



the zonal mean zonal wind should be smaller than a critical value for vertical propagation. This theory also suggest that large scale waves (zonal wave number = 1, 2, 3) are more likely to propagate upwards because their associated critical wind speeds are higher. Studies by (Matsuno, 1970; Limpasuvan and Hartmann, 2000; Hu and Tung, 2002; Dickinson, 1969) not only confirmed this theory but also stressed the importance of vertical shear of the zonal mean zonal wind as well as the vertical gradient of the buoyancy frequency for vertical propagation of large scale waves.

Matsuno (1970) introduced the refractive index for stationary planetary waves as a diagnostic tool for studying the influence of the background zonal flow on planetary wave propagation. According to linear wave theory planetary waves, away from the source regions, tend to propagate toward the region of large positive refractive index squared. The existence of Rossby waves are prohibited where the refractive index squared is small or negative, which can happen if the zonal mean zonal wind is easterly, or westerly exceeding the critical wind speed.

The refractive index of Rossby waves as a diagnostic tool provides a framework in which the dynamical forcing of the stratosphere by tropospheric waves can be investigated. However, as shown by Li et al. (2007) the traditional analysis of the refractive index squared makes it difficult, if not impossible, to study the climatological state of the background flow for propagation of planetary waves. In calculating the climatology of the refractive index squared, the problem arises from averaging a time series that could consist of positive and negative values that may cancel each other and hence makes the interpretation of climatologies of this quantity difficult. Another weakness of the refractive index is that it is somewhat vague. Randel (1988) pointed out that, while using the refractive index as a diagnostic tool one should not overemphasize the details, since it is a qualitative guide. For instance Smith (1983) found that planetary waves can only propagate when and where the refractive index squared is positive and very large or avoid the region of large negative values of the refractive index. The vagueness arises from vague expressions such as very large positive and very large

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negative values of the refractive index which demonstrates the arbitrariness of the classic time mean diagnostic.

Here we attempt to address the modeling of such vagueness which has not been previously addressed. We present an algorithm based on fuzzy logic theory which addresses the above-mentioned vagueness and provides an estimate of the favorability of atmospheric background condition for planetary wave propagation as a function of latitude and altitude. Any diagnostic tool should be consistent with the general knowledge about stationary Rossby wave propagation condition (Table 1). The first and second criterion of the Table 1 are the most important findings of the seminal papers of Charney and Drazin (1961) and Matsuno (1970). They made a great contribution on the understanding of the propagation of planetary scale disturbances from the troposphere into the stratosphere. Eliassen and Palm (1961) based on the wave-mean flow interaction theorem showed that the planetary waves also have a strong influence on the zonal mean zonal wind. Matsuno (1970) and Charney and Drazin (1961) argue that only ultra-long waves (wave numbers 1–3) have the capability to propagate from the troposphere into the middle atmosphere. The criterion 3 expresses that the jet maxima blocks the planetary wave propagation and penetration through the jet maxima is prohibited Karoly and Hoskins (1982). The study of Chen and Robinson (1992) shows that the key parameter that controls the planetary wave propagation is the tropopause which acts like a valve for the vertical wave propagation from the troposphere into the stratosphere. Furthermore the study of Hu and Tung (2002) and Li et al. (2007) indicated that the large positive vertical shear of zonal wind at the tropopause height tends to enhance wave propagation (criterion 4).

Chen and Robinson (1992) and Hu and Tung (2002) have discussed the importance of vertical shear of zonal mean zonal wind on the vertical propagation of Rossby waves. Chen and Robinson (1992) showed that penetration of planetary waves from the troposphere into the stratosphere is sensitive to small changes in the vertical shear of zonal wind near the tropopause height. Hu and Tung (2002) identified that a positive vertical shear of zonal wind enhances wave propagation across the tropopause. Similarly large

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negative shear of zonal wind tends to trap the planetary waves in the troposphere and hence less is left to penetrate into the stratosphere. Any diagnostic tool that attempts to provide a climatology of stationary Rossby wave propagation conditions should reflect this theory. In fact, we try to develop an algorithm that is capable of demonstrating the enhancing influence of positive vertical shear of zonal wind and impeding influence of negative vertical shear of zonal wind on stationary Rossby wave propagation from the troposphere to the stratosphere.

Figure 1 shows the climatology of the zonal mean zonal wind and the vertical shear of zonal mean zonal wind ( $\text{m s}^{-1} \text{ km}^{-1}$ ) for the Northern Hemisphere winter months. Northern Hemisphere winter months include December, January and February (DJF) and Southern Hemisphere winter months include June, July and August (JJA). Due to the larger meridional temperature gradient between the tropics and mid latitudes, the magnitude of the wind shear between 20–40° N is about four times stronger than the vertical shear at higher latitudes. Regardless of magnitude, it is evident that it is positive in the troposphere and negative in the stratosphere in this latitude band. At tropopause heights of these regions, where the sign changes, we expect to see a discontinuity in the Rossby wave propagation as discussed by Hu and Tung (2002). We will show that our new diagnostic is consistent with this theory while both the time mean of refractive index squared and the probability of positive refractive index introduced by Li et al. (2007) cannot capture this characteristic. In this study, we focus on the vertical propagation of the planetary wave, as there are also many studies using refractive index studying the horizontal propagation of the planetary waves.

## 2 Data and method

In the current study we used daily mean zonal wind and temperature from the National Center for Environmental Prediction-National Centre for Atmospheric Research (NCEP-NCAR) Kalnay et al. (1996) to calculate the refractive index of Rossby waves for 50 winters (1961–2010) of both Northern and Southern Hemispheres. The refractive

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index for stationary planetary waves is defined as:

$$n_{k,l}^2(y, z) = \left( \frac{N^2}{f^2 \cos^2(\phi)} \right) \left[ \frac{\overline{q_\phi}}{\overline{u}} - \left( \frac{k}{a} \right)^2 - \left( \frac{\pi l}{2a} \right)^2 - \left( \frac{f \cos(\phi)}{2NH} \right)^2 \right] \quad (1)$$

where

$$\overline{q_\phi} = \cos(\phi) \left[ \frac{2\Omega}{a} \cos(\phi) - \frac{1}{a^2} \frac{\partial}{\partial \phi} \left[ \frac{\partial}{\partial \phi} (\overline{u} \cos(\phi)) \right] - \frac{f^2}{\rho_0} \left[ \frac{\partial}{\partial z} \left( \rho_0 \frac{\partial \overline{u}}{\partial z} \right) \right] \right] \quad (2)$$

is the meridional gradient of the zonal mean potential vorticity which is a fundamental quantity in Planetary wave dynamics and the stability of the zonal mean flow Andrews et al. (1987). Here  $H$ ,  $k$ ,  $l$ ,  $\rho_0$ ,  $f$ ,  $N^2$ ,  $a$ ,  $\Omega$  and  $\phi$  are the scale height, zonal and meridional wavenumbers, air density, Coriolis parameter, buoyancy frequency, the Earth's radius and rotation frequency and latitude respectively (Andrews et al., 1987; Matsuno, 1970).

Figures 2 and 3 show the time Mean Refractive Index Squared (MRIS, in the plots weighted with the Earth radius squared) of 50 winters for Northern and Southern Hemispheres respectively. The dependence of the MRIS on the zonal ( $k = 1, 2, 3$ ) and meridional wavenumbers ( $l = 1, 2, 3$ ) is visible in both figures. It can be seen that the multi-year average of MRIS gives unsatisfactory results. For instance, for  $(k, l) = (1, 1)$  very high values of the refractive index squared are found in high latitudes of the troposphere and the lower stratosphere. Moreover, in most areas of mid and high latitudes of the troposphere alternating positive and negative values of the refractive index squared leads to a noisy structure and makes the interpretation very difficult. The problem originates from overlapping of positive and negative values in the time-series and results in a reduction of climatological information. Such features of the time mean are also discussed by others (Mukougawa and Hirooka, 2004; Li et al., 2007). Too high values of MRIS northward of 75° N in the lower stratosphere are not consistent with criterion 3 in Table 1, because the strong jet is expected to block wave penetration from the

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5 troposphere to the stratosphere. The MRIS is also not able to capture the meridional wavenumber dependency on the wave propagation conditions (criterion 2 in Table 1). For example in the Southern Hemisphere, the difference between time mean of for wave (2, 1), (2, 2) and (2, 3) in the stratosphere (above 100 hPa) is small, suggesting  
 5 no considerable influence from the meridional wavenumbers on the vertical propagation of planetary waves from the troposphere to the stratosphere.

### 3 Probability of positive refractive index squared

Li et al. (2007) introduced the frequency distribution of days with negative refractive index squared as an alternative metric to describe how planetary waves can propagate.  
 10 Figure 4 shows the PDFs of positive refractive index squared for Northern Hemisphere winter time expressed as the percentage of days with positive  $n_{k,l}^2(y, z)$  for wave (1, 1), (1, 2) and (1, 3). By comparing to the time mean of the same waves we conclude that this quantity is capable of describing the required wave properties better than the time mean of  $n_{k,l}^2(y, z)$ . However, it results in high values of probability between 20–40° N in  
 15 the lower and middle stratosphere. This might be an over-optimistic result, because it is due to small positive values at these locations that exist throughout the winter season. In this respect the climatology of probability of positive refraction index squared does not meet the criterion 4 in Table 1.

Further evidence to show the importance of  $\frac{\partial}{\partial z}\bar{u}$  for vertical propagation of Rossby  
 20 waves can be provided by calculating the normalized vertical component of the Eliassen–Palm (EP) flux. Figure 5 shows that the normalized vertical component of EP flux has a discontinuity at the tropopause, indicating that upward penetration of waves is suppressed by the negative values above tropopause heights as suggested by Hu and Tung (2002). Sensitivity of  $n_{k,l}^2(y, z)$  to  $\bar{u}$  can be studied by comparing the  
 25 values of  $a^2\frac{\overline{q\phi}}{\bar{u}}$  and  $a^2\frac{\overline{q\phi}}{10}$ . Figure 6 shows the climatology of  $a^2\frac{\overline{q\phi}}{\bar{u}}$  and  $a^2\frac{\overline{q\phi}}{10}$  for DJF in the Northern Hemisphere. The subpolar maxima of  $a^2\frac{\overline{q\phi}}{\bar{u}}$  at the troposphere are not

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related to values at these regions, since by taking away the  $\bar{u}$ , the maxima is shifted to subtropics (25–40° N). This infers that small values of  $\bar{u}$  rather than  $\frac{\partial}{\partial z}\bar{u}$  at subpolar regions cause the maxima of  $n_{k,l}^2(y, z)$  at these regions.

### 4 Probability of favorable propagation condition for Rossby waves

5 A long standing issue in the interpretation of  $n_{k,l}^2(y, z)$  is its vagueness. As suggested by Matsuno (1970), large waves tend to propagate in regions of positive refractive index  $n_{k,l}^2(y, z)$  while they may be refracted or absorbed where  $n_{k,l}^2(y, z) < 0$ . Here (in the light of fuzzy sets and logic), we attempt to address the modeling of such vagueness. Fuzzy logic is a mathematical method for answering questions with imprecise  
 10 information (such as very large or very small refractive index), it deals with reasoning that is approximate rather than fixed and precise. The basic approach is to assign a value between zero and one to describe the area between the upper and lower limit. In classical logic everything is either true or false. However, in fuzzy logic truth is a matter of degree (Zadeh, 1965; Novak et al., 1999).

15 Here we assume that instead of each of the individual  $n_{k,l}^2(y, z, t)$  contributing equally to the final  $n_{k,l}^2(y, z)$  some  $n_{k,l}^2(y, z, t)$  contribute more than others. In this way, we distinguish between small positive and very large positive values to let very large positive values influence the final result more than small positive values. In this way classes or sets whose boundaries are not sharp will be introduced. We introduce  $\mu_{Ro}(y, z, t)$  as the Rossby wave MVF which provides mPDF and estimate the probability of favorable propagation condition of Rossby wave  $Pr_{Ro}(y, z)$ , as a function of latitude and height. We also provide the physical basis of the proposed method. For a detailed discussion of MVF, see the Appendix.

25 The advantage of our analysis over the traditional analysis of the refractive index is that without any reduction in the information due to cancellation of negative and positive

values of the refractive index squared, we estimate the likeliness for planetary waves to propagate from one region to another at any time, altitude and latitude.

In the Fig. 7 the black curve shows the MVF used in the calculation of favorable propagation condition of Rossby waves. For the negative  $n_{k,l}^2(y,z,t)$  region (part a) this function suggest that the rate of attenuation is very high and therefore wave propagation is prohibited in this region. Since our method is still based upon the linear wave theory, we assume a linear relationship between the magnitude of the  $n_{k,l}^2(y,z,t)$  and the probability of favorable propagation conditions for positive  $n_{k,l}^2(y,z,t)$  in a way that the higher the values of the  $n_{k,l}^2(y,z,t)$  the chances of propagation for the Rossby waves increases linearly (part b). Large values of the  $n_{k,l}^2(y,z,t)$  occur near the critical line where zonal mean zonal wind approaches zero ( $\bar{u} < 0.5 \text{ m s}^{-1}$  in this study). This region is also not favorable for Rossby wave propagation since at this region the linear wave theory breaks down and waves start to break and the waves are absorbed (part c). In the study of Li et al. (2007) the effect of the critical line on Rossby wave propagation is neglected since all the positive values of the  $n_{k,l}^2(y,z,t)$  are regarded as though small and very large positive values of the  $n_{k,l}^2(y,z,t)$  are equally favorable places for wave propagation. In fact very high values of the  $n_{k,l}^2(y,z,t)$  are not necessarily favorable conditions for the Rossby wave propagation. In this study the  $n_{k,l}^2(y,z,t)$  higher than 600 is considered as the critical line region, obtained from the climatology of the refractive index when  $\bar{u} < 0.5 \text{ m s}^{-1}$ . As we will show, this function gives us an improved picture of planetary wave propagation condition in climatologies. Higher values of  $Pr_{\text{Ro}}(y,z)$  provide a window of opportunity for planetary waves to propagate at any latitude and height. Likewise, smaller values of  $Pr_{\text{Ro}}(y,z)$  demonstrate the places where Rossby waves propagate away from these regions. The sensitivity of  $Pr_{\text{Ro}}(y,z)$  values to the shape of the MVF function is discussed in Appendix.

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## 5 Results and discussions

Figure 8 demonstrates the climatology of probability of favorable propagation condition of Rossby waves for zonal wavenumbers ( $k = 1, 2, 3$ ) and meridional wavenumbers ( $l = 1, 2, 3$ ) for the Northern Hemisphere winter season. The most common feature for all waves are their rather large probability to propagate in the troposphere (below 200 hPa) in winter season. It is also evident that the most favorable propagation condition is in the lower troposphere of the mid-latitude region. The values of Fig. 8 are independent of Rossby wave generation and explain how the waves, when generated, would propagate given the structure of the mean flow. However the regions of highly favorable Rossby wave propagation and source region for wave generation (asymmetries at the surface, land-sea contrasts, and sea surface temperature asymmetries) are coincident. It is also clear that longer waves have more opportunity to penetrate to the stratosphere.

Karoly and Hoskins (1982) by using ray tracing technique from geometrical optics and wave propagation in a slowly varying medium, showed that wave rays which are parallel to the group velocity vector tend to refract toward large refractive index squared. They also found that Rossby waves have a tendency to propagate along greater circles and most of the upward propagation of Rossby waves will be refracted toward the equator (even if the refractive index squared were positive at all height in their study). Similar to this theory, we also found a channel or waveguide of large probability of favorable propagation condition for Rossby waves. The strong westerlies act as a waveguide of Rossby waves and direct them vertically through the tropopause and allow them to penetrate to higher altitudes from their source region (troposphere). These areas are south of  $40^\circ \text{ N}$  in winter of the Northern Hemisphere for large waves and are indicated by  $Pr_{\text{Ro}}(y,z) > 50\%$ .

The study of Karoly and Hoskins (1982) also revealed that Rossby waves tend to propagate on the edges of strong westerlies and avoid penetrating through the jet. This fact is also clear in our results, where north of  $60^\circ \text{ N}$  and above 200 hPa, the probability

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of favorable condition for Rossby waves show relatively smaller values, comparing to similar altitude ranges between 30 and 50° N. The same climatologies as Fig. 8 are presented in Fig. 9 for the Southern Hemisphere. Similar to the Northern Hemisphere, all large scale waves have a rather large chance to propagate in the troposphere in winter. It can be seen that the larger the waves, the probability of favorable condition for them to propagate upward are larger.

Figure 10 demonstrates the differences between probability of positive refractive index (calculated by PDFs) and probability of favorable propagation condition of Rossby waves (calculated by mPDFs) for Northern Hemisphere wintertime for wave (1, 1), (1, 2) and (1, 3). The maximum difference is found at 20–40° N of the middle and upper troposphere which can reach to 50 %. This unsatisfactory result of the probability of positive refractive index is due to small positive values at these places which are consistent throughout the winter season. The area of maximum difference between  $Pr_{Ro}(y, z)$  and probability of positive refractive index remains the same for all wavenumbers at both Northern and Southern Hemispheres (not shown).

The most important difference between the Northern and Southern Hemisphere occurs in the high latitudes of the stratosphere, where in the Northern Hemisphere, zonal wavenumber = 1 has a good opportunity to propagate upward ( $Pr_{Ro}(y, z) > 40\%$ ), while in the Southern Hemisphere it has a rather poorer chance to propagate upward. This is consistent with the theoretical explanation of the vertical propagation of Rossby waves from the troposphere to the stratosphere by Charney and Drazin (1961). The zonal mean zonal wind should be weaker than a critical strength for upward propagation of Rossby waves. The strong stratospheric winter polar vortex of the Southern Hemisphere will block and reflect wave activity. The critical strength depends on the scale of the wave and is not a function of the background zonal regime.

The most important information which is lost from the time mean of  $n_{k,l}^2(y, z)$  is the role of meridional wavenumbers on the wave propagation conditions. For instance in the Southern Hemisphere, the difference between the time mean of  $n_{k,l}^2(y, z)$  for wave (2, 1), (2, 2) and (3, 3) in the stratosphere (above 100 hPa) is not large which is one of

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the unsatisfactory results of time mean of  $n_{k,l}^2(y, z)$ . It is only in the light of  $Pr_{Ro}(y, z)$  values that we can understand the impact of meridional wavenumbers on the wave propagation in the stratosphere. Note that, at the same latitude range of the Southern Hemisphere,  $Pr_{Ro}(y, z)$  values are as high as 45 % for wave (2, 1) in mid-latitudes of stratosphere, while the  $Pr_{Ro}(y, z)$  values reach to less than 5 % for wave (3, 3).

## 6 Usefulness and appropriateness of $Pr_{Ro}(y, z)$

In order to test the appropriateness of the  $Pr_{Ro}(y, z)$  in climatological studies of stationary planetary wave propagation, we further investigate the sensitivity of the  $Pr_{Ro}(y, z)$  to different zonal flow regimes in the stratosphere. Following Castanheira and Graf (2003), we constructed two data sets based upon the strength of the westerlies in the lower stratosphere (50 hPa) at 65° N. According to the Charney and Drazin (1961) criterion, if the background flow is westerly and smaller than the latitude and wave number dependent critical Rossby velocity, the planetary waves can penetrate from the troposphere into the stratosphere, otherwise wave reflection occurs and tropospheric flow may be modified. SVR is identified when  $\bar{u}_{50}(65^\circ \text{N}) > 20 \text{ m s}^{-1}$  and WVR is considered when  $0 < \bar{u}_{50}(65^\circ \text{N}) < 10 \text{ m s}^{-1}$ , where  $\bar{u}_{50}(65^\circ \text{N})$  is the 50 hPa zonal mean zonal wind at 65° N. The  $20 \text{ m s}^{-1}$  threshold reflects the critical Rossby velocities ( $20 \text{ m s}^{-1}$ ) for ZWN = 1 for a climatological Northern Hemisphere zonal wind profile.

Table 2 demonstrates the periods of different polar vortex regimes that last for at least 30 consecutive days in DJF. Since in DJF the stratospheric flow consists of strong westerlies (in the absence of vertical wave propagation), the number of SVR events are more than WVR events. The results of  $n_{k,l}^2(y, z)$  and  $Pr_{Ro}(y, z)$  for WVR and SVR for wave (1, 1) are presented in Fig. 11. It is found that in comparison to climatologies (Fig. 8) both WVR and SVR show similar patterns. However, the waveguide at mid latitudes is much narrower in SVR than WVR. In addition, the average values of  $Pr_{Ro}(y, z)$  in the stratosphere are greater in WVR than SVR. These results show that

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planetary waves have more chance to penetrate and force the stratosphere in WVR than SVR. In other words, values of  $Pr_{Ro}(y, z)$  are sensitive to stratospheric westerlies and are consistent with the general knowledge about planetary wave propagation from the troposphere to the stratosphere. An enhancement of wave propagation northward of  $70^\circ$  N in the lower stratosphere and a slight reduction in the favorability of wave propagation between  $50$ – $70^\circ$  N in the stratosphere are found. On the other hand it can be seen that due to the high level of noisiness the interpretation of the difference of  $n_{k,l}^2(y, z)$  between WVR and SVR is very difficult. Since the highest difference in the favorability of wave propagation between WVR and SVR occurs northward of  $50^\circ$  N in the stratosphere, we further calculate the difference in the vertical component of EP flux between WVR and SVR in this region (Fig. 12). An enhancement of vertical EP flux is obtained northward of  $65^\circ$  N in the lower stratosphere during WVR while a decrease in this quantity is obtained southward of this region in the middle and upper stratosphere. By comparing the differences of  $n_{k,l}^2(y, z)$ ,  $Pr_{Ro}(y, z)$  and vertical component of EP flux during WVR and SVR, it can be seen that the pattern of differences between  $Pr_{Ro}(y, z)$  and vertical component of EP flux are similar. Therefore, based upon this analyses, we suggest that this diagnostic tool can be useful for studying wave-mean flow interaction.

## 7 Conclusions

Climatological values of the time mean of the refractive index squared derived from 50 winters (1961–2010) of both Northern and Southern Hemispheres are calculated to show several problematic features of this important quantity in climatologies. In order to improve these unsatisfactory results, we introduced probability density functions (PDFs) of positive refractive indices as a function of zonal and meridional wave numbers. We also compared this quantity with a modified set of PDFs (mPDFs) and demonstrate their superior performance compared to the climatological mean of refractive indices and the original PDFs. Without any reduction in the information,  $Pr_{Ro}(y, z)$  estimates the likeliness for stationary Rossby waves to propagate from one region to

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another at any time, altitude and latitude in a climatological sense. The higher the  $Pr_{Ro}(y, z)$  the easier it is for planetary waves to propagate. Smaller values of  $Pr_{Ro}(y, z)$  demonstrate the places where Rossby waves are absorbed or reflected from these regions. It is also found that by using this quantity one can easily study the climatological effect of meridional wavenumbers on stationary Rossby waves propagation without the difficulty of the interpretation of the noisy structure of the time mean refractive index. Our diagnostic tool is also capable of demonstrating the enhancing influence of positive vertical shear of zonal wind and impeding influence of negative vertical shear of zonal wind on stationary Rossby wave propagation from the troposphere to the stratosphere. This diagnostic tool successfully shows that for WVR there is more space for the vertical propagation of Rossby waves from the troposphere to the stratosphere. In contrast, SVR tend to block and reflect vertical propagation of stationary Rossby waves. Since our diagnostic tool is consistent with the theoretical understanding of vertical propagation of Rossby waves from the troposphere to the stratosphere, therefore we suggest that this diagnostic tool has the capacity to be used in assessing planetary wave propagation conditions in climate models.

## Appendix

The probability of favorable propagation condition of Rossby waves  $Pr_{Ro}(y, z)$  can be written as:

$$Pr_{Ro}(y, z) = \frac{\sum_{t=1}^n \mu_{Ro}(y, z, t)}{\sum_{t=1}^n t} \times 100 \quad (\text{A1})$$





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**Table 1.** A summary of known facts about stationary Rossby wave propagation. Any diagnostic tool that attempts to provide a climatology of stationary Rossby wave propagation conditions should be consistent with these criteria.

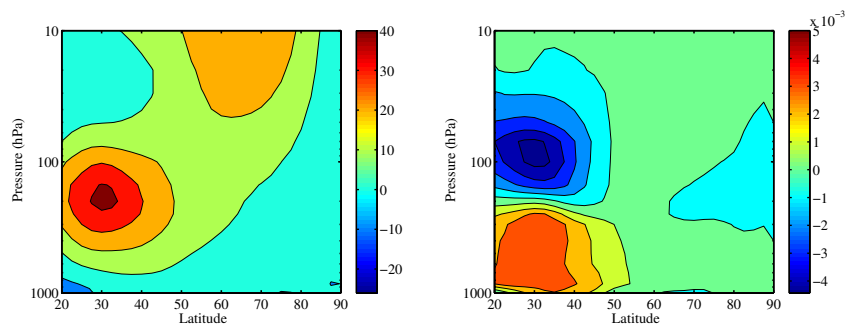
1	For all stationary Rossby waves the most favorable propagation conditions are in the lower troposphere of the mid-latitude region. Upper troposphere and lowermost stratosphere of mid-latitude regions are also favorable for Rossby wave propagation.	Matsuno (1970); Charney and Drazin (1961)
2	For large scale waves (horizontal and meridional wave numbers 1 to 3) the probability to propagate vertically is highest.	Matsuno (1970); Charney and Drazin (1961)
3	Rossby waves tend to propagate on the edges of strong westerly winds and avoid penetrating through the jet maxima. Therefore, the strong stratospheric polar night jet of the Southern Hemisphere in the winter will block and reflect large scale waves.	Karoly and Hoskins (1982)
4	Strong vertical shear (positive) is likely to enhance the vertical propagation of waves.	Chen and Robinson (1992)

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**Table 2.** Periods of polar vortex regimes lasting for at least 30 consecutive days in DJF; left: Strong Vortex Regime. Right: Weak Vortex Regime.

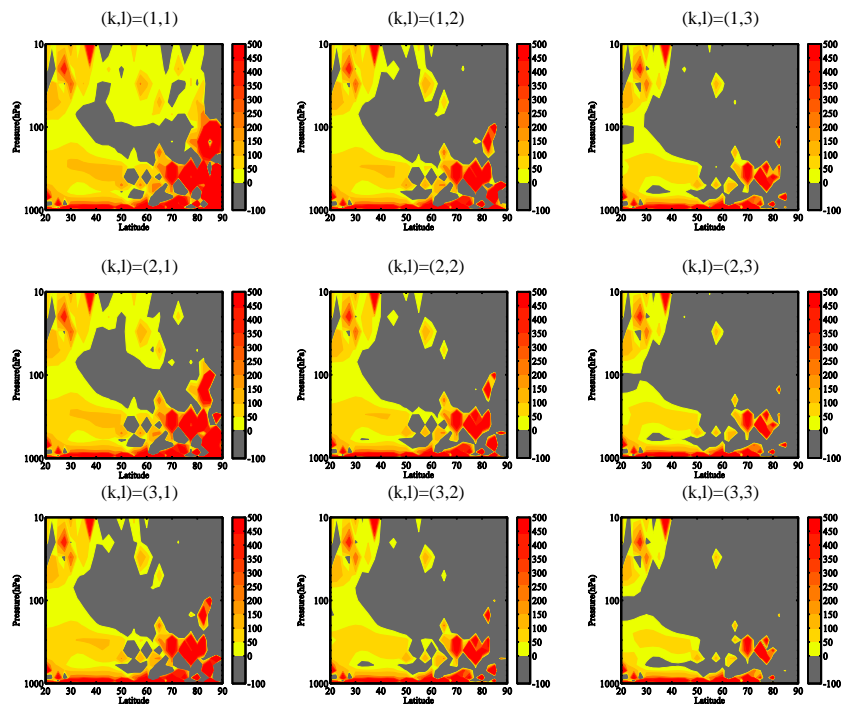
Strong Vortex Regime (SVR)		Weak Vortex Regime (WVR)	
Starting date	Ending date	Starting date	Ending date
20 Dec 1961	20 Feb 1962	20 Dec 1968	27 Jan 1969
24 Dec 1963	28 Feb 1964	28 Dec 1984	13 Feb 1985
3 Jan 1967	28 Feb 1967	09 Dec 1998	11 Jan 1999
1 Dec 1975	28 Feb 1976	2 Jan 2004	28 Feb 2004
1 Dec 1987	14 Jan 1988		
16 Dec 1988	17 Feb 1989		
17 Dec 1989	28 Feb 1990		
1 Dec 1991	18 Jan 1992		
5 Dec 1992	11 Feb 1993		
1 Dec 1994	18 Jan 1995		
7 Dec 2004	21 Feb 2005		
30 Dec 2006	26 Feb 2007		
23 Dec 2007	13 Feb 2008		

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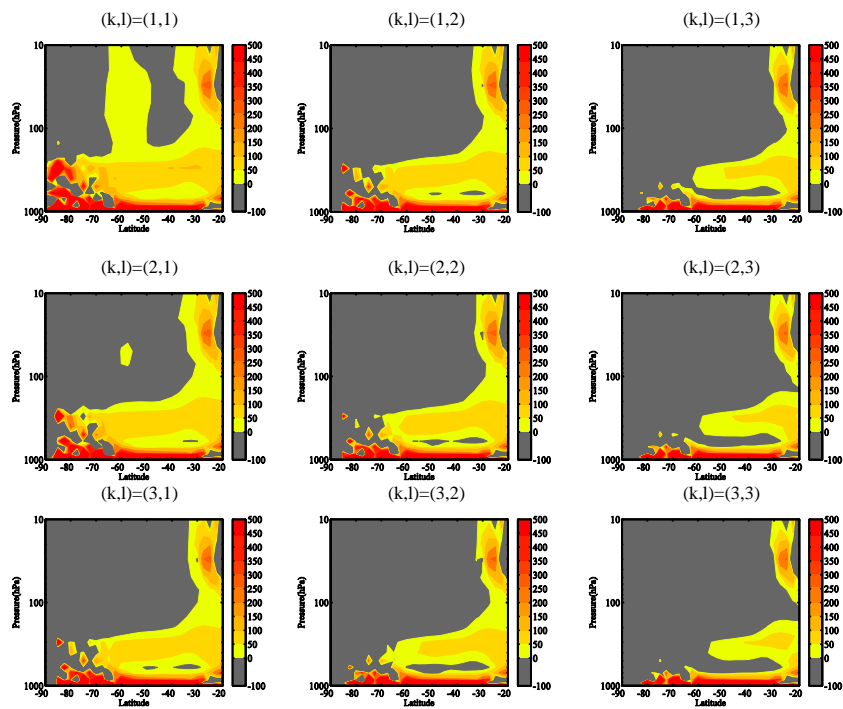
**Figure 1.** Climatology of the zonal mean zonal wind (left) in and the vertical shear of zonal mean zonal wind (right) for the Northern Hemisphere during DJF. The units are  $\text{m s}^{-1}$  for zonal mean zonal wind and  $\text{m s}^{-1} \text{ km}^{-1}$  for the vertical shear of zonal mean zonal wind respectively.

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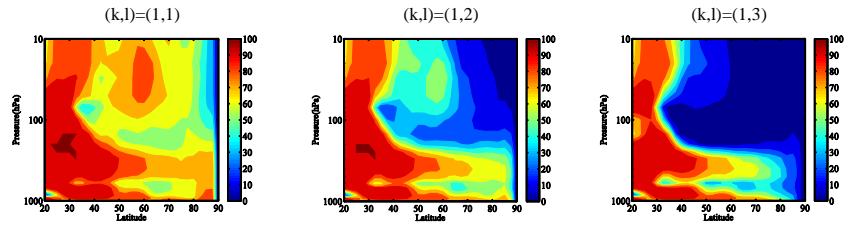
**Figure 2.** Climatology of refractive index squared ( $a^2 n_{k,l}^2(y, z)$ ) of 50 winters (1961–2010) in the Northern Hemisphere. Regions with negative  $a^2 n_{k,l}^2(y, z)$  are shaded with gray color.

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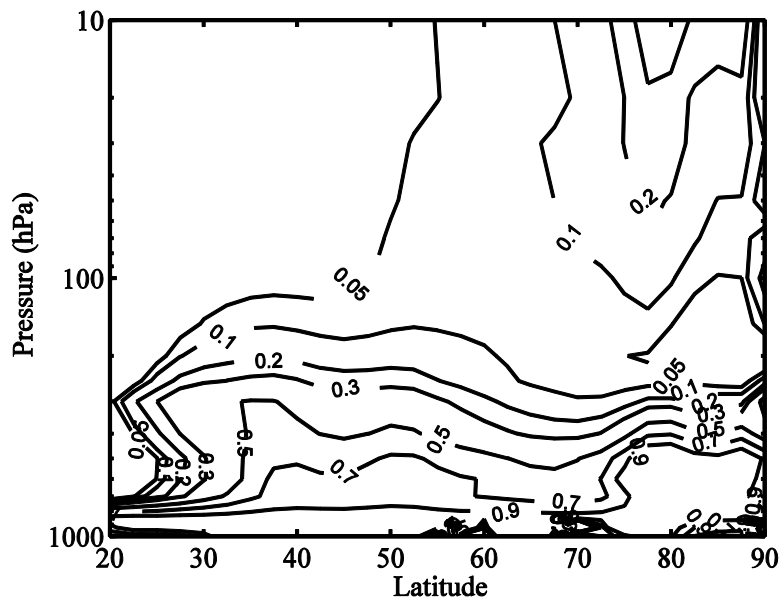
**Figure 3.** Climatology of refractive index squared ( $a^2 n_{k,l}^2(y, z)$ ) of 50 winters (1961–2010) in the Southern Hemisphere. Regions with negative  $a^2 n_{k,l}^2(y, z)$  are shaded with gray color.

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**Figure 4.** Probability of positive refractive index squared for Northern Hemisphere wintertime for wave (1, 1), (1, 2) and (1, 3).

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**Figure 5.** Climatology of vertical component of EP flux normalized by vertical component of EP flux at 850 hPa for DJF at Northern Hemisphere. Discontinuity of this quantity at the tropopause heights indicates the strong suppression of wave penetration from troposphere into the stratosphere at lower stratosphere.

32312

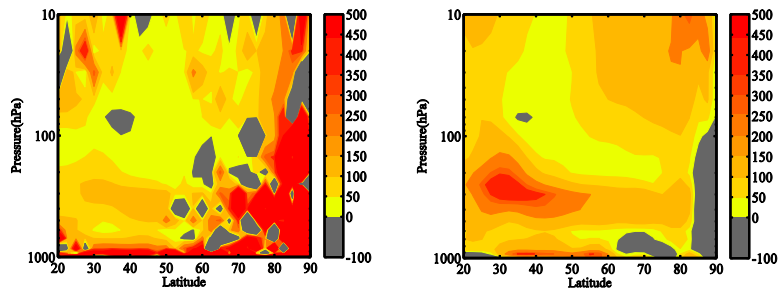


Figure 6. Climatology of  $a^2 \frac{q_\phi}{u}$  (left) and  $a^2 \frac{q_\phi}{10}$  (right) for DJF in the Northern Hemisphere.

32313

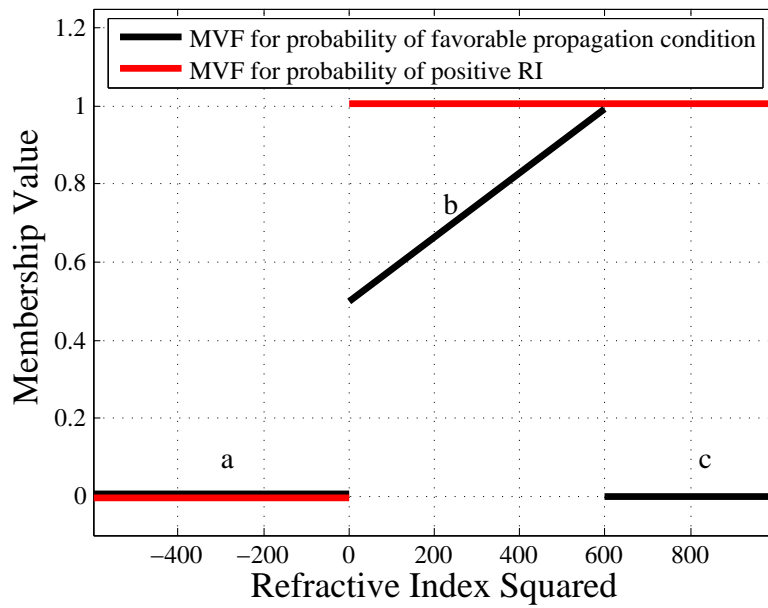
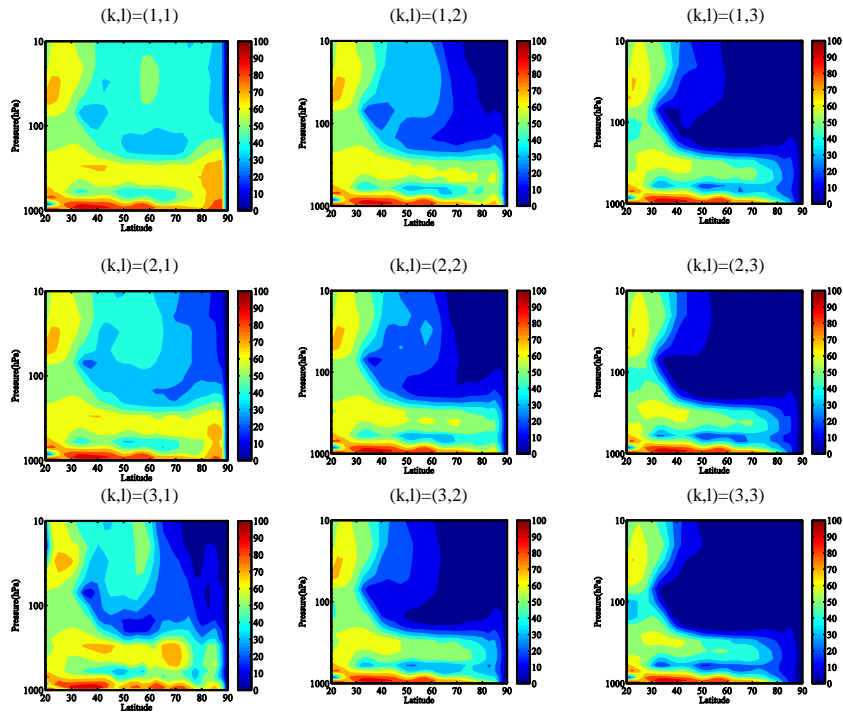


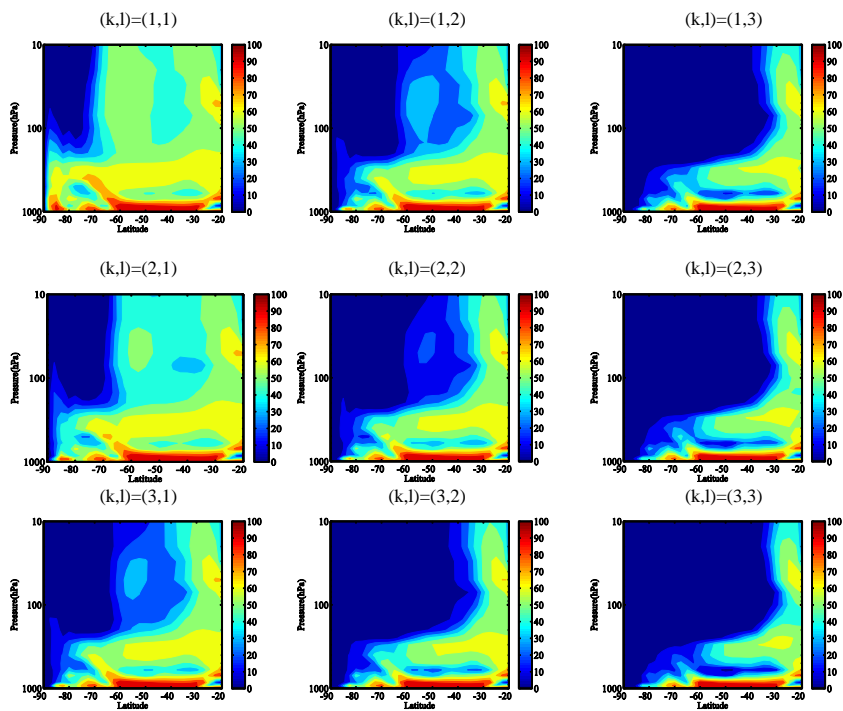
Figure 7. MVF used in the calculation of favorable propagation condition of Rossby waves (black curve). Red lines show MVF for calculating probability of positive refractive index which are used by Li et al. (2007). In their study the effect of the critical layer (part c) is not considered.

32314



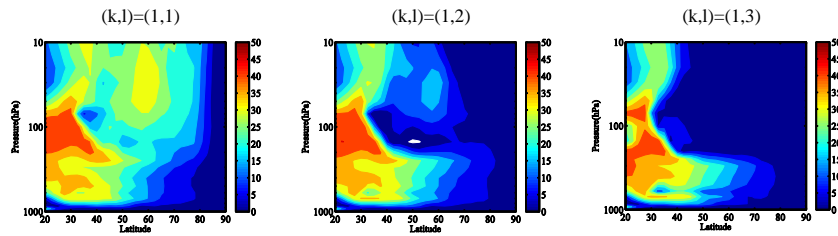
**Figure 8.** Probability of favorable propagation condition for Rossby waves derived from 50 winters (1961–2010) in the Northern Hemisphere. The higher the values, it is convenient for planetary waves to propagate to that regions. In contrast, planetary waves tend to propagate away from regions of low values of this quantity.

32315



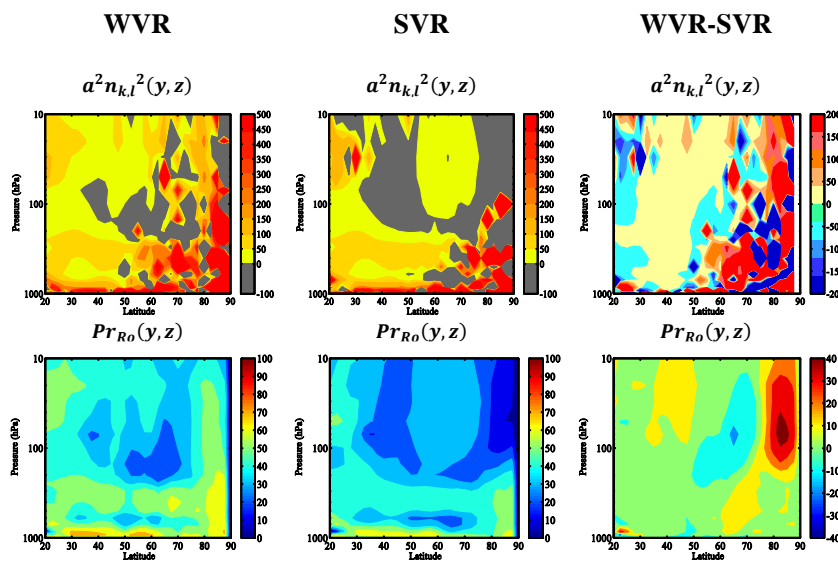
**Figure 9.** The same as Fig. 8 but for Southern Hemisphere wintertime.

32316



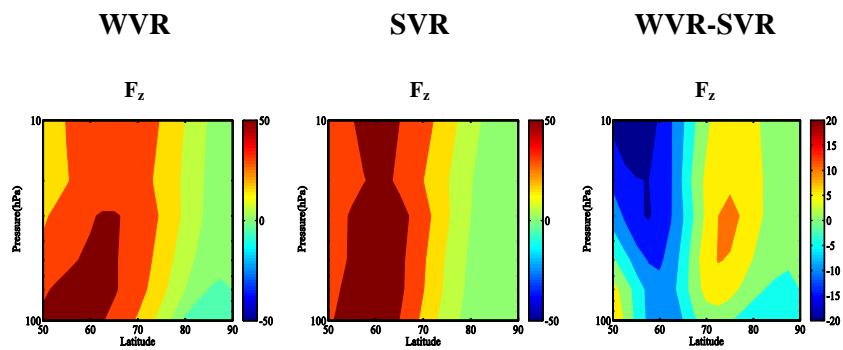
**Figure 10.** The differences between the probability of positive refractive index squared and the probability of favorable propagation condition of stationary Rossby waves.

32317



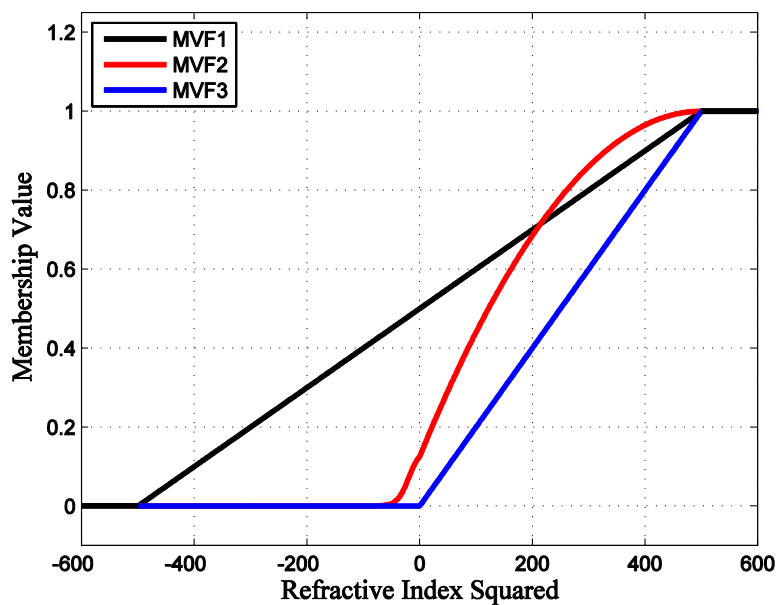
**Figure 11.**  $a^2 n_{k,l}^2(y, z)$  (first row) and  $Pr_{Ro}(y, z)$  (second row) during WVR and SVR.

32318



**Figure 12.** Same as Fig. 11 but for the vertical component of EP flux. The values are divided by  $10^5$ . Since the highest differences in the  $n_{k,l}^2(y, z)$  and  $P_{R_0}(y, z)$  between WVR and SVR are in the high latitude stratosphere the vertical component of EP fluxes are shown in this region.

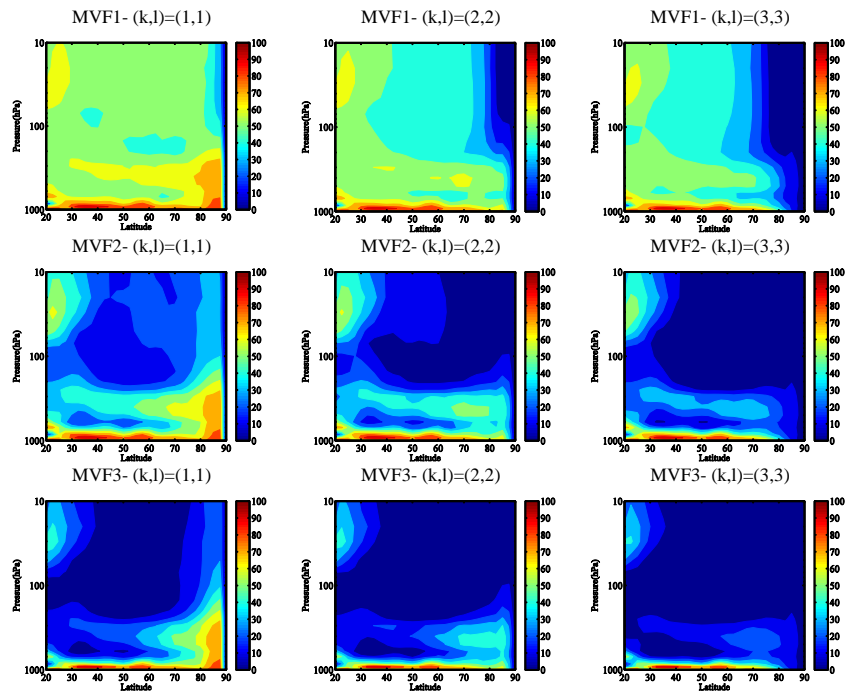
32319



**Figure 13.** Shape of three MVFs that are used to calculate the values of  $P_{R_0}(y, z)$ .

32320





**Figure 14.** Probability of favorable propagation condition for Rossby waves derived from 50 winters (1961–2010) in the Northern Hemisphere based on different MVF values described in Fig. 13.