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Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009)

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Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Abstract

New burned area datasets and top-down constraints from atmospheric concentration measurements of pyrogenic gases have decreased the large uncertainty in fire emissions estimates. However, significant gaps remain in our understanding of the contribution of deforestation, savanna, forest, agricultural waste, and peat fires to total global fire emissions. Here we used a revised version of the Carnegie-Ames-Stanford-Approach (CASA) biogeochemical model and improved satellite-derived estimates of area burned, fire activity, and plant productivity to calculate fire emissions for the 1997–2009 period on a 0.5° spatial resolution with a monthly time step. For November 2000 onwards, estimates were based on burned area, active fire detections, and plant productivity from the Moderate Resolution Imaging Spectroradiometer (MODIS) sensor. For the partitioning we focused on the MODIS era. We used burned area estimates based on Tropical Rainfall Measuring Mission (TRMM) Visible and Infrared Scanner (VIRS) and Along-Track Scanning Radiometer (ATSR) active fire data prior to MODIS (1997–2000) and Advanced Very High Resolution Radiometer (AVHRR) derived estimates of plant productivity during the same period. Average global fire carbon emissions were 2.0 Pg yr^{-1} with significant interannual variability during 1997–2001 (2.8 Pg yr^{-1} in 1998 and 1.6 Pg yr^{-1} in 2001). Emissions during 2002–2007 were relatively constant (around 2.1 Pg yr^{-1}) before declining in 2008 (1.7 Pg yr^{-1}) and 2009 (1.5 Pg yr^{-1}) partly due to lower deforestation fire emissions in South America and tropical Asia. During 2002–2007, emissions were highly variable from year-to-year in many regions, including in boreal Asia, South America, and Indonesia, but these regional differences cancelled out at a global level. During the MODIS era (2001–2009), most fire carbon emissions were from fires in grasslands and savannas (44%) with smaller contributions from tropical deforestation and degradation fires (20%), woodland fires (mostly confined to the tropics, 16%), forest fires (mostly in the extratropics, 15%), agricultural waste burning (3%), and tropical peat fires (3%). The contribution from agricultural waste fires was likely a lower bound because our approach for measuring

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



burned area could not detect all of these relatively small fires. For reduced trace gases such as CO and CH₄, deforestation, degradation, and peat fires were more important contributors because of higher emissions of reduced trace gases per unit carbon combusted compared to savanna fires. Carbon emissions from tropical deforestation, degradation, and peatland fires were on average 0.5 Pg C yr⁻¹. The carbon emissions from these fires may not be balanced by regrowth following fire. Our results provide the first global assessment of the contribution of different sources to total global fire emissions for the past decade, and supply the community with an improved 13-year fire emissions time series.

1 Introduction

Over the last decade, the role of fire in shaping the environment and atmosphere has been increasingly appreciated (e.g., Langmann et al., 2009; Bowman et al., 2009). Fire is one of the most important disturbance agents in terrestrial ecosystems on a global scale and is widely used by humans to manage and transform land for many purposes, especially in tropical and subtropical ecosystems. Fires contribute significantly to the budgets of several trace gases and aerosols (Andreae and Merlet, 2001) and are one of the primary causes of interannual variability in the growth rate of several trace gases, including the greenhouse gases CO₂ and CH₄ (Langenfelds et al., 2002).

In many regions, pre-industrial levels of fire activity may have been comparable to or even higher than contemporary levels (Pyne, 1982; Marlon et al., 2008). In deforestation regions, however, humans are known to have increased fire activity (Fearnside, 2005; Schultz et al., 2008; Field et al., 2009). Fire activity has also increased in more remote regions due to humans (e.g., Mollicone et al., 2006). In addition, climate change may lead to more frequent and intense fires if drought conditions in areas with abundant fuel loads become more severe (Kasischke et al., 1995; Westerling et al., 2006).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[I◀](#)[▶I](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

To understand how fires influence and interact with the Earth system, quantitative information on emissions and a breakdown of emissions into different sources is required. This breakdown of emissions is especially important to quantify the extent to which fires contribute to the build-up of atmospheric CO₂ since only deforestation fires, fires in drained peatlands, and fires from other areas that have had increasing levels of disturbance are a net source of CO₂ to the atmosphere. In many other areas, CO₂ emissions from fires are balanced by carbon uptake during regrowth on decadal timescales. For other trace gases and aerosols, this distinction is less important but a breakdown into categories is useful to better characterize non-CO₂ anthropogenic climate forcing and to understand how human activities are affecting atmospheric chemistry. For example, all fires contribute to emissions of methane (CH₄), but the amount of CH₄ released per unit biomass combusted varies greatly between different fire types. Peat fires, for example, may emit almost ten times more CH₄ per unit biomass combusted than fires in savannas (Yokelson et al., 1997; Andreae and Merlet, 2001; Christian et al., 2003).

Seiler and Crutzen (1980) made the first global estimates of fire emissions, which subsequently have been refined and updated (e.g., Crutzen and Andreae, 1990; Galanter et al., 2000; Andreae and Merlet, 2001 based on unpublished data from Yevich). Hao et al. (1996) used climatological information to better understand the temporal distribution of emissions, while Schultz et al. (2002) and Duncan et al. (2003) improved the understanding of the spatial and temporal distribution of fires, as well as their interannual variability using satellite information on fire activity (ATSR) and/or aerosol optical depths from the Total Ozone Mapping Spectrometer (TOMS). In addition to improving contemporary estimates, long-term time series during the 20th century have been constructed, primarily with the aim of understanding changes in ecosystems, the carbon cycle, and atmospheric chemistry (Mouillot et al., 2006; Schultz et al., 2008; Mieville et al., 2010).

The early approaches to estimate global fire emissions (Seiler and Crutzen, 1980; Crutzen and Andreae, 1990) de facto estimated the contribution from different sources,

because estimates were based on biome-averaged fuel loads and fire frequency (savanna and forest fires) and estimates of per-capita clearing rates in combination with population densities to estimate deforestation and shifting agriculture emissions. In these studies a clear distinction was made between deforestation fires where forest is removed permanently and shifting agriculture where the clearing of forest is followed by a few years of production, after which the forest is allowed to regrow. If we exclude wood burning (which is not assessed in this study) then emissions estimates from Seiler and Crutzen (1980) were 2.6 Pg C yr^{-1} (range of 1.7–3.5), with agricultural waste burning estimated to be the largest source of fire carbon emissions (33%), followed by shifting agriculture (29%), savanna (21%), deforestation (12%), fires in temperate areas (4%) and fires in boreal areas (1%). For the tropics, Crutzen and Andreae (1990) later revised the emissions estimates for deforestation and savanna fires upwards, with savanna burning becoming the main source of emissions.

More recently, global satellite-derived burned area information has become available (Grégoire et al., 2003; Simon et al., 2004; Giglio et al., 2006). These datasets have been used in combination with biogeochemical or dedicated fuel load models to estimate emissions (Hoelzemann et al., 2004; Ito and Penner, 2004; van der Werf et al., 2006; Jain et al., 2006). These studies point towards a fire carbon loss estimate between $1\text{--}3 \text{ Pg C yr}^{-1}$ (excluding biomass burned for domestic purposes such as cooking) with considerable uncertainty and large interannual variability (Rander-son et al., 2005). The partitioning between different sources drew less attention in these new global studies because it did not lie at the core of the calculations, although the very different fuel consumption estimates for different sources were accounted for. Other studies have focused more on a particular sector; Yevich and Logan (2003), for example, calculated emissions from the burning of biofuels and agricultural waste and suggested that the latter source was substantially smaller ($\sim 0.2 \text{ Pg C yr}^{-1}$) than in earlier estimates.

Besides improvements in quantifying global fire emissions, our understanding of the multifaceted role of fire in the Earth system is improving from studies quantifying the

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



different ways fires influence climate. Randerson et al. (2006), for example, showed that the impact of (increasing) boreal forest fires on climate warming may be limited or even result in regional cooling because the negative forcing from increased surface albedo following a fire due to higher snow exposure offsets positive forcings from greenhouse gas emissions and the deposition of black carbon on snow. In the tropics, aerosol emissions from fires have been shown to influence the radiation budget at regional scales (Duncan et al., 2003). Climate modeling studies suggest these aerosols may lengthen or intensify periods of drought in the Amazon (Zhang et al., 2008) and in Indonesia (Tosca et al., 2010). At a global scale, changing levels of fire emissions influence 8 out of the 13 radiative forcing terms identified in the IPCC 4th Assessment (Bowman et al., 2009).

Improvements in emissions estimates are also necessary to calibrate and/or validate prognostic fire modules in dynamic global vegetation models and climate-carbon models for the period they overlap with satellite observations to make better predictions about future fire activity (e.g., Thonicke et al., 2001; Arora and Boer, 2005; Kloster et al., 2010). These models can also take advantage of a better understanding of the drivers of fires based on new satellite information. Archibald et al. (2009), for example, showed that tree cover density, rainfall over the last 2 years, and rainfall seasonality explained more than half of the variability in burned area in Southern Africa. In addition, interannual variability in precipitation rates controls part of the variability in fire-driven deforestation rates from year to year, with strongest relations in Equatorial Asia where annual variability in precipitation is highest (Le Page et al., 2008; van der Werf et al., 2008).

During the last years, several new burned area datasets have been developed at 500 m or 1 km resolution (Roy et al., 2008; Plummer et al., 2006; Tansey et al., 2007; Giglio et al., 2009). Comparisons with Landsat-based burned area have reduced uncertainties, but large differences persist between the different approaches (Roy et al., 2009; Giglio et al., 2010). Ideally, these moderate resolution burned area datasets would be combined with fuel-load modeling at the same resolution to improve estimates

of emissions. This would provide an added benefit from the perspective of validating fuel loads and fuel consumption with ground measurements. However, with the exception of Ito and Penner (2004), who built a dedicated 1-km fuel model, most global modeling frameworks are based on global biogeochemical models that were developed at coarser spatial resolutions because of conceptual and data constraints. Many of these models, for example, also are used to estimate the net carbon balance of terrestrial ecosystems. This requires additional model complexity, including tracking the flow of carbon through multi-decadal vegetation, litter, and soil carbon pools, representing the age dynamics of different forest types, and capturing climate change effects on primary production and ecosystem respiration. On regional scales, fire-dedicated research has employed native resolution satellite data; see for example Hely et al. (2007) for savanna regions in Southern Africa. Working also at relatively high resolution (1 km), Wiedinmyer et al. (2006) combined data on fire activity, land cover, and literature-derived fuel load and combustion completeness to estimate emissions for North America, while more recently Chang and Song (2010) combined MODIS burned area with fuel statistics and combustion completeness based on the Normalized Difference Vegetation Index (NDVI) to estimate emissions for tropical Asia at 500 m resolution. For the global scale, however, combining native resolution burned area and a coarser resolution biogeochemical model for fuel may provide a useful interim solution until these biogeochemical models can run at the native resolution of the satellite data. As an example, Lehsten et al. (2009) used 1 km burned area to drive a 1° dynamic vegetation model with 100 subgrid elements in Africa to account for stochastic processes.

Here we used new burned area estimates and an improved biogeochemical model at 0.5° spatial resolution and a monthly time step to investigate global patterns of fire emissions. Our main objectives were to refine the Global Fire Emissions Database (GFED) and to quantify the relative contributions of deforestation, savanna, forest, agricultural, and peat fires to global contemporary fire emissions. We used MODIS data on burned area and active fires (Giglio et al., 2010), land cover characteristics (Friedl

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

et al., 2002; Hansen et al., 2003), and plant productivity (Myneni et al., 2002) to study fires over the November 2000–2009 period. Information from other sensors (TRMM-VIRS and ATSR) for burned area (Giglio et al., 2010), and AVHRR for plant productivity (Tucker et al., 2005) was used to extend our estimates back in time starting in 1997.

2 Methods and datasets

2.1 Introduction

The work presented here builds on our earlier work in which we combined information on fire activity (Giglio et al., 2003; Giglio et al., 2006) with global biogeochemical modeling (van der Werf et al., 2003; Randerson et al., 2005; van der Werf et al., 2006). The model we used was originally derived from the satellite-driven Carnegie Ames Stanford Approach (CASA) model (Potter et al., 1993; Field et al., 1995; Randerson et al., 1996). In this methods and datasets section we start with a brief overview of the model structure including minor modifications (Sect. 2.2), describe the major input datasets used to drive the model (Sect. 2.3), explain the major changes we made to the model (Sect. 2.4), and we conclude with a description of our approach for assessing uncertainties (Sect. 2.5).

2.2 Modeling overview

CASA calculates carbon “pools” for each grid cell and time step based on carbon input from net primary productivity (NPP) and carbon emissions through heterotrophic respiration (R_h), fires, herbivory, and fuelwood collection. The CASA version used here had a $0.5^\circ \times 0.5^\circ$ grid and a monthly time step. NPP was calculated based on satellite-derived estimates of the fraction of available photosynthetically active radiation ($fAPAR$) absorbed by plants:

$$NPP = fAPAR \times PAR \times \varepsilon_{(T,P)} \quad (1)$$

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



where PAR is photosynthetically active radiation, and ε is the maximum light use efficiency (LUE) that is downscaled when temperature (T) or moisture (P) conditions are not optimal. NPP was delivered to living biomass pools (leaves and roots for herbaceous vegetation, and leaves, roots, and stems for woody vegetation) following the Hui and Jackson (2005) allocation scheme with more NPP delivered to leaves and stems when mean annual precipitation (MAP) was high while larger amounts of NPP were delivered to roots when MAP is low (ter Steege et al., 2006). By introducing this partitioning scheme we captured 87% of the variability in biomass density in the Amazon (Fig. 1), based on a biomass density assessment combining forest inventory plots with satellite data (Saatchi et al., 2007). Since the main NPP drivers (f APAR and incoming solar radiation at the surface) were relatively uniform over the Amazon and other tropical forest areas, the original partitioning based on fixed fractions of NPP would yield little spatial variability in biomass density estimates.

Carbon in the living biomass pools was transferred to litter pools depending on turnover rates and satellite-derived changes in f APAR (Randerson et al., 1996) and subsequently decomposed based on turnover times regulated by temperature and soil moisture conditions (Potter et al., 1993). Other loss pathways include herbivory based on empiric relations between NPP and herbivore consumption (McNaughton et al., 1989) and fuelwood collection based on national fuelwood use statistics and population densities (following van der Werf et al., 2003; see Fig. S1). Although fuelwood collection and combustion is calculated internally, we do not further discuss or present these emissions because more comprehensive analyses are available (e.g., Yevich and Logan, 2003); the module is included to more realistically simulate spatial variability in fuel availability for other types of fires.

For each grid cell and month, fire carbon emissions were then based on burned area, tree mortality, and the fraction of each carbon pool combusted (combustion completeness, CC). Each carbon pool was assigned a unique minimum and maximum CC value with the fine fuels (leaves, fine litter) having relatively high values while coarse fuels (stems, coarse woody debris) having lower values (Table 1). The actual

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

combustion completeness was then scaled linearly based on soil moisture conditions with CC closer to the minimum value under relatively moist conditions, and vice versa (see van der Werf et al., 2006 for more details). Burned area and tree mortality will be discussed further below.

2.3 Main driver datasets

Key datasets for our model were burned area, active fires, and $fAPAR$, which are described below. Additional datasets used to drive the model are summarized in Table 2, and the main changes we made to the model are summarized in Table 3.

2.3.1 Burned area and active fires

We used the Giglio et al. (2010) burned area time series that is based on four satellite data sets. At the core lies a 500 m burned area mapping algorithm based on a burn-sensitive vegetation index, with dynamic thresholds aided by active fires applied to MODIS imagery (Giglio et al., 2009). Over 90% of the area burned over 2001–2009 was mapped this way. Local and regional scale relationships between MODIS active fires and burned area were used to map remaining areas in the MODIS era, while a mix of VIRS (Giglio et al., 2003) and ATSR (Arino et al., 1999) active fire data were used to map pre-MODIS burned area in a similar way. Several corrections were made to arrive at a consistent, long-term burned area dataset, see Giglio et al. (2010) for more details. This new burned area data set compared well to independent burned area estimates for North America, as well as to subsets of burned area derived from Landsat in tropical regions.

The burned area dataset includes an uncertainty assessment as well as information on the partitioning of burned area over different land cover classes and fractional tree cover bins within the 0.5° grid cell. For this, the MOD12Q1 land cover map for 2001 (Friedl et al., 2002) at 1 km resolution in combination with the University of Maryland (UMD) land cover classification scheme, and the MOD44 vegetation continuous fields

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(VCF; fraction tree, herbaceous, and bare cover; Hansen et al., 2003) for 2004 was used. The distribution of burned area over land cover and fraction tree cover (FTC) is shown in Fig. 2. In addition, the fraction of burned area that occurred on tropical peatlands in Indonesia as well as Malaysian Borneo was obtained using the *Terrestrial Ecoregions of the World* map (Olson et al., 2001; <http://www.worldwildlife.org/science/data/item1875.html>) sampled to a 500 m resolution grid. To partition burned area over the various land cover types for the pre-MODIS era and when the 500 m burned area maps were not available, we used a monthly climatology based on mapped burned area during the MODIS era instead of information derived from active fires. This was done to avoid inconsistencies. For example, the ATSR nighttime detection will give a smaller weight to those fires exhibiting a more pronounced diurnal cycle compared to MODIS.

2.3.2 *f*APAR

Our approach to estimating *f*APAR for the full study period was to take advantage of the sophisticated MODIS radiative transfer algorithms for calculating *f*APAR (Myneni et al., 2002) and the longer time series of NDVI observations from AVHRR (Tucker et al., 2005). We obtained MOD15 data from collection 5 at 4 km monthly resolution, including quality assurance meta data (QA) from the Boston University web site (<http://cliveg.bu.edu/modismisr/index.html>). The 4 km monthly product was produced by averaging the 1 km monthly product. The monthly 1 km pixels values were derived from the 8 day 1 km MODIS *f*APAR (native resolution). The QA for the 4 km product was calculated as the fraction of monthly 1 km pixels that were judged high quality relative the total number of pixels within a 4 km pixel (16). The QA for each 1 km monthly pixel was specified as high quality if at least one of the 4 eight-day pixels used the main algorithm. If none of the 4 inputs were from the main algorithm then the value of *f*APAR for that month was set to the maximum value and the QA to low quality. We only used 4 km pixels with $\geq 75\%$ high quality (main algorithm) to calculate the mean within the 0.5° aggregation.

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



To extend the f APAR time series back to 1997 we obtained GIMMS (Global Inventory Monitoring and Modeling Study) NDVI available at biweekly 8 km resolution, which we aggregated to monthly, 0.5° resolution. We then derived f APAR for each month (m) and year (y) of 1997–1999 (as well as January and February 2000) and 0.5° land grid cell (i) as:

$$fAPAR_{m,y,i} = \left(\frac{dM}{dG} \right)_i \Delta G_{m,y,i} + M_{m,i} \quad (2)$$

where $\left(\frac{dM}{dG} \right)_i$ is the slope of the linear correlation between MODIS f APAR monthly anomalies and GIMMS NDVI monthly anomalies (calculated separately for each month over the 2001–2008 period of overlap), $\Delta G_{m,y,i}$ is the GIMMS NDVI anomaly and $M_{m,i}$ is the mean MODIS seasonal cycle. This equation was only used in those grid cells where the p -value derived from the linear correlation was below 0.05. Otherwise the climatology was used. The high correlation of the MODIS and GIMMS anomalies justified this simple approach (see Fig. S2).

2.4 Modifications made to the modeling framework

While the overall structure of our modeling framework did not undergo major changes, new input datasets and several smaller modifications led to substantial changes in our estimation of fire carbon emissions. We modified the models so that for the period from November 2000 onwards, it now exclusively uses MODIS data for burned area, active fire detections, vegetation productivity (f APAR), land cover classification, and fractional tree cover estimates. Data from the VIRS and ATSR sensors were used to extrapolate fire information back in time (Giglio et al., 2010).

Besides these changes to input data, two major modifications were made. First, we adjusted the burned area estimates to better account for fire-driven deforestation and used these estimates to calculate deforestation fire emissions as a separate class within each grid cell (Sects. 2.4.1, 2.4.2). Second, the sub-grid cell information on the partitioning of burned area according to land cover type and fraction tree cover bin

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



was used to better estimate the contribution of different sources, and to partition total burned area within the 0.5° grid cell into herbaceous and woody burned area. Besides deforestation fires this included savanna fires, woodland fires, forest fires, agricultural fires, and peat fires. Savanna fires were further separated into grassland and savanna fires on one hand, and woodland fires on the other. These steps allowed, amongst others, for better estimates of trace gas emissions and aerosols (Sect. 2.4.5).

2.4.1 Deforestation rates

Active fire observations may be more successful in capturing fire activity in tropical high tree cover regions than burned area datasets (Roy et al., 2008). In addition, the number of times an active fire is observed in the same grid cell yields information on the fuel load and type of burning; fires in savanna and grassland areas burn rapidly with near-complete combustion of existing fuels, so if a fire is detected in a grid cell it rarely burns in the same grid cell during the consecutive overpass (Giglio et al., 2006). Deforestation fires, however, may burn over longer time periods before fuels are depleted. More fires are observed in the same location when forest is replaced with agriculture that requires near-complete removal of biomass than when land use following deforestation is pastureland (Morton et al., 2008). To better predict deforestation fire extent, we therefore combined burned area and active fire detections as a proxy for the area cleared by fire in deforestation regions. We first separated burned area for the 0.5° grid cell into area burned in wooded and in herbaceous (see Sect. 2.4.3). We then assumed that the cleared area was the product of the wooded burned area and fire persistence.

This proxy was calculated for each 0.5° grid cell and for each month, with the fire persistence averaged over all 1 km observations within the 0.5° grid cell. The proxy was used only in the humid tropical forest biome based on the WWF ecoregions map (Olson et al., 2001). Although empirical, it compared reasonably well to independent assessments of deforestation rates. Our approach captured about 49% of the variability in country-level deforestation rates over 2000–2005 when compared with a deforestation

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



assessment based on the hybrid use of Landsat and MODIS data (Hansen et al., 2008). Total pan-tropical deforestation rates based on our proxy were about 82% of those from Hansen et al. (2008). The state-level comparison against Landsat-derived PRODES (Programa de cálculo do desflorestamento da Amazônia) deforestation estimates for the Brazilian Amazon (<http://www.obt.inpe.br/prodes/>) from the Brazilian National Space Research Institute (INPE) was more favorable; we captured about 78% of the variability and also 82% of the total deforestation over the 2001–2006 period (Fig. 3). Cleared area was used to adjust the VCF fields over time; starting from the year 2004 (the base year for the VCF product) backwards in time the cleared fraction was added to the fraction tree cover while it lowered the fraction herbaceous cover, and vice versa for 2004 onwards.

In addition to the use of fire persistence in the deforestation rate assessment, it was also used to amplify combustion completeness and fire-induced tree mortality in deforestation zones (Fig. 4). Specifically, we set the combustion completeness so it ranged from its “normal” value (based on plant moisture content or soil moisture within the range defined in Table 1) to 1, and fire-induced tree mortality from 80% to 100% based on the fire persistence with the minimum value set at a fire persistence of 1 and maximum values defined when fire persistence was 4 (the 95th percentile of cleared area weighted persistence).

2.4.2 Deforestation emissions

Deforestation emissions were calculated based on biomass density from the forested fraction of each grid cell and deforestation rates (Sect. 2.4.1). The fate of the deforested land was tracked using a new sub-grid cell class representing land that had been deforested. Carbon pool density in this class was based on the carbon pool density of the forested fraction, with the combusted fraction subtracted. In case the grid cell underwent multiple deforestation events over the study period, the carbon pools of the deforested part of the grid cell were based on an area-weighted average of the previously and newly deforested fractions. NPP allocation in the deforested fraction was

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



treated the same as for herbaceous cover. R_h in deforested grid cells usually exceeded NPP due to decomposition of remaining forest carbon pools in the grid cell, with the effect larger if the combustion completeness was low.

Combining the deforestation rates with biomass density estimates in the wooded fraction of the grid cell, we found that on average $12 \pm 5 \text{ kg C m}^{-2}$ burned for Southern hemisphere South America. These estimates were near the upper bound of field measurements (Kaufmann et al., 1995; Guild et al., 1998) while those for Central America ($9 \pm 3 \text{ kg C m}^{-2}$ burned) and Northern hemisphere South America ($10 \pm 5 \text{ kg C m}^{-2}$ burned) were closer to average measurements, although still on the high side. The difference in modeled fuel consumption between these three regions was mostly due to higher fire persistence that boosted our combustion completeness in Southern hemisphere South America compared to other regions. In areas outside tropical America, fire persistence was lower and so were our fuel consumption estimates; about $5 \pm 3 \text{ kg C m}^{-2}$ burned for Africa and $7 \pm 4 \text{ kg C m}^{-2}$ burned for Central Asia. Only in Equatorial Asia was fuel consumption comparable to tropical America ($10 \pm 6 \text{ kg C m}^{-2}$ burned), this was likely caused by our inability to separate increased fire persistence due to repetitive burning of aboveground material from increased fire persistence due to incessant burning of peatlands. In other words, the high fuel consumption in Equatorial Asia may be a consequence of the co-existence of forests and peat soils, especially in deforestation areas where drainage canals expose peat soils to fire and oxidation during the deforestation process.

For the Southern Amazon, our fuel consumption estimates resembled those found by a related modeling approach focusing on the state of Mato Grosso, Brazil, which highlighted the possibility of high combustion completeness and thus high fuel consumption in these areas (DeFries et al., 2008; van der Werf et al., 2009). Especially if forests are replaced with large-scale agriculture, such as for soy plantations, all aboveground biomass and even part of the belowground biomass will be combusted through repeated burns (Morton et al., 2006). Our average fuel consumption estimates for Southern hemisphere South America were close to those calculated for conversions

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



to soy plantations (van der Werf et al., 2009), despite 2000–2005 trends indicating higher deforestation rates for cattle ranching in Amazonia where complete combustion of forest biomass is not required. However, very large deforestation events (>500 ha) accounted for the majority of deforested area in the Brazilian Amazon in recent years (Walker et al., 2009), indicating a trend towards mechanization of the clearing process (and higher combustion completeness) regardless of post-clearing land use for pasture or soy. It is important to note that our fuel consumption estimates are annual means and may include multiple deforestation fires in the same area during a single dry season. This makes it challenging to compare our estimates with literature values, mostly based on a single event.

2.4.3 Partitioning of non-deforestation fires

A novel aspect of this work was to separate deforestation fires from other types of fires. In addition, we used the partitioning of 500 m burned area over the different land cover types within the 0.5° grid to separate the non-deforestation fires in several sources. The model tracked woody and herbaceous vegetation separately within each grid cell. Because woody fuels are an order of magnitude larger than herbaceous fuels, separating these two sources should provide better emissions estimates. In previous versions we applied the same amount of burned area to both fractions, with a mortality scalar based on fraction tree cover to ensure that tree mortality was low in open savanna ecosystems and increased with increasing tree cover density. Here, however, we used the partitioning of burned area maps over 5% fraction tree cover bins (Fig. 2) to separately estimate the woody and herbaceous burned area within each 0.5° grid cell (Fig. 4). The amount of burned area (BA) was distributed over tree cover bins (TC) each sized 5 percent point (*i*) apart and we calculated the fraction of the total burned area occurring in the wooded part of the grid cell as:

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$$\text{Fraction of tree cover weighted burned area (TCWBA)} = \frac{\sum_{i=1}^{i=20} \text{BA}_i \times \text{TC}_i}{\sum_{i=1}^{i=20} \text{BA}_i} \quad (3)$$

The resulting average fraction of burned area occurring in woody fuel types is shown in Fig. S3. For each grid cell we thus had herbaceous and woody burned area estimates that were used to drive the sub-grid cell carbon flux calculations, with the most important difference that in the woody carbon flux calculations wood and coarse woody debris pools were included. More details on changes in the carbon model are described below; here we focus on how we partitioned the emissions into different sources (Fig. 4). The simplest partitioning within our model framework would be grassland fires (herbaceous) versus forest fires (woody). However, most land cover types consist of a mixture of herbaceous and woody plant functional types, such as savannas, where trees and grasses are interspersed over the landscape. We therefore based the partitioning of non-deforestation and non-peat fires into savanna, woodland, forest, and agricultural fire emissions on the partitioning of burned area within different landcover types defined by the MODIS MOD12Q1 product (Friedl et al., 2002) using the UMD classification scheme.

To calculate agricultural waste (AGW) emissions, we multiplied the herbaceous emissions with the fraction of total herbaceous burned area occurring in agricultural areas (class 12 in the UMD land cover classification). Another class of emissions that was solely derived from either the herbaceous or woody emissions (versus the mixture) were fires in wooded areas outside the humid tropics, but still containing evergreen broadleaf forest. We confined our deforestation assessment to humid tropical forests defined in spatial extent by the WWF ecoregions map (Olson et al., 2001). Fire emissions from trees that occurred in grid cells containing evergreen broadleaf forest but outside the humid tropical forest domain were here included as deforestation (or degradation) emissions to separate them from deforestation and degradation fires within the

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



humid tropical forest biome, and to be able to assign them a different emission factor. However, this distinction is somewhat arbitrary; below we will refer to deforestation and degradation emissions to cover all non-savanna or agricultural fires occurring in the tropical forest domain irrespectively of whether they caused permanent land use changes (deforestation) or were, for example, escaped fires (degradation).

We next calculated the fraction of emissions associated with forest fires. In the boreal region, according to the UMD classification a large fraction of the burned area were observed in savanna-type ecosystems more likely resembling forests with relatively low tree cover; we therefore labeled fires in shrublands and woody savannas (class 7–9) as forest fires in this region. We defined the boreal region as all land with below zero mean annual temperature and more than 100 mm yr^{-1} mean annual precipitation. The precipitation threshold was included so that high latitude arid grassland areas such as found in Mongolia were treated as grassland. Boreal forests were unique in that emissions included burning in forest, shrubland, and wood savanna classes. In other regions, forest emissions were based only on burned area that occurred within the forest classes (Fig. 4).

The remainder of emissions stemmed from grasslands and savannas, with the latter ranging from open savannas to woodlands. To distinguish grassland and open savannas from woodlands, we separated these two sources based on the dominant source of emissions; if herbaceous emissions dominated then we labeled them savanna fires, otherwise woodland fires (Fig. 4).

2.4.4 Additional changes (tree mortality, combustion completeness, leaf litterfall)

Tree mortality (M_w) was modeled similar to earlier model versions as a function of fractional tree cover so that savanna-type ecosystems had only 1% mortality which started to increase when tree cover exceeded 30% to reach the maximum of 60%

Global fire emissions (1997–2009)

G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[I◀](#)[▶I](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

mortality in areas with more than 70% tree cover following:

$$M_w = 0.01 + 0.59 / \left(1 + e^{(25 \times (0.50 - \text{TCWBA}))} \right) \quad (4)$$

While in earlier version this mortality scalar was fixed in each grid cell based on the fraction tree cover in the grid cell, here we used the sub-grid cell information to model mortality more dynamically, allowing it to change over time. Specifically, we used a burned area weighted estimate of tree cover (TCWBA) that changed for each time step (Eq. 3). Two region-specific modifications to Eq. (4) were made; we scaled the mortality in deforestation regions to values between 80 and 100% based on fire persistence (as described in Sect. 2.4.1) and (similar to earlier modeling versions) applied a fixed mortality of 60% in forested regions in temperate and boreal regions (based on a mean annual temperature threshold of below 15 °C) where tree cover density was often far below 70%. We made this modification in recognition that stand-replacing crown fires often occur in many temperate and boreal forests. A map of mean fire-induced tree mortality is shown in Fig. S4. Although in some areas of the boreal forest mortality can approach 100%, particularly in areas with moderate and severe fires, we applied a 60% mortality to reflect the observation that within burn perimeters there are often many areas that are incompletely burned, or even entirely unburned.

The key fuel component in the boreal region is the soil, which is most often the major source of emissions. This is also the case for Equatorial Asia, most importantly in Indonesia (Page et al., 2002). Organic soil burning was modeled in a similar fashion as combustion completeness; we set a minimum and maximum value (0 and 15 cm for the boreal region, 0 and 50 cm for Equatorial Asia), which was then scaled based on soil moisture conditions (from both the current and the previous month).

In North America, organic soil burning had a mean depth of 8 ± 3 cm during 1997–2009 and with this parameterization our fuel consumption estimates agreed with the $0.8\text{--}3 \text{ kg C m}^{-2}$ dominant range (and outliers to 5 kg C m^{-2}) found in recent literature (DeGroot et al., 2007; DeGroot et al., 2009; Boby et al., in press). Average depth of burning in Indonesia (30 ± 8 cm) was similar to results from a large-scale assessment

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



of depth of burning in Borneo using LIDAR measurements (Ballhorn et al., 2009), resulting in a burning depth of 33 ± 18 cm.

In addition, we modified the leaf litterfall parameterization; in previous versions the amount of leaves and grasses decreased only slightly after the growing season. This led to a larger than desired build-up of leaves, and thus to an overestimation of fuel, especially in areas dominated by herbaceous fuels such as savannas. By lowering the turnover time of leaves to 6 months and modifying other parts of the algorithm, the leaf litterfall component, the leaf pool build-up and its depletion following the growing season performed better. Average fuel consumption estimates for savanna-dominated regions (Africa and Australia) were about half of those previously found (Table 4 versus Table 4 in van der Werf et al., 2006).

Measurements of Savadogo et al. (2007) in savanna-woodlands in West Africa showed that grazing may lower fuel loads by 50% compared to areas without grazing. They also found significant differences in fuel loads between annual and perennial grasses. Although our model includes grazing based on a global relation between plant productivity and herbivory, fine-scale differences like these cannot be reproduced due to the relatively coarse spatial resolution of our model. Fuel loads between different treatments varied between 170 and 450 g C m^{-2} (Savadogo et al., 2007). In the half degree grid cell encompassing their study region, modeled minimum fuel loads were 200 g C m^{-2} , based on one wet season fuel build-up. Maximum fuel loads were 550 g C m^{-2} when fires were excluded in our model, which was somewhat larger than observed in the field. In savanna areas of Northern Australia, Williams et al. (1998) performed a landscape-scale experiment where fuel loads were found to range between 75 and 650 g C m^{-2} with most fires burning in areas with $100\text{--}200 \text{ g C m}^{-2}$ of fuel. For Australia as a whole, we found that most fires burned between less than 100 g C m^{-2} of fuel but with a substantial amount of burning in the $100\text{--}400 \text{ g C m}^{-2}$ of fuel (Figs. 5, 6). In the area where Williams et al. (1998) performed their measurements average fuel consumption was about 250 g C m^{-2} of fuel while maximum fuel consumption (reached when fires were excluded for 5 years) was 600 g C m^{-2} of fuel.

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



While far from exhaustive, these comparisons are encouraging.

2.4.5 Trace gas emissions

Our modeling framework calculated carbon fluxes. Emission factors (EF) were then used to translate the fire carbon loss to trace gas and aerosol emissions. EFs have been measured in most fire-prone biomes, compiled by Andreae and Merlet (2001) and updated annually (Andreae, personal communication). We used separate EFs from this database for fires in (1) tropical forests, (2) savannas and grasslands, (3) extratropical forests, and 4) agricultural areas. The EFs were based on the mean of the measurements for each species within each of the 4 biomes described above. For tropical peat burning only one campaign made measurements to our knowledge (Christian et al., 2003), with EFs for reduced species about twice as high as those for tropical forest fires. Deforestation and degradation fires in the non-humid tropics received the average EF of savanna and deforestation fires because they represent a mixture of these fire types, and we applied the same EF to woodland fires. EFs were reported per kilogram dry matter burned; based on mass balance equations of the EFs ($\text{CO}_2 + \text{CO} + \text{CH}_4$) we used a dry matter carbon content of $\sim 48\%$ to translate calculated carbon to dry matter emissions. EF used for several species are given in Table 5.

2.4.6 Spin-up

We spun up CASA for 250 years based on average input data from 1997–2009 so that carbon release (R_h , fires, herbivory, fuelwood collection) matched input (NPP), indicating that all carbon pools were in steady state. Because of the long turnover rates of slowly decomposing soil pools, these pool sizes were tuned prior to the spin up to match measured carbon densities (Batjes et al., 1996) by adjusting the turnover times of the slow soil pools by a single scalar in each 0.5° grid cells. Fire return times for forests were based on the mean fire interval for each basis region (Fig. 7) and for each 10% fraction tree cover bin to create region-specific and to some extent

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



ecosystem-specific average fire return times (Fig. S5), trading space for time (Chuvieco et al., 2008). This was done to create realistic fuel loads in areas that saw frequent fires during the study period but had possibly little fire activity in early periods.

2.5 Uncertainty

5 While the burned area assessment underwent a formal uncertainty assessment, a similar approach to estimate uncertainties in fuel loads, combustion completeness, and emission factors is not yet possible due to a lack of ground truth data. However, to get an indication of the spatial variability in uncertainties in carbon emissions we have propagated the uncertainties in the burned area estimates through our model in a Monte Carlo situation using best-guess estimates of model parameter uncertainty (Table 6) following French et al. (2004) and Jain et al. (2007), although the latter used error propagation instead of a Monte Carlo simulation.

10 Specifically, we attributed best-guess uncertainties to several parameters used to calculate emissions (Table 6). Normally-distributed uncertainties for the light use efficiency (scaling directly to biomass density), burned area, combustion completeness, and burning depth into organic soil were used. We performed 2000 runs based on spin-up data from the main run and with the biomass pools adjusted with the change in LUE (which led to a linear change in biomass density), and then ran 1997–2009 and changed the other parameters independently. We focused the uncertainty analysis on carbon emissions and not all errors were included; we did not, for example, include errors in the fractional tree or land cover that were used in several places in our model.

15 The uncertainty we assigned to biomass was based on the comparison with Amazonian biomass (Fig. 1). Specifically, we used the square root of the mean of the squared residuals of the comparison. Since the scatter increases with biomass density (heteroskedasticity), the standard deviation was applied as a scaling factor (here the light use efficiency) instead of an absolute value. For herbaceous fuels, the same standard deviation was used but we doubled the value to account for additional uncertainties such as the amount of grazing and our inability to accurately determine the time since

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



the previous fire (one of the key factors in herbaceous fuel amount) due to our relatively coarse model set-up. Standard deviations for combustion completeness and depth of burning were chosen somewhat arbitrary as half of their respective ranges. For depth of burning into organic soil in boreal regions, for example, the standard deviation was set to 7.5 cm. These values are relatively large and should account for the substantial uncertainty here, which probably exceeds uncertainties in combustion completeness, although no formal assessment can be done.

For combustion completeness, burned area, and the depth of burning into organic soil, we truncated the distributions to avoid physically unrealistic scenarios. For example, if the combustion completeness would exceed unity or the depth of burning would be negative, these would be cut-off by 1 and 0, respectively. Therefore uncertainties in some areas are not necessarily normally-distributed, and the mode of the Monte Carlo runs may not line up with the values we report as our best estimates

3 Results

3.1 Emissions

3.1.1 Global overview

Average carbon emissions over 1997–2009 were 2.0 Pg C yr^{-1} with considerable inter-annual variability, especially over the 1997–2001 period (Figs. 8, 9, Table 7). Emissions in the peak fire year 1998 (2.8 Pg yr^{-1}) were 78% higher than those in 2001 (1.6 Pg yr^{-1}). From 2002 through 2007, emissions were relatively constant from year to year on a global scale. Regionally, however, large variations occurred but high fire years in some regions cancelled low fire years in other regions. In 2006 for example, emissions in Southern hemisphere South America were relatively low while in Equatorial Asia emissions were higher than in any other year except 1997. In 2007 the

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



reverse occurred with high emissions in Southern hemisphere South America and low emissions in Equatorial Asia. In 2008, almost all regions experienced below average emissions with the notable exception of boreal Asia, leading to a relatively low fire year globally (1.7 Pg yr^{-1}). This situation persisted in 2009, although boreal Asia was now also low and even though emissions in Equatorial Asia increased somewhat, 2009 was the year with lowest emission over our study period (1.5 Pg yr^{-1}).

Over half of the global carbon emissions were from Africa (Table 7), with emissions from Africa south of the equator (28%) somewhat exceeding those from Northern hemisphere Africa (24%). South America accounted for 15% of global carbon emissions, mostly from Southern hemisphere South America. On average, Equatorial Asia was the fourth most important region (10%) with its relative contribution growing considerably during El Niño years. In 1997, for example, we estimated that emissions from Equatorial Asia contributed to 40% of global emissions. According to our estimates, the boreal region accounted for 9% of total global carbon emissions with emissions from boreal Asia almost 2.5 times as high as those from boreal North America and comparable to emissions from Australia. While the Scandinavian countries and Finland were not included in the boreal region in our assessment, emissions here were negligible compared to boreal North America and Asia.

When translating our estimated carbon emissions to emissions of trace gases (Andreae et al., 2001; Andreae, personal communication) the role of savanna regions like Africa and Australia diminished while the role of forest and deforestation fires in areas with higher woody fuel loads increased (Fig. 10). This is because of more complete oxidation and thus reduced production of CH_4 , CO and other reduced trace gases in grass fuels compared to fuels from forests, shrublands, and peatlands. Africa, for example, accounted for 72% of global burned area, 52% of global carbon emissions, 44% of CO emissions and 36% of CH_4 emissions (Fig. 10). On the other hand Southeast and Equatorial Asia had only 2.5% of global burned area but due to high fuel loads including peat that emit more reduced trace gases per unit carbon combusted it accounted for 22% of CO and 32% of CH_4 emissions. Since the surface area of Equatorial Asia

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



is much smaller than any of our other regions, the emissions density is highest in this region (Fig. 11). This is exacerbated by the large interannual variability in this region.

3.1.2 Partitioning between fire types

By far the largest contributors to the global fire carbon loss estimates were fires in savannas and grasslands, on average 44% over 2001–2009 (Figs. 12, 13). 23% of fire carbon emissions were net carbon emissions (likely not compensated for by regrowth) either due to deforestation, degradation, or peat fires, most importantly in Southern hemisphere South America (37% of all deforestation fires) and Equatorial Asia (all tropical peat fires and 19% of all deforestation fires). Since the partitioning in the MODIS era was assumed most reliable, the results we present were for 2001–2009 (Fig. 12, Table S1). However, if we were to consider the full 1997–2009 period, the role of peat fires would increase from 3 to 5% of total global emissions due to high peat fire emissions in 1997. The major regions contributing to net fire carbon emissions were Southern hemisphere South America and Equatorial Asia. While deforestation emissions in Southern hemisphere South America were substantially larger than those in Equatorial Asia, total net emissions in Equatorial Asia exceeded those in Southern hemisphere South America because of the important role of peat burning.

Forest fires were a dominant contributor to emissions in the boreal regions (96% in boreal North America and 81% in boreal Asia) while agricultural waste fires were important in Europe, Middle East, and Central Asia. Almost all fires in Central Asia north of the Black, Caspian, and Aral Seas were in agricultural areas (Fig. 13). Since these agricultural fires may not leave a clear burned area signal in the 500 m data used here, emissions estimates in these fires are likely conservative because we detected only a fraction of the actual agricultural fire activity. Part of this underestimate may be compensated for because our framework does not explicitly include harvest, leading to an overestimation of fuel loads.

On a global scale, peat fire emissions were most variable from year to year as measured by the coefficient of variation (CV, standard deviation divided by the mean, Fig. 8).

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Their CV of 1.76 was about a factor 5 larger than the CV of deforestation and degradation, and forest fires (0.37 and 0.36). Agricultural waste burning and savanna fires were relatively constant on a global scale with a CVs of 0.18 and 0.11, respectively. Regionally, deforestation fires also showed more variability, especially in Northern hemisphere South America (CV of 1.37) and Equatorial Asia (1.10). The main deforestation region (Southern hemisphere South America) had a CV of 0.64. Agricultural waste burning and savanna fires were not only relatively constant on a global scale; also regional variations were modest compared to other sources with the exception of Australia (CV of 0.29 for savanna fires).

3.1.3 Uncertainties

Results of the Monte Carlo simulation indicated that globally, uncertainties were around 20% (1σ) for annual estimates during the MODIS era (2001–2009) and somewhat higher during the years before when burned area was derived from ATSR and VIRS hot spots (Fig. 14). Regionally, uncertainties were highest in the boreal regions and Equatorial Asia where organic soil burning occurs (Fig. S6). One factor that had a major impact on the spatial distribution of the uncertainties was whether mapped burned area was available, or whether burned area estimates were derived from fire hot spot–burned area relations. For the latter, uncertainties were much higher. Because uncertainties were often higher than the absolute burned area and because negative burned area estimates were truncated at 0, the mode of the Monte Carlo runs was higher than the estimates reported throughout the paper (Fig. S6).

3.2 Fuel consumption

Fuel consumption (reported here as g C per m^2 of area burned) broadly followed biome distributions with low biomass density biomes such as grasslands and savannas burning less fuel than high biomass density types such as forests (Fig. 6). Fuel consumption in our model depended on biomass availability, tree mortality, and combustion

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

completeness (Fig. S7). The relatively abrupt incline in fuel consumption when moving from savannas to woodlands to tropical forest was for a large part related to increased tree mortality. In areas with low tree cover density the tree mortality was low (1%) so fuels consisted mainly of litter and (dead) grass, while in tropical forest area up to 100% of the biomass was subject to fire. In addition, the combustion completeness was increased in deforestation areas when fire persistence was high; this led to higher fuel consumption in the Southern Amazon, for example, as compared with Africa. High fuel loads were also found in peatland areas in Indonesia, most notably the Southern part of Borneo and northwestern Sumatra. On average, Equatorial Asia had by far the highest mean fuel consumption (Table 4), because the majority of fires burned in peatlands and forested areas.

In savannas, the areas that had lower fire return times, higher precipitation rates, or had larger tree densities showed the highest fuel consumption. High precipitation rates lead in general to more fuels, causing the gradient of fuel consumption observed in Africa and Australia (Fig. 6). There is evidence that since increased precipitation rates are often accompanied by increased tree cover density, the amount of grass available to burn decreases. In our model this was compensated for by an increase in fire-induced tree mortality. Savannas in South America had considerably higher fuel consumption than those in Africa, which in turn consumed more carbon than most of the fires in Australia. Because in Southern hemisphere Africa comparatively more fires were detected in woodlands than in Northern hemisphere Africa, average fuel consumption here was higher (448 g C m^{-2} in Southern hemisphere Africa vs. 377 g C m^{-2} in Northern hemisphere Africa). Although fuel consumption was somewhat skewed (exacerbated by the log scale in Fig. 5), median fuel consumption was close to the mean; most fires in Northern hemisphere Africa burned between 200 and 300 g C m^{-2} , and between 300 and 400 g C m^{-2} in Southern hemisphere Africa.

In boreal forests most fires burned between 2 and 5 kg C m^{-2} (Figs. 5, 6). When considering all fires, average fuel consumption was 2662 g C m^{-2} in boreal North America and 1979 g C m^{-2} in BOAS (Table 4). Less than 10% of these fuels were from the

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



burning of standing trees and leaves; most was related to burning of the duff layer and organic soil (Table 4). Because of a larger share of fires burning in areas with lower fuel loads (e.g., in grasslands and agriculture), average fuel consumption in boreal Asia was lower than in boreal North America, see Figs. 5, 6.

4 Discussion

4.1 Source contributions

Deforestation, degradation, and peatland fires

We found that on average, about a quarter of total carbon emissions from fire (or 0.5 Pg C yr^{-1}) contributed to the build-up of atmospheric CO_2 because they were likely not or only partly balanced by regrowth. This analysis was confined to the tropics (Fig. 13, S8). Roughly 3/5th of this stemmed from deforestation and degradation fires inside the humid tropical forest biome, 1/5th from deforestation and degradation in the non-humid tropical forest biome, and the remaining 1/5th from tropical peat burning (Table S1). While crucial for interpreting the atmospheric CO_2 signal, the uncertainty in this net carbon emissions estimate is large. There are reasons to believe that our estimates are conservative. First, the approach used to estimate deforestation rates captured about 80% of total deforestation rates reported by other studies. Whether the remaining 20% is due to deforestation or degradation not involving fires (e.g., logging) or whether our simple approach underestimated fire-driven deforestation rates is unknown. Second, we focused on the tropical forest biome for deforestation. In addition, net emissions may occur in areas outside this biome. Hansen et al. (2010), for example, found higher rates of forest loss in boreal and temperate regions than in the tropical forest biome over 2000–2005. These may be partly compensated for by regrowth or forest thickening in other regions, but may also indicate net forest loss. In our analysis, some of the woodland areas, especially in Southern Africa, had fire return intervals

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



that did not appear sustainable. This points towards degradation that is not included in our approach because it is not trivial to separate fires that are part of a natural cycle (or sustainable human-driven fire intervals) from areas that see degradation. Case studies that investigate the fate of the burned areas at higher spatial resolution could be more successful in understanding these dynamics over time. Multi-decadal high resolution time series of burned area are also crucial for detecting these types of changes in fire regimes.

Fire emissions estimates in the deforestation regions have been used in the past to complement other approaches to calculate deforestation emissions, for example those based on reported deforestation rates, biomass density inventories, and deforestation biomass loss trajectories (e.g., Houghton et al., 2003). The fire emissions were doubled to account for emissions other than fires, for example from the respiration of leftover plant materials or soil carbon following deforestation (e.g., Olivier et al., 2005). Following this approach, total deforestation estimates (including peat carbon emissions) would be 1 Pg C yr^{-1} , supporting earlier satellite-based deforestation estimates for the 1990s (DeFries et al., 2002; Achard et al., 2004). However, besides the substantial uncertainties in our deforestation fire flux estimate, the doubling of the estimate to arrive at total deforestation carbon emissions is prone to large errors and more research is needed to lower uncertainties. One feature that is likely robust are the interannual variations, supporting for example the downward trend in deforestation emissions over the past few years in Brazil (e.g., Nepstad et al., 2009). One exception was found in the year 2007; while PRODES reported deforestation rates were below average, emissions for Southern hemisphere South America were higher than in any other year in our record (see also Gloudemans et al., 2009). While the exact causes remain unclear and may change from year to year (e.g., Torres et al., 2010), precipitation rates may explain part of the variability (Aragão et al., 2008; Le Page et al., in press). This precipitation-fire-deforestation link is more pronounced in Equatorial Asia (Field et al., 2009; van der Werf et al., 2008), where emissions increased again in 2009 due to El Niño conditions after two very low fire years, but not to levels seen during earlier El Niño

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

events in 1997, 2002, and 2006. Combining Southern hemisphere South America and Equatorial Asia, emissions from these two important deforestation regions resulted in below-average deforestation fire emissions in the latter part of our record.

While the sum of deforestation, degradation, and peat fire emissions accounted for about a quarter of total carbon emissions, for CH₄ these sources were key with 44% of total CH₄ emissions of 20 Tg CH₄ yr⁻¹. This was mostly due to the large emission factor for CH₄ in peat areas, which were three times as high as the emission factor for tropical deforestation fires, and almost 10 times that of savanna and grassland fires.

Agricultural waste burning

Another source of emissions for which our estimates are likely conservative are fires in agricultural areas, for example due to the burning of leftover crop residues. Most of these fires are small in comparison to the 500 m grid cell for which burned area was calculated. In addition, they often follow harvest, leaving a smaller drop in the vegetation index that was used to calculate burned area. Hence, we probably underestimate the area burned in agricultural waste burning areas. Fuel loads may compensate for part of this; since harvest is not explicitly included in the model (although the drop in *f*APAR will translate into a transfer of biomass from leaves to litter) we will likely overestimate fuel loads. On average, we found emissions from this source to be 55 Tg C yr⁻¹, which is factor 4 lower than found by Yevich and Logan (2003). The spatial distribution of emissions closely resembles those found by Korontzi et al. (2006), which is not surprising because the datasets used are similar. Noteworthy is that Korontzi et al. (2006) found that between 8 and 11% of fire hot spots were observed in agricultural areas over 2001–2003; combining this with average emissions of 1.9 Pg C yr⁻¹ we found during these years results in a similar estimate as found by Yevich and Logan (2003) if we assume that each fire hot spot represents an equal amount of emissions. While hot spots integrate the effect of burned area and fuel consumption, they likely do not scale linearly with emissions (e.g., van der Werf et al., 2006), but these independent

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



assessments provide some confidence in the amount of carbon burned in the agricultural waste process.

Savanna and woodland burning

By far the largest source of carbon emissions were fires in grasslands and savanna fires (44%), especially when woodland fires are included (total of 60%). This is not surprising because of the prominent role of savanna-dominated Africa and Australia in the global budget. The distinction between savanna and grassland on one hand, and woodland on the other hand was somewhat arbitrary (based on whether herbaceous or woody fuels were the dominant source of emissions). In earlier studies (e.g., Crutzen and Andreae, 1990) this distinction was not made. The combined estimates of savanna and woodland emissions between this study and the estimate dating 20 years back in time agree to within 5% (Fig. 15). The approach used to derive this estimate, however, was very different from the ones used in the past. While Crutzen and Andreae (1990) assumed that all savanna areas in the tropics burned with frequencies ranging from every year to every 4 years based on data from Menaut et al. (1991) and consumed on average 55 g C m^{-2} , our estimates indicate that even the most frequently burning savanna areas (those receiving between 1000 and 1500 mm of rain annually), burned on average only once every 4 years during 1997–2009. Consequently, fuel consumption was higher and because we also take woodland burning into account, this compensated for the lower estimates of areas burned.

Forest fires

Although fuel loads were high in forests, globally the role of forest fires (excluding deforestation fires and woodland burning) was relatively modest; about 15% of total carbon emissions was due to the burning of forests. Our model did not separate ground from crown fires and thus fuel consumption in boreal North America and boreal Asia was relatively similar. There are indications though that the fire regime in boreal North

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



America is more characterized by crown fires while ground fires are more prevalent in boreal Asia (Harden et al., 2000; Wooster and Zhang, 2004). Even though the aboveground fuel component is relatively small, not including these dynamics is an additional source of uncertainties.

4.1.1 Comparison against GFED2

Our revised fire carbon loss estimate was on average 13% lower than found previously (van der Werf et al., 2006) when taking the 1997–2008 period into account, see Fig. 16 (GFED2 was not updated after 2008). The decrease in CO emissions was also 13% (Fig. S9), but regionally changes in C and CO were different. Trace gas calculations based on the modeled carbon fluxes were impacted in two contrasting ways. First, in GFED2 we assumed a dry matter carbon content of 45% while mass balance equations indicate that 48% is more appropriate. The calculated carbon emissions thus represent 6% less dry matter emissions to which emission factors are subsequently applied to derive trace gas fluxes. This has the effect of reducing trace gas emissions. However, here we included peat-specific emission factors for tropical peatland fires for CO and CH₄, which were double those used previously when they were based on deforestation emission factors. In addition, the deforestation and degradation fires in the non-humid tropical forest biome as well as woodland fires received higher emission factors than in GFED2. These two types of adjustments more or less compensated each other on a global scale.

Regionally, differences are larger than the global comparison suggests, but interannual variability was relatively similar in both versions. The only regions where emissions increased substantially were boreal North America (+43%), and the Middle East (including the Sahara, +126%). This was mostly due to increases in burned area (11 and 90%, respectively) and, for boreal North America, higher fuel consumption (39%). The higher fuel consumption in boreal North America was primarily the result of 1) an increase in the fraction of area burned in forest, and 2) the inclusion of soil burning in the herbaceous fraction of grid cells occurring in forested regions. Specifically, tree

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



cover density is relatively low in the forests of boreal North America and previously we assumed that fires burning in the herbaceous fraction of these forests were grassland fires that only burned aboveground fuels. It is, however, more likely that this herbaceous fraction is due to larger distances between trees (i.e., more open taiga areas) and that the fires also consume the soil organic matter in the areas between trees.

All areas dominated by grassland and savanna fires had a decrease in fire emissions due to the modifications we made to better represent senescence of herbaceous fuels, most importantly lowering the turnover time of leafs. The decrease in emissions in Southern hemisphere Africa was small (−3%), mostly because lower average fuel consumption (−41%) was partly compensated for by a 38% increase in burned area. Burned area also increased in Northern hemisphere Africa (Northern hemisphere Africa, 20% more burned area) and Australia (6% more burned area), but not enough to offset the decrease in fuel consumption. The 22% decrease in emissions in Central Asia was solely due to lower fuels; burned area did not change substantially here.

Other areas that saw a reduction in estimates of fire carbon emissions included temperate North America and Central America, as well as Northern hemisphere South America, mostly due to decreased burned area and in the case of temperate North America also a reduction in fuel consumption. Burned area also decreased significantly in Europe leading to a 66% reduction in carbon emissions. The decreases in Southeast (−31%) and Equatorial Asia (−17%) were also largely due to reduced levels of burned area, that was partly compensated for by higher fuel consumption in Equatorial Asia. Including peat-specific emission factors resulted in an increase in CO emissions in Equatorial Asia (9%).

Because of the large range in fuel consumption, changes in burned area do not necessarily translate directly into changes in emissions; it depends where those changes occur. Our current version is for a large part driven by mapped 500 m burned area (in the MODIS era, where more than 90% of the total burned area was mapped directly) while in GFED2 the model depended solely on a regression tree relationship between

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



fire hot spots and burned area. The improved burned area led to higher quality predictions that better matched observations across the continental US, Canada, and Alaska (Giglio et al., 2010). Together with the other improvements described above (summarized in Table 3) we have reasons to believe that our current approach is more reliable.

4.1.2 Comparison against other bottom-up estimates

Few global estimates of fire emissions exist that employ new burned area datasets; most assessments were based on pilot burned area studies for the year 2000; GBA2000 (Gregoire et al., 2003) and GLOBSCAR (Simon et al., 2004). These burned area estimates have been used in different ways to estimate emissions. Hoelzemann et al. (2004) coupled GLOBSCAR burned area and ATSR fire hot spots in burned area data gaps with biomass density estimates from the Lund-Potsdam-Jena (LPJ) dynamic global vegetation model (DGVM) at a 0.5° spatial resolution to estimate total emissions of 1.7 Pg C yr^{-1} in the year 2000. For the same year and using the GBA2000 burned area dataset, Ito and Penner (2004) estimated emissions from open fires of 1.4 Pg C yr^{-1} combining satellite-derived vegetation information and literature-based values of combustion completeness and relations between vegetation types and fuel loads, most at the native (1 km) burned area resolution. Earlier versions of our modeling work estimated emissions in 2000 of 2.7 Pg C yr^{-1} (GFED1, van der Werf et al., 2004) and 2.0 Pg C yr^{-1} (GFED2, van der Werf et al., 2006).

The substantial differences between these studies was due in part to differences in burned area data; Jain et al. (2007) employed three different burned area datasets in their Integrated Science Assessment Model (ISAM) to find that CO emissions ranged between 320 and $390 \text{ Tg CO yr}^{-1}$ with much larger regional differences than the range of global burned area estimates would suggest. These CO estimates are substantially lower than earlier inventory methods; Galanter et al. (2000), for example, estimated CO emissions from savanna and forest burning of $554 \text{ Tg CO yr}^{-1}$. Estimates for the year 2000, however, are probably not representative for mean annual emissions because emissions from several important fire areas (including tropical Asia and America) were

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



below average (van der Werf et al., 2004). Globally, our results now indicate that emissions for the year 2000 were 1.7 Pg C yr^{-1} while the average for 1997–2009 was 2.0 Pg C yr^{-1} . When taking the full uncertainty range into account, our estimates still overlap with those mentioned above with the exception of van der Werf et al. (2004), which produced substantially higher estimates. Interannual variability, however, is still comparable to the results from van der Werf et al. (2004) supporting the conclusion that fires in Indonesia, Central and South America, and the boreal region contributed substantially to the high CO_2 and CH_4 growth rates during the 1997–1998 El Niño.

Much progress has recently been achieved in estimating emissions at regional scales. Lehsten et al. (2009) estimated fire emissions for 2001–2005 from Africa based on L3JRC burned area data (Tansey et al., 2008) and the SPITFIRE fire model embedded in the LPJ-Guess DGVM. They estimated that over this time period on average $195 \times 10^4 \text{ km}^2 \text{ yr}^{-1}$ burned, emitting $723 \pm 70 \text{ Tg C yr}^{-1}$. We found that over the same time period $247 \times 10^4 \text{ km}^2 \text{ yr}^{-1}$ burned which emitted on average $1031 \text{ Tg C yr}^{-1}$. Average fuel consumption for the continent was thus comparable in these approaches, $417 \text{ g C m}^{-2} \text{ yr}^{-1}$ (this study) vs. $371 \text{ g C m}^{-2} \text{ yr}^{-1}$ (Lehsten et al., 2009), although the spatial distribution of burned area varied substantially.

Besides Africa, the boreal region and Australia may be the mostly studied fire regions. Several authors have estimated emissions for the boreal region, with large differences partly related to the complex fuel composition where belowground fuels may comprise the largest fraction of the fuels (e.g., Amiro et al., 2001; DeGroot et al., 2007). Both our previous work and the emission estimates presented here fall in the range of previous emissions estimates (Kasischke et al., 2005; Soja et al., 2004) but uncertainties will remain large because several parameters (most importantly the depth of burning into organic soil) remain difficult to validate. Our emissions estimates for North America (including Mexico) are substantially lower than those from Wiedinmyer et al. (2006), especially for the US.

While both inventory and satellite-based estimates described above use the product of burned area, fuel loads, and combustion completeness to calculate emissions

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(Seiler and Crutzen, 1980), studies using satellite-measured fire radiative power (FRP) do not require these input datasets – each with substantial uncertainty – but directly integrate FRP to total fire radiative energy (FRE), which is in turn related to total emissions (Wooster et al., 2002). Roberts and Wooster (2007) used geostationary Spinning Enhanced Visible and Infrared Imager (SEVIRI) and calculated emission for Southern Africa for July–October 2004 of 259–339 Tg dry matter (DM) burned, or 127–166 Tg C assuming a DM carbon content of 48%. For the same region and time period, GFED2 emissions were 387 Tg C while GFED3 estimates were 404 Tg C. FRE-based estimates using MODIS observations are also substantially lower than our estimates; Ellicott et al. (2009) estimated DM emissions averaged over 2001–2007 for all of Africa of 716 Tg DM yr⁻¹ (351 Tg C yr⁻¹) and 305 Tg DM yr⁻¹ for South America (149 Tg C yr⁻¹). For Africa, GFED2 estimates were 1175 Tg C year while GFED3 estimates were 1018 Tg C yr⁻¹, exceeding those of Ellicott et al. (2009) by factors of 3.3 and 2.9, respectively. The difference was somewhat smaller for South America; while Ellicott et al. (2009) estimated emissions of 149 Tg C yr⁻¹, the GFED2 based estimate was 373 Tg C yr⁻¹ (2.5 times higher) and GFED3 estimates 334 Tg C yr⁻¹ (2.2 times higher). Thus, there remains a large discrepancy between biogeochemical model estimates combining burned area and fuel consumption and FRE/FRP approaches, with the latter having a factor of 2–3 lower emissions.

4.1.3 Comparison against regional top-down assessments

Since trace gas and aerosol emissions from fires are transported into the atmosphere, and many of these gases are measured by satellite, atmospheric measurements may be used to validate bottom-up estimates if the transport and chemical pathways are modeled correctly. If this approach is used, emission factors translating carbon or dry matter emissions to trace gas or aerosol emissions are used. Since these emission factors are sometimes poorly known and since their spatial and temporal variability is hardly taken into account, care should be taken translating findings of top-down assessments directly to carbon or dry matter emissions, although these shortcomings

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



are probably less important for the study of interannual variability (e.g., GlouDEMANS et al., 2009).

Pfister et al. (2005) estimated that 30 ± 5 Tg CO was emitted from fires in Alaska and Canada between June and August 2004, which was significantly higher than estimates based on GFED2 (16 Tg CO) but in line with our new results (30 Tg CO), mostly due to somewhat higher burned area and higher CO emission factors. TurquetY et al. (2007) found comparable emissions. Another regional assessments employing satellite-based CO focused on Indonesia and surrounding countries between 2000 and 2006 and indicated that GFED2 overestimated emissions from Sumatra but that emissions from Borneo matched reasonably well with the distribution of CO columns observed in the atmosphere (van der Werf et al., 2008). In this area, carbon emissions in GFED3 are somewhat lower than those in GFED2 for 2001–2006 but due to the inclusion of specific peat emission factors, the CO emissions are somewhat higher GFED3 for Borneo, and lower for Sumatra. In Borneo, however, GFED3 indicates that 2002 and 2006 emission peaks were of similar magnitude while several atmospheric indicators (CO, aerosols) indicate that 2006 was higher (Logan et al., 2008; Tosca et al., 2010).

Focusing on Africa, Chevallier et al. (2009) used MOPITT and the LMDZ-INCA chemistry transport model to estimate emissions with GFED2 CO emissions as a-priori input for 2000–2006. While Northern hemisphere African CO concentrations appeared in agreement with those seen from MOPITT, those from the Southern hemisphere were found to be too low, on average 26%. GFED3 CO emissions are 5% higher in southern hemisphere Africa than GFED2 and likely more in agreement with the results from Chevallier et al. (2009), but the 22% decrease in Northern hemisphere Africa will likely reduce the agreement between modeled and observed CO fluxes.

In short, we expect that overall the newly calculated emissions lead to better predictions of atmospheric trace gas and aerosol burdens. However, improvements in some regions will be accompanied by deterioration in other regions, similarly to the situation when we replaced version 1 with version 2 (Stavrakou et al., 2009).

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



4.2 Uncertainties

By combining new burned area with improved biogeochemical modeling to better model fuel loads we have attempted to reduce uncertainties in global fire emissions estimates. However, our estimates are still uncertain to about $\sim 20\%$ (1σ) on global, annual scales for carbon emissions, and higher for trace gas and aerosol emissions or when smaller regions or shorter time windows are considered. In addition, the uncertainties in the pre-MODIS era were up to 25% higher. Several factors contribute to these uncertainties, which are described further below.

4.2.1 Burned area

Global-scale moderate resolution burned area maps as used here have improved over the last few years and compare favorably against high-resolution case studies in temperate and boreal forest regions, as well as in savanna regions in Southern Africa (Giglio et al., 2009; Giglio et al., 2010). In wooded savannas, burned area may be somewhat underestimated (Roy et al., 2009) while burned or cleared area in deforestation regions has not undergone a formal validation. Other uncertainties arise from the native resolution of the burned area product, in our case 500 m. Here we have assumed that these grid cells burn completely when flagged as burned. In reality, however, this is often not the case. Early season burns, for example, are patchier than late season fires in Australia (e.g., Russell-Smith and Edwards, 2006). The patchiness thus clearly leads to an error in estimates. Part of this overestimation of area burned due to including unburned patches may be compensated for by the underestimation of area burned in grid cells where only a small fraction of the grid cell burned, leaving little signal to detect burned area. Errors may also be expected from fires in tropical forest areas where persistent cloud cover and small fires lead to obscure the burned area signal (Roy et al., 2008). By combining burned area with fire persistence, we have made the first attempt to distinguish deforestation fires from other types of fires, and at

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



the same time increase burned area estimates here to better capture reported cleared or deforested areas.

4.2.2 Fuel loads

Although uncertainty in burned area datasets is still substantial and the highest quality datasets still do not capture all burned area (Roy et al., 2009), some are of substantially higher quality than earlier attempts to characterize burned area around the world indicating that fuel loads, combustion completeness, and emission factors are now becoming the most uncertain components of emissions estimates at a global scale (Hoelzemann et al., 2004). This ordering has existed on regional scales for many years where higher resolution (e.g., Landsat) data has been available (French et al., 2004). Our model captured biomass density estimates in the Amazon (Saatchi et al., 2007), fuel build-up in savanna regions of Africa (Savadoغو et al., 2007) and Australia (Williams et al., 1998), and estimated fuel consumption in boreal America in agreement with field measurements and higher spatial resolution studies, see Sects. 2.2–2.4. However, discrepancies remained and variability in reported fuel consumption is large. Part of the reason may be related to scale; measurements are often made on small plots and less frequently extrapolated over the area within a fire perimeter. This scaling is crucial for making realistic comparisons with global-scale models, especially in heterogeneous landscapes.

Savanna and grassland fuel loads in GFED2 may have been too high (Roberts et al., 2008; Ellicott et al., 2009). By improving the scalar that determines the transfer of leaves to litter our modeled fuel loads are now lower than in previous work and are in better agreement with measurements from several field studies. Fuel loads are still higher than observed in some areas though; the mode fuel consumption in Africa was 300–400 g C m⁻² burned while most field studies indicate fuel loads of about 200 g C m⁻² burned or even lower. The model, however, does a reasonable job in simulating fuel build-up over the first few years after a fire. The model bias may have several causes. First, field campaigns may focus on frequently burning regions that have little

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



time to accumulate fuel. For example, it is often stated that most savannas burn every year or every other year. Our burned area data, however, indicate that the majority of the landscape exposed to fire in Africa had a fire return time between 2 and 5 years with only 14% burning annually. For Australia, fire return times were mostly between 5 and 10 years, and 2% of the total area burned annually. Second, if fires burn preferentially in a certain part of the grid cell then our approach will overestimate fuel loads. For example, if we find that on average 50% of a grid cell burns we assume the whole grid cell has a fire return time of 2 years and each year 50% of the whole grid cell is burned. Another possibility though is that 50% of the grid cell burns annually and the other 50% does not burn. In the latter case emissions are lower than in the former because fuel loads are smaller. If this is the case, our approach could overestimate emissions by ~30%. In reality the error is smaller because its magnitude decreases with increasing fire return time. We also expect to have reduced these errors moving from GFED2 to GFED3 by increasing the spatial detail by a factor 4. Finally, fuel consumption in Africa showed much spatial variability, with increasing values towards the equator. For Southern Africa, most measurements were made in the more arid and less productive areas (e.g., South Africa, Zambia) while satellite data indicates that areas further north (Angola, D.R. Congo) see more fire activity. Since parts of these areas are woodlands, they boost the average fuel consumption estimates.

4.2.3 Combustion completeness and emission factors

Both combustion completeness and emission factors vary to a large extent based on geographical and meteorological conditions as well as fuel composition (e.g., Kortontzi et al., 2003; Shea et al., 1996; Andreae and Merlet, 2001). While our combustion completeness values as well as depth of burning in peatlands were scaled based on soil moisture conditions, this should be seen as a simple approach to incorporate spatial and regional variability and more work is needed to better represent this variability. This may be feasible for emission factors because of a relatively large body of research providing ground measurements, but improving the representation of depth of

Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



burning into peat and organic soil may be difficult because of limited availability of observations and concurrent meteorology, although new approaches hold considerable promise (Ballhorn et al., 2009; Bobby et al., 2010). Validating combustion completeness is problematic as well; in savanna and grassland areas the error may be small because fine fuels are almost always observed to burn complete. In deforestation regions, however, the variability is large due to differences in land management and post-fire land use that will be difficult to model and validate with the currently available information.

4.3 Future developments

Several of the uncertainties in our estimates are related to landscape heterogeneity within our 0.5° modeling grid. Although we have made a distinction between deforestation fires, peat fires, fires occurring in the herbaceous parts of the grid cell, and fires occurring in the wooded parts of the grid cells, our approach is a first attempt and other key parameters (e.g., $fAPAR$) is still taken as the average of the 0.5° grid cell ignoring variability. In addition, we cannot distinguish the fire history within a grid cell; if for example a grid cell burns for 50% on average each year it makes a substantial difference whether 50% of the area burns every year or whether the whole grid cell burns every two years.

Modeling at higher spatial resolution may partly overcome some of these issues. Ideally, results are based on the native resolution of important datasets (burned area, $fAPAR$) that are available at 500 m resolution, so this may be feasible from a technical point of view. An added benefit is that this may allow for better validation with local measurements. Geo-location issues, especially when combining input datasets from different platforms (Hyer et al., 2009), and realistically extrapolating other input datasets (e.g., climate data) to the finer resolution are issues that have to be dealt with.

Besides improving spatial resolution and the quality of input datasets, there is a clear need to better understand the partitioning of combusted biomass into trace gases. While a large number of emission factor measurements used for this partitioning have revealed the broad-scale differences in emission factors between biomes (Andreae

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



and Merlet, 2001; Janhall et al., 2009), a clear description of spatial and temporal variability in emission factors is not yet included in large-scale modeling efforts, although variability may be large in some areas (e.g., Ward et al., 1996, Korontzi et al., 2003). Multi-species satellite measurements or continuous fire plume measurements in important biomass burning regions are needed to better understand the rules governing the partitioning.

5 Conclusions

We have updated our Global Fire Emissions Database (GFED) using improved satellite input data and made several modifications to our modeling framework. These modifications include introducing deforestation and degradation in the model, enabling the partitioning of fire carbon emissions into different sources, and adding an uncertainty analysis. The main findings can be summarized as follows:

- Global fire emissions averaged over 1997–2009 were 2.0 Pg C yr^{-1} according to our modeling framework with important contributions from Africa (52%), South America (15%), Equatorial Asia (10%), the boreal region (9%), and Australia (7%).
- By far the largest contributor (44%) to fire carbon emissions were fires in savannas and grasslands, with another 16% emitted from woodland fires. Forest fires, mostly confined to temperate and boreal regions accounted for 15%. We estimated that about 23% of fire carbon emissions stemmed from tropical deforestation, degradation, and fires burning in tropical peatland areas. These fire emissions of approximately 0.5 Pg C yr^{-1} were probably a net source of CO_2 to the atmosphere because they may not be compensated for by regrowth. This estimate does not include carbon fluxes from decomposition of remaining plant material following deforestation fires, and is confined to the tropics. Agricultural areas contributed about 3% of the global total carbon emissions from fire, but were likely conservative. These numbers represent averages during 2001–2009

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

- due to the need for MODIS observations; the role of peat burning increased when considering the 1997–2000 period due to high emissions in 1997 in Indonesia.
- Emissions were highest in 1998 (2.8 Pg C yr^{-1}) with almost all fire regions experiencing above normal fire activity, followed by 1997 (2.7 Pg C yr^{-1}) mostly due to emissions from Indonesia. The years 2000, 2001, 2008, and 2009 had lower amounts of fire activity with mean emissions of $\sim 1.6 \text{ Pg C yr}^{-1}$. From 2002 to 2007, fire emissions were relatively constant around 2.0 Pg C yr^{-1} . Within this period, positive anomalies in some regions cancelled negative anomalies, most notably in 2006 and 2007 when fire emissions in South America were low and high, respectively, while fire emissions in Equatorial Asia showed the reverse pattern. The decrease in the last 2 years was partly due to decreasing deforestation and degradation emissions.
 - While Africa was by far the largest source of fire CO_2 emissions, emissions of reduced trace gases like CO and CH_4 in tropical America and tropical Asia were almost as high due to a larger share of deforestation and peat fires emitting higher amounts of reduced trace gases per unit biomass combusted. Deforestation, degradation, and tropical peat fires caused almost half of global CH_4 emissions.
 - Emissions estimates are more reliable than our previous attempts to characterize global fire emissions, but uncertainties remain and are in the order of 20% (1σ) on global, annual scales. They are higher in the boreal region as well as in Equatorial Asia due to difficulties in estimating fuel consumption in organic soils. In addition, deforestation-fire carbon emissions are uncertain mostly due to uncertainties in cleared or burned area estimates. Since the burned area assessments used here may underestimate the amount of burned area in wooded savanna and agricultural areas, and our approach to model deforestation fires may also underestimate cleared areas, our estimates are likely conservative.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Future fire emissions estimates could take better advantage of native resolution burned area estimates, which would enable a better representation of heterogeneity in the landscape and enable more useful comparisons with ground measurements. Due to lack of spatially-explicit input datasets necessary to estimate fuel load, difficulties in estimating depth of burning in boreal and tropical peat areas, and unreliable burned area estimates in deforestation areas, it is unlikely that in the near future uncertainties will be reduced to levels of uncertainty associated with fossil fuel emissions.

Gridded $0.5^{\circ} \times 0.5^{\circ}$ monthly burned area, emissions estimates, the C4 fraction of CO_2 emissions, and biospheric fluxes (NPP and R_h) are provided to the scientific community for large-scale research through <http://www.falw.vu/~gwerf/GFED/> and will be updated frequently.

Supplementary material related to this article is available online at:
<http://www.atmos-chem-phys-discuss.net/10/16153/2010/acpd-10-16153-2010-supplement.pdf>.

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**Global fire emissions
(1997–2009)**

 G. R. van der Werf et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


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(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**Global fire emissions
(1997–2009)**G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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**Global fire emissions
(1997–2009)**

 G. R. van der Werf et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


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**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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Global fire emissions (1997–2009)

G. R. van der Werf et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


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**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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G. R. van der Werf et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


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**Global fire emissions
(1997–2009)**

 G. R. van der Werf et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[I◀](#)
[▶I](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


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- 30

Global fire emissions (1997–2009)

G. R. van der Werf et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[I◀](#)[▶I](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Table 1. Minimum and maximum combustion completeness (CC, unitless) or maximum burn depth for different fuel types (cm).

Fuel type	CC _{min}	CC _{max}	Max. burn depth (cm)
Leaves	0.8	1.0	–
Stems	0.2	0.4	–
Fine leaf litter	0.9	1.0	–
Coarse woody debris	0.4	0.6	–
Boreal organic soil	–	–	15
Tropical peat organic soils	–	–	50

Global fire emissions
(1997–2009)

G. R. van der Werf et al.

Table 2. Data sets used to drive the CASA-GFED modeling framework.

Variable	Role in CASA	Data product name	Source	Product resolution	Reference
Precipitation	Soil moisture, impacting NPP, F_{h1} , combustion completeness	Global Precipitation Climatology Project (GPCP) version 1.1	Multi-satellite and rain gauges	1° × 1°	Huffmann et al. (2001)
Temperature	Soil moisture, impacting NPP and F_{h1}	Climatology: IIASA IAV: GISTEMP	Station data Station data	0.5° × 0.5° 2° × 2°	Leemans and Cramer (2001) Hansen et al. (1999)
f APAR	NPP calculation	2000–2009: MOD15 1997–1999: GIMMSg anomalies with MODIS climatology	MODIS AVHRR	1 × 1 km 8 × 8 km	Myneni et al. (2002) Tucker et al. (2005)
Solar radiation	NPP	GISS, ISCCP-FD		280 km	Zhang et al. (2004)
Vegetation continuous fields (2001 onwards)	NPP allocation, mortality, partitioning of burned area	MOD44	MODIS	500 × 500 m	Hansen et al. (2003)
Land cover classification	Partitioning of burned area	MOD12 with UMD classification	MODIS	500 × 500 m	Friedl et al. (2002)
Burned area	Emissions calculation	GFED3 burned area	MODIS	0.5° × 0.5°	Giglio et al. (2010)
Burned area derived from fire hot spots	Emissions calculation when MODIS 500 m maps were unavailable	GFED3 burned area	MODIS, VIRS, ATSR	0.5° × 0.5°	Giglio et al. (2010)
Fire hot spots (2001–2009 climatology used pre-2001)	Deforestation rates	MOD14	MODIS	1 × 1 km	Giglio et al. (2003)

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Global fire emissions (1997–2009)

G. R. van der Werf et al.

Table 3. Improvements in biogeochemical modeling framework for GFED3.

Parameter	Description of modification	Impact on model estimates
Burned area	We now primarily use 500 m burned area maps from MODIS during 2001–2009 instead of regional relationships between fire hot spots and burned area	Estimates of burned area significantly improved in North America (Giglio et al., 2010); fire emissions estimates are no longer impacted by regional variations in fire hot spot to burned area relationships.
Spatial resolution	Increased from 1° to 0.5°	Factor of 4 increase in spatial resolution; smaller errors due to heterogeneity in landscape
Leaf senescence	Reduced carry-over of leaves during the dry season to the following wet season in herbaceous vegetation	Decreased biomass in herbaceous fuels, more in line with measurements (e.g., Savadogo et al., 2007; Williams et al., 1998)
NPP allocation	Changed from a fixed to a dynamic allocation based on mean annual precipitation (Hui and Jackson, 2006)	Better representation of spatial variability in aboveground biomass in highly productive ecosystems as compared with Saatchi et al. (2007)
Sub-grid cell information on burned area distribution over land cover and fractional tree cover bins	Changed from uniform distribution of burned area to herbaceous and woody fuel classes to dynamic distribution based on sub-grid cell information	Improved representation of spatial and temporal variability in fuel type burning and mortality rates; better ability to apply emission factors
Deforestation rates	Previous calculation based solely on burned area changed to combine burned area in wooded ecosystems and fire hot spot persistence (Morton et al., 2008; Roy et al., 2008)	Ability to separately estimate deforestation emissions; fuel loads in deforestation regions are no longer impacted by other fire activity in the grid cell (e.g., agricultural maintenance fires)
Deforestation emissions	Newly introduced; deforestation emissions based on clearing rates in the wooded fraction of the grid cell	New insights into deforestation fire activity; ability to track deforested land through time to calculate emissions from respiration (forthcoming work)
Uncertainty	Assessment of uncertainties	Monte Carlo approach provides insight into the spatial and temporal variability in global fire emissions.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


Global fire emissions
(1997–2009)

G. R. van der Werf et al.

Table 4. 1997–2009 area-averaged fire return time, NPP, fuel consumption, and combustion completeness for different regions.

Region ¹	Area (Mkm ²)	Fire return time (Yr)	NPP g C m ⁻² yr ⁻¹	Fuel consumption (g C m ⁻²)				Combustion completeness (–)		
				Standing	Surface	Soil	Total	Standing (burned)	Standing (all)	Surface
BONA	11.25	550	235	270	488	1904	2662	0.38	0.23	0.69
TENA	7.94	540	388	219	407	0	627	0.40	0.17	0.75
CEAM	2.71	197	674	682	803	4	1489	0.45	0.22	0.79
NHSA	3.02	139	1001	455	551	2	1007	0.56	0.20	0.81
SHSA	14.79	72	796	668	634	9	1311	0.55	0.29	0.82
EURO	5.41	827	400	202	462	3	667	0.43	0.21	0.80
MIDE	12.03	1363	35	30	169	0	198	0.59	0.16	0.93
NHAF	14.73	12	366	108	269	0	377	0.60	0.09	0.86
SHAF	9.92	8	627	108	340	0	448	0.55	0.07	0.83
BOAS	15.28	236	257	157	398	1424	1979	0.34	0.21	0.70
TEAS	18.25	130	205	50	142	61	253	0.40	0.30	0.84
CEAS	6.63	94	545	800	650	8	1459	0.47	0.29	0.80
EQAS	2.61	130	1213	2937	1181	5382	9500	0.49	0.47	0.77
AUST	7.98	15	238	53	206	0	259	0.69	0.09	0.88

¹ See Fig. 7 for the list of region abbreviations.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Global fire emissions (1997–2009)

G. R. van der Werf et al.

Table 5. Emission factors used for different fire types, in g specie per kg dry matter burned.

	Deforestation ¹	Savanna and Grassland ¹	Woodland ²	Extratropical forest ¹	Agricultural waste burning ¹	Peat fires ³
Carbon ⁴	489	476	483	476	440	563
CO ₂	1626	1646	1636	1572	1452	1703
CO	101	61	81	106	94	210
CH ₄	6.6	2.2	4.4	4.8	8.8	20.8
NMHC	7.00	3.41	5.21	5.69	11.19	7.00
H ₂	3.50	0.98	2.24	1.78	2.70	3.50
NO _x	2.26	2.12	2.19	3.41	2.29	2.26
N ₂ O	0.20	0.21	0.21	0.26	0.10	0.20
PM _{2.5}	9.05	4.94	7.00	12.84	8.25	9.05
TPM	11.8	8.5	10.2	17.6	12.4	11.8
TC	6.00	3.71	4.86	8.28	6.19	6.00
OC	4.30	3.21	3.76	9.14	3.71	4.30
BC	0.57	0.46	0.52	0.56	0.48	0.57
SO ₂	0.71	0.37	0.54	1.00	0.40	0.71

¹ Based on Andreae and Merlet (2001) and Andreae (personal communication, 2009).

² Based on the average of the savanna and grassland, and deforestation emission factor. The same emission factor was applied to deforestation and degradation emissions outside the humid tropical forest biome.

³ Based on Christian et al. (2003) for CO₂, CO, and CH₄; other species based on deforestation fires.

⁴ Dry matter carbon content based on carbon content in CO₂, CO, and CH₄ emission factors.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Global fire emissions (1997–2009)

G. R. van der Werf et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Table 6. Reported and best-guess uncertainties (1σ) for various parameters influencing fire emission estimates. We used a Monte Carlo simulation with 2000 runs to analyze the impact of uncertainties on estimated fire carbon emissions. Distributions for individual variables were truncated to avoid physically unrealistic scenarios, such as negative depth of burning values.

Parameter	Uncertainty
Burned area	Reported standard deviation (Giglio et al., 2010)
Deforested area	Reported burned area standard deviation $\times 2$
Woody biomass	22% ¹
Herbaceous biomass	44% ²
Tree mortality	25%
Depth of soil burning	50% of range
Combustion completeness	50% of range

¹ Based on a comparison of Amazon biomass with data from Saatchi et al. (2007)

² Double the uncertainty of woody biomass due to more factors impacting herbaceous biomass that may not be accurately represented at high resolutions, such as time since last fire, grazing, etc.

Global fire emissions
(1997–2009)

G. R. van der Werf et al.

Table 7. Annual emissions estimates (Tg C yr⁻¹) over 1997–2009 for different regions.¹

Region ²	Year													Mean	Contribution (%)
	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009		
BONA	19	116	36	14	5	69	60	139	66	50	40	49	44	54	2.7
TENA	2	8	11	12	6	10	9	4	6	11	20	13	8	9	0.5
CEAM	14	60	14	27	9	13	28	8	27	20	14	14	19	20	1.0
NHSA	20	51	14	19	17	9	54	26	13	11	25	13	13	22	1.1
SHSA	201	412	298	137	143	231	214	327	459	241	572	194	91	271	13.4
EURO	4	6	3	9	5	2	5	3	5	4	7	2	2	4	0.2
MIDE	1	2	2	1	2	2	2	2	1	2	3	1	2	2	0.1
NHAF	581	586	511	532	428	479	506	407	532	442	441	445	362	481	23.9
SHAF	514	682	534	514	514	483	597	579	621	548	533	578	544	557	27.7
BOAS	42	338	85	141	103	191	333	16	48	96	46	165	66	128	6.4
TEAS	57	31	18	37	33	49	43	25	27	35	35	40	31	36	1.8
CEAS	65	187	160	56	40	91	69	166	87	83	165	64	106	103	5.1
EQAS	1069	184	33	21	70	285	71	109	123	368	21	25	101	191	9.5
AUST	118	112	182	146	186	153	128	155	89	147	122	78	136	135	6.7
Global	2705	2775	1901	1665	1561	2066	2118	1966	2105	2059	2043	1681	1524	2013	100.0

¹ Annual estimates for other trace gases, as well as the contribution of different fire types, can be found on <http://www.falw.vu/~gwerf/GFED/>

² See Fig. 7 for the list of region abbreviations.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



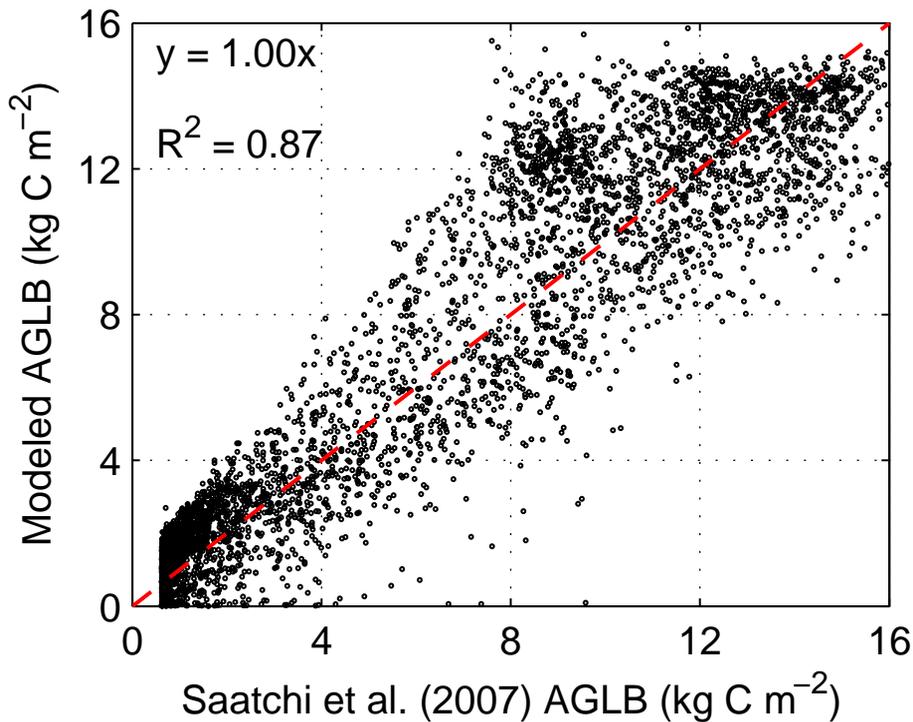


Fig. 1. Comparison of CASA predicted aboveground live biomass (AGLB; leaves and stems) with estimates from Saatchi et al. (2007) based on plot measurements and remote sensing metrics. Red dashed line indicates 1:1 slope, regression forced through origin.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

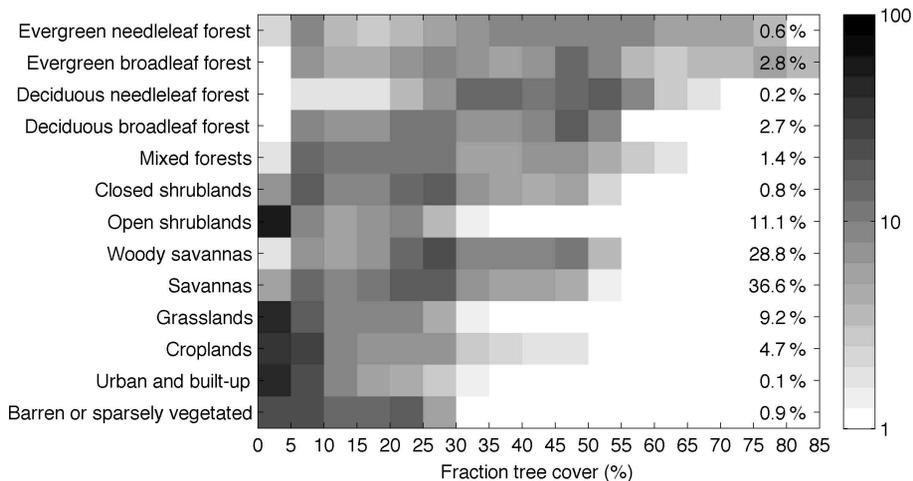


Fig. 2. Percent of burned area in each land cover class (MOD12Q1, UMD land cover classification) and fractional tree cover bin (MOD44, 5% bins). Numbers on the right denote the average contribution of each land cover type to global burned area over 2001–2009.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

◀ ▶

◀ ▶

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



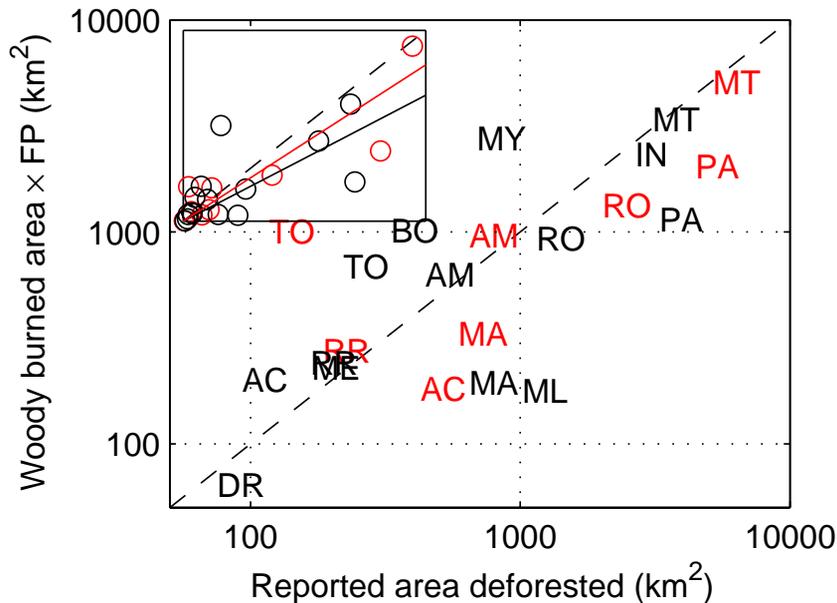


Fig. 3. Comparison of modeled deforested area with estimates from PRODES for states in the Brazilian Amazon (red, total for 2001–2006) and from Hansen et al. (2008) for tropical forest countries (black, total for 2000–2005). The dashed black line depicts the 1:1 slope. Modeled deforestation rates were based on a metric combining burned area in woody vegetation types with fire persistence (FP). Note the log scale; inset in top left shows the same data on a linear scale. The red solid line indicates the linear fit with PRODES data (slope of 0.75; $R^2=0.78$, $n=9$), and black solid line shows the linear fit with Hansen et al. (2008) data (slope of 0.66; $R^2=0.49$; $n=15$), both forced through origin. Modeled fire-driven deforestation rates were 82% of the total rates from the independent estimates. Abbreviations are DR (Democratic Republic of Congo), ME (Mexico), BO (Bolivia), MY (Myanmar), ML (Malaysia), IN (Indonesia), and the Brazilian states of AC (Acre), RR (Roraima), AM (Amazonas), RO (R ndonia), PA (Par ), and MT (Mato Grosso).

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



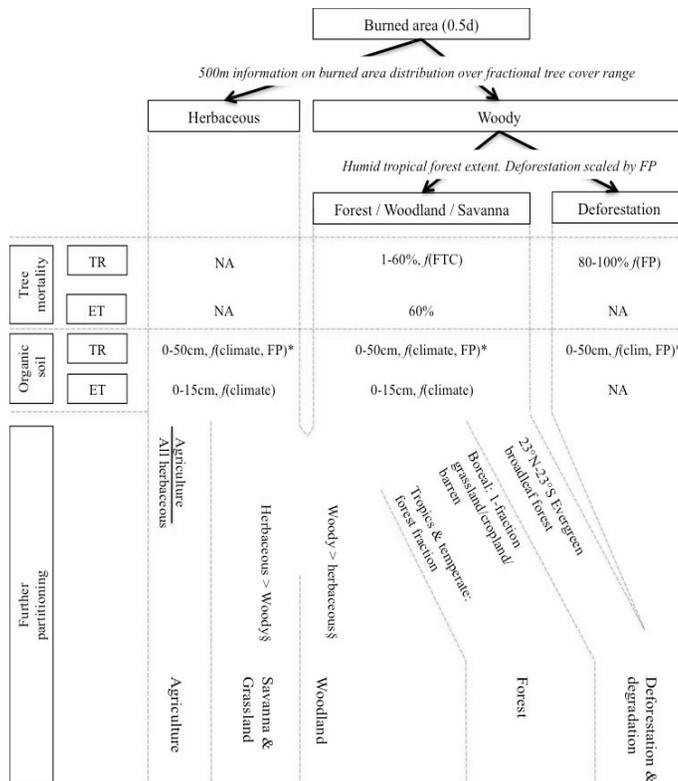


Fig. 4. Flowchart of our approach to partition burned area and emissions into different fire types. FP is fire persistence, TR tropics, and ET extratropics. The tropical “peat” class is not shown, but is based on the fraction of burned area detected in tropical peatlands and the TR organic soil burning scheme (*). Note that fires in savannas and grasslands, woodlands, and forests burn both herbaceous and woody fuels, while agricultural fires burn only herbaceous fuels, and deforestation fires burn only woody fuels. Agricultural and forest fraction of emissions were subtracted before woodland and savanna and grassland burning were calculated.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

◀ ▶

◀ ▶

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Global fire emissions
(1997–2009)

G. R. van der Werf et al.

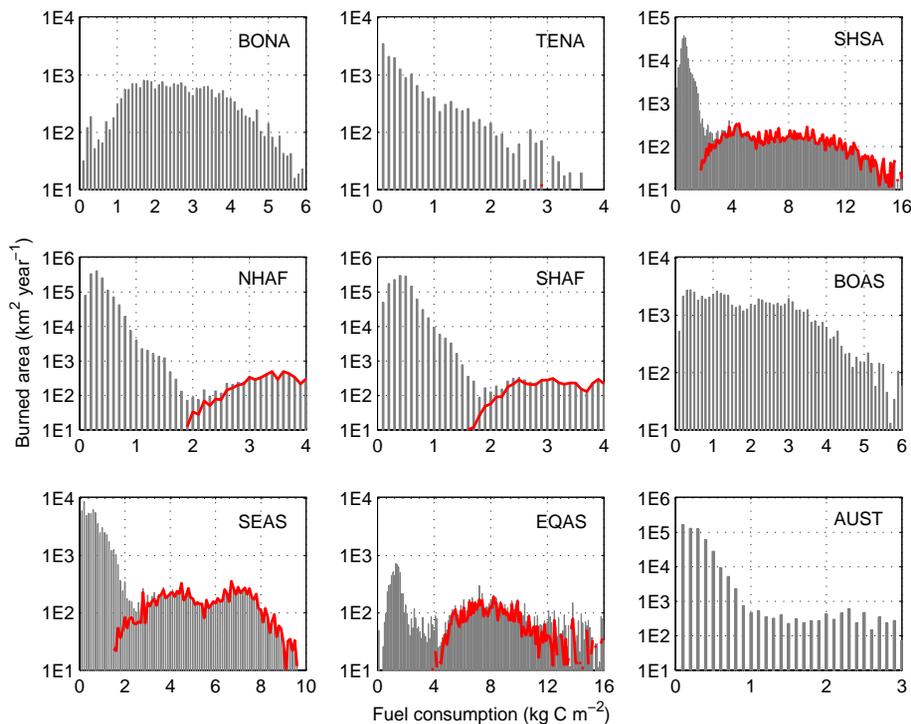


Fig. 5. Frequency distribution of fuel consumption in different regions (see Fig. 7 for the list of region abbreviations), with deforestation fires marked in red. Note the logarithmic y-axis scale and the different x-axis scales for each plot. Each bar represents 0.1 kg C m^{-2} of burned area, averaged over 1997–2009, centered upon its mean.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

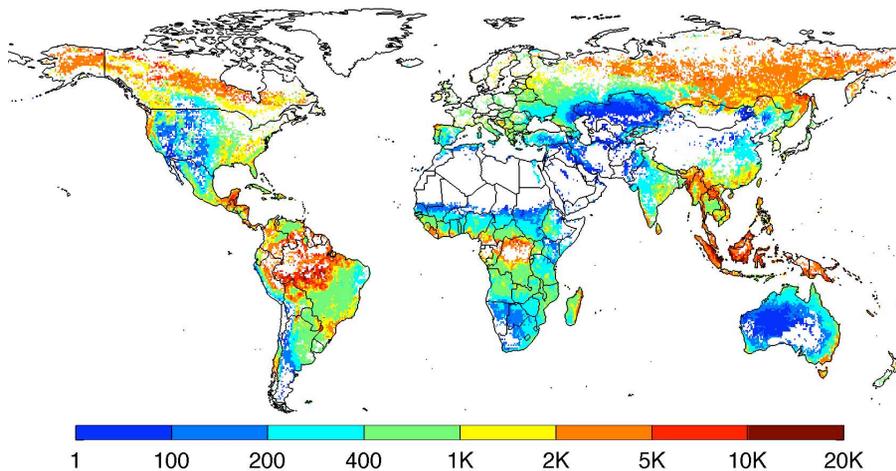
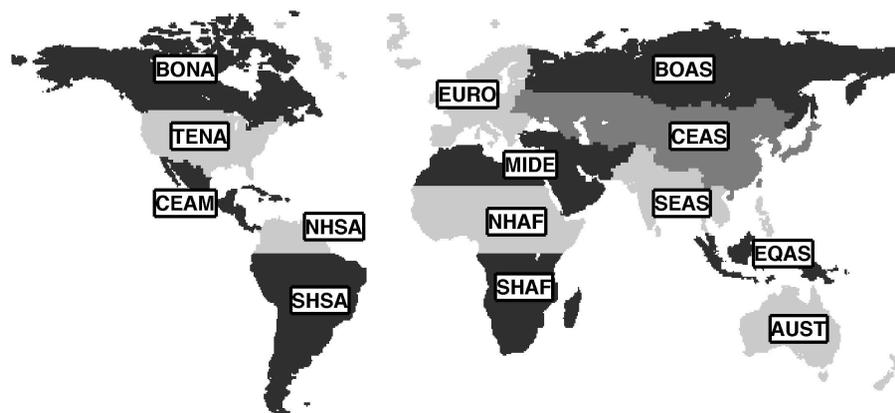


Fig. 6. Fuel consumption (g C m^{-2} of area burned), averaged over 1997–2009.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Global fire emissions
(1997–2009)

G. R. van der Werf et al.



BONA	Boreal North America	NHAF	Northern Hemisphere Africa
TENA	Temperate North America	SHAF	Southern Hemisphere Africa
CEAM	Central America	BOAS	Boreal Asia
NHSA	Northern Hemisphere South America	CEAS	Central Asia
SHSA	Southern Hemisphere South America	SEAS	Southeast Asia
EURO	Europe	EQAS	Equatorial Asia
MIDE	Middle East	AUST	Australia and New Zealand

Fig. 7. Map of the 14 regions used in this study, after Giglio et al. (2006) and van der Werf et al. (2006).

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[I◀](#)
[▶I](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


Global fire emissions (1997–2009)

G. R. van der Werf et al.

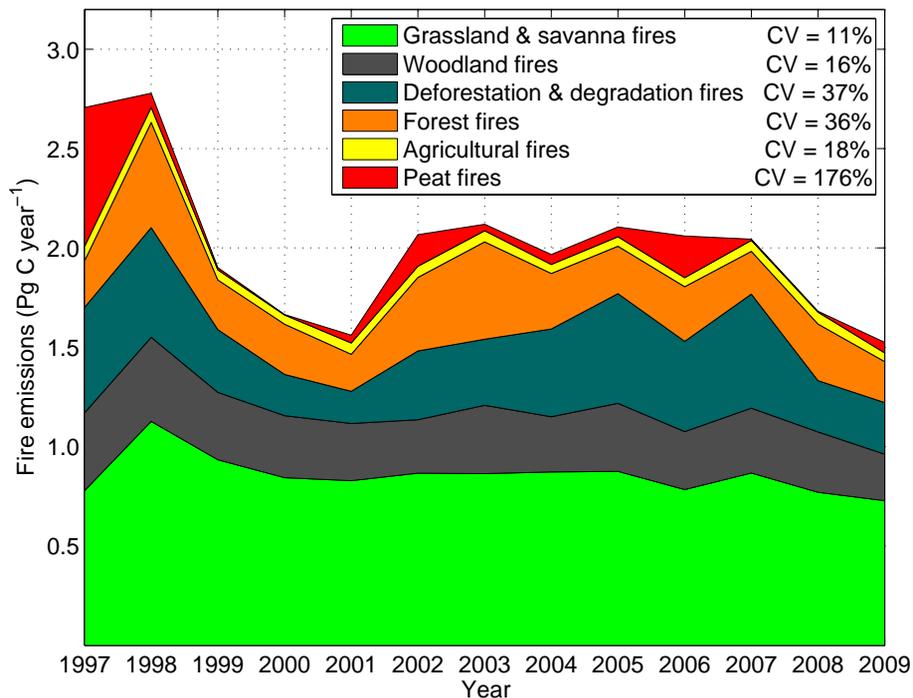


Fig. 8. Cumulative annual carbon emissions from different fire types and their coefficient of variation (CV) during 1997–2009.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

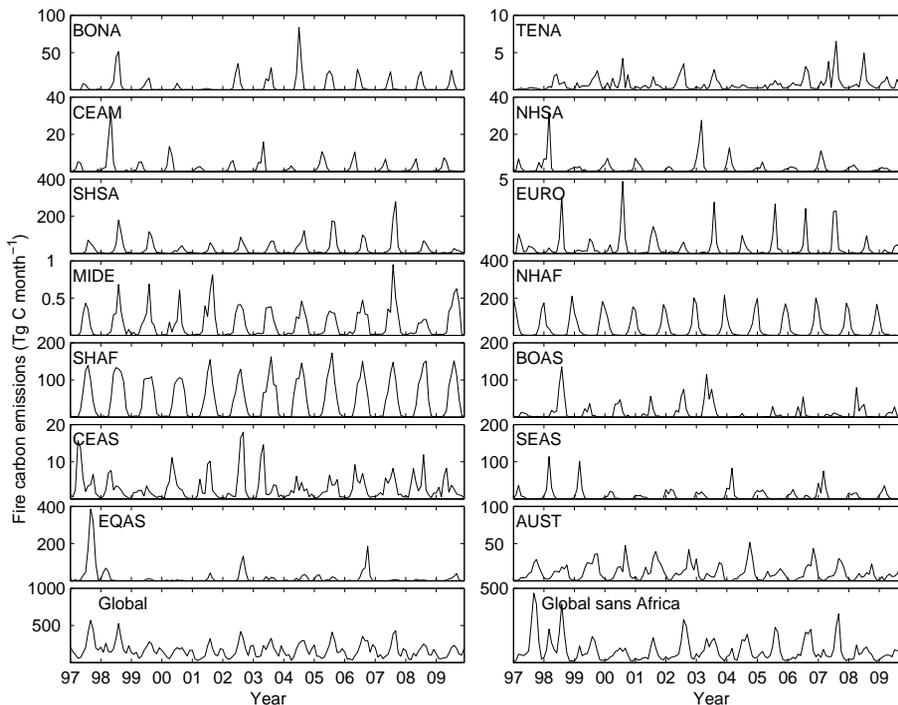


Fig. 9. Monthly fire emissions estimates (Tg C month^{-1}) over 1997–2009 for different regions (Fig. 7), as well as the global total with and without African emissions (bottom). Note the different y-axis scales for each plot.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



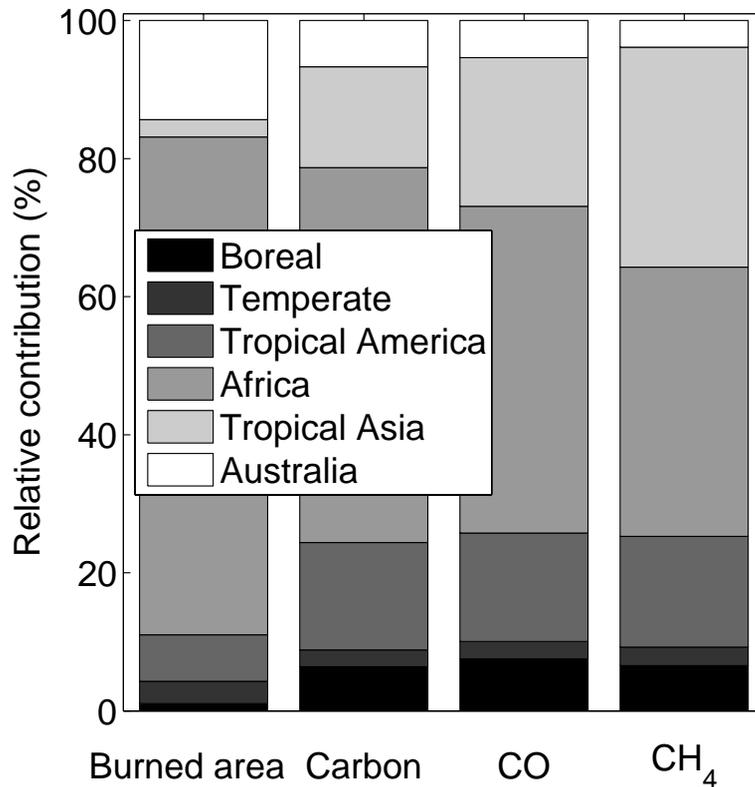


Fig. 10. Relative contribution (%) from different regions to 1997–2009 average global total burned area and fire emissions of carbon, CO, and CH₄. The different regions were composed of BONA and BOAS (Boreal), TENA, EURO, CEAS (Temperate), CEAM, NHSA, and SHSA (Tropical America), MIDE, NHAF, and SHAF (Africa), SEAS and EQAS (Tropical Asia), and AUST for Australia.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

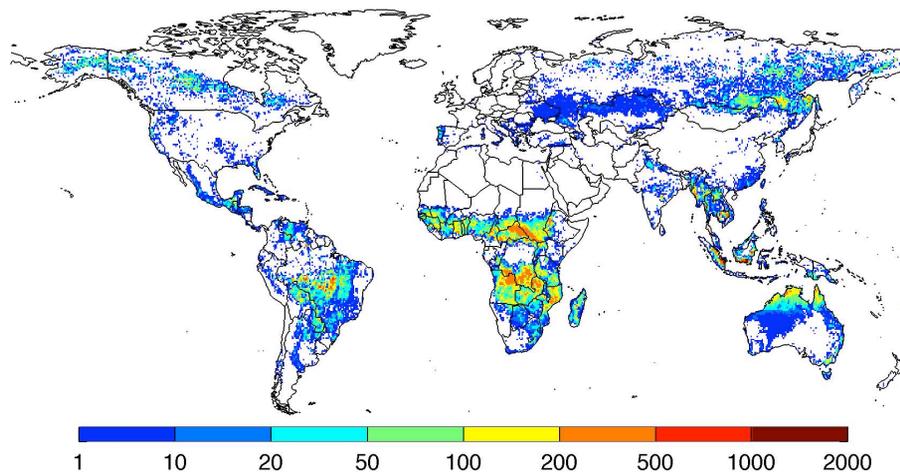


Fig. 11. Mean annual fire carbon emissions ($\text{g C m}^{-2} \text{yr}^{-1}$), averaged over 1997–2009. This quantity is the product of the fuel consumption (e.g., Fig. 6) and the burned area within the grid cell, divided by the total area of the grid cell.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[I◀](#)[▶I](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Global fire emissions
(1997–2009)

G. R. van der Werf et al.

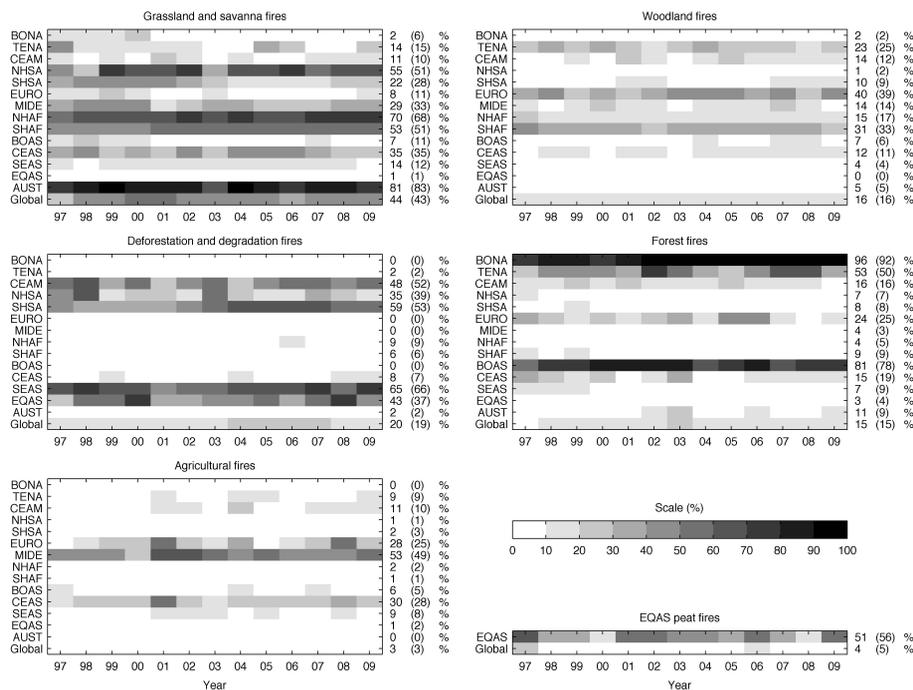


Fig. 12. Percent contribution of different fire types to annual fire emissions in each region. Absolute values are given in Table S1. Average percentages are given for the 2001–2009 MODIS period as well as for the full 1997–2009 time period (in parenthesis).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Global fire emissions
(1997–2009)**

G. R. van der Werf et al.

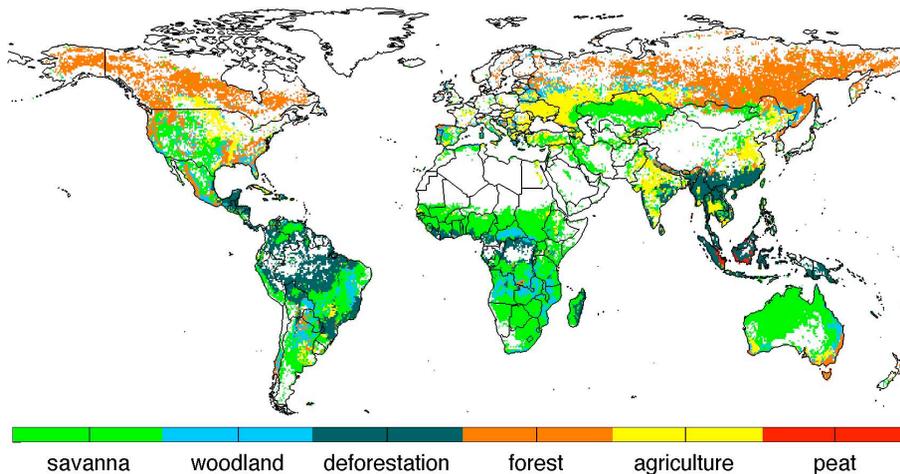


Fig. 13. Dominant fire type in each 0.5° grid cell based on carbon emissions. Savanna fires include grassland fires; deforestation includes degradation. Woodland and savanna fires were separated based on the relative contributions from woody or herbaceous fuels to total emissions, respectively.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

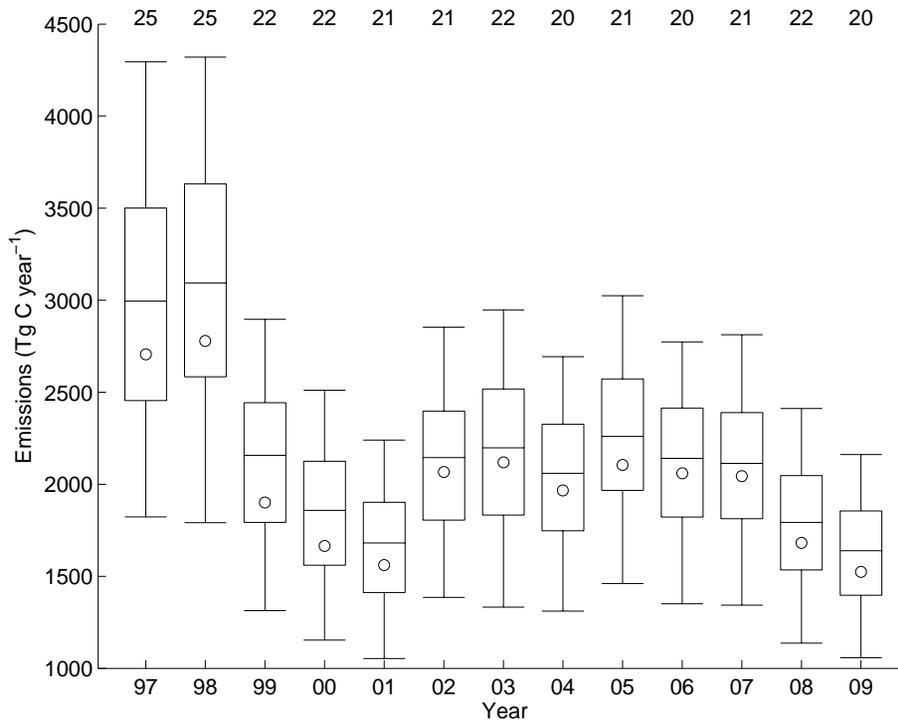


Fig. 14. Annual uncertainties expressed as the 5th, 25th, 50th, 75th, and 95th percentiles of 2000 runs in a Monte Carlo set-up. Circles denote the estimates reported throughout the paper, which do not necessarily align with the 50th percentiles due to truncation of several parameters in the Monte Carlo simulations. Numbers on the top of the plot indicate 1σ uncertainties in annual fire emissions in percent of the median estimate from Monte Carlo simulations, assuming a Gaussian distribution. Note that uncertainties are larger on regional and monthly scales (Fig. S6).

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



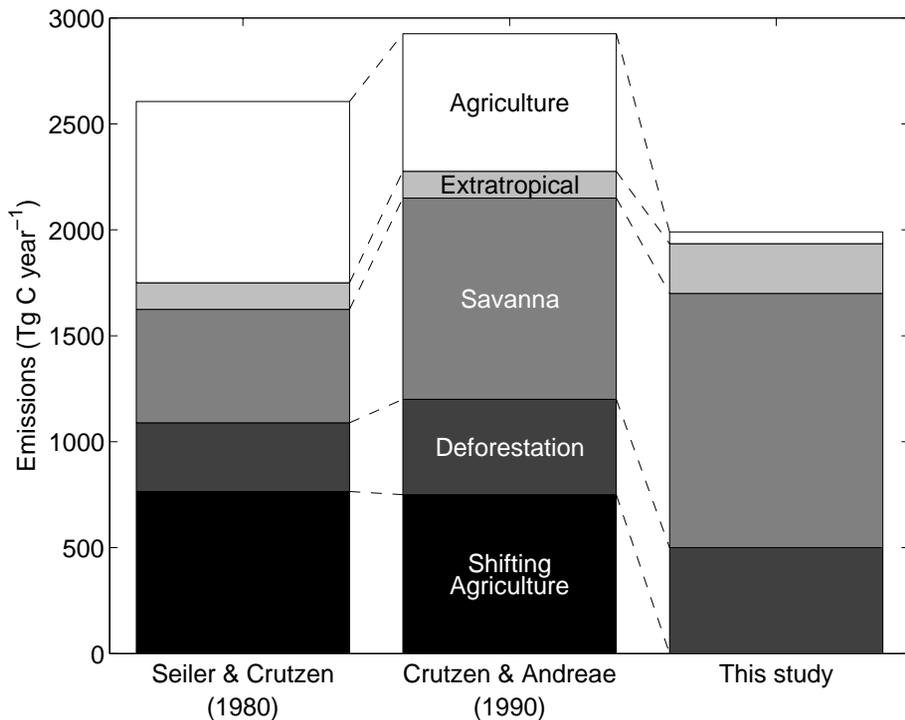


Fig. 15. Partitioning of total fire carbon emissions according to Seiler and Crutzen (1980), Crutzen and Andreae (1990), and this study. We used the extratropical assessment of Seiler and Crutzen (1980) to complete the tropical assessment of Crutzen and Andreae (1990). Note that we did not separate emissions from shifting agriculture from other types of deforestation in this study; our “deforestation” class therefore includes all fire types in tropical forest regions.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Global fire emissions (1997–2009)

G. R. van der Werf et al.

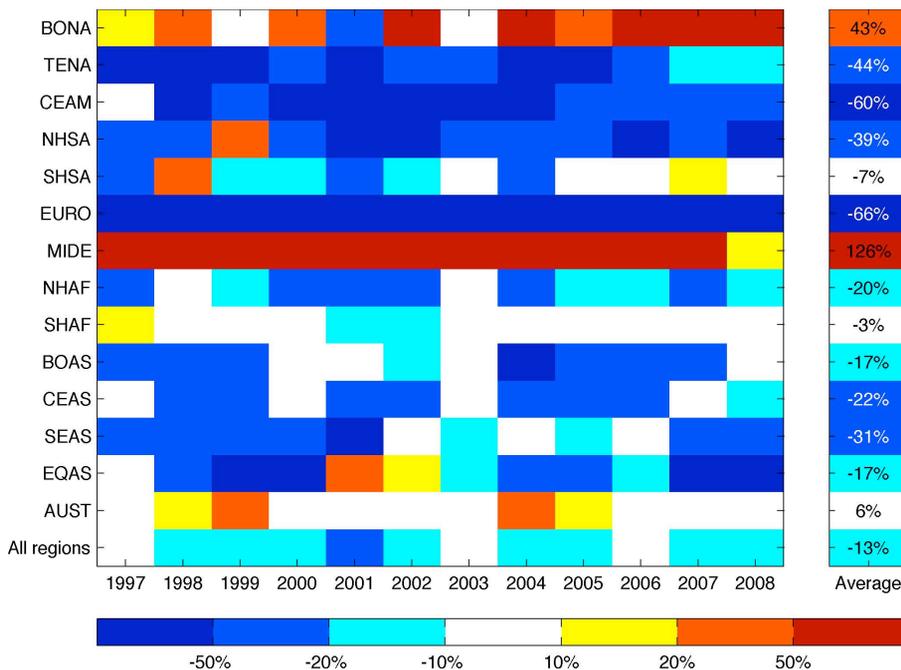


Fig. 16. Differences in fire carbon emissions estimates between GFED3 and GFED2, as a percent of GFED2 estimates. Positive numbers indicate GFED3 is higher than GFED2 and vice versa.

Title Page

Abstract	Introduction
Conclusions	References
Tables	Figures

◀
▶

◀
▶

Back
 Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

