

**PTR-TOF-MS
measurements and
MEGAN 2.1 model
results**

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Undisturbed and disturbed above canopy ponderosa pine emissions: PTR-TOF-MS measurements and MEGAN 2.1 model results

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Abstract

We present the first eddy covariance flux measurements of volatile organic compounds (VOCs) using a proton-transfer-reaction time-of-flight mass-spectrometer (PTR-TOF-MS) above a ponderosa pine forest in Colorado, USA. The high mass resolution of the PTR-TOF-MS enabled the identification of chemical sum formulas. During a 30 day measurement period in August and September 2010, 649 different ion mass peaks were detected in the ambient air mass spectrum (including primary ions and mass calibration compounds). Eddy covariance with the vertical wind speed was calculated for all ion mass peaks. On a typical day, 17 ion mass peaks including protonated parent compounds, their fragments and isotopes as well as VOC-H⁺-water clusters showed a significant flux with daytime average emissions above a reliable flux threshold of 0.1 mg compound m⁻² h⁻¹. These ion mass peaks could be assigned to seven compound classes. The main flux contributions during daytime (10:00–18:00 LT) are attributed to the sum of 2-methyl-3-buten-2-ol (MBO) and isoprene (50%), methanol (12%), the sum of acetic acid and glycolaldehyde (10%) and the sum of monoterpenes (10%). The total MBO + isoprene flux was composed of 10% isoprene and 90% MBO.

There was good agreement between the light and temperature dependency of the sum of MBO and isoprene observed for this work and those of earlier studies. The above canopy flux measurements of the sum of MBO and isoprene and the sum of monoterpenes were compared to emissions calculated using the Model of Emissions of Gases and Aerosols from Nature (MEGAN 2.1). The best agreement between MEGAN 2.1 and measurements was reached using emission factors determined from site specific leaf cuvette measurements. While the modelled and measured MBO + isoprene fluxes agree well the emissions of the sum of monoterpenes is underestimated by MEGAN 2.1. This is expected as some factors impacting monoterpene emissions, such as physical damage of needles and branches due to storms, are not included in MEGAN 2.1.

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After a severe hailstorm event, 22 ion mass peaks (attributed to six compound classes plus some unknown compounds) showed an elevated flux for the two following days. The sum of monoterpene emissions was 4–23 times higher compared to emissions prior to the hailstorm while MBO emissions remained unchanged. If one heavy storm occurs at this site every month we calculate that the monthly monoterpene emissions (in mg compound m⁻²) would be underestimated by 40 % if this disturbance source is not considered.

1 Introduction

A great variety of Volatile Organic Compounds (VOCs) is continuously emitted into the atmosphere by pyrogenic, anthropogenic and biogenic sources. Due to their central role in tropospheric ozone chemistry and their ability to form aerosols, they can influence local and global climate (Koppmann, 2007). Globally, total biogenic VOC sources are considered to be approximately ten times larger than the sum of anthropogenic emissions (Muller, 1992; Olivier et al., 1999; Guenther, 2002; Lamarque et al., 2010). While terpenoids such as hemiterpenes, monoterpenes and sesquiterpenes are an important fraction of biogenic VOCs (BVOCs), other compounds such as oxygenated compounds (methanol, acetone, acetaldehyde) also exhibit significant emissions from the biosphere (Fowler et al., 2009 and references therein). In the last decade, BVOC emissions from the biosphere have been measured using proton-transfer-reaction mass-spectrometer (PTR-MS) instruments, which allow monitoring a selected set of ion mass peaks that are related to different compounds based on disjunct eddy covariance flux measurements (e.g., Karl et al., 2002). The recent development of a proton-transfer-reaction time-of-flight mass-spectrometer (PTR-TOF-MS) (Jordan et al., 2009; Graus et al., 2010) enables measurements of the entire mass spectrum within a fraction of a second, allowing 10 Hz eddy covariance measurements over a wide mass range. Additionally, the high mass resolution of the time-of-flight mass spectrometer allows for separation of isobaric compounds. So far this instrument has been tested

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for flux measurements above grassland (Müller et al., 2010; Ruuskanen et al., 2011; Bamberger et al., 2011) and above a citrus plantation (Park et al., 2013). Ruuskanen et al. (2011) and Park et al. (2013) report fluxes (emissions and depositions) for up to 30 (27) different ion mass peaks corresponding to a dozen different VOC species. Müller et al. (2010) and Park et al. (2013) evaluated eddy covariance measurements by PTR-TOF-MS by comparing with the well-established disjunct eddy covariance technique using PTR-MS. Both studies showed good agreement between the two techniques.

Ponderosa pine forest BVOC fluxes of monoterpenes, 2-methyl-3-buten-2-ol (MBO), methanol, ethanol, acetaldehyde and acetone have been previously measured at a site in the California Sierra Nevadas (Schade et al., 1999; Schade and Goldstein, 2001; Lee et al., 2005). The magnitude and type of compounds emitted by plants is complex depending on the species, the age and the health of the vegetation. Effects of short time variations in temperature, moisture and light levels (Guenther et al., 1995; Kesselmeier and Staudt, 1999; Niinemets et al., 2010a) are better understood. Studies have shown that temperature is the main driver for monoterpene emissions, while both temperature and light are important for MBO emissions. These studies took place in ponderosa pine forests using above canopy relaxed eddy accumulation flux measurements, (Schade and Goldstein, 2001) and leaf level measurements (Harley et al., 1998). These observations are typically used to develop emission models such as the Model of Emissions of Gases and Aerosols from Nature (MEGAN), which uses leaf area index (LAI), light, temperature, humidity, wind speed, soil moisture and average canopy environmental conditions of the past 24 to 240 h to calculate emissions of VOCs from ecosystems to the atmosphere (Guenther et al., 2006, 2012).

Additionally, abiotic and biotic stress factors such as drought, herbivore infestation, severe storms, mechanical wounding, air pollutants, etc. can initially enhance VOC emissions of plants although chronic stress may eventually reduce emissions (Niinemets, 2010; Niinemets et al., 2010b; Holopainen and Gershenson, 2010; Duhl et al., 2013). Most emission measurements have focused on VOC emissions of intact and undamaged plants; however, studies by Schade and Goldstein (2003), Räisä-

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nen et al. (2008) and Kim et al. (2011) demonstrated enhanced monoterpene fluxes after mechanical wounding on the ecosystem scale. A few studies also reported enhanced mixing ratios of monoterpenes and sesquiterpenes after severe storms (Helmig et al., 1998; Bouvier-Braun et al., 2009; Haase et al., 2011; Bamberger et al., 2011).

5 Holzinger et al. (2006) reported increased monoterpene emissions on days following rain events. While temperature and light dependent emissions of VOCs are well understood, the prediction of emissions induced by stress factors is complicated (Grote and Niinemets, 2008). The development of emission algorithms to account for ecosystem disturbance in biogenic emission models requires the availability of representative leaf,
10 plant and ecosystem scale measurements (Arneth and Niinemets, 2010). Niinemets et al. (2010b) highlight the importance of measurements in order to reduce model uncertainties related to stress, environmental variability, seasonality and foliage developmental stage.

In this paper we present above canopy flux measurements of light and temperature driven emissions of BVOCs from a ponderosa pine forest during 30 days in August/September 2010. We compare these to emissions calculated with the MEGAN 2.1 model using default parameters and values based on land cover and enclosure measurements at the same site. We also report enhanced BVOC emissions after a severe hailstorm, which damaged needles and branches. Our canopy flux measurements
20 serve as a starting point to evaluate the magnitude of these stress driven emission events.

2 Methods

2.1 Field site

VOC eddy covariance measurements were conducted at the Manitou Forest Observatory located in the US Forest Service Manitou Experimental Forest near Woodland
25 Park, Colorado, USA (2300 m elev., lat. 39°6' 0" N, long. 105°5' 30" W) as part of the

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Bio-Hydro-Atmosphere Interactions of Energy, Aerosols, Carbon, H₂O, Organics and Nitrogen (BEACHON)-Rocky Mountain Organic Carbon Study (ROCS) campaign in summer 2010. The field site has been described in detail elsewhere (see Kim et al., 2010; DiGangi et al., 2011). An extensive overview on VOC concentration measurements at this site has been summarized by Kaser et al. (2013). Ponderosa pine trees about 100 yr in age and with an average height of 18.5 m dominate BVOC emissions within the flux footprint. The canopy is open and of varying density. The ground surface is covered by grasses, sage, forbs and exposed cryptogammic soils. The typical leaf area index (LAI) for tree canopies is 3 and the tree cover fraction is 0.38 resulting in a landscape average LAI at this site of ≈ 1.14 .

Here we present 30 days of VOC flux measurements between 3 August and 8 September 2010. VOC fluxes were not measured between 25 August–28 August and 31 August–2 September, when the PTR-TOF-MS was used for other measurements.

2.2 Instrumentation and calibration

VOC measurements with 10 Hz time resolution allowing the calculation of eddy covariance fluxes were conducted on a walk-up chemistry tower at 25.1 m height. A 35 m long unheated teflon line (OD: 3/8 in., 20 SLPM) was mounted 20 cm below the sonic anemometer. VOC volume mixing ratios were measured by a PTR-TOF-MS located at the bottom of the tower in a seacontainer laboratory. Details on the instrument, calibration and instrument performance can be found in Kaser et al. (2013) and will be mentioned here only briefly. The PTR-TOF-MS developed at the University of Innsbruck (Graus et al., 2010) was operated under standard conditions at a drift tube temperature of 60 °C, 580 V drift voltage and 2.3 mbar drift tube pressure corresponding to an E/N of 125 Td (E being the electric field strength and N the gas number density; 1 Td = 10^{-17} V cm²). The PTR-TOF-MS was calibrated weekly by dynamic dilution of VOCs using a quantitative gas standard provided by the US National Oceanic and Atmospheric Administration (NOAA, Boulder, CO, USA) containing methanol, acetonitrile, acetaldehyde, acetone, methyl vinyl ketone, limonene, MBO, pyrrole, benzene, toluene

signature of hydrometeors impacting the sensor head surface and, for solid particles such as hail, provides a total count of hailstone impacts. Also, a laser ranging sensor (Jenoptik, Inc. SHM30 laser snow depth sensor) was also deployed at the site to measure winter snowfall but also can measure hail accumulations on the ground. The sensor suggests that approximately 1.4 cm of hail accumulated on the ground during the main hail event on 4 August 2010 presented below.

2.3 Flux calculation

Eddy covariance fluxes (30 min values) were calculated for 649 ion mass peaks from the 10 Hz PTR-TOF-MS signals and the 10 Hz sonic anemometer data according to Karl et al. (2002) and Ruuskanen et al. (2011). Three dimensional wind data were recorded on a second computer. The two computer clocks drifted relative to each other during the measurement period. The time stamps of the two datasets were corrected accordingly. The three dimensional wind vector was rotated on an 5 hourly basis according to Wilczak et al. (2001). To detect fluxes among the 649 ion mass peaks, we developed a screening method based on the delay times that occur from the covariance analysis of the individual ion mass peaks and the corresponding vertical wind speed during 30 min time intervals. In order to obtain ion mass peaks that show a significant covariance between the vertical wind and their volume mixing ratio, we allowed the flux calculation routine to choose a maximum or minimum of the covariance peak within a delay time window of max. ± 20 s. The physically meaningful delay time window (0 to 10 s) was defined as the maximum difference between the two computer clocks used for wind and VOC measurements and the maximum delay time caused by the 35 m long sampling line. We used these delay times to determine ion mass peaks exhibiting a significant flux or no flux. For this purpose the fraction of half hour values where the delay time was found in the physically meaningful delay time window between 0 to 10 s was used to attribute all ion mass peaks to an individual flux grade. This fraction is plotted in Fig. 1. For example the delay time of the exact ion mass peak m/z 69.0704 ($C_5H_9^+$) which corresponds to the sum of isoprene (protonated isoprene) plus

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MBO (protonated MBO minus H₂O) was found 70 % of the time within the 0 to 10 s time window.

Ion mass peaks where the above described fraction was smaller than 35 % do not exhibit a significant covariance between vertical wind and volume mixing ratio and are depicted by red circles. These ion mass peaks exhibit no flux or fluxes that are close to or below the limit of detection for most of the time and are classified as grade C flux. Blue circles depict ion mass peaks for which a physically meaningful covariance between vertical wind and volume mixing ratio can be found for at least 35 % of all measured half hour values. We rank these as grade B fluxes. Green circles were defined as ion mass peaks where a physically meaningful delay time was found for more than 50 % of all measured half hour flux values. We rank these as grade A flux ion mass peaks. As a second cut-off criteria we defined a flux detection limit of > 0.1 mg compound m⁻² h⁻¹.

3 Results

3.1 Meteorology

Figure 2 depicts a typical diurnal cycle for temperature and photosynthetically active radiation (PAR) during the VOC flux measurement period. During the 30 days of flux measurements average air temperatures reached a maximum of 23 °C at 1 p.m. (Mountain standard time, MST) and a minimum of 9 °C at 6 a.m. (MST).

On average, PAR reaches a maximum of 1700 μmol m⁻² s⁻¹ at 1 p.m. (MST). Typically days were sunny with some cloudiness as well as rainfall during some afternoons. This is visible in the average diurnal PAR cycle, which has a slight dip in the afternoon. On 4 August the site experienced a severe hailstorm with hailstones of 1–3 cm in diameter, which caused significant damage to needles and branches. During the two days following the hail storm (Fig. 2, blue curves) temperature and PAR reached the typi-

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cally observed midday maxima of 23 °C and 2000 $\mu\text{mol m}^{-2} \text{s}^{-1}$, respectively, but both decreased rapidly in the afternoon.

Flux data during the hailstorm and the following two days are treated in Sect. 3.5. Estimates of emission factors and discussion of typical fluxes in Sect. 3.3 are based on data after 7 August when the effect of the hailstorm had decreased.

Generally the weather changed from some cloudiness and rain during the beginning of the measurement period to dry and sunny at the end of the campaign. Precipitation events with more than 2 mm water occurred only on 27 July (hail), 4 August (hail), 5 August (rain) and 9 August (rain). Maximum daytime temperatures remained similar over the measurement period but night time temperatures decreased gradually from 10 °C to 4 °C due to the approaching autumn.

3.2 Footprint and quality control

The flux footprint was calculated using the 90 % flux recovery value according to Hsieh et al. (2000). The canopy height was 18.5 m. The zero plane displacement, 12.7 m, was defined as 2/3 of the average canopy height.

For the momentum roughness height, we used 8 % of the canopy height. The ratio between z_m (measurement height above the zero plane displacement) and the Obukov length (L) was used to distinguish between stable ($z_m/L > 0$) and unstable ($z_m/L < 0$) atmospheric conditions. The calculated footprint was < 900 m during unstable conditions (45 % of the measurement time) and < 2500 m for stable conditions (55 %). Figure 3 shows the footprint in more detail for unstable atmospheric conditions. Data points are grouped along the main valley wind-direction coming from the north (positive axis: 270° to 90°) and from the south (negative axis: 90° to 270°).

Quality control for the flux data included some criteria that are mass independent (1–4) as well as mass dependent (5–6). Data were removed if (1) the wind came from east (35°–145° for unstable conditions, 10°–170° for stable conditions) to eliminate any contribution from a lightly traveled highway \approx 500 m to the east of the field site; if (2) the

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calculated footprint was larger than 1850 m (occurring during stable conditions) as the vegetation changes slowly outside this radius; if (3) 30 out of 18 000 wind data points per half hour are outside 5σ of the mean; if (4) the third rotation angle exceeded $\pm 10\%$; if (5) 30 out of 18 000 VOC volume mixing ratios per half hour are outside 5σ of the mean; if (6) the stationarity test exceeded 60 % (Foken and Wichura, 1996). Depending on the individual compound between 33–55 % of all half-hour flux values passed these quality criteria (Table 1). For MBO we calculated a daytime limit of detection of $0.15 \text{ mg m}^{-2} \text{ h}^{-1}$ and an advection flux of $0.07 \text{ mg m}^{-2} \text{ h}^{-1}$.

3.3 Light and temperature driven emissions

Ion mass peaks showing a significant correlation with the vertical wind component were determined as described in Sect. 2.3. After filtering out all ion mass peaks that (1) are related to primary ions (including those masked by a water flux interference) or (2) do not pass the delay time criterion described in Sect. 2.3 or (3) have an average daytime flux $< 0.1 \text{ mg m}^{-2} \text{ h}^{-1}$, we obtain 14 grade A flux ion mass peaks (black) and four grade B flux ion mass peaks (grey). Table 1 lists these ion mass peaks by descending flux magnitude grouped by likely compounds. As an example, MBO and isoprene show a flux on the protonated MBO parent mass peak (m/z 87.0811), the protonated isoprene and MBO fragment mass peak (m/z 69.0704) and the fragment mass peak (m/z 41.0400) as well as the isotopes (m/z 88.0824, m/z 70.0729). After unravelling isotopic patterns and fragments, we obtain seven different compounds or compound classes emitted by this ecosystem during daytime.

Diurnal cycles for all 7 compounds showing a significant flux are depicted in Fig. 4. The emissions are dominated by the sum of MBO and isoprene, accounting for 50 % of the daytime emissions. When using a PTR-TOF-MS operated in H_3O^+ mode the fragment of protonated MBO exhibits the same ion mass peak as protonated isoprene. Therefore the sum of MBO and isoprene is given. Using the calibrated fragmentation pattern of MBO and comparing it to the ratio of m/z 69.0704 and m/z 87.0811 measured in ambient air it is possible to estimate the isoprene flux at this site. During

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the measurements in summer 2010 10% of the total MBO + isoprene flux can be attributed to isoprene. Other BVOC fluxes include methanol (11%), the sum of acetic acid and glycolaldehyde (10%), the sum of monoterpenes (10%), the sum of acetone and propanal (7%), ethanol (5%) and acetaldehyde (5%). Formaldehyde flux measurements published by DiGangi et al. (2011) contribute 2% of the total observed BVOC emission with a mean daytime emission of $0.08 \text{ mg m}^{-2} \text{ h}^{-1}$.

As shown by Harley et al. (1998) and Schade et al. (2001) MBO emissions are light (Eq. 1) and temperature (Eq. 2) dependent, and these dependencies are well described by the following expressions.

$$C_L = \frac{\alpha \cdot C_{L_1} \cdot L}{\sqrt{1 + \alpha^2 \cdot L^2}} \quad (1)$$

C_L is a scalar that accounts for the emission affected by PAR, α and C_{L_1} are empirical coefficients and L is PAR.

$$C_T = \frac{E_{\text{opt}} \cdot C_{T_2} \cdot e^{C_{T_1} \cdot x}}{C_{T_2} - C_{T_1} \cdot (1 - e^{C_{T_2} \cdot x})}, \quad x = \frac{1}{T_{\text{opt}}} - \frac{1}{T} \quad (2)$$

C_T accounts for the emission influenced by temperature. E_{opt} is the emission capacity at the temperature T_{opt} , T the temperature, R the ideal gas constant and C_{T_1} , C_{T_2} the empirical coefficients.

To calculate light dependencies, data from a narrow temperature range (e.g. 23–27°C) was chosen and the remaining data was binned into $100 \mu\text{mol m}^{-2} \text{ s}^{-1}$ light steps. Similarly for the temperature dependence, only data for which PAR exceeded $1000 \mu\text{mol m}^{-2} \text{ s}^{-1}$ was taken and binned into 1°C temperature windows.

A comparison of the light dependence of our measurements between 23–27°C with previous measurements and MEGAN 2.1 is shown in Fig. 5 and summarized in Table 2. The Harley et al. (1998) light curve represents leaf level measurements; MBO

Emission of most compounds discussed so far (except ethanol) also show some light dependence. Whether this is due to stomatal conductance or due to a direct light effect cannot be distinguished with these measurements. The results of a linear fit through data in a temperature range of 23–27 °C binned in 100 $\mu\text{mol m}^{-2} \text{s}^{-1}$ light windows are given in Table 6.

3.4 Comparison to MEGAN 2.1 model

The observed above canopy emissions of the sum of MBO and isoprene and the sum of monoterpenes are compared with MEGAN 2.1 model results using the landscape average emission factors based on four different approaches: MEGAN 2.1 Southern Rockies Conifers (M21SRC, MEGAN 2.1 average value for the region including ponderosa pine woodlands and nearby mixed conifer forests), MEGAN 2.1 ponderosa pine (M21PP, MEGAN 2.1 value for ponderosa pine woodlands), site specific measurements of ponderosa pine (SSLC, measured at MFO using a leaf cuvette), site specific measurements of ponderosa pine (SSBE, measured at MFO using a branch enclosure). Landscape average emission factors are the weighted average of species-specific emission factors for all plant species in a landscape. The SSBE and SSLC emission factors use the species composition and vegetation cover fraction observations obtained with ground surveys at the Manitou Forest site. The difference between them is that the SSBE estimate is based on branch enclosure (Teflon bag) measurements and the SSLC estimate is based on leaf cuvette (enclosure using a modified LiCOR 6400) measurements. All of the trees sampled by branch enclosure or leaf cuvette were within the flux tower footprint. The M21SRC and M21PP emission factors use the same species-specific emission factors (based on literature values including Guenther et al., 1994; Harley et al., 1998; Schade et al., 1999, 2001; Karl et al., 2002) and the same needle leaf tree cover fraction based on Moderate Resolution Imaging Spectroradiometer (MODIS) satellite land cover estimates (see Guenther et al., 2012). The difference is in the species composition estimated for the Manitou Forest site by the two approaches. The M21SRC estimate uses averaged species composition data

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agree with the measurements. After DOY 240 all emission factor approached agree similarly well with an R^2 of 0.68 and slopes between 0.57–0.77.

3.5 The hailstorm event

On 4 August a severe hailstorm impacted the site. Hailstones of about 1–3 cm in diameter caused damage to needles and branches. As a result, 22 ion mass peaks showed an elevated flux signal. All ion mass peaks are given in Table 8; they are grouped into their likely compound classes and sorted by descending amount of emission. Assignments of exact ion mass peaks to compounds/compound classes were made using the work of Kim et al. (2010) and references therein. Especially interesting is that we also observe first order oxidation products of monoterpenes such as pinonaldehyde and nopinone exhibiting an elevated flux after the hailstorm. At this point we cannot determine whether this results from increased in-canopy overall oxidation rates or if this is due to direct emission from the wounded branches and needles.

The sum of monoterpenes shows by far the highest emission with a maximum of $4.75 \text{ mg m}^{-2} \text{ h}^{-1}$. This corresponds to a 4 times higher maximum flux as observed on the day before the hailstorm. If we use MEGAN 2.1 to calculate the regular (temperature and light) dependent flux on the day after the hailstorm and compare this to the measurements the observed flux is ~ 23 times higher. This is justified as the measurements on 3 August (day before the hailstorm) might be still influenced from a hailstorm on 27 July. Figure 9 shows the maximum of emission before and after the hailstorm (3 August and 5 August) for all compounds that show a flux during the 30 day average as well as those which are influenced by the hailstorm. The fluxes of MBO + isoprene as well as the flux of the sum of monoterpenes are also compared to the expected flux calculated by MEGAN 2.1 based on the emission factors calculated for this work (see Sect. 3.4, Table 7). Elevated monoterpene concentrations and fluxes after mechanical wounding have been previously observed (Bouvier-Brown et al., 2009; Haase et al., 2011; Bamberger et al., 2011; Holzinger et al., 2006). New in this study is that a natural event such as a hailstorm can cause such a strong mechanical damage to enhance

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monoterpene emissions by a factor of 4–23 as well as the enhancement of several other compounds such as cymene, camphor, nopinone, pinonaldehyde and the sum of sesquiterpenes. Some compounds such as the sum of sesquiterpenes as well as the flux measured on m/z 105.070 show a maximum of elevation only two days after the hailstorm indicated in Fig. 9.

Figure 10 shows the concentrations (upper panel) and fluxes (lower panel) of the sum of monoterpenes (red) and the sum of MBO and isoprene (blue) for the whole measurement period. After the hailstorm on 4 August (DOY 216) monoterpene volume mixing ratios and fluxes are elevated. In contrast, the sum of MBO + isoprene is not elevated after the hailstorm. Some monoterpenes are stored in pools and due to the wounding of the pools monoterpenes are released in higher amounts. MBO + isoprene are not stored, but emitted directly after their production. Higher emissions of stored compounds could be induced by mechanical wounding of the trees.

Figure 11 shows the cumulative flux calculated from the measurements (blue) and the expected flux assuming only temperature and light dependencies using MEGAN 2.1 (red) adjusted to the measurements at the end of the measurement campaign when we assume no more influence of the hailstorm (after DOY 240). This indicates that if monoterpene emissions are calculated based on temperature and light dependence only then the cumulative flux over one month is underestimated by $\sim 40\%$.

4 Summary and conclusion

Eddy covariance PTR-TOF-MS measurements above a ponderosa pine forest in Colorado, USA revealed fluxes of 17 ion mass peaks that can be grouped into seven compound classes. The sum of MBO and isoprene is the dominant flux, contributing 50% to the total observed flux. The other compounds showing a significant daily flux are methanol, the sum of acetic acid and glycolaldehyde, the sum of monoterpenes, the sum of acetone and propanal, and ethanol.

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Temperature and light dependence of MBO+isoprene are found to be in good agreement with previous measurements. Eddy covariance measurements of MBO + isoprene and the sum of monoterpenes are compared to four different emission factor estimation approaches available for MEGAN 2.1 at this site. The emission factor obtained from site specific leaf cuvette emission measurements and tree cover surveys agree best with the above canopy flux measurements.

The wounding of branches and needles due to a severe hailstorm triggered the emission of monoterpenes and other compounds such as camphor, pinonaldehyde, nopinone and sesquiterpenes. The storm increased the monthly emissions of monoterpenes by 40 % compared to model results assuming only temperature and light dependent emission. These results stress the importance of long term above canopy flux measurements to increase the chance to capture irregularly occurring events such as hailstorms, which can significantly influence the budget of VOCs emitted from a needle tree forest.

Acknowledgements. This work was financially supported by the Austrian Science Fund (FWF) under the project number L518-N20. The National Center for Atmospheric Research is operated by the University Corporation for Atmospheric Research under sponsorship from the National Science Foundation. Lisa Kaser is a recipient of a DOC-fORTE-fellowship of the Austrian Academy of Science. The authors would like to thank Ionimed/Ionicon Analytik in Innsbruck for providing the liquid calibration unit (LCU) and the BEACHON ROCS science team that assured the success of this measurement campaign.

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Table 1. Volatile organic compounds showing an emission flux during the period of 5 August to 8 September 2010. Grade A fluxes (bold) and grade B fluxes (italic) are listed with sum formula and proposed compound.

Molecular formula	Likely compound	Parent ion (m/z)	Fragment (m/z)	Isotopologue (m/z)	Water cluster (m/z)	Flux ($\text{mg m}^{-2} \text{h}^{-1}$) ^a	Percent data used ^b	Calibration method	Sensitivity (ncps/ppbv)
$\text{C}_9\text{H}_{10}\text{O}-\text{H}^+$	MBO and isoprene	87.0811	69.0704 41.0400	<i>88.0824</i> <i>70.0729</i>		1.84	40	Gas standard	13.5
$\text{CH}_4\text{O}-\text{H}^+$	Methanol	33.0336			<i>51.0438</i>	0.42	47	Gas standard	11.7
$\text{C}_2\text{H}_2\text{O}-\text{H}^+$	acetic acid and glycolaldehyde	<i>61.0287</i>	<i>43.0180</i>			0.36	36	LCU	7.4
$\text{C}_{10}\text{H}_{16}-\text{H}^+$	sum of monoterpenes	137.134	81.0706 <i>95.0866</i>	138.137 <i>82.0738</i>		0.35	33	Gas standard	13.1
$\text{C}_3\text{H}_6\text{O}-\text{H}^+$	acetone and propanal	59.0495				0.27	39	Gas standard	19.7
$\text{C}_2\text{H}_6\text{O}-\text{H}^+$	Ethanol	47.0469				0.17	39	LCU	2.5
$\text{C}_2\text{H}_4\text{O}-\text{H}^+$	Acetaldehyde	45.0337				0.17	50	Gas standard	17.8

^a Daytime (10:00–18:00 LT) average flux ($\text{mg compound m}^{-2} \text{h}^{-1}$).

^b Percentage of datapoints used after quality control.

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Table 2. Comparison of the empirical coefficients from the light dependent MBO and isoprene emission following Eq. (1).

	Harley et al. (1998)	Schade et al. (2001)	This work	MEGAN 2.1
A	0.0011	0.0011	0.0011	0.0007
C_{L1}	1.44	1.37	1.36	1.74

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Table 4. Comparison of the empirical coefficients from the temperature dependent MBO and isoprene emission following Eq. (2).

	Harley et al. (1998)	Schade et al. (2001)	This work	MEGAN 2.1
E_{opt}	1.54	1.45	1.29	1.32
C_{T1}	67	131	128	134
C_{T2}	209	154	149	156
T_{opt}	312.3	312	312	312

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Compound ($T_{\text{ref}} = 303 \text{ K}$)	F_{ref} [$\text{mg m}^{-2} \text{ h}^{-1}$]	B	r^2	Q_{10}
Sum of monoterpenes*	0.50	0.12	0.85	3.32
Methanol	0.94	0.13	0.81	3.67
Acetone + propanal	0.63	0.15	0.92	4.48
Acetaldehyde	0.33	0.12	0.95	3.32
Acetic acid + glycolaldehyde	0.91	0.15	0.94	4.48
Ethanol	0.28	0.07	0.86	2.01

* For the sum of monoterpenes only data after DOY 230 was used as the data before this day was significantly affected by the hailstorm.

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Compound ($T_{\text{ref}} = 303 \text{ K}$)	Slope	Intercept	r^2
Sum of monoterpenes*	1.7×10^{-4}	0.15	0.59
Methanol	2.5×10^{-4}	0.15	0.87
Acetone + propanal	1.5×10^{-4}	0.11	0.53
Acetaldehyde	6.8×10^{-5}	0.11	0.67
Acetic acid + glycolaldehyde	2.5×10^{-4}	0.09	0.70

* For the sum of monoterpenes only data after DOY 230 was used as the data before this day 9 was significantly affected by the hailstorm.

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Table 7. Summary of emission factor approaches. The sum of monoterpenes includes: myrcene, sabinene, limonene, 3-carene, trans- β -ocimene, β -pinene, α -pinene and other monoterpenes.

Emission factor ($\mu\text{g compound m}^{-2} \text{h}^{-1}$)	M21SRC	M21PP	SSLC	SSBE	This study
MBO	5300	9000	23 000	15 000	21 485
Sum of monoterpenes	1862	1650	1820	1370	2395*

* For monoterpene emissions after DOY 240 where model results imply that there is no more influence from the hailstorm and assuming the same partitioning of the different monoterpenes as measured by SSLC.

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Table 8. List of ion mass peaks that show an elevated flux after the hailstorm on 4 August. If possible they were grouped into compound classes and sorted by the amount of maximum emission after the hailstorm. The remaining ion mass peaks are listed at the end of the table by decreasing molecular weight.

Molecular formula	Likely compound	Exact masses (m/z)
$C_{10}H_{16}-H^+$	Sum of monoterpenes	137.134 (parent ion) 138.137 (isotope) 81.0706 (fragment) 82.0738 (isotope) 95.0866 (fragment)
$C_{10}H_{14}-H^+$	p-cymene	135.118 (parent ion) 136.13 (isotope) 93.0707 (fragment)
$C_{10}H_{16}O-H^+$	e.g. camphor	153.128 (parent ion) 154.132 (isotope)
$C_{10}H_{16}O_2-H^+$	Pinonaldehyde	169.124 (parent ion) 151.113 (fragment) 109.105 (fragment) 107.086 (fragment) 99.081 (fragment)
$C_9H_{14}O-H^+$	Nopinone	139.117 (parent ion) 121.098 (fragment)
$C_{15}H_{24}-H^+$	sum of sesquiterpenes	205.198 (parent ion) 95.0866 (fragment)
$C_8H_8-H^+$	e.g. styrene	105.070 (parent ion)
Several still unassigned ion mass peaks that also had elevated fluxes after the hailstorm:		
$C_7H_{10}O_3-H^+$	Unknown	143.069
$C_7H_{10}O_2-H^+$	Unknown	127.077

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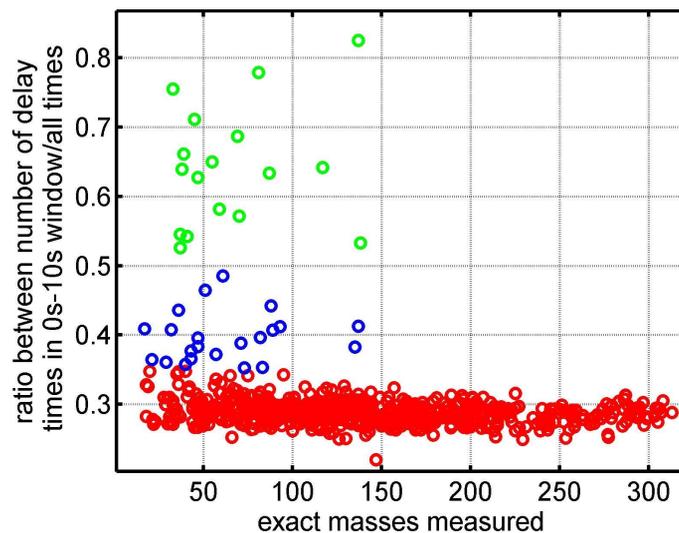


Fig. 1. Fraction of physically plausible delay times as a function of recorded exact ion mass peaks. Green circles (classified as grade A flux) indicate ion mass peaks that were recorded more than 50% of the time having a physically plausible delay time in the range of 0 to 10 s. Blue circles are grade B fluxes and red circles represent the majority of ion mass peaks that did not show a significant flux during the measurement campaign lasting from 4 August to 8 September 2010.

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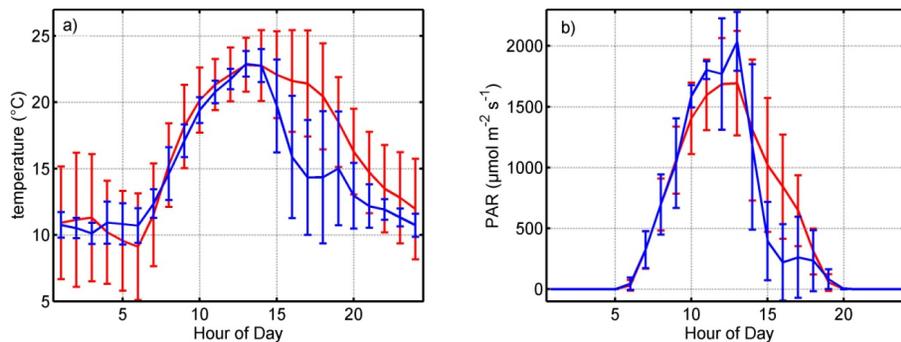


Fig. 2. Diurnal cycles of temperature **(a)** and PAR **(b)** during the 30 day period of VOC flux measurements (red) and on the day of the hailstorm and the following two days (blue).

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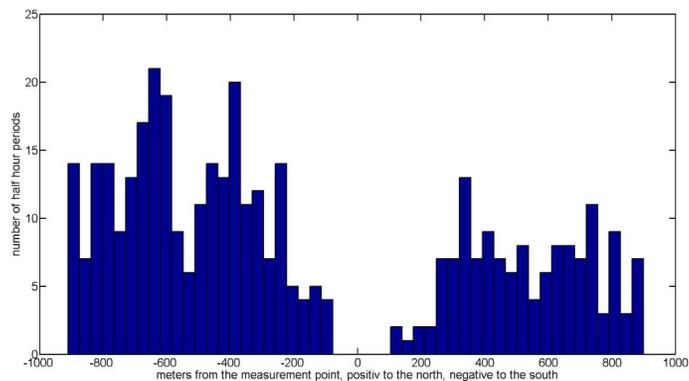


Fig. 3. All half hour values for the 90 % flux recovery during the unstable atmospheric conditions lie within 900 m. Negative values indicate the distance in meters from the flux tower to the south, positive from the tower to the north.

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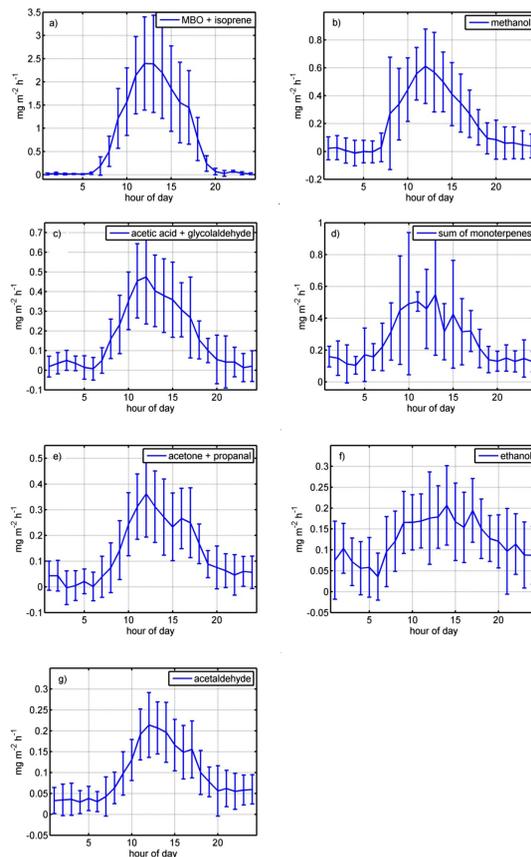


Fig. 4. (a–g) show averaged diurnal cycles of the 7 VOC species observed to have a significant flux.

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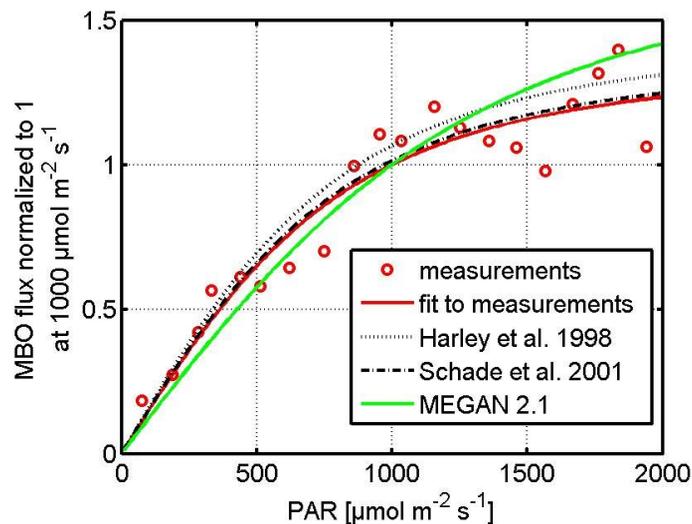
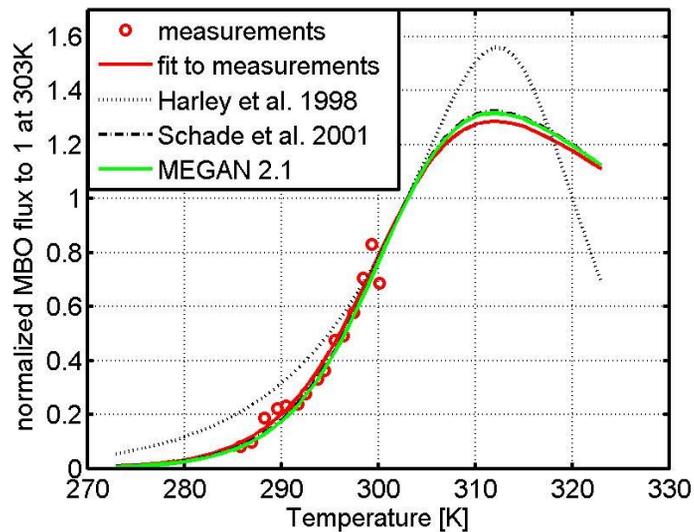


Fig. 5. Light dependence of measurements between 23–27°C for the sum of MBO and isoprene emissions. Results from this study (red-solid) are compared to earlier leaf-level results by Harley et al. (1998) (black-dot) and canopy-scale results of Schade et al. (2001) (black-dash-dot). Red circles represent our measurements averaged into $100 \mu\text{mol m}^{-2} \text{s}^{-1}$ bins.

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**Fig. 6.** Temperature dependence of MBO emissions.[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

PTR-TOF-MS
measurements and
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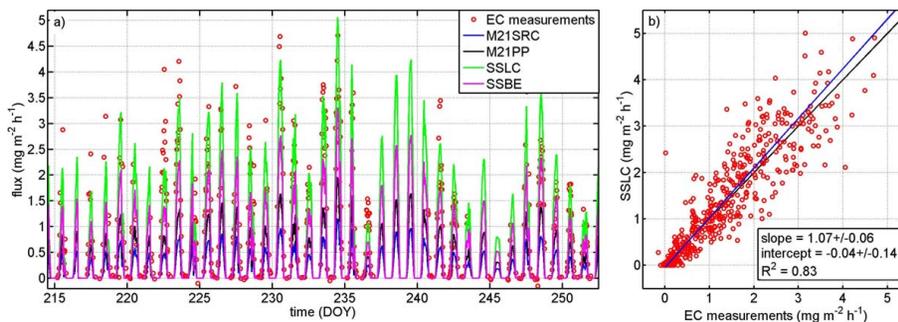


Fig. 7. (a) Emissions for MBO + isoprene (modelled) and MBO + isoprene (measured): red circles indicate the measurements, solid lines the four available emission types. **(b)** Scatterplot between PTR-TOF-MS EC flux measurements of MBO + isoprene and MEGAN 2.1 results using the SSLC emission factor.

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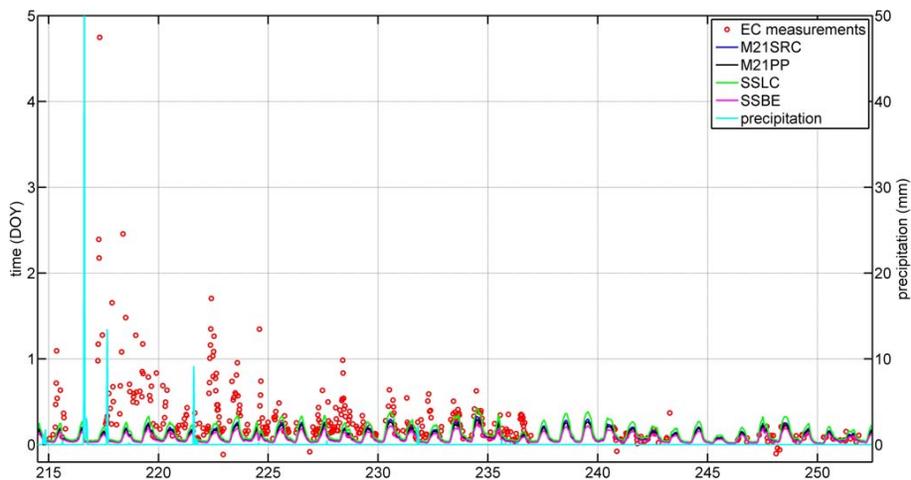


Fig. 8. Emissions for the sum of monoterpenes: red circles indicate the measurements, solid lines the four available emission factor approaches used in MEGAN 2.1 to model the emissions. Precipitation in mm is given for references (cyan line).

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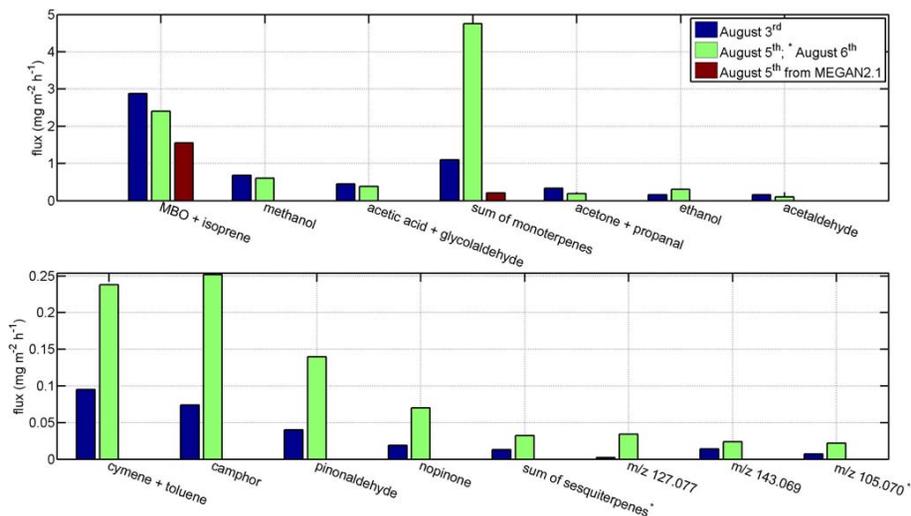


Fig. 9. Comparison of the fluxes on the day after the hailstorm with measured fluxes on the day before the hailstorm as well as the calculated flux from MEGAN 2.1 using emission factors calculated from this work (see Table 7).

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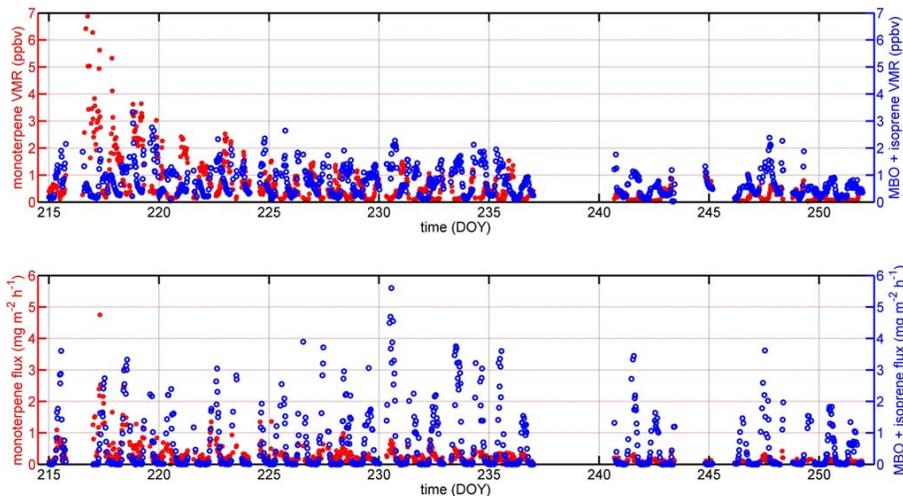


Fig. 10. Volume mixing ratios of monoterpenes (red) and the sum of MBO + isoprene (blue) are shown for the entire measurement period in the upper panel. The lower panel shows fluxes for the same species and time. The hailstorm occurred on 4 August (DOY 216).

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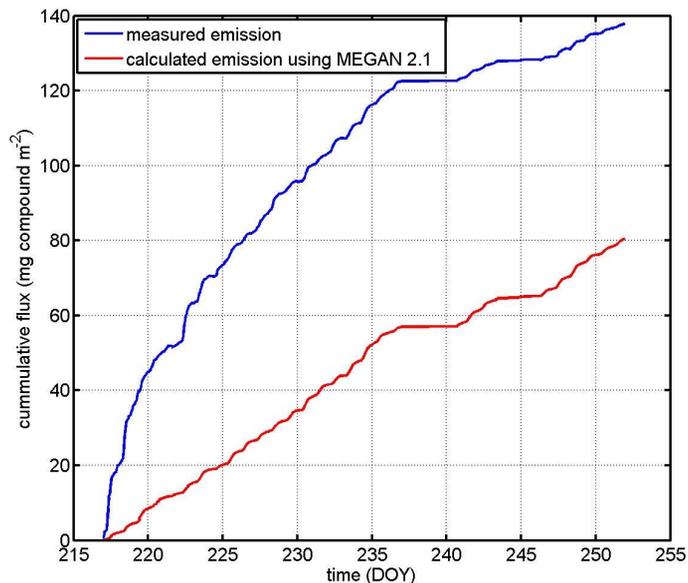


Fig. 11. Cumulative flux calculated from the 30 day measurement period (blue) and cumulative flux based on MEGAN 2.1 assuming only temperature and light dependence (red).

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