

**Hygroscopic growth
of iodic acid solution
droplets**

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**Glass formation and unusual hygroscopic
growth of iodic acid solution droplets with
relevance for iodine oxide particles in the
coastal marine boundary layer**

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Abstract

Iodine oxide particles are known to nucleate in the marine boundary layer where gas phase molecular iodine and organoiodine species are produced by macroalgae. There has been some debate over the chemical identity of these particles. Hygroscopic measurements have been used to infer that they are composed of insoluble I_2O_4 , while elemental analysis of laboratory generated particles suggests soluble I_2O_5 or its hydrated form iodic acid, HIO_3 ($I_2O_5 \cdot H_2O$). In this paper we explore the response of super-micron sized aqueous iodic acid solution droplets to varying humidity using both Raman microscopy and single particle electrodynamic traps. These measurements reveal that the propensity of an iodic acid solution droplet to crystallise is negligible on drying to $\sim 0\%$ relative humidity (RH). On applying mechanical pressure to these droplets they shatter in a manner consistent with an ultra-viscous liquid or a brittle glass, but subsequent water uptake between 10 and 20% RH causes their viscosity to reduce sufficiently that the cracked droplets flow and merge. The persistence of iodic acid solution in an amorphous state, rather than a crystalline state, suggests they will more readily accommodate other condensable material and are therefore more likely to grow to sizes where they may serve as cloud condensation nuclei. On increasing the humidity to $\sim 90\%$ the mass of the droplets only increases by $\sim 20\%$ with a corresponding increase in radius of only $\sim 6\%$, which is remarkably small for a highly soluble material. We suggest that the small growth factor of aqueous iodic acid solution droplets is consistent with the small aerosol growth factors observed in field experiments.

1 Introduction

In seaweed-rich coastal regions there is a marked increase in the number of nucleation mode particles observed during daytime low tides when macroalgal species are exposed to air and emit large quantities of molecular iodine (McFiggans et al., 2010; Saiz-Lopez et al., 2011). Particle production is driven by biogenic emissions of iodine

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vapours that rapidly undergo photochemical processing to produce gas phase iodine oxides such as IO and OIO which undergo further reactions leading to higher oxides which can condense to form new particles (Saiz-Lopez et al., 2011; O'Dowd and Hoffmann, 2005). Under suitable conditions these particles may grow, through the condensation of other condensable vapours (Mäkelä et al., 2002; Saunders et al., 2010), to sufficiently large sizes that they can serve as cloud condensation nuclei (McFiggans et al., 2010). At present it is unclear if iodine production from open ocean sources can be sufficiently strong to lead to significant particle formation over the open ocean (Saiz-Lopez et al., 2011; Mahajan et al., 2010).

There has been some debate in the literature over the most likely composition of iodine oxide particles with insoluble I_2O_4 and the soluble I_2O_5 being identified as the most likely candidates (Jimenez et al., 2003; O'Dowd and De Leeuw, 2007). Ultrafine particles sampled during low tide particle "bursts" at Mace Head in Ireland were reported to have a characteristic hygroscopic (diameter) growth factor (based on electrical mobility diameters, $D_{final}/D_{initial}$) of less than 1.1 at an RH of 90% (Väkevä et al., 2002), ruling out sulfate or sea salt aerosol which are known to have much greater growth factors. Mäkelä et al. (2002) showed that the ultrafine particles sampled at low-tide periods at the same coastal location contained iodine of biogenic origin. Iodine oxide particles generated in the laboratory through a range of processes also have small hygroscopic growth factors consistent with the field observations (Saunders et al., 2010; Jimenez et al., 2003; McFiggans et al., 2004). In general, a small growth factor indicates a weakly soluble material or an insoluble material internally mixed with a small amount of soluble material (Swietlicki et al., 2008; Väkevä et al., 2002). Since I_2O_5 is highly soluble (263.2 g per 100 g of water at 298 K, Kumar et al., 2010), Jimenez et al. (2003) concluded that I_2O_4 was the most likely composition of iodine oxide particles. However, elemental analysis of particles generated photochemically from I_2 in the presence of O_3 revealed a composition consistent with I_2O_5 (Saunders and Plane, 2005). I_2O_4 exists as an insoluble crystalline solid which slowly disproportionates to I_2 and HIO_3 when immersed in water at room temperature, or even just exposed to humid

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air for a prolonged period (Daehlie and Kjekshus, 1964). Hence, if I_2O_4 forms in the humid marine boundary layer it may later convert to iodic acid solution (Saiz-Lopez et al., 2011).

The physical state of coastal atmospheric iodine oxide particles is an unresolved issue. Aerosol particle can exist in a liquid, crystalline solid, amorphous solid or amorphous semi-solid state. Amorphous substances by definition have no long range order, in contrast to crystalline materials, and an amorphous material is said to become glassy when its viscosity exceeds 10^{12} Pa s (Angell, 2002). At such extreme viscosities an amorphous material will have physical properties consistent with that of a solid. At lower viscosities materials can exist as ultra-viscous liquids, rubbers or gels (Mikhailov et al., 2009). Gels are defined as a two-phase mixture in which one phase is a solid matrix while the other retains some mobility within the matrix. Rubbers form when intermolecular bonding between molecules forms a complex cross linked matrix. Ultra-viscous liquids will exhibit viscous flow on long timescales, but if subjected to a shear force on a short time scale they can shatter. An example of an ultra-viscous liquid is pitch which shatters when hit with a hammer, but flows on a decadal timescale with an estimated viscosity of 10^8 Pa s (Edgeworth et al., 1884). On cooling, rubbers, gels and ultra-viscous liquids can form a glass at the glass transition temperature. In addition, water serves as a plasticiser in many amorphous materials and reduction of relative humidity may also cause an amorphous material to become semi-solid or glassy (Mikhailov et al., 2009).

It has been suggested that aerosol particles can become glassy at the very low temperatures found at the tropopause (Murray, 2008; Zobrist et al., 2008) and it was later shown that this can influence cirrus cloud formation through heterogeneous ice nucleation (Murray et al., 2010b). It has also been shown that on drying atmospherically relevant aqueous droplets at temperatures relevant for much of the troposphere that they can become semi-solid or even glassy rather than crystallising (Mikhailov et al., 2009; Tong et al., 2011; Bodsworth et al., 2010) and that the rate of exchange of water with air is greatly reduced at low humidity (Zobrist et al., 2011; Tong et al., 2011).

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Recent work also shows that secondary organic aerosol may exist in an amorphous solid state (Virtanen et al., 2010; Koop et al., 2011). The existence of aerosol particles in an amorphous solid or semi-solid state is important because their growth through uptake of condensable vapours will be reduced relative to a liquid droplet since diffusion into their interior is limited (Shiraiwa et al., 2011). The behaviour of organic solutions contrasts strongly with typical atmospherically relevant inorganic solutions, such as NaCl or $\text{NH}_4(\text{SO}_4)_2$, which readily crystallise at low RH (Martin, 2000; Onasch et al., 1999; Pant et al., 2006). However, Mikhailov et al. (2009) also discussed a number of inorganic solution aerosol species which tend not to exhibit prompt crystallisation, such as NH_4HSO_4 and NH_4NO_3 , and it has been suggested that the freeze concentrated solution in frozen inorganic solution droplets is glassy (Murray and Bertram, 2008, 2007). Also, delayed water transport in amorphous particles of MgSO_4 at RHs below 40 % has been recently reported, with MgSO_4 persisting in an amorphous gel state to 0 % RH without crystallisation (Li et al., 2011).

Previous work suggests that iodic acid droplets may become very viscous at low humidity (Kumar et al., 2010; Ivanov et al., 1985). Kumar et al. (2010) quantified the deliquescence RH (DRH) of crystalline I_2O_5 and HIO_3 as well as studying the effect of the drying of iodic acid solution droplets (note that I_2O_5 and HIO_3 share a common aqueous phase, which we refer to as an iodic acid solution). The crystalline solids were shown to have DRH values in excess of 80 % below 300 K, but when the micrometre sized solution droplets were gradually dried out to 0 % RH they did not appear to crystallise when viewed with an optical microscope. Kumar et al. (2010) suggested that the droplets may not have crystallised instead becoming extremely viscous even to the extent of possibly forming a glass at low RH. This is intriguing because it suggests that if iodine oxide particles are either produced in an aqueous liquid state or as crystalline particles that have undergone deliquescence in the humid marine boundary layer, they will persist in an amorphous (liquid or solid) state.

In this paper, we use a Raman microscope coupled to a temperature and humidity controlled stage to explore the physical state of iodic acid solution droplets. We also use

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two types of electrodynamic balances to quantify the hygroscopic growth of aqueous iodic acid solution droplets and probe the rate at which they lose water.

2 Experimental

2.1 Raman microscope and humidity controlled stage

5 A diagram of the humidity controlled stage coupled to the Raman microscope system is shown in Fig. 1. Crystalline HIO_3 particles (99.5 %, Sigma-Aldrich) with dimensions of several 10s of micrometres were mounted on a silanised glass cover slip. Similar cover slips have been used in previous nucleation studies and shown not to catalyse nucleation (Murray et al., 2010a, b, 2011a; Knopf and Lopez, 2009; Dymarska et al., 10 2006; Broadley et al., 2012). The cell was constructed of Teflon and sealed on the base with an O-ring with a window (cover slip) on top through which the sample could be viewed. The humidity in the nitrogen flowing over the particles was controlled by passing nitrogen gas through a bubbler immersed in a temperature controlled bath and diluted using mass flow controllers (MKS 1179) to produce the desired humidity. 15 The humidity of the gas downstream of the cell was recorded with a dew/frost point hygrometer (General Eastern D-2 sensor). The temperature of the stage was set using a re-circulating chiller and the temperature of the sample was determined from the average of two T-type thermocouples embedded in the stage immediately below the sample. In these experiments the temperature was set to 294 ± 0.5 K. The temperature of the stage and dew point of the gas were used to derive the RH of the gas above the sample according the parameterisations of Murphy and Koop (Murphy and Koop, 20 2005).

The humidity and temperature controlled stage was mounted on the automated x-y translation stage of the Raman system which allowed the stage to move independently of the objective and Raman system. Raman microscopy enabled us to optically view 25 particles under magnification and then probe individual particles by focusing a laser

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5 onto them. The resulting inelastic Raman emission was then collected and analysed using a spectrometer. The Raman spectrum provides a means of unambiguously identifying chemical compounds and determining phase and phase changes (Reid et al., 2007). Our experiments were performed with a 514 nm laser coupled to a Renishaw
inVia Raman microscope system equipped with a Leica microscope and Olympus 20× long working distance objective.

2.2 Electrodynamic balance (EDB) instruments

10 Two separate electrodynamic balances (EDBs) were used in these experiments: one with a double ring electrode design and the second with a concentric cylindrical electrode design. The description and standard operation of the double ring EDB has been published previously (Hargreaves et al., 2010). Briefly, an aqueous iodic acid (from the same source as used in Raman experiments) solution droplet was placed into the reservoir of a droplet-on-demand dispenser (GeSiM, A-010-006, 30 μm orifice). The dispenser was positioned directly ~2 mm above the induction electrode, to which a
15 +175 V DC voltage is applied, and ~12 cm above the double ring electrodes of the EDB, to which an AC potential (2 kV(0-P), 50–200 Hz) was applied. The field generated by the induction electrode imparted a mirror charge onto the droplet as it was ejected from the dispenser. The presence of this net charge enabled the droplet to be levitated by the AC field generated by the ring electrodes. A 632 nm He-Ne laser was used to illuminate the droplet.
20

The size of the levitated droplet was estimated using two different methods. The relative mass of the droplet was measured by monitoring the amplitude of the DC offset to the AC field required to hold the droplet at a fixed position. The RH in the EDB was controlled by varying the relative proportion of dry and humidified nitrogen gas flows with a net flow rate held constant at 200 sccm. To generate the humidified airflow, nitrogen was passed through a 2 l bubbler filled with pure water (18 MΩ cm). The mixture of humidified and dry air was passed through an electrostatic precipitator (ESP) prior to the EDB. The temperature and relative humidity of the airflows entering and leaving
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the EDB were measured using two Honeywell probes (HIH-4602-A/C Series). The RH was then ramped from <20 % to >90 % then back to <20 % over a period of ~24 h, with an hour given to equilibrate at each 10 % increase in RH and two hours given to equilibrate each 10 % decrease in RH.

5 A second electrodynamic balance, based on concentric cylindrical electrodes, was used to observe mass transfer into dry conditions from droplets of known size and composition. A detailed description of the technique and method utilized is described by Davies et al. (2012). A key benefit of this technique is the ability to accurately determine initial size and composition, allowing an estimate of the dry size. The elastic scattering of the laser by the droplet in the angular range of 42° to 50° was collected with a CCD and the angular separations between the peaks in the fringe pattern were used to estimate the radius of the droplet (Glantschnig and Chen, 1981). Each measured scattering pattern was compared with Mie theory simulations and the best fit size was obtained for each frame. This estimation results in an error in absolute radius of less than 50 nm, which is a significantly more sensitive measure of changes in size than the DC offset. The timescales accessible with the instrument range from sub-100 ms to many hours, with a resolution down to 10 ms, allowing both the initial rapid evaporation and the slow approach to equilibrium to be readily measured.

3 Results and discussion

3.1 Raman study of HIO₃ solution droplets

20 Crystalline HIO₃ particles were ground with pestle and mortar and deposited directly onto the silanised glass slide. The slide was then mounted on the RH and temperature controlled stage. The temperature of the cell was maintained at 294 ± 0.5 K throughout the experiment and the RH was initially set to 48.5 %. The Raman spectrum and accompanying image of the crystal are shown in Fig. 2a. The RH was increased (or decreased) in steps and the particles allowed to equilibrate for at least 10 min before a

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Raman spectrum was recorded. No change in the particles was observed until 82.5 % RH when the image of the droplet clearly shows water uptake and an aqueous droplet forming around a crystal (Fig. 2c). The corresponding spectrum is still dominated by the features of the crystalline material, but a weak water peak can be seen to appear centred at about 3400 cm^{-1} . On increasing the RH to 83.5 % (Fig. 2d) the crystal can be seen to have fully dissolved in the image and the Raman spectrum has altered substantially. The peaks associated with IO_3 below 1000 cm^{-1} (see Table 1) all become much broader and a clear OH stretch signal appears around 3400 cm^{-1} which is indicative of liquid water. Kumar et al. (2010) determined that the DRH of HIO_3 at 294 K was $84 \pm 1\%$, which is consistent with the RH at which we observe the solid to aqueous transition. In Table 1, we identify and compare the bands in the Raman spectra of the aqueous HIO_3 solution droplets and show they are consistent with Raman spectra of iodic acid solutions in the literature (Durig et al., 1965).

A selection of the spectra recorded when drying a deliquesced aqueous HIO_3 droplet is shown in Fig. 3. Other droplets behaved in the same way. The most obvious change in the Raman spectra on reducing RH is that the OH bands at around 3400 cm^{-1} decreased in intensity as water evaporated from the droplets. Even though the water content was clearly reduced on drying and a supersaturated solution results, crystallisation did not occur and the droplet remained in an amorphous state in which water was retained. After leaving the droplet overnight at 0% RH, the droplet (or any other droplets) did not crystallise (Fig. 3e).

In the spectra at very low RH (Fig. 3c to e), the water signal around 3400 cm^{-1} appears to be distorted. The band at 3400 cm^{-1} is indicative of water molecules hydrogen bonded to other water molecules in a liquid environment, whereas the Raman signal from water molecules bound to a cation such as H_3O^+ or H_5O_2^+ forms a continuum on the low frequency side of the water band (Giguere and Guillot, 1982). The hydronium ion itself has a Raman signal which is orders of magnitude less in intensity than that of water (Giguere and Guillot, 1982). The continuum on the low frequency side of the

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3400 cm⁻¹ peak in Fig. 3c to e is consistent with many of the water molecules in this solution being protonated.

On increasing the RH, the intensity of the OH band increased again, further confirming that the droplet remained in a supersaturated aqueous state even at 0% RH without a first order phase change (i.e. crystallisation) occurring. This spectroscopic information is summarised in Fig. 4 which shows the ratio of the integrated intensities of the OH stretch peak, indicative of water, centred at between 2900 and 3800 cm⁻¹ and the iodate peaks between 500–1000 cm⁻¹ as a function of RH. This ratio is a measure of the amount of H₂O hydrogen bonded to other H₂O in the droplets rather than the total amount of water in the droplet (for the reasons discussed above). Nevertheless, this ratio is a useful measure of the water content as a function of RH. Figure 4 clearly shows that iodine acid solution droplets absorb and loose water in a reversible manner as RH is cycled and that even at 0% RH a small, but significant, quantity of water remained in the droplets.

In order to probe the physical state of the droplets at low RH we applied a mechanical force to them. This involved pressing the flat end of a metal rod onto them by hand while maintaining a flow of 0% RH over the droplets. Images of the droplets in Fig. 5 reveal that many of the droplets shattered or cracked when subject to this mechanical force. These droplets behaved in a manner consistent with a glassy solid or an ultra-viscous liquid such as pitch. The effect of increasing RH on these cracked droplets is illustrated in Fig. 6. When the RH reaches about 16% RH the sharp edges begin to become rounded indicating increased mobility and at 18% RH the individual fragments of droplets merged. This indicates that water serves as a plasticiser: i.e. as water is absorbed into the amorphous solid (or ultra-viscous liquid) it reduces its viscosity to the point where it behaves like a liquid with smaller fragments merging to form larger units.

These results may also provide an alternative explanation for growth factors well below unity for iodine oxide particles generated under dry conditions. Jimenez et al. (2003) and Saunders et al. (2010) observed a substantial reduction in the mobility

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diameter of the aerosol generated under dry conditions when exposed to RH of greater than 20% RH. Under dry conditions, the aerosol particles were shown to have fractal morphologies (Jimenez et al., 2003; Saunders and Plane, 2006) and it was suggested that water uptake into capillaries would cause the fractals to collapse and form more dense particles (Saunders et al., 2010; McFiggans et al., 2004), hence resulting in an apparent growth factor of less than 1. Alternatively, we suggest here that the particles generated under dry conditions were amorphous solids. When the RH was increased, it is possible that water adsorbed and served as a plasticiser allowing the material to flow and coalesce into a more compact morphology in a similar manner to our cracked droplets which can be seen to flow on increasing RH in Fig. 6.

3.2 Electrodynamic balance measurements of the rate of droplet evaporation

The time-dependence of the size of an aqueous iodic acid droplet injected and trapped in dry nitrogen gas is shown in Fig. 7 as a function of time. These measurements were made with the concentric cylindrical electrode trap. The initial evaporation of water from a droplet of initial size $21.5\ \mu\text{m}$ and a concentration of $50.4\ \text{g l}^{-1}$ occurs at a rate close to that of pure water (Fig. 7a). On approach to the dry size the evaporation rate decreases sharply, consistent with formation of an amorphous solid state in which diffusion and evaporation are slow. Slow mass transfer is measured over $\sim 33\ \text{h}$ (Fig. 7b) indicating that the droplets did not come to equilibrium during this period. Based on the density of crystalline iodic acid ($4.63\ \text{g cm}^{-3}$, Lide, 2009) and the density of the initial solution (Kumar et al., 2010), we estimate the dry size of the iodic acid particle to be $4.95 \pm 0.15\ \mu\text{m}$. In addition, a regular elastic scattering pattern persists at all times, even after $\sim 33\ \text{h}$ (Fig. 7c), indicating sphericity is maintained throughout evaporation. The slow release of water over an extended period of time has been observed in glassy sucrose solutions (Tong et al., 2011; Zobrist et al., 2011); hence, these measurements are consistent with the iodic acid solution droplet existing in an amorphous solid state at very low humidity.

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3.3 Electrodynamic balance measurements of hygroscopic growth

The mass growth curve for an iodic acid solution droplet is shown in Fig. 8a. The mass growth is only around 20 % on increasing the RH from below 15 % to 90 %. This is a remarkably small uptake of water for a species that is highly soluble in water (0.26 moles per mole of water or 308 g/100 g of water at 25 °C, Lide, 2009). The mass of an aqueous sodium chloride particle, for example, increases by about 700 % on increasing the RH to 90 % from 0 % (see Fig. 8b). We can estimate the diameter growth of the iodic acid solution droplet if we assume that the density of supersaturated iodic acid solutions does not change over this range: a 20 % change in mass would correspond to an increase in diameter of only 6 %. This value is consistent with the small diameter growth factors observed in the field (Väkevä et al., 2002) and under humid conditions in the laboratory (Jimenez et al., 2003; McFiggans et al., 2004). Hence, the hygroscopic growth measurements in the field and laboratory are consistent with the aerosol being composed of aqueous iodic acid solutions. This does not rule out the possibility that iodine oxide particles in nature are I₂O₄. However, given that the elemental composition measurements of laboratory generated material indicate an I₂O₅ composition (the anhydrous crystalline form of iodic acid) and that it is known I₂O₄ reacts in water to produce iodic acid, we suggest iodine oxide particles in the marine boundary layer are composed of aqueous iodic acid in a liquid, amorphous solid or semi-solid state.

The reason why iodic acid solution droplets should have a weak growth compared to other soluble salts is unclear. In solution, iodic acid is thought to form polymeric structures which lead to strongly non-ideal properties (Goldman et al., 1974; Ivanov et al., 1985). It has been suggested that the formation of polymers of iodic acid with bound water may also be responsible for glass formation in concentrated solutions (Ivanov et al., 1985). Hence, the formation of polymers may be related to the very strong dependence of water activity on solute concentration and its unusual hygroscopic growth. MgSO₄ is also known to form polymeric structures in concentrated solutions and water transfer into and out of concentrated MgSO₄ solutions is retarded when it forms an

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amorphous solid at low RH (Li et al., 2011; Zhang and Chan, 2000), although aqueous MgSO_4 droplets differ from HIO_3 droplets in that they still grow substantially on increasing RH to 90 % (Tang et al., 1997; Zhang and Chan, 2000). Further work is needed to understand the role of polymerisation in aqueous HIO_3 , as well as other atmospherically relevant solutions, and its influence on hygroscopic properties.

In Fig. 9 we have referenced the mass growth data to the direct measurements of water activity and solute concentration previously reported by Kumar et al. (2010). The data are referenced at 83.8 % RH. The data for greater RH are in agreement within experimental uncertainty and the general trend at lower RH is also consistent with the previous bulk measurements. This indicates that the droplets retain $\sim 13 \pm 3$ wt % water (i.e. 87 ± 3 wt % HIO_3) at low RH, consistent with an amorphous solid or semi-solid state. Ivanov et al. (1984) estimated that aerosol with more than 83 wt % HIO_3 (17 wt % H_2O) were glassy based on whether solutions of varying concentration froze when immersed in liquid nitrogen. Their simple test is remarkably close to our value of 87 ± 3 wt %. The presence of water in these droplets is corroborated by the presence of the distorted OH peak in the Raman spectra in Fig. 3. The general trend of increasing mass with increasing RH is reflected in the increase in OH stretch signal in the Raman study (Fig. 4), hence the Raman spectra and EDB results are in qualitative agreement.

4 Summary and conclusions

Our laboratory experiments show that iodic acid solution droplets have a very low propensity to crystallise and therefore probably would not do so in the atmosphere. This inhibition of efflorescence is most likely related to their high viscosity at low humidity. Physically probing the droplets and measurements of slow loss of water over many hours indicate that aqueous iodic acid droplets exist in a semi-solid or glassy state below ~ 15 % RH.

In addition, we also show that the hygroscopic growth of iodic acid solution droplets is very small, which is surprising considering HIO_3 and I_2O_5 are highly soluble. Such

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droplets grow by only 20% between 20 and 90% RH and we determine the hygroscopicity parameter (κ), as defined by Petters et al. (2007), for iodine acid droplets is 0.024. This value is within the range commonly associated with weakly soluble organic materials and secondary organic aerosol particles (Petters and Kreidenweis, 2007).

Hygroscopicity parameter values for highly soluble inorganic species normally fall in the range of 0.5 to 1.4. Iodine oxide particles in the marine boundary layer and in the laboratory also have very small growth factors. This has led previous authors to conclude that iodine oxide particles must be insoluble (Jimenez et al., 2003; Väkevä et al., 2002) and Jimenez et al. (2003) therefore went on to suggest the composition of iodine oxide particles was the insoluble I_2O_4 rather than the highly soluble I_2O_5 or HIO_3 . Here we have shown that the very small hygroscopic growth factors of nucleation mode particles observed in the coastal marine boundary layer (Väkevä et al., 2002) are consistent with these aerosol being composed of aqueous iodine acid.

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- Angell, C. A.: Liquid fragility and the glass transition in water and aqueous solutions, *Chem. Rev.*, 102, 2627–2650, 2002.
- 5 Bodsworth, A., Zobrist, B., and Bertram, A. K.: Inhibition of efflorescence in mixed organic-inorganic particles at temperatures less than 250 K, *Phys. Chem. Chem. Phys.*, 12, 12259–12266, doi:10.1039/c0cp00572j, 2010.
- Broadley, S. L., Murray, B. J., Herbert, R. J., Atkinson, J. D., Dobbie, S., Malkin, T. L., Condliffe, E., and Neve, L.: Immersion mode heterogeneous ice nucleation by an illite rich powder representative of atmospheric mineral dust, *Atmos. Chem. Phys.*, 12, 287–307, doi:10.5194/acp-12-287-2012, 2012.
- 10 Chan, C. K., Kwok, C. S., and Chow, A. H. L.: Study of hygroscopic properties of aqueous mixtures of disodium fluorescein and sodium chloride using an electrodynamic balance, *Pharmaceutical Research*, 14, 1171–1175, doi:10.1023/a:1012146621821, 1997.
- Chen, L. Y., Jeng, F. T., Chen, C. C., and Hsiao, T. C.: Hygroscopic behavior of atmospheric aerosol in taipei, *Atmos. Environ.*, 37, 2069–2075, doi:10.1016/S1352-2310(03)00071-2, 2003.
- 15 Cohen, M. D., Flagan, R. C., and Seinfeld, J. H.: Studies of concentrated electrolyte-solutions using the electrodynamic balance .1. Water activities for single-electrolyte solutions, *J. Phys. Chem.*, 91, 4563–4574, doi:10.1021/j100301a029, 1987.
- 20 Daehlie, G., and Kjekshus, A.: Iodine oxides:Part i. On $i_2O_3 \cdot SO_3$, $i_2O_3 \cdot 4SO_3 \cdot H_2O$, $i_2O_3 \cdot SeO_3$ and i_2O_4 , *Acta Chemica Scandinavica*, 18, 144–156, 1964.
- Davies, J. F., Haddrell, A. E., and Reid, J. P.: Time-resolved measurements of the evaporation of volatile components from single aerosol droplets, *Aerosol Sci. Technol.*, 46, 666–677, doi:10.1080/02786826.2011.652750, 2012.
- 25 Durig, J. R., Bonner, O. D., and Breazeale, W. H.: Raman studies of iodic acid and sodium iodate, *J. Phys. Chem.*, 69, 3886–3892, doi:10.1021/j100895a041, 1965.
- Dymarska, M., Murray, B. J., Sun, L. M., Eastwood, M. L., Knopf, D. A., and Bertram, A. K.: Deposition ice nucleation on soot at temperatures relevant for the lower troposphere, *J. Geophys. Res.-Atmos.*, 111, D04204, doi:10.1029/2005JD006627, 2006.
- 30 Edgeworth, R., Dalton, B. J., and Parnell, T.: The pitch drop experiment, *Eur. J. Phys.*, 5, 198–200, 1984.

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- Giguere, P. A. and Guillot, J. G.: Raman spectrum of hydronium (H_3O^+) ions in aqueous acids, *J. Phys. Chem.*, 86, 3231–3233, doi:10.1021/j100213a031, 1982.
- Glantschnig, W. J. and Chen, S.-H.: Light scattering from water droplets in the geometrical optics approximation, *Appl. Opt.*, 20, 2499–2509, 1981.
- 5 Goldman, S., Bates, R. G., and Robinson, R. A.: Osmotic coefficients and activity coefficients of iodic acid at high concentrations, *J. Solut. Chem.*, 3, 593–602, 1974.
- Hargreaves, G., Kwamena, N. O. A., Zhang, Y. H., Butler, J. R., Rushworth, S., Clegg, S. L., and Reid, J. P.: Measurements of the equilibrium size of supersaturated aqueous sodium chloride droplets at low relative humidity using aerosol optical tweezers and an electrodynamic balance, *J. Phys. Chem. A*, 114, 1806–1815, doi:10.1021/jp9095985, 2010.
- 10 Ivanov, A. A., Kirilenko, I. A., Azarova, L. A., and Vinogradov, E. E.: The electrical conductivities of solutions of the glass forming system $\text{HIO}_3\text{-H}_2\text{O}$, *Russian Journal of Inorganic Chemistry*, 29, 1209–1211, 1984.
- Ivanov, A. A., Kirilenko, I. A., Azarova, L. A., and Vinogradov, E. E.: Properties and structure of solutions in the $\text{HIO}_3\text{-H}_2\text{O}$ glass forming system, *Russian Journal of Inorganic Chemistry*, 30, 603–605, 1985.
- 15 Jimenez, J. L., Bahreini, R., Cocker, D. R., Zhuang, H., Varutbangkul, V., Flagan, R. C., Seinfeld, J. H., O'Dowd, C. D., and Hoffmann, T.: New particle formation from photooxidation of diiodomethane (CH_2I_2), *J. Geophys. Res.-Atmos.*, 108, 4318, doi:10.1029/2002jd002452, 2003.
- 20 Knopf, D. A. and Lopez, M. D.: Homogeneous ice freezing temperatures and ice nucleation rates of aqueous ammonium sulfate and aqueous levoglucosan particles for relevant atmospheric conditions, *Phys. Chem. Chem. Phys.*, 11, 8056–8068, doi:10.1039/b903750k, 2009.
- 25 Koop, T., Bookhold, J., Shiraiwa, M., and Pöschl, U.: Glass transition and phase state of organic compounds: Dependency on molecular properties and implications for secondary organic aerosols in the atmosphere, *Phys. Chem. Chem. Phys.*, 13, 19238–19255, doi:10.1039/c1cp22617g, 2011.
- 30 Kumar, R., Saunders, R. W., Mahajan, A. S., Plane, J. M. C., and Murray, B. J.: Physical properties of iodate solutions and the deliquescence of crystalline I_2O_5 and HIO_3 , *Atmos. Chem. Phys.*, 10, 12251–12260, doi:10.5194/acp-10-12251-2010, 2010.

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Li, K.-K., Wang, F., Zeng, G., Reid, J. P., and Zhang, Y.-H.: Probing the time scale for bulk equilibration and mass transport of water in amorphous inorganic aerosol, *J. Phys. Chem. B*, 115, 14397–14403, doi:10.1021/jp209379f, 2011.

Lide, D. R.: *Crc handbook of chemistry and physics*, 90th ed., CRC, Boca Raton, Fla., London, 2009.

Mahajan, A. S., Plane, J. M. C., Oetjen, H., Mendes, L., Saunders, R. W., Saiz-Lopez, A., Jones, C. E., Carpenter, L. J., and McFiggans, G. B.: Measurement and modelling of tropospheric reactive halogen species over the tropical Atlantic Ocean, *Atmos. Chem. Phys.*, 10, 4611–4624, doi:10.5194/acp-10-4611-2010, 2010.

Mäkelä, J. M., Hoffmann, T., Holzke, C., Vakeva, M., Suni, T., Mattila, T., Aalto, P. P., Tapper, U., Kauppinen, E. I., and O'Dowd, C. D.: Biogenic iodine emissions and identification of end-products in coastal ultrafine particles during nucleation bursts, *J. Geophys. Res.-Atmos.*, 107, 8110, doi:10.1029/2001jd000580, 2002.

Martin, S. T.: Phase transitions of aqueous atmospheric particles, *Chem. Rev.*, 100, 3403–3453, doi:10.1021/cr990034t, 2000.

McFiggans, G., Coe, H., Burgess, R., Allan, J., Cubison, M., Alfarra, M. R., Saunders, R., Saiz-Lopez, A., Plane, J. M. C., Wevill, D., Carpenter, L., Rickard, A. R., and Monks, P. S.: Direct evidence for coastal iodine particles from *Laminaria* macroalgae – linkage to emissions of molecular iodine, *Atmos. Chem. Phys.*, 4, 701–713, doi:10.5194/acp-4-701-2004, 2004.

McFiggans, G., Bale, C. S. E., Ball, S. M., Beames, J. M., Bloss, W. J., Carpenter, L. J., Dorsey, J., Dunk, R., Flynn, M. J., Furneaux, K. L., Gallagher, M. W., Heard, D. E., Hollingsworth, A. M., Hornsby, K., Ingham, T., Jones, C. E., Jones, R. L., Kramer, L. J., Langridge, J. M., Leblanc, C., LeCrane, J.-P., Lee, J. D., Leigh, R. J., Longley, I., Mahajan, A. S., Monks, P. S., Oetjen, H., Orr-Ewing, A. J., Plane, J. M. C., Potin, P., Shillings, A. J. L., Thomas, F., von Glasow, R., Wada, R., Whalley, L. K., and Whitehead, J. D.: Iodine-mediated coastal particle formation: an overview of the Reactive Halogens in the Marine Boundary Layer (RHAMBLE) Roscoff coastal study, *Atmos. Chem. Phys.*, 10, 2975–2999, doi:10.5194/acp-10-2975-2010, 2010.

Mikhailov, E., Vlasenko, S., Martin, S. T., Koop, T., and Pöschl, U.: Amorphous and crystalline aerosol particles interacting with water vapor: conceptual framework and experimental evidence for restructuring, phase transitions and kinetic limitations, *Atmos. Chem. Phys.*, 9, 9491–9522, doi:10.5194/acp-9-9491-2009, 2009.

Hygroscopic growth of iodic acid solution droplets

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- Murphy, D. M. and Koop, T.: Review of the vapour pressures of ice and supercooled water for atmospheric applications, *Q. J. Roy. Meteorol. Soc.*, 131, 1539–1565, 2005.
- Murray, B. J.: Inhibition of ice crystallisation in highly viscous aqueous organic acid droplets, *Atmospheric Chemistry and Physics*, 8, 5423–5433, 2008.
- 5 Murray, B. J. and Bertram, A. K.: Strong dependence of cubic ice formation on droplet ammonium to sulfate ratio, *Geophys. Res. Lett.*, 34, L16810, doi:10.1029/2007GL030471, 2007.
- Murray, B. J. and Bertram, A. K.: Inhibition of solute crystallisation in aqueous $\text{h}^+ \text{-nh}_4^+ \text{-so}_4^{2-} \text{-h}_2\text{o}$ droplets, *Phys. Chem. Chem. Phys.*, 10, 3287–3301, doi:10.1039/b802216j, 2008.
- Murray, B. J., Broadley, S. L., Wilson, T. W., Bull, S. J., Wills, R. H., Christenson, H. K., and
10 Murray, E. J.: Kinetics of the homogeneous freezing of water, *Phys. Chem. Chem. Phys.*, 12, 10380–10387, 10.1039/c003297b, 2010a.
- Murray, B. J., Wilson, T. W., Dobbie, S., Cui, Z., Al-Jumur, S. M. R. K., Mohler, O., Schnaiter, M., Wagner, R., Benz, S., Niemand, M., Saathoff, H., Ebert, V., Wagner, S., and Karcher, B.: Heterogeneous nucleation of ice particles on glassy aerosols under cirrus conditions, *Nature Geosci.*, 3, 233–237, 2010b.
- 15 Murray, B. J., Broadley, S. L., and Morris, G. J.: Supercooling of water droplets in jet aviation fuel, *Fuel*, 90, 433–435, 2011a.
- Murray, B. J., Broadley, S. L., Wilson, T. W., Atkinson, J. D., and Wills, R. H.: Heterogeneous freezing of water droplets containing kaolinite particles, *Atmos. Chem. Phys.*, 11, 4191–4207, doi:10.5194/acp-11-4191-2011, 2011b.
- 20 O'Dowd, C. D. and De Leeuw, G.: Marine aerosol production: A review of the current knowledge, *Philos. T. Roy. Soc. A*, 365, 1753–1774, doi:10.1098/rsta.2007.2043, 2007.
- O'Dowd, C. D. and Hoffmann, T.: Coastal new particle formation: A review of the current state-of-the-art, *Environ. Chem.*, 2, 245–255, doi:10.1071/en05077, 2005.
- 25 Onasch, T. B., Siefert, R. L., Brooks, S. D., Prenni, A. J., Murray, B., Wilson, M. A., and Tolbert, M. A.: Infrared spectroscopic study of the deliquescence and efflorescence of ammonium sulfate aerosol as a function of temperature, *J. Geophys. Res.-Atmos.*, 104, 21317–21326, 1999.
- Pant, A., Parsons, M. T., and Bertram, A. K.: Crystallization of aqueous ammonium sulfate particles internally mixed with soot and kaolinite: Crystallization relative humidities and nucleation rates, *J. Phys. Chem. A*, 110, 8701–8709, doi:10.1021/jp060985s, 2006.
- 30

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- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7, 1961–1971, doi:10.5194/acp-7-1961-2007, 2007.
- Pope, F. D., Tong, H.-J., Dennis-Smith, B. J., Griffiths, P. T., Clegg, S. L., Reid, J. P., and Cox, R. A.: Studies of single aerosol particles containing malonic acid, glutaric acid, and their mixtures with sodium chloride. II. Liquid-state vapor pressures of the acids, *J. Phys. Chem. A*, 114, 10156–10165, doi:10.1021/jp1052979, 2010.
- Reid, J. P., Meresman, H., Mitchem, L., and Symes, R.: Spectroscopic studies of the size and composition of single aerosol droplets, *Int. Rev. Phys. Chem.*, 26, 139–192, doi:10.1080/01442350601081899, 2007.
- Saiz-Lopez, A., Plane, J. M. C., Baker, A. R., Carpenter, L. J., von Glasow, R., Gomez Martín, J. C., McFiggans, G., and Saunders, R. W.: Atmospheric chemistry of iodine, *Chem. Rev.*, 112, 1773–1804, doi:10.1021/cr200029u, 2011.
- Saunders, R. W. and Plane, J. M. C.: Formation pathways and composition of iodine oxide ultra-fine particles, *Environ. Chem.*, 2, 299–303, doi:10.1071/en05079, 2005.
- Saunders, R. W. and Plane, J. M. C.: Fractal growth modelling of i_2O_5 nanoparticles, *J. Aerosol Sci.*, 37, 1737–1749, doi:10.1016/j.jaerosci.2006.08.007, 2006.
- Saunders, R. W., Kumar, R., Gomez Martin, J. C., Mahajan, A. S., Murray, B. J., and Plane, J. M. C.: Studies of the formation and growth of aerosol from molecular iodine precursor, *Zeitschrift Fur Physikalische Chemie-International Journal of Research in Physical Chemistry & Chemical Physics*, 224, 1095–1117, doi:10.1524/zpch.2010.6143, 2010.
- Shiraiwa, M., Ammann, M., Koop, T., and Pöschl, U.: Gas uptake and chemical aging of semisolid organic aerosol particles, *P. Natl. Acad. Sci. USA*, 108, 11003–11008, doi:10.1073/pnas.1103045108, 2011.
- Swietlicki, E., Hansson, H. C., Hameri, K., Svenningsson, B., Massling, A., McFiggans, G., McMurry, P. H., Petaja, T., Tunved, P., Gysel, M., Topping, D., Weingartner, E., Baltensperger, U., Rissler, J., Wiedensohler, A., and Kulmala, M.: Hygroscopic properties of submicrometer atmospheric aerosol particles measured with h-tdma instruments in various environments – a review, *Tellus B*, 60, 432–469, doi:10.1111/j.1600-0889.2008.00350.x, 2008.
- Tang, I. N., Tridico, A. C., and Fung, K. H.: Thermodynamic and optical properties of sea salt aerosols, *J. Geophys. Res.*, 102, 23269–23277, 1997.
- Tong, H.-J., Reid, J. P., Bones, D. L., Luo, B. P., and Krieger, U. K.: Measurements of the timescales for the mass transfer of water in glassy aerosol at low relative humidity and

ambient temperature, *Atmos. Chem. Phys.*, 11, 4739–4754, doi:10.5194/acp-11-4739-2011, 2011.

Väkevää, M., Hämeri, K., and Aalto, P. P.: Hygroscopic properties of nucleation mode and aiten mode particles during nucleation bursts and in background air on the west coast of ireland, *J. Geophys. Res.-Atmos.*, 107, 8104, doi:10.1029/2000jd000176, 2002.

Virtanen, A., Joutsensaari, J., Koop, T., Kannosto, J., Yli-Pirila, P., Leskinen, J., Makela, J. M., Holopainen, J. K., Poschl, U., Kulmala, M., Worsnop, D. R., and Laaksonen, A.: An amorphous solid state of biogenic secondary organic aerosol particles, *Nature*, 467, 824–827, 2010.

Zhang, Y. H. and Chan, C. K.: Study of contact ion pairs of supersaturated magnesium sulfate solutions using raman scattering of levitated single droplets, *J. Phys. Chem. A*, 104, 9191–9196, doi:10.1021/jp0013659, 2000.

Zobrist, B., Marcolli, C., Pedernera, D. A., and Koop, T.: Do atmospheric aerosols form glasses?, *Atmos. Chem. Phys.*, 8, 5221–5244, doi:10.5194/acp-8-5221-2008, 2008.

Zobrist, B., Soonsin, V., Luo, B. P., Krieger, U. K., Marcolli, C., Peter, T., and Koop, T.: Ultra-slow water diffusion in aqueous sucrose glasses, *Phys. Chem. Chem. Phys.*, 13, 3514–3526, doi:10.1039/c0cp01273d, 2011.

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12, 7879–7908, 2012

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Table 1. Assignment of bands in Raman spectra of aqueous HIO_3 (cm^{-1}).

This Work	12 M Aqueous Solution (Durig et al., 1965)	Description
814	814	
768	771	IO_2 (HO· IO_2) assym. stretch
–	737	
644	638	IO stretch
333	324	IO_2 deformation

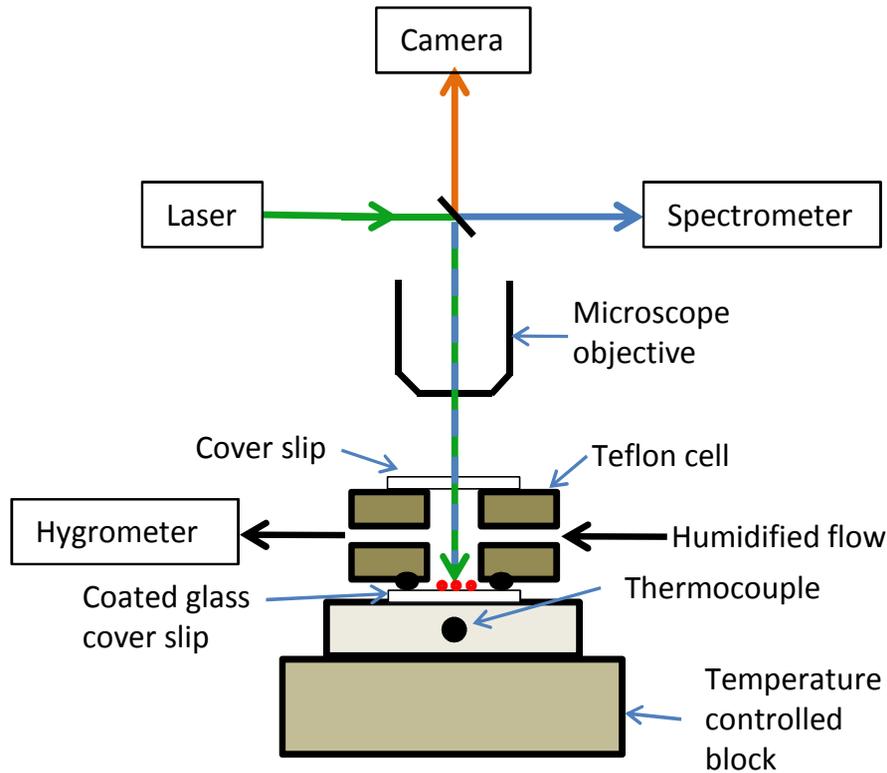


Fig. 1. Diagram of the humidity controlled stage coupled to a Raman microscope system used in this study. Not to scale.

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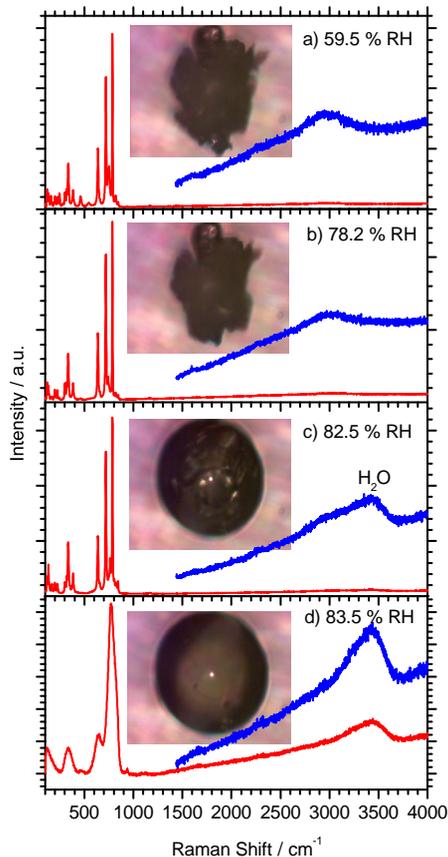


Fig. 2. Raman spectra of an initially crystalline HIO_3 particle on increasing RH. The blue spectra are the region above $\sim 1500 \text{ cm}^{-1}$ expanded for clarity. Deliquescence occurs around 82.5% RH (**c–d**) and is clearly evident in changes in both the image of the droplets and the Raman spectra. The width of the microscope images is $160 \mu\text{m}$.

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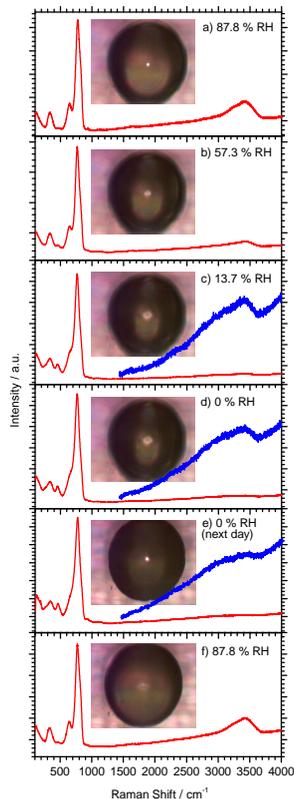


Fig. 3. Raman spectra of deliquesced HIO_3 on decreasing RH (**a–d**) and then increasing RH again on the following day (**e, f**) after leaving the droplets at 0% RH overnight. The spectra are for the same droplet as in Figure 2, but tests demonstrated that all droplets (~ 30 droplets) on the slide responded in a similar manner. The OH stretch band associated with water ($\sim 3400 \text{ cm}^{-1}$) decreases dramatically on reducing RH, but the droplet does not revert back to its crystalline form. The blue expanded spectra are included only in cases where the OH band is weak.

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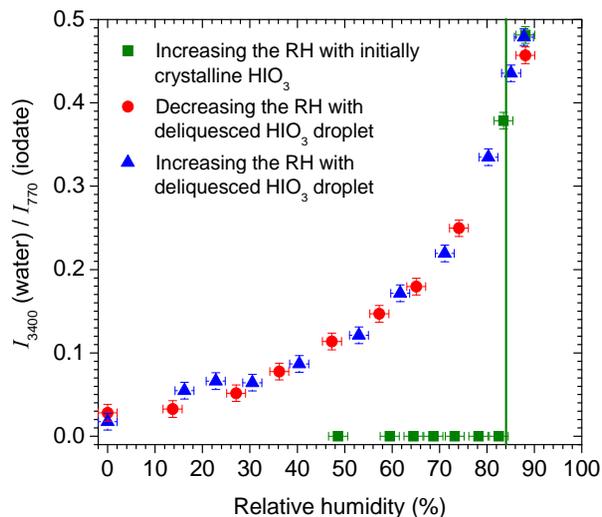


Fig. 4. Intensity ratio of peaks roughly centred at 3400 cm^{-1} (I_{3400} , OH stretch, i.e. water), and 770 cm^{-1} (I_{770} , iodate) as a function of RH determined from spectra such as those shown in Fig. 2 and Fig. 3. The vertical green line indicates the expected deliquescence point (Kumar et al., 2010) and is in good agreement with the RH at which HIO₃ clearly takes up water to become an aqueous liquid droplet. The estimated uncertainty on our RH measurement is 2%. The intensities I_{3400} and I_{770} were determined by integrating the background subtracted signal between $2900\text{--}3800\text{ cm}^{-1}$ and $500\text{--}1000\text{ cm}^{-1}$, respectively, for the liquid droplets. A value of zero was taken when the spectra were dominated by crystalline peaks (up to 82.5% RH). All these measurements were done with a single particle/droplet as the RH was initially increased around a dry crystalline particle until it had fully deliquesced (green), and then as the RH was reduced to 0% RH (red), and then on the following day as it was increased again (blue).

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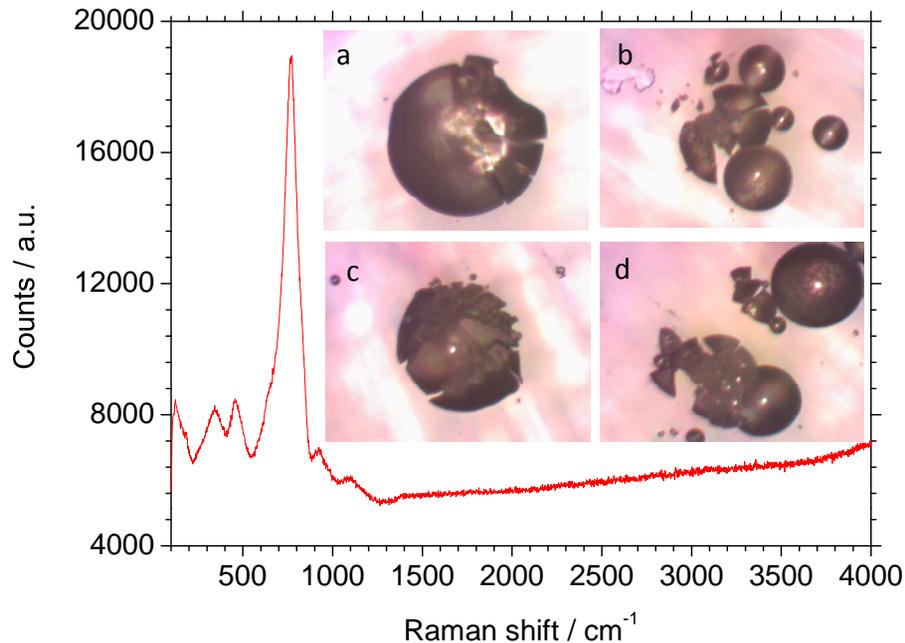


Fig. 5. Raman spectrum and images of “cracked” aqueous HIO₃ droplets at 0% RH. For this test, deliquesced droplets were dried out to 0% RH gradually over the course of ~1.5 h and were then subjected to mechanical force using the flat end of a metal rod. The droplets shattered in much the same way as one would expect an ultra viscous liquid, such as pitch, or a brittle glass to shatter. The spectrum corresponds to **(b)**.

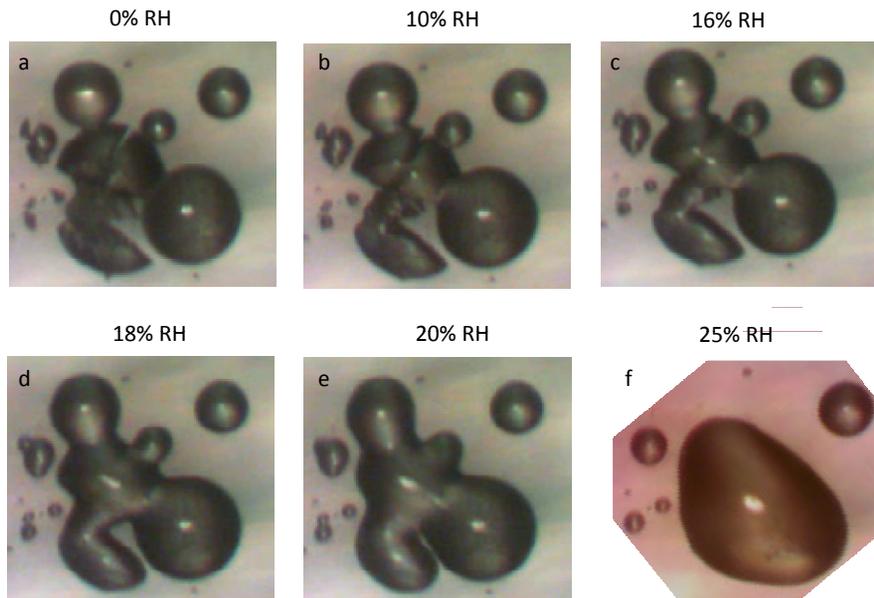


Fig. 6. Sequence of images of a cracked droplet gradually taking up water and flowing to reform a single droplet on increasing RH. Image **(f)** was taken with a different camera, hence the difference in orientation.

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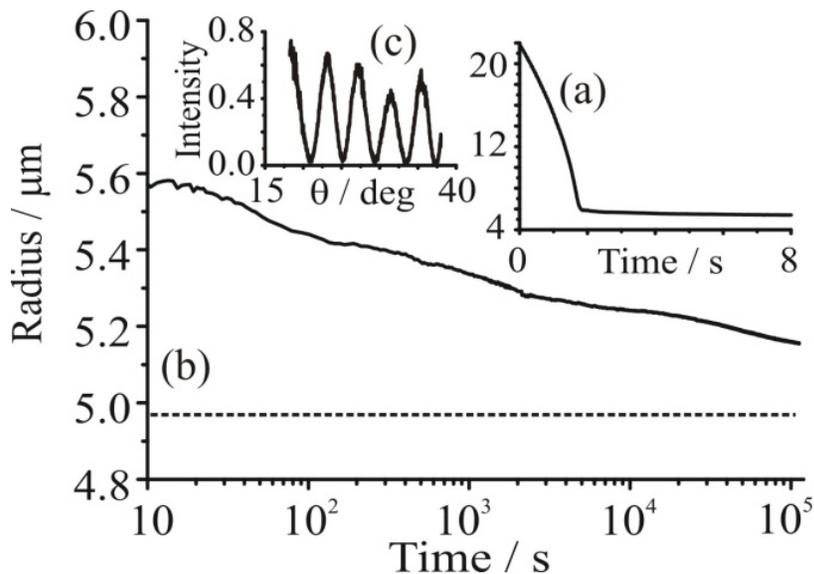


Fig. 7. (a) Initial loss of water on injection of a dilute aqueous HIO_3 droplet into a dry atmosphere within the concentric cylindrical electrodynamic balance. The gray dashed horizontal line indicates the estimated dry size for a crystal of HIO_3 ($4.95 \pm 0.15 \mu\text{m}$). (b) indicates the size decrease over ~ 33 h. (c) is a representative phase function observed after 70 000 s indicating the droplet remained approximately spherical.

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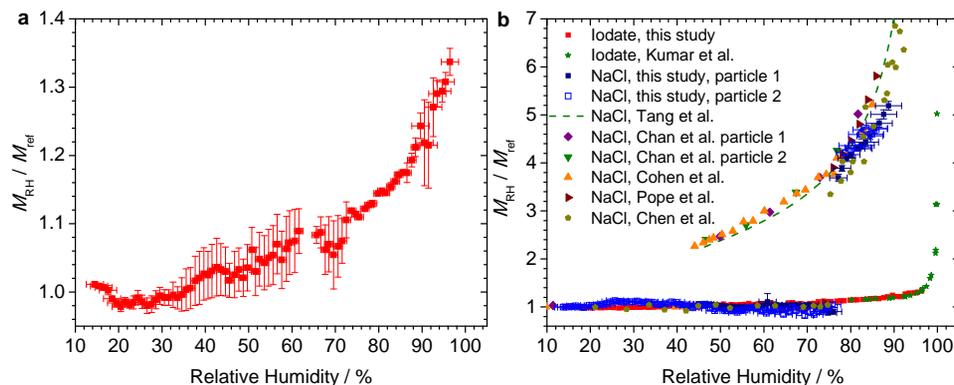


Fig. 8. Relative change in mass of droplets suspended in the electrodynamic balance on varying RH. **(a)** The relative change in mass of iodic acid solution droplets and **(b)** comparison of the iodic acid solution data with the growth of sodium chloride. Our measurements of NaCl hygroscopic growth are compared to literature data (Chan et al., 1997; Cohen et al., 1987; Pope et al., 2010; Chen et al., 2003; Tang et al., 1997). Water activity literature data for bulk iodic acid solutions of known composition (Kumar et al., 2010) are referenced to the growth factor at 83.81 % RH.

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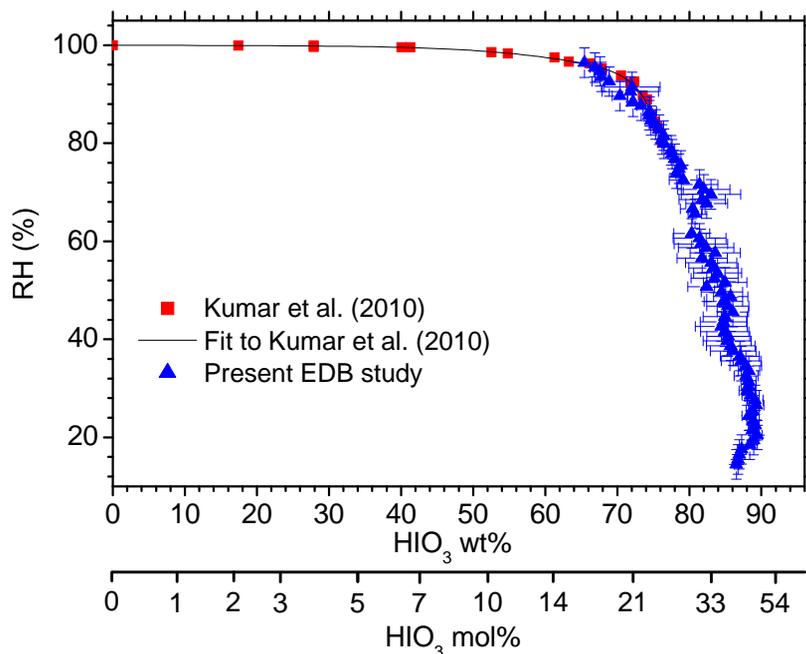


Fig. 9. Variation of solute mass concentration as a function of RH. Direct measurements of water activity of solutions of known composition were taken from Kumar et al. (2010) (red points and line) and the data for lower water activities is determined using the measured growth factor. The electrodynamic balance data is referenced to the water activity data at 83.81 % RH.

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