florida: 20–24 ²	contiguous U.S.: 16 ²	Central Canada: 12–18 ²
Langford et al. (2004)	NLDN: 12 Sep 2002			Colorado: 15
Ridley et al. (2004)	NLDN: 29/16 July 2002	Florida: 20/26		
Chowdhuri et al. (2005) this paper	NLDN: 1997–1998 and 1999–2000 LINET: only CG strokes <10 kA	SE U.S.: 20 ² Subtropical Brazil (18 Feb 2005): 19 ³	Mid-latitude Germany (June–July 2005): 19–24	Central/NW U.S.: 16/18 ² Tropical Brazil (Jan–Feb 2005): 19–27

¹ Here mean value independent of polarity.

² Here median value instead of mean value.

³ Here mean value for negative CG and IC strokes. No separation possible, since strokes occurred along LINET periphery.

14874

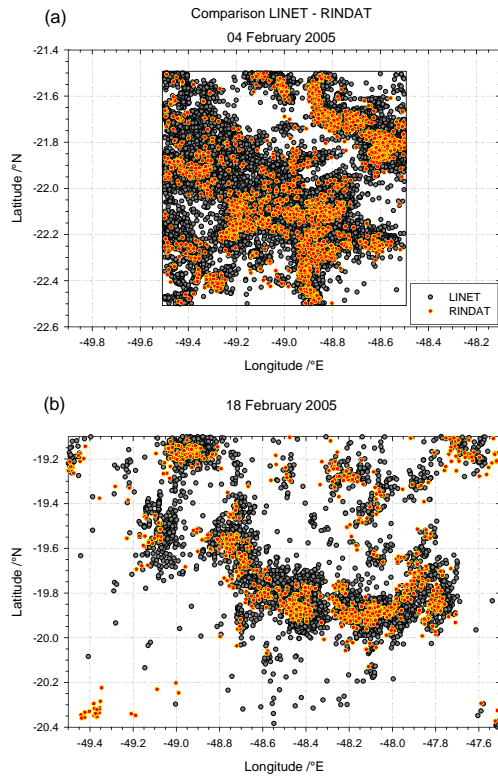


Fig. 1. Horizontal distributions of RINDAT and LINET strokes for the **(a)** 4 February 2005 in the centre area of the LINET detection network and for the **(b)** 18 February 2005 along the northern periphery, 00:00 UTC–24:00 UTC.

14875

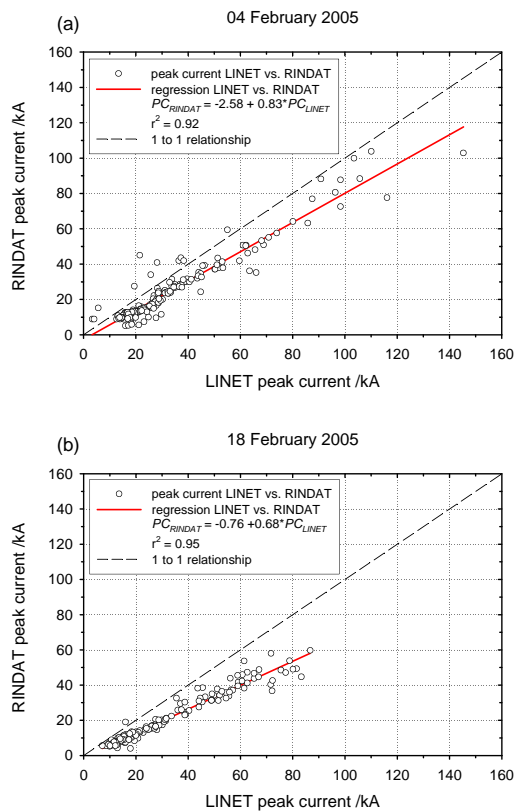


Fig. 2. Correlation between LINET and RINDAT peak currents (PC) for 222 and 173 selected strokes on **(a)** 4 and **(b)** 18 February 2005, respectively, in the areas shown in Fig. 1a–b.

14876

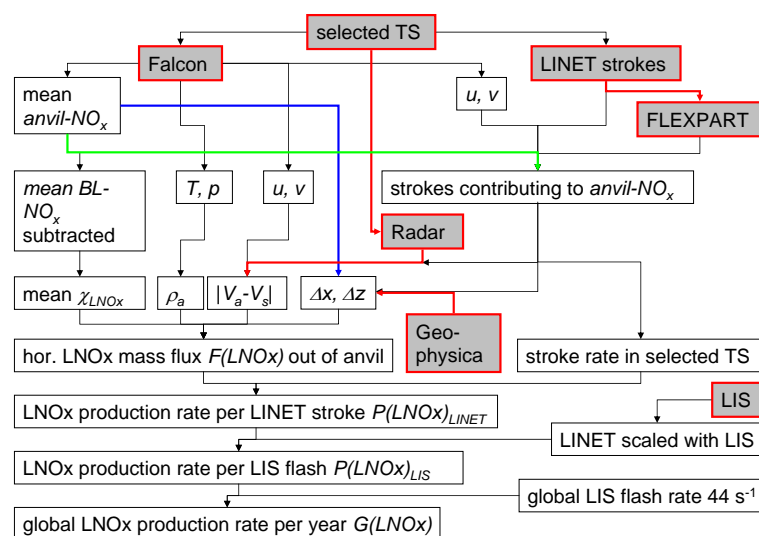


Fig. 3. Flow chart showing the introduced method to estimate the annual global LNO_x production rate from TROCCINOX field measurements (Falcon, Geophysica, LINET, Radar) combined with LIS data and model simulations (FLEXPART) of the selected thunderstorms (TS) all indicated with grey background. Different line colours are used to avoid misunderstanding of the flow direction at line intersections.

14877

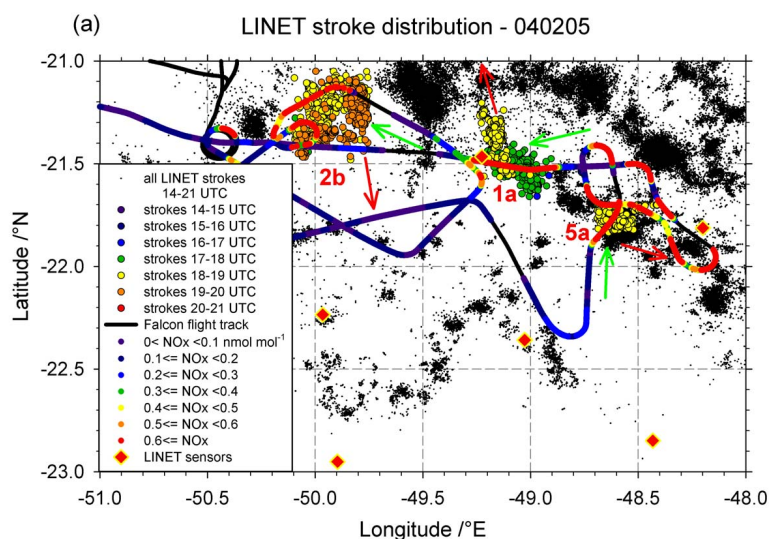


Fig. 4a. Horizontal distributions of LINET strokes on (a) 4 and (b) 18 February 2005. All strokes registered before the Falcon penetration within the selected thunderstorm systems on 4 February (tropical: labelled 1a, 5a and 2b) and 18 February (subtropical) are coloured. Falcon flight paths and NO_x mixing ratios are superimposed (colour/grey scale). The red arrows indicate the direction of the storm motion and the green arrows the main wind direction in the anvil outflow. In addition, the positions of the 6 LINET sensors listed in Table 1 are indicated in (a).

14878

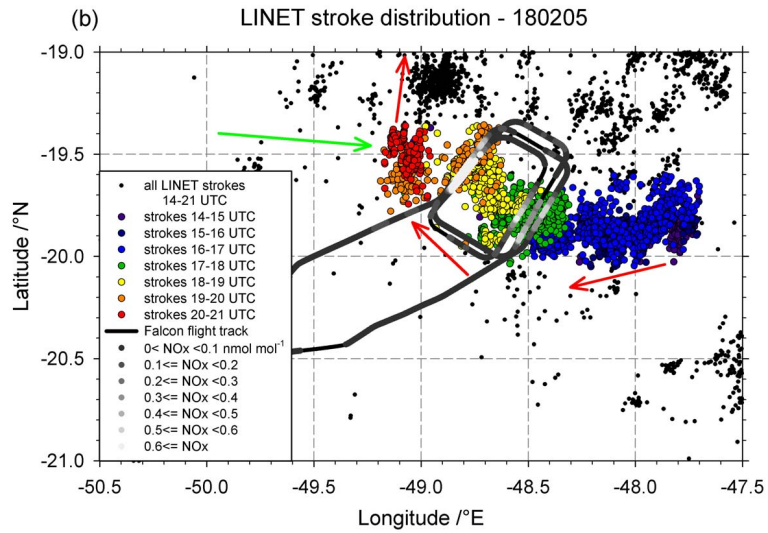


Fig. 4b. Continued.

14879

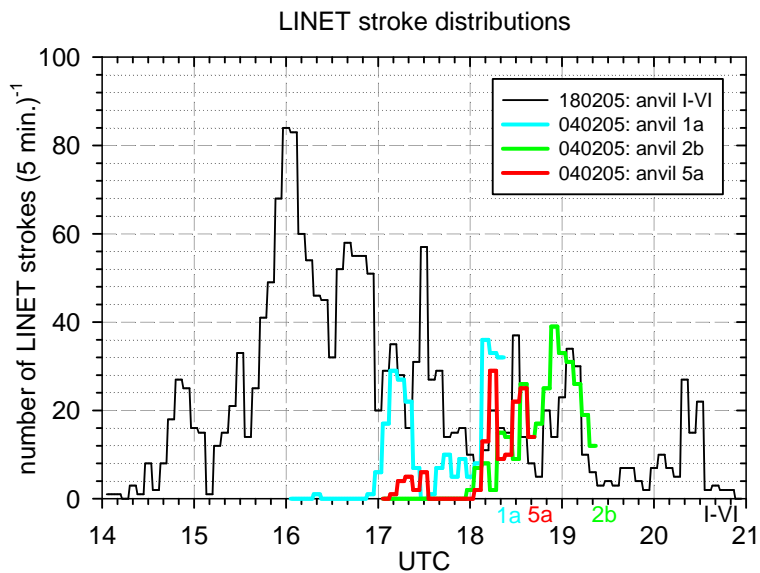


Fig. 5. Time series of LINET stroke rates for the selected thunderstorms (only strokes with peak currents ≥ 10 kA considered). On 4 February 2005, stroke rates in the investigated thunderstorms (active) are shown from storm initiation until penetration (tropical: labelled 1a, 5a and 2b). On 18 February 2005, the stroke rate within the selected thunderstorm system (sub-tropical) is shown from storm initiation until decay. The repeated penetration started first at 20:34 UTC (labelled I-VI) when the lightning activity decayed.

14880

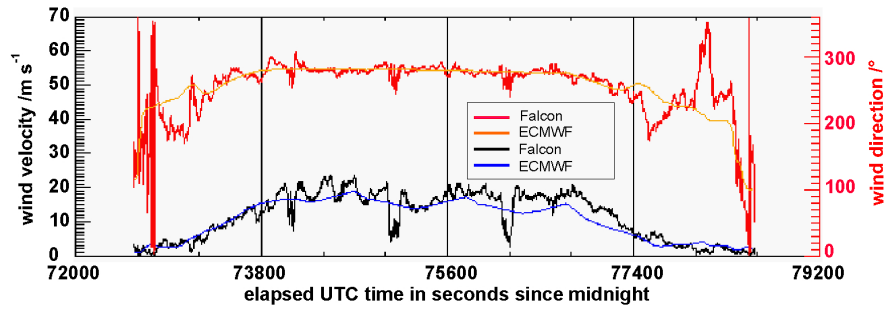


Fig. 6. Comparison between Falcon and ECMWF wind velocity (in black and blue, respectively) and wind direction (in red and orange, respectively) for the flight on 18 February 2005.

14881

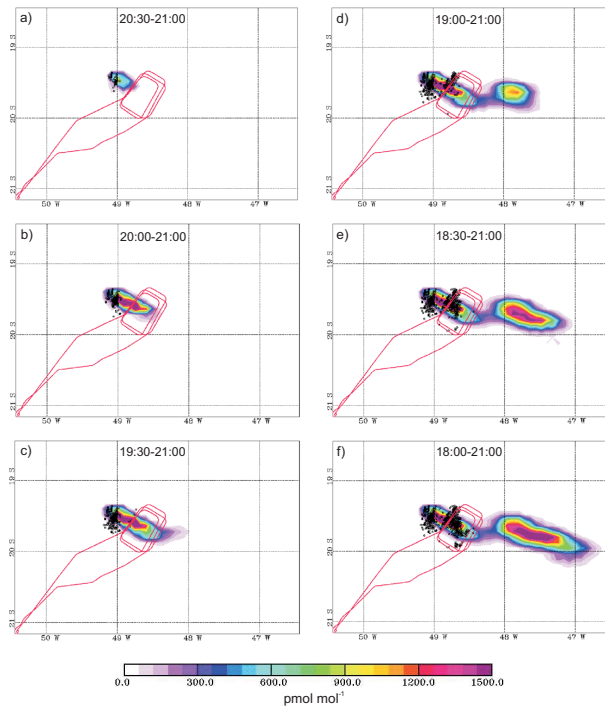


Fig. 7. FLEXPART lightning tracer simulations (NO_x at 10 km in pmol mol^{-1}) for the 18 February 2005 subtropical thunderstorm system considering all LINET strokes (black dots) after (a) 20:30 UTC, (b) 20:00 UTC, (c) 19:30 UTC, (d) 19:00 UTC, (e) 18:30 UTC, (f) 18:00 UTC and until 21:00 UTC. The Falcon track is superimposed in red.

14882

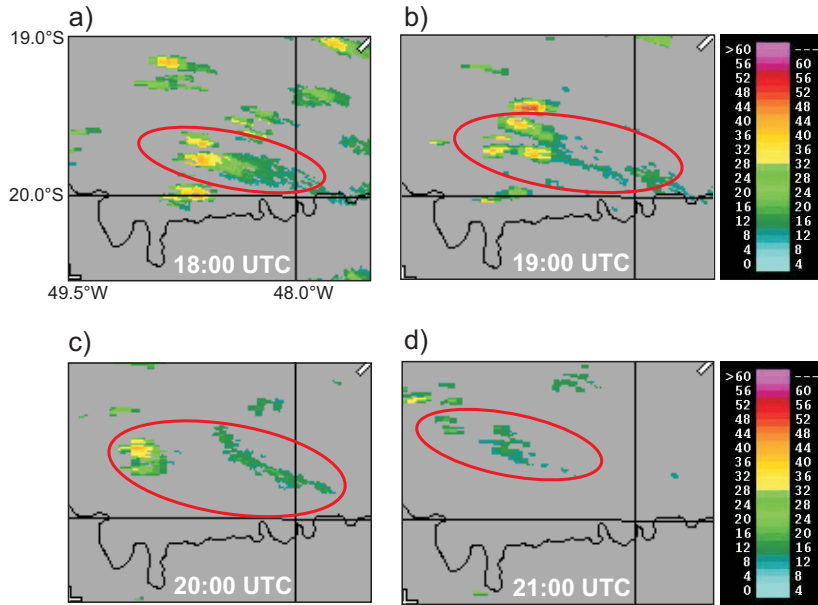


Fig. 8. Radar reflectivity as PPI scan (PPI_SURVEILLANCE operational product) in dBZ units measured at 0.0° elevation by the Bauru radar (22.4° S, 49.0° W) for the 18 February 2005 subtropical thunderstorm system (marked in red) at (a) 18:00 UTC, (b) 19:00 UTC, (c) 20:00 UTC and (d) 21:00 UTC.

14883

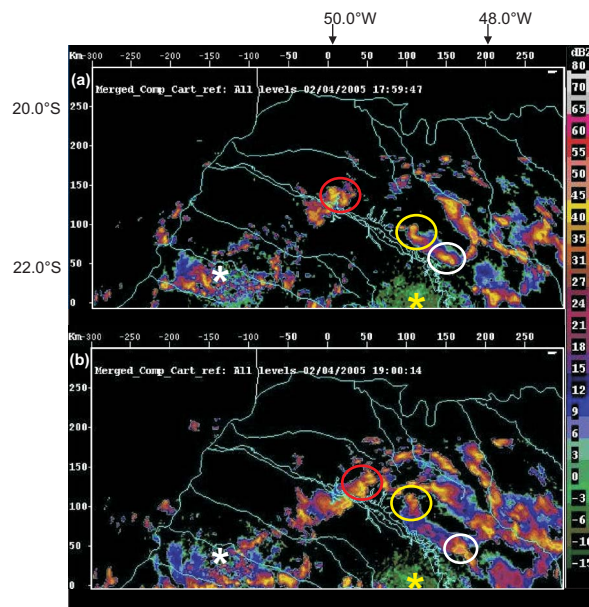


Fig. 9. Vertical maximum of the radar reflectivity (max CAPPI frame, unit dBZ). Composite from the Bauru (22.4° S, 49.0° W) and Presidente Prudente (22.1° S, 51.4° W) radars for the 4 February 2005 thunderstorms at (a) 17:59:47 UTC, and (b) 19:00:14 UTC. Selected thunderstorms are marked (tropical: yellow circle anvil 1a, white circle anvil 5a and red circle anvil 2b). The change in location of each of these circles between (a) and (b) indicates the storm motion, which can be compared with the storm motion indicated by arrows in Fig. 4a. The Bauru and Presidente Prudente radar sites are indicated with a yellow and white *, respectively. The left and upper axes give the distance in km, and the latitude and longitude are indicated in (a). (TITAN Software, installed at IPMet in collaboration with NCAR.)

14884

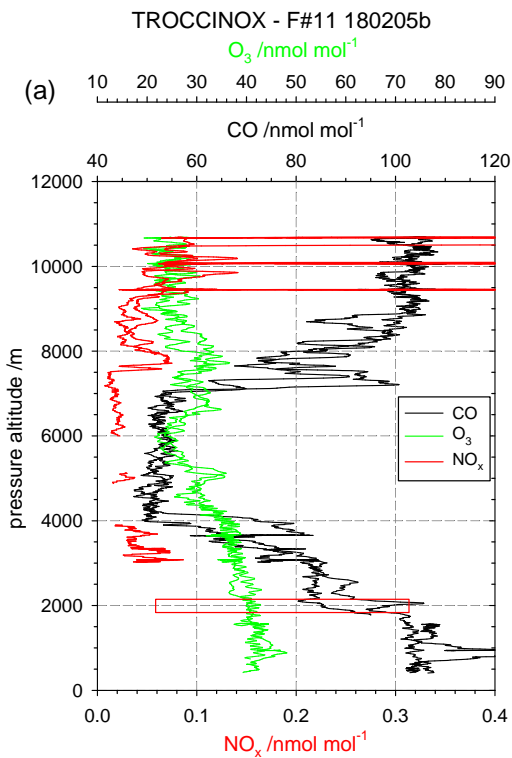


Fig. 10a. (a) Vertical profiles for CO, O_3 , and NO_x mixing ratios from the Falcon flight on 18 February 2005. The red box at 2 km altitude indicates the top of the mixed layer.

14885

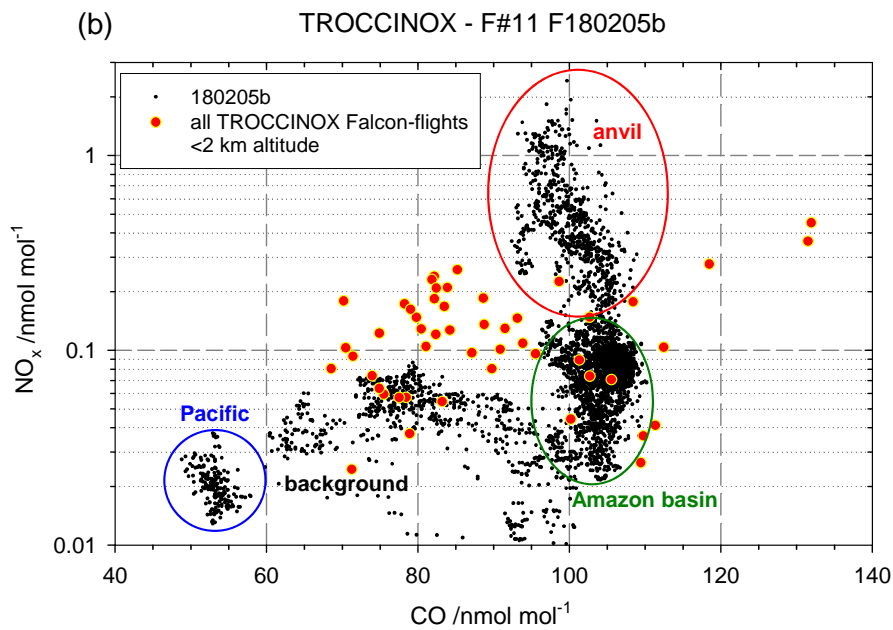


Fig. 10b. (b) Correlation plot for NO_x and CO for the same flight (black dots), and superimposed data from all TROCCINOX Falcon-flights in the boundary layer (<2 km) (red-yellow dots).

14886

TROCCINOX - F#11 180205b

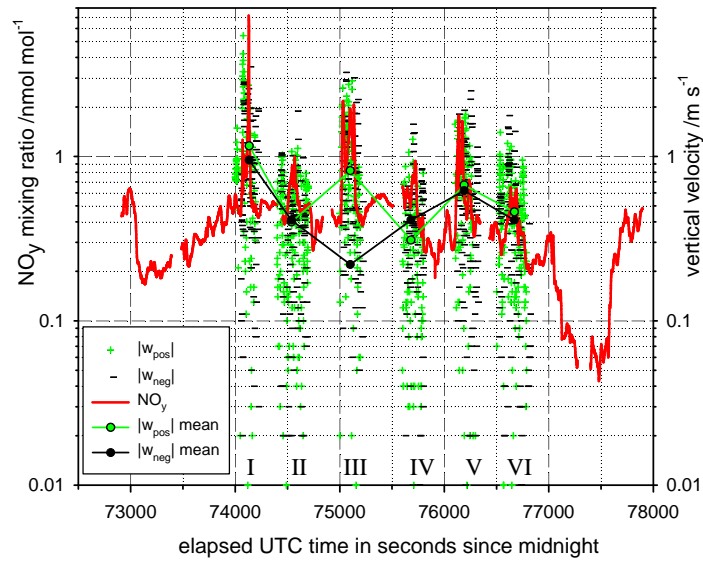


Fig. 11. Time series of NO_y mixing ratio and absolute vertical velocity for the Falcon flight on 18 February 2005. The anvil penetrations are labelled I–VI.

14887

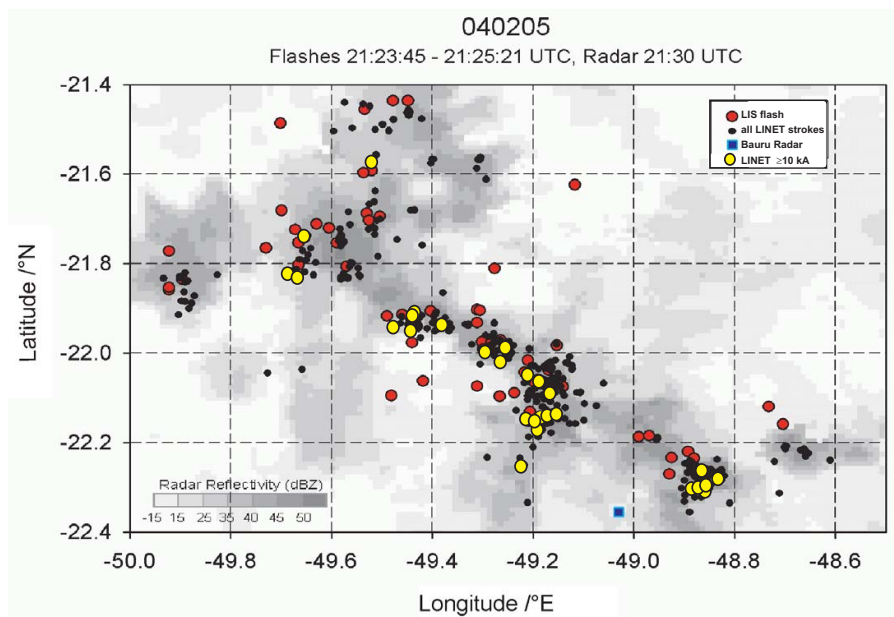


Fig. 12. Vertical maximum of the radar reflectivity (max CAPPI frame, unit dBZ) on 4 February 2005 at 21:30 UTC measured by the Bauru (22.4° S, 49.0° W) and Presidente Prudente (22.1° S, 51.4° W) radars. Superimposed are the horizontal distributions of LIS flashes (in red) and LINET strokes (black and yellow) for the time period 21:23:45 UTC–21:25:21 UTC

14888

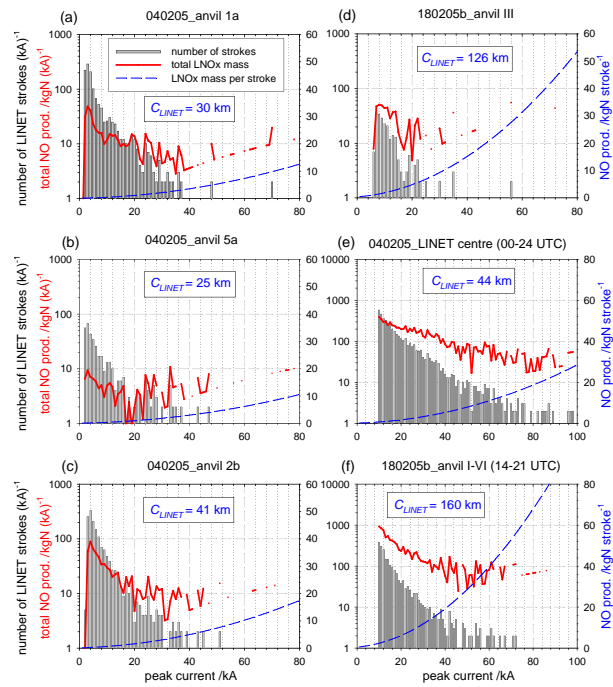


Fig. 13. Frequency distribution of LINET strokes (vertical grey bars) as a function of peak current for the selected tropical and subtropical thunderstorm systems on 4 and 18 February 2005, respectively (see Table 3). Superimposed is the laboratory result by Wang et al. (1998) modified for LINET strokes (blue dashed line) according to Table 3 (C_{LINET}) and the estimated total amount of LNOx mass produced per 1 kA LINET stroke interval for the selected thunderstorm systems (red line), see mass estimates in Table 3.

14889

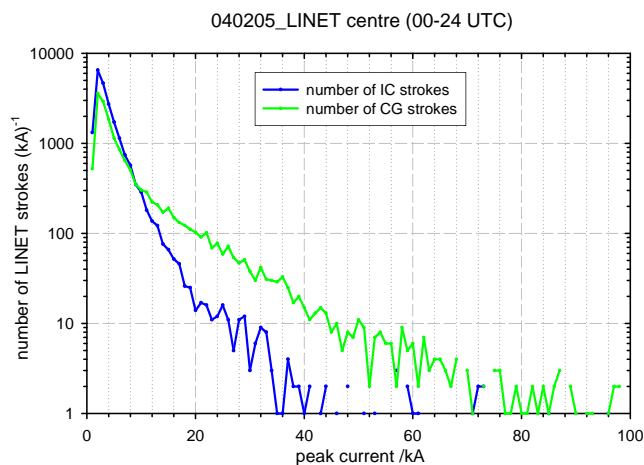


Fig. 14. Frequency distributions of LINET strokes as a function of peak current separated for IC (blue) and CG (green) strokes on 4 February 2005.

14890

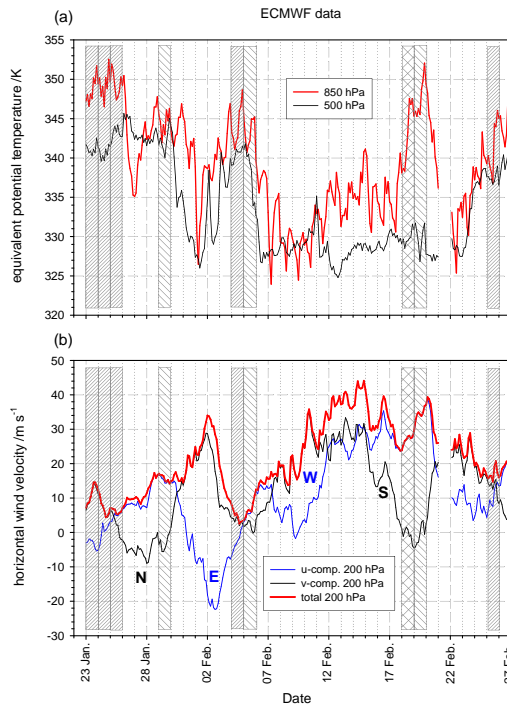


Fig. 15. Time series of **(a)** equivalent potential temperature and **(b)** horizontal wind velocity (separated in N, E, W, and S components) at different levels for the period of the TROCCINOX field phase in 2005 (22° S, 49° W). Days with registered LINET strokes in the LINET centre area (criterion: at least 40 strokes ≥ 10 kA, see Table 4b) are marked with grey patterns (tropical days: dense tilted, transition tropical/subtropical: less dense tilted, subtropical: crossed).

14891

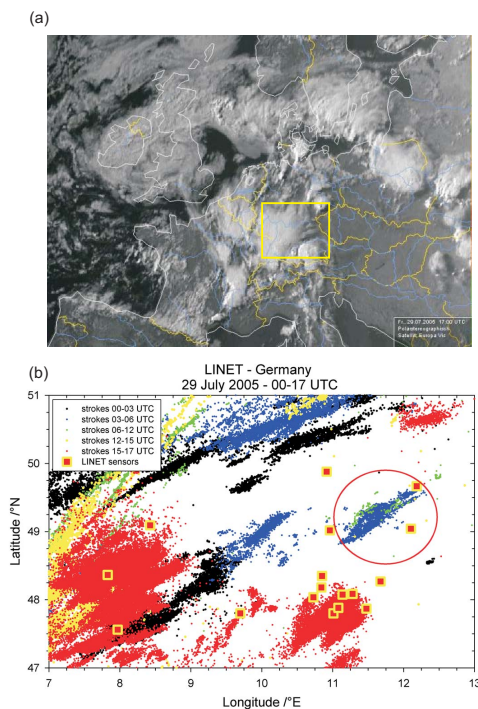


Fig. 16. Visible satellite image (Meteosat, DWD) for 29 July 2005 at 17:00 UTC showing the cloud distribution over Europe. Over southern Germany (mid-latitude) thunderstorms with high cloud tops (white) are visible. The yellow box indicates the area shown in **(b)**; horizontal distribution of LINET strokes for the 29 July 2005, 00:00 UTC–17:00 UTC. In the LINET centre area an isolated thunderstorm system was selected for further investigations (red circle).

14892

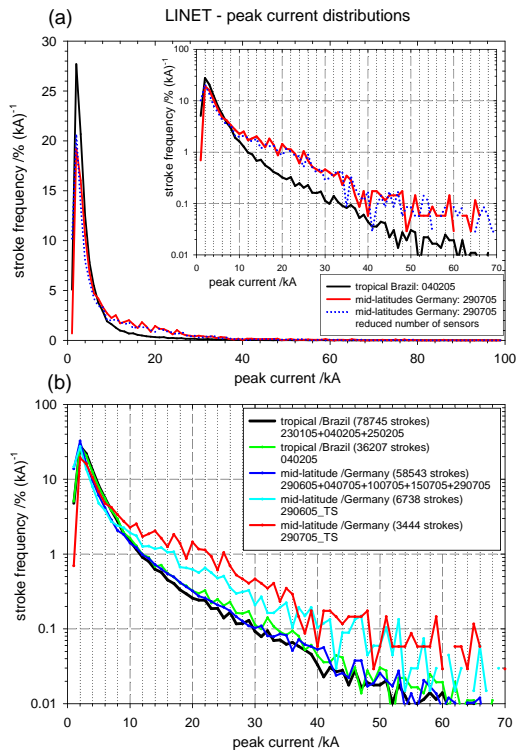


Fig. 17. Frequency distributions of LINET strokes (in the LINET centre area) as a function of peak current for different datasets given in colour (for comparison all curves were normalised). (c) Estimated amount of NO produced per 1 kA stroke interval (based on the Wang et al. 1998 relationship) for the dataset in (a) considering different values for C_{LINET} .

14893

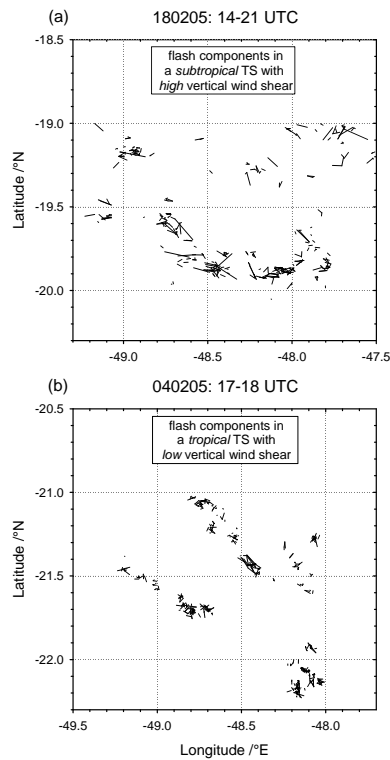


Fig. 18. Horizontal distributions of RINDAT flash components (several VLF sources in a flash connected to a line) for the selected (a) 18 February subtropical thunderstorm (TS) (14:00 UTC–21:00 UTC) and the (b) 4 February tropical thunderstorm (17:00 UTC–18:00 UTC).

14894