



An important
mechanism
sustaining the
atmospheric “water
tower”

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An important mechanism sustaining the atmospheric “water tower” over the Tibetan Plateau

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Received: 27 April 2014 – Accepted: 22 June 2014 – Published: 10 July 2014

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Published by Copernicus Publications on behalf of the European Geosciences Union.

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Abstract

The Tibetan Plateau (TP), referred to as the “roof of the world” is also known as the “world water tower”, because it contains a large amount of water resources and ceaselessly transports these waters to its surrounding areas. However, it is not clear how these waters are being supplied and replenished. In particular, how plausible hydrological cycles can be realized between tropical oceans and the TP. In order to explore the mechanism sustaining the atmospheric “water tower” over the TP, the relationship of a “heat source column” over the plateau and moist flows in the Asian summer monsoon circulation is investigated, here we show that the plateau’s thermal structure leads to dynamic processes with an integration of two couples of lower convergences and upper divergences, respectively, over the plateau’s southern slopes and main platform, which relay moist air in two ladders up to the plateau. Similarly to the CISK (Conditional Instability of the Second Kind) mechanism of tropical cyclones, the elevated warm-moist air, in turn, forces convective weather systems, hence building a water cycle over the plateau. An integration of mechanical and thermal TP-forcing is revealed in relation to the Asian summer monsoon circulation knitting a close tie of vapor transport from tropical oceans to the atmospheric “water tower” over the TP.

1 Introduction

It has long been known that the Tibetan Plateau (TP) as the third pole and “the world water tower”(Xu et al., 2008; Qiu, 2008) plays an important and special role in global climate and energy/water cycle. In particular, due to its elevated land surface and thus enhanced sensible heating, the TP becomes a unique heat source, nonexistent in any other part of the world (Flohn, 1957; Yeh et al., 1957; Yanai et al., 1992; Webster et al., 1998; Wu and Zhang, 1998; An et al., 2001; Sugimoto and Ueno, 2010). From its topographic structure, we know that the TP possesses steep slopes with dramatic rising of land surfaces on its south and east rims. Over the plateau, however, the TP extends

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into north and west extensively in a relatively flat fashion, thus being presented as an oversized “mesa”, although there are large mountains over the TP triggering convective cloud formations. In the boreal summer, this massive “mesa” is strongly heated by solar radiation. One of the consequences of this thermal structure is its virtual functionality serving as an “air pump”, which attracts warm and moist air from low-latitude oceans up to the north into the Asian continent (Wu et al., 1997, 2012). During boreal winter, this flow pattern reverses with the TP’s cooling source (Ding, 1994). Hence, the TP’s role in the world’s largest monsoon system is explained.

Furthermore, classic studies (Flohn, 1957; Yeh et al., 1957; Luo and Yanai, 1984; Wu and Zhang, 1998; Yanai et al., 1992; Hahn and Manabe, 1975; Webster et al., 1998; Xu et al., 2010; Ye and Gao, 1979) also indicate that the rising warm and moist air from the tropical oceans tends to be deflected predominantly to the right (carried along the midlatitude westerlies), once encountered with the sharply elevated plateau. The deflected warm and moist air forms the well-known “south-westerly monsoonal flows”, transporting water vapor down to the southeastern China, plausibly explaining the abundant water resources in these areas (Xu et al., 2010, 2012; Zhao and Chen, 2001) (see the small rectangle in the low reach of Yangtze River Basin in upper panel of Fig. 1).

However, many environment resource surveys (Lu et al., 2005; Yao et al., 2012; Qiu, 2008) confirm that the TP itself contains a large amount of water resources, in the forms of snowpacks, glaciers, lakes, rivers, and aquifers (the large rectangle over the TP in upper panel of Fig. 1). The TP region contains one of the richest water resources and constitutes one of the densest hydrological systems in the world. Xu et al. (2008) identified the role of TP as the world water tower, and elucidated how a hydrological cycle is completed over the plateau and its surrounding areas and how atmosphere is able to supplement and reinforce the water that has been continuously transported away from the TP. These studies certainly indicate that despite the fact that a large amount of water vapor is deflected to the southeast China, there must be appreciable

correlation coefficients of the TP-heat source column to U-, V- and W-components of vector for wind and vapor transport flux.

3 Results

3.1 Elevated heat and wet islands over the TP

5 The upper panel in Fig. 2, respectively, depicts the vertical distribution in zonal differences of air temperature and specific humidity around the TP, and these differences are calculated respectively by subtracting the zonal means of air temperature and specific humidity in the Northern Hemisphere from their regional values for summer (June, July and August) averaged over 2000–2009. A “warm-wet island” elevated in the middle troposphere over the TP is identified from the positive differences of air temperature and humidity over the TP (upper panel of Fig. 2). This heat island over the massive TP exceeds that of any urban agglomerations in the world in both intensity and area.

10 A high total solar irradiance of 1688 W m^{-2} , 23 % higher than the solar constant was observed over the TP (Lu et al., 1995), as the plateau absorbs a large proportion of solar radiation. Because the TP is a high solar radiation region with the value exceeding the solar constant in the world, air temperatures over the TP could be $4 \sim 6^\circ\text{C}$ and even up to 10°C higher than its surrounding atmosphere at the same altitude in summer (Yeh and Chen, 1992). The high solar radiation on the TP could result in a strong sensible heat exchange in the surface layers. A good positive correlation between surface air temperature and vertical velocity at 500 hPa over the TP (lower panel of Fig. 2) reflects an important role of the surface sensible heating and its vertical transfer in building the heat and wet islands over the TP. The surface heating from the Plateau could trigger the air ascent for the vertical water vapor transport up to the free troposphere. Even if the surface heat fluxes from the plateau have a negligible impact on the South Asian summer monsoon circulation strength (Boos and Kuang, 2010), they could greatly impact the convective precipitation over the TP. As shown in the upper

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panel of Fig. 2 for the vertical structures of the elevated heat and wet islands, a heat source column reaching the upper troposphere over the TP could be visualized from the distribution of positive temperature differences with two high cores, respectively, within near-surface layers and between 200 and 400 hPa (upper panel of Fig. 2). Due to a monotonic decrease in surface sensible heating with the increasing elevation, the “hollow heat island” with a warm core at 200–400 hPa could be dominated by the latent heating released from the convective cloud and precipitation processes over the TP in association with the vertical structure of air vapor in the wet island over the TP (upper panel of Fig. 2).

The elevated land surface with a strong radiative heating could make the massive TP “mesa” more favorable for initiating a large number of convective cells. These convective cells over the plateau often give rise to precipitation over the TP and its surroundings in the boreal summer (Xu et al., 2012; Sugimoto and Ueno, 2010). In fact, the annual occurrences of convective clouds (cumulonimbus) over the TP are observed with 2.5 times of the regional mean over the other areas of China (Xu et al., 2002), and the TP region is regarded as a high frequency center of cumulonimbus or mesoscale convective systems (MCSs) in China (Sugimoto and Ueno, 2012), which is also confirmed by the satellite-observed convective clouds over the TP (Fig. 6) in the plateau low vortex region (upper panel of Fig. 4).

3.2 Processes of water vapor transport upward the TP

Based on the differences of temperature and humidity at a given pressure level of the atmosphere over the TP and over adjacent non-elevated areas in boreal summer, the vertical structures of heat source column and wet island on the TP are characterized in Fig. 2 (upper panel) with the particularly surprising “hollow heat island” between 200 and 400 hPa in the shape of “warm core” and “mushroom cloud” (high zonal air temperature deviation) over the TP. The vertical structure of the elevated wet island over the TP can also confirm that the TP prevents dry and cool extratropical air from “ventilating” the moist and warm tropics and subtropics (upper panel of Fig. 2). It is

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particularly interesting that the TP “hollow heat island” structure is similar to the warm core of Typhoon-CISK (Conditional Instability of the Second Kind) process (Charney and Eliassen, 1964; Smith, 1997) in the company of the elevated wet island (upper panel of Fig. 2) and the meridional circulation with strong convections (left upper panel of Fig. 3). The “CISK-like process” relaying warm-moist air up to the TP in two ladders is identified between two couples of tropospheric lower convergences (LC) and upper divergences (UD) corresponding to (1) the LC in the South Asian monsoon regions and the UD over the southern TP-slopes as well as (2) the LC on the TP main platform and the UD in the middle and upper troposphere over herein (left upper panel of Fig. 3).

The strength of “heat source column Q1” could be represented by the atmosphere column integration of apparent heat source Q_1 over the TP-region. The middle panel of Fig. 3 presents the correlations of the TP heat source column strength Q1 over the TP, respectively, to the divergences in the contours as well as to the W- and V-wind components in the correlation vectors at the vertical sections around the TP averaged in July of 2000–2009, indicating that the air ascent motions induced by the TP heating are profound over the TP during the summer monsoon period. The large topography of TP with the “hollow heat island” can force a water vapor pump with the strong upward air flows. A meridional circulation produced by the thermal effect of “hollow heat island” and the mechanical impact of the TP-topography can not only result in the Asian summer monsoon circulations and also enhance the water vapor transport from the oceans crossing the Asian monsoon areas up to the TP (see two dotted rectangles in middle panel of Fig. 3). The strong divergences of the South Asian High in the upper troposphere are collocated with the near-surface convergences associated with the plateau low vortex, which is a favorable pattern for vertical circulation enforcing a strong water vapor uplift over the TP (left upper and middle panels of Fig. 3; upper panel of Fig. 4). The TP surface sensible heat and the latent heat release from the convective cloud and precipitation may maintain the vertical circulation driving the vapor transport up into the atmospheric “water tower” over the TP (lower panel of Fig. 2; Figs. 3–5). A water vapor pump with cloud convective activities is motivated in the near-surface air convergences

over the TP, driven by the plateau heating (upper panel of Fig. 4; Fig. 6). The atmospheric “water tower” is set up by the air pump forced with the TP heating (Xu et al., 2008).

A coupling of two “dynamic pumps” with the CISK-like mechanism, contiguous horizontally but staggered vertically, are revealed with the cooperative interaction of the “heat source column” and the elevated wet islands over the roof of the world (middle panel of Fig. 3). This interaction could be achieved with a positive feedback, when the forcing effect of the “heat source column” drives the water vapor flows climbing up the TP in the vertical motion, in turn, and the phase changes of water vapor to clouds and precipitation in the moist convection release latent heating intensifying the “heat source column” and especially the “warm core” in the upper troposphere associated with the South Asian High (Sugimoto and Ueno, 2012). The “heat source column” could enhance convergences at lower levels and divergences at upper levels in the troposphere for pushing the moist air up the TP (middle panel of Fig. 3; Fig. 4). There could be a mutual feedback between the UD on the southern plateau slopes and the LC on the TP-platform through the dynamical interaction of the horizontally contiguous UD and LC (right upper and middle panels of Fig. 3). The UD over the southern TP-slopes and the LC on the TP-platform could be contributed by the water vapor flow acceleration at the inflection point between the southern slopes and the platform with the mechanical TP-impact on the LC and the air pump on the platform (upper and middle panels of Fig. 3).

The two ladders of “CISK-like process” over the South Asian summer monsoon region and the TP-platform knit a close tie of vapor transport from tropical oceans to the atmospheric “water tower” over the TP (Fig. 3). The South Asian summer monsoon precipitation is produced in the first ladder of air vapor transport toward the TP atmosphere, which could be attributed to the TP-topographical block at the steep southern slopes with less thermal impact (Boos and Kuang, 2010). The second ladder resulting in convective cloud precipitation over the large TP-platform with less terrain obstacles for water vapor flows is dominantly controlled by thermal forcing of the “hollow heat

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convective clouds bring the precipitation over the TP and its surroundings (Xu et al., 2012; Sugimoto and Ueno, 2010). To further clarify the atmospheric “water tower” over the TP in the Asian water cycles, Fig. 6 presents the spatial distribution of total cloud cover over the TP and its surrounding area in July 2008.

During the Asian summer monsoon period, the dense cloud covers exist over the regions from the Bengal Bay, Asian monsoon regions to the southern TP (Fig. 6). As characterized with the correction vectors of the column heat source over the TP to the moisture transport over and around the TP in middle panel of Fig. 3, two convergence zones of moisture transport fluxes ($\nabla \cdot qV < 0$) are found in two ladders over the plateau’s southern slopes and main platform during the moisture transport from the oceans up to the TP, resulting in these regions of cloud covers in Fig. 6. It is noteworthy that the high cloud amounts are zonally concentrated between the steep southern plateau edges and the shear line of the plateau low vortex over the TP (upper panel of Fig. 4; Fig. 6) with the monthly mean cloud cover fractions up to 90 %, which could be formed from the “CISK-like mechanism” for building the TP’s atmospheric “water tower” (Fig. 3). Over the large TP platform with relatively plain terrain, the monthly mean cloud covers of around 45 % are mostly observed on the central-eastern region with the less cloud covers over the northwestern TP, depending on the moisture transport across the TP. The plateau low vortex over the TP and the southward air flows with less moisture on the north of the shear line could lead to the less cloud covers in the northwestern platform of TP (upper panel of Fig. 4).

The observed cloud distribution over the TP confirms that the “CISK-like mechanism” is an important mechanism sustaining the atmospheric “water tower” over the TP. Connecting with the cloud and precipitation in the atmospheric “water tower”, the plausible hydrological cycles could be realized between tropical oceans and the TP.

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The present analyses clearly indicate that the TP presents itself as a “warm-wet island”. The surface heating over the plateau leads to a low-pressure center, causing flow convergence at low levels of the plateau and triggers vertical motion subsequently. This convective system will result in plateau clouds and precipitation, which would explain abundant water storage over the TP and its surrounding regions.

The classic Asian summer monsoon theory elucidated an “air pump” mechanism in relation to the TP. The warm-moist air from the low-latitude oceans is drawn toward the plateau by this air pump. The analyse on relationship of the “heat source column” over the TP and warm-moist air transport in the present study further suggests a CISK-like mechanism on water vapor suction up the plateau. An appreciable portion of warm-moist air converges at the foot of the south rim of the plateau. The convergence of the warm-moist air ascends along the plateau’s slope and diverges at about the altitude of the plateau top. This divergence flow enforces the convergence at the heated low-pressure center over the TP and feeds in the convective system with warm-moist air, which results in the clouds and precipitations for the atmospheric water tower over the TP.

These dynamic and thermodynamic processes depict a coupling of two CISK type systems, both with convergence at low levels and divergence at upper levels, but the systems are horizontally contiguous as well as vertically staggered. The two systems display a mutually supportive mechanism (Fig. 7). It is this coupling that ladders the moist air up to the plateau, which may an important mechanism sustaining the atmospheric “water tower” over the TP.

In this study, the mean climate of air vapor transport to the TP is only investigated based on the summertime averages over the past years. The two ladder “CISK-like mechanism” is identified as a key process sustaining the atmospheric “water tower” over the TP, which can be used to further study the inter-annual variability of the atmospheric heat source and moisture budget over the TP. There is an important role

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- Lu, L., Zhou, G., and Zhang, Z.: Direct and global solar radiaions in the region of Qomolangma during the summer 1992, *Acta Energiæ Solaris Sinica*, 16, 229–233, 1995 (in Chinese).
- Luo, H. and Yanai, M.: The large-scale circulation and heat sources over the Tibetan Plateau and surrounding areas during the early summer of 1979. Part II: Heat and moisture budgets, *Mon. Weather Rev.*, 112, 966–989, 1984.
- 5 Qiao, Q. and Zhang, Y.: Synoptics in Tibetan Plateau, Chinese Meteorology Press, Beijing, 15–18, 1994 (in Chinese).
- Qiu, J.: China: the third pole, *Nature*, 454, 393–396, 2008.
- Smith, R. K.: On the theory of CISK, *Q. J. Roy. Meteor. Soc.*, 123, 407–418, 1997.
- 10 Sugimoto, S. and Ueno, K.: Formation of mesoscale convective systems over the eastern Tibetan Plateau affected by plateau-scale heating contrasts, *J. Geophys. Res.*, 115, D16105, doi:10.1029/2009JD013609, 2010.
- Sugimoto, S. and Ueno, K.: Role of mesoscale convective systems developed around the Eastern Tibetan Plateau in the eastward expansion of an upper tropospheric high during the monsoon season, *J. Meteorol. Soc. Jpn.*, 90, 297–310, doi:10.2151/jmsj.2012-209, 2012.
- 15 Sugimoto, S., Ueno, K., and Sha, W. M.: Transportation of water vapor into the Tibetan Plateau in the case of a passing synoptic-scale trough, *J. Meteorol. Soc. Jpn.*, 86, 935–949, 2008.
- Webster, P. J., Magaña, V. O., Palmer, T. N., Shukla, J., Tomas, R. A., Yanai, M., and Yasunari, T.: Monsoons: processes, predictability, and the prospects for prediction, *J. Geophys. Res.*, 103, 14451–14414, doi:10.1029/97JC02719, 1998.
- 20 Wu, G. X. and Zhang, Y.-S.: Tibetan Plateau forcing and the timing of the monsoon onset over South Asia and the South China Sea, *Mon. Weather Rev.*, 126, 913–927, 1998.
- Wu, G. X., Li, W. P., Guo, H., Liu, H., Xue, J. S., and Wang, Z. Z.: Sensible heat driven air-pump over the Tibetan Plateau and its impacts on the Asian Summer Monsoon. Collections on the Memory of Zhao Jiuzhang, edited by Ye, D. Z., Chinese Science Press, Beijing, 116–126 (in Chinese), 1997.
- 25 Wu, G. X., Liu, Y., He, B., Q. Bao, Duan, A., and Jin, F.-F.: Thermal controls on the Asian summer monsoon, *Scientific Reports*, 2, 404, doi:10.1038/srep00404, 2012.
- Xu, X., Bian, L., Li, S., Ding, Y., Zhou, M., and Chen, J.: A comprehensive physical pattern of land-air dynamics and thermal structure on the Qinghai-Xizang Plateau, *Sci. China Ser. D*, 45, 577–594, 2002.
- 30 Xu, X., Lu, C., Shi, X., and Gao, S.: World water tower: an atmospheric perspective, *Geophys. Res. Lett.*, 35, L20815, doi:10.1029/2008GL035867, 2008.

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Xu, X., Shi, X., and Lu, C.: Theory and Application for Warning and Prediction of Disastrous Weather Downstream from the Tibetan Plateau, Novinka Science Publishers, Inc., New York, 2012.

Yanai, M. and Johnson, R. H.: Impacts of cumulus convection on thermodynamic fields, in: *The Representation of Cumulus Convection in Numerical Models of the Atmosphere*, edited by: Emanuel, K. A. and Raymond, D. J., AMS Monograph, 24, 39–62, 1993.

Yanai, M. and Tomita, T.: Seasonal and interannual variability of atmospheric heat sources and moisture sinks as determined from NCEP–NCAR reanalysis, *J. Climate*, 11, 463–482, 1998.

Yanai, M., Li, C. F., and Song, Z. S.: Seasonal heating of the Tibetan Plateau and its effects on the evolution of the Asian summer monsoon, *J. Meteorol. Soc. Jpn.*, 70, 319–351, 1992.

Yao, T., Thompson, L., Yang, W., Yu, W., Gao, Y., Guo, X., Yang, X., Duan, K., Zhao, H., Xu, B., Pu, J., Lu, A., Xiang, Y., Kattel, D. B., and Joswiak, D.: Different glacier status with atmospheric circulations in Tibetan Plateau and surroundings, *Nature Clim. Change*, 2, 663–667, 2012.

Ye, D. Z. and Gao, Y. X.: *Meteorology of the Qinghai-Xizang Plateau*, Chinese Science Press, Beijing, 1979.

Yeh, T. C. and Chen, B.: *Global Change Research in China (Part 2)*, Chinese Earthquake Press, Beijing, 9–111, 1992 (in Chinese).

Yeh, T. C., Luo, S. W., and Chu, P. C.: The wind structure and heat balance in the lower troposphere over Tibetan Plateau and its surrounding, *Acta Meteorol. Sin.*, 28, 108–121, 1957.

Zhao, P. and Chen, L.: Climatic features of atmospheric heat source/sink over the Qinghai-Xizang Plateau in 35 years and its relation to rainfall in China, *Sci. China Ser. D*, 44, 858–864, 2001.

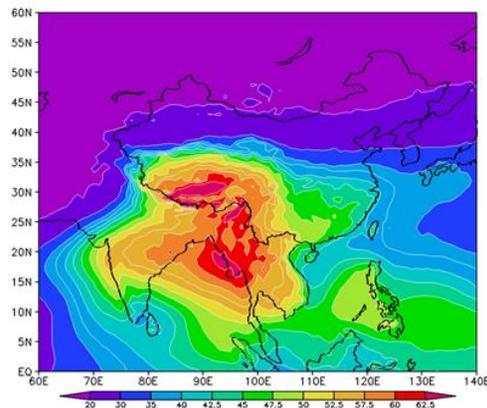
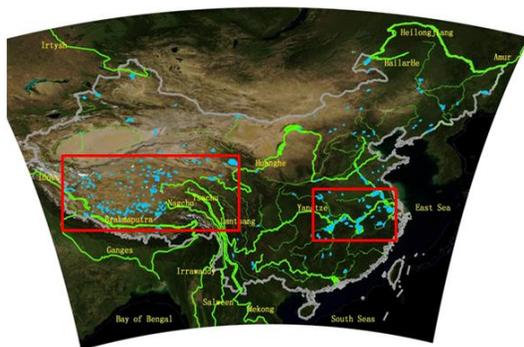


Figure 1. Geographical distribution of water sources in glaciers, snowpacks, rivers and lakes over China with white, yellow and light blue colors respectively. Two major lake groups are marked by two red rectangles in the TP and Eastern China (upper panel). Column vapor content ($10^{-2} \text{ g cm}^{-2}$) over 500 hPa in summer averaged over 2000–2009 (lower panel).

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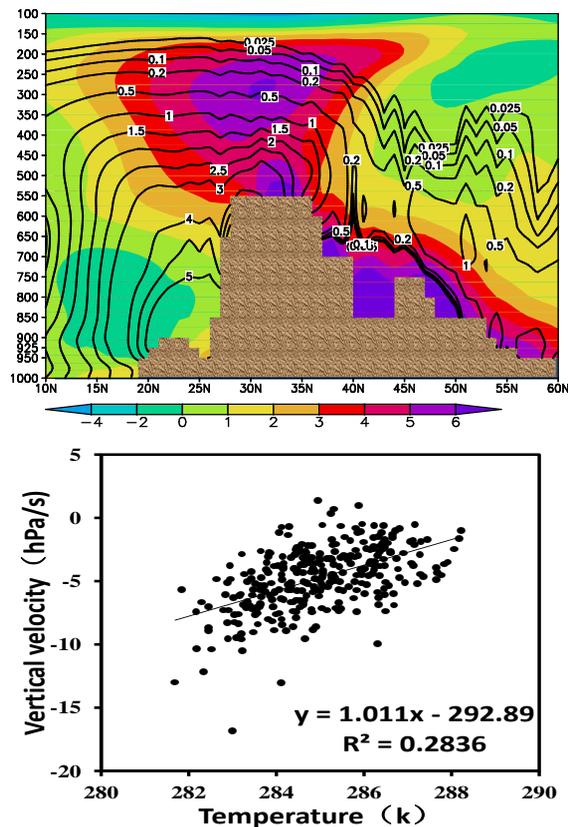


Figure 2. Vertical sections of the temperature ($^{\circ}\text{C}$; filled contours) and specific humidity (g kg^{-1} ; contour lines) differences relative to the zonal means along 93–94° E in summer averaged over 2000–2009. The plateau section is marked with soil color (upper panel). A scatter plot of surface air temperature and vertical velocity at 500 hPa in the TP region in July of 2000–2009 (lower panel).

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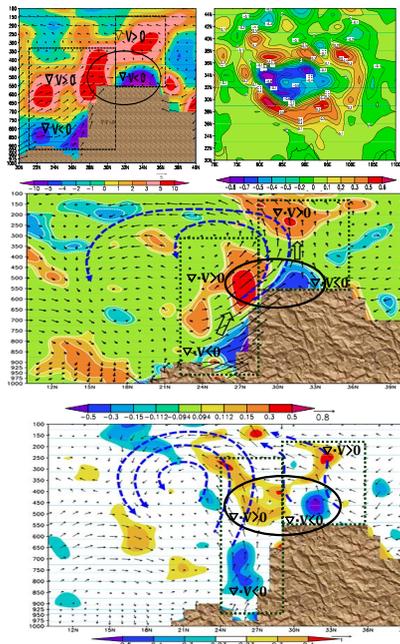


Figure 3. Vertical sections of wind vectors and divergences (filled contours) for summer averaged over 2000–2009 along 93–94° E (left upper panel); distribution of summertime 500 hPa divergence averaged over 2000–2009 (right upper panel). Vertical sections of the daily correlations of TP heat source column Q_1 to the divergences (filled contours) and the correction vectors of Q_1 to V- and W-wind components in July of 2000–2009 along 93–94° E with the meridional circulations and the uplifting vapor transport denoted by blue dash lines and black arrows, respectively (middle panel). Vertical sections of the lag-correlations of TP heat source column Q_1 at 10 prior days to divergences and the lag-correlation vectors in the meridional circulations in July of 2000–2009 along 93–94° E (lower panel). In all panels, two couples of lower convergences (LC) and upper divergences (UD) are denoted with $\nabla \cdot V < 0$ and $\nabla \cdot V > 0$ in two dotted rectangles and the interaction of LC in the TP and UD over the southern slopes in the black ovals. The plateau section is marked with soil color.

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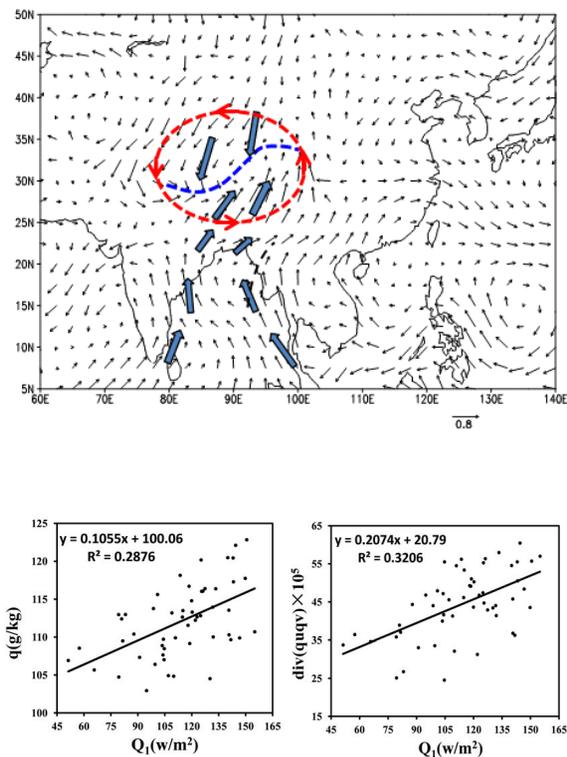


Figure 4. Correlation vectors of the TP heat source column strength Q_1 to the horizontal moisture transport flux components over June, July and August of 2000–2009. A shear line between southward and northward air flows (light blue arrows) and the plateau low vortex over the TP are marked with dash lines in blue and red (upper panel). Correlations of the heat source strength Q_1 , total water vapor q and net vapor transport flux divergence (div) in the TP air column in summer of 2000–2009 in scatter plots of Q_1 to q (left lower panel) and div (right lower panel).

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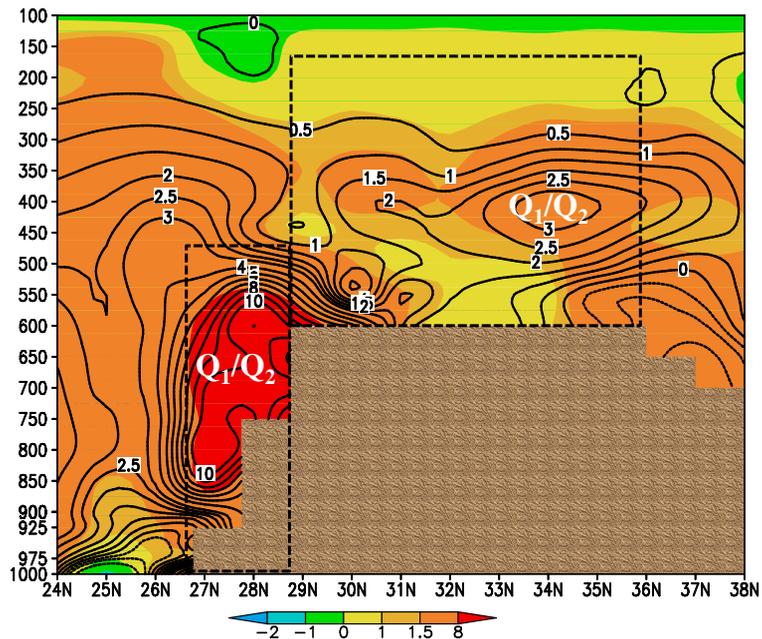


Figure 5. The vertical distributions of apparent heat source Q_1 (filled contours) and apparent moisture sink Q_2 (w m^{-2}) averaged between 85°E and 100°E in summer over 2000–2009. The Q_1/Q_2 in two dash rectangles are produced with two ladders of CISK-like process respectively over the TP’s southern slopes and main platform. The plateau section is marked with soil color.

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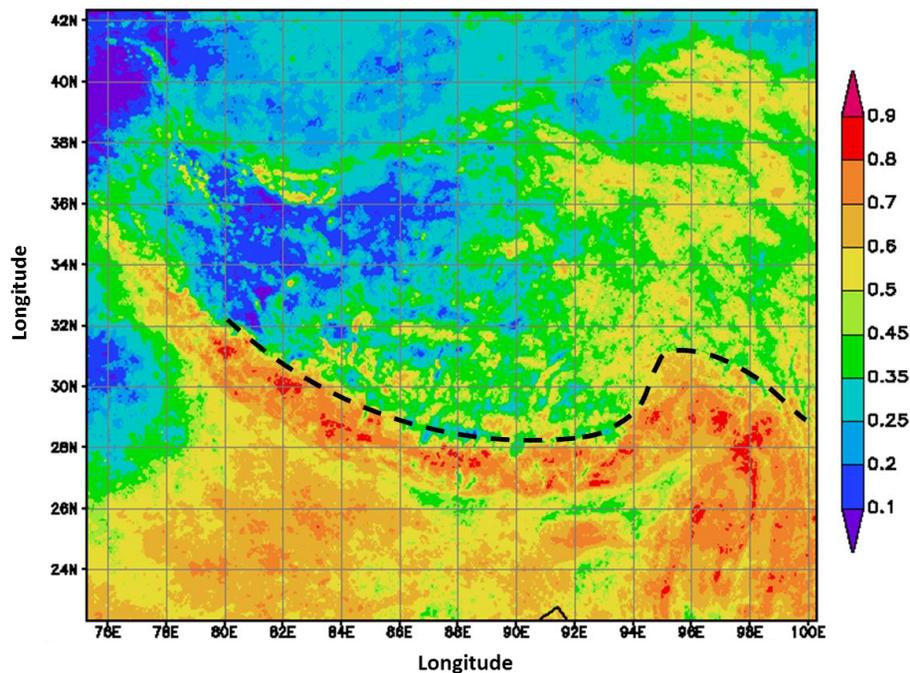


Figure 6. Horizontal distribution of cloud cover fraction in July 2008 derived from the Chinese meteorological satellite FY-2F. The black dash line separates the high and low amounts of cloud cover over the TP.

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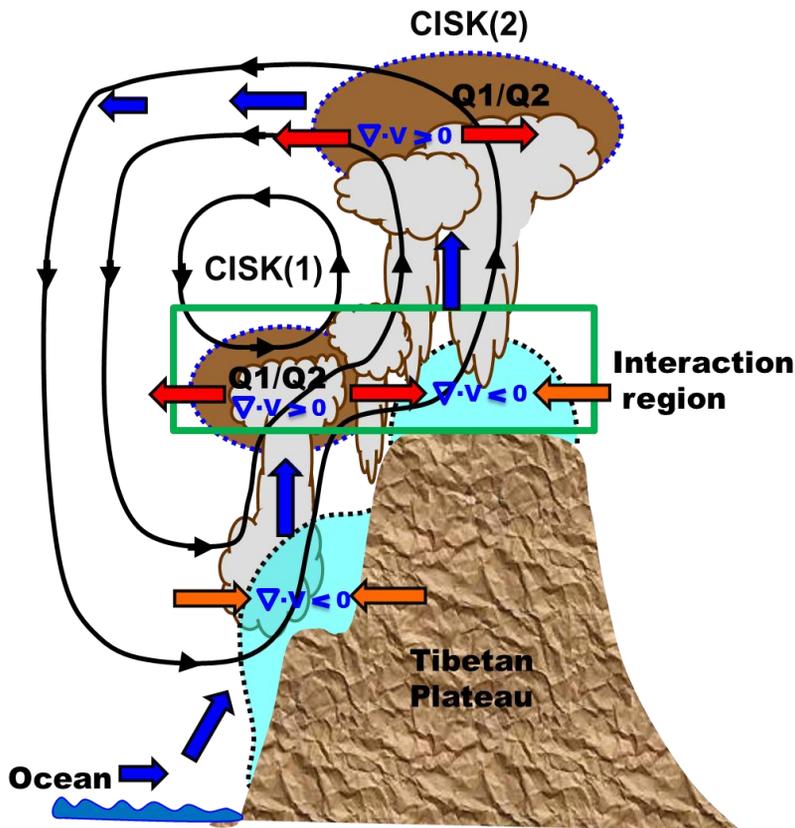


Figure 7. A diagram of the summary on two ladders of CISK-like processes with two couples of heat source Q_1 and moisture sink Q_2 over the TP’s southern slopes and main platform in forcing water vapor flows climbing up the TP, which is marked with soil color.